Preliminary Report

On

Magnitude and Frequency of Floods From Small Watersheds In Semi-Arid Areas

To

U.S. Bureau of Public Roads

By Richard A. Schleusener, George L. Smith, and Nobu Yotskura

From

Civil Engineering Section Colorado State University Fort Collins, Colorado

CER59RAS39



Department of Atmospheric Science

PRELIMINARY REPORT

ON

MAGNITUDE AND FREQUENCY OF FLOODS FROM SMALL WATERSHEDS IN SEMI-ARID AREAS

TO

U. S. BUREAU OF PUBLIC ROADS

BY

RICHARD A. SCHLEUSENER, GEORGE L. SMITH, AND NOBU YOTSUKURA

îrom

CIVIL ENGINEERING SECTION COLORADO STATE UNIVERSITY FORT COLLINS, COLORADO

CER59RAS39

TABLE OF CONTENTS

Chapter		Page
	ABSTRACT	vi
ı.	INTRODUCTION	1
	Objectives	1
	History	1
II.	ESTIMATES OF RUNOFF FROM PHYSIOGRAPHIC	
	PARAMETERS	1
	Preliminary Studies	1
	Collection of Data	1
	Analysis	3
	Relations Between Short-Term and	
	Long-Term Floods	3
	Parameters Used in Multiple Correlation	4
	Refined Estimates	6
	Collection of Basic Data	6
	Relations Between Floods of Short-Term	
•	and Long-Term Frequency	6
	Graphical Correlation	10
ш.	BACKGROUND STUDIES	13
	Outside Contacts	13
	Conferences	14
	Precipitation Studies	14
	Runoff Study	17
IV.	SUMMARY OF RESULTS	18
	Presentation of Results	18
	Area of Application	18
	Estimates of Q_{10} , and Q_{40}	
	from Physiographic Parameters	18
	Use of Fig. 14 for Estimates of Peak Rates	
	of Runoff	18
	Geographical Limitations	18
	Limitation on Size of Contributing Watershed .	18
	Example of Use	18

TABLE OF CONTENTS - Continued

Chapter		Page
_	Limitations and Precautions	19
	Limitations in Basic Data	19
	Errors of Estimate	19
	Recommended Maximum and Minimum Values	
	of Q ₁₀ as a Function of Area	19
	Summary of Results of Related Studies	20
	Characteristics of Precipitation Associated	
	with Annual Maximum Flood Events	20
	Estimates of Clock-Hourly Precipitation from	
	Precipitation Amounts of Longer	
	Duration	20
	Utilization of Weather Radar to Provide	
	Increased Area Coverage of Rain-	
	fall Events	20
	Correlation of Precipitation Factors with	
	Physiographic Factors	20
•	Seasonal Distribution of Annual Maximum	
	Flood Events	22
	LIOOC MAGNES	22
V	DISCUSSION	22
	Research Needs	22
	Needs for Additional Records From Small	
	Watersheds	22
	Effect of Diversions on Peak Rates of Runoff	23
	Increased Importance of Precipitation Data	
	for Small Watersheds	23
	Techniques for Representation of Frequency Distribution	23
		24
	Plans for Future Studies	25
	Acknowledgments	26
	References	20
vi.	APPENDICES	
	1. Identification of Stations Used in the Study.	
	2. Gumbel Plots from Individual Stations.	
	3. List of persons contacted During the Study.	
	4. Use of "Relative Wetness" Parameters for	
	Estimates of Clock-hourly Precipitation	
	In Eastern Colorado.	
	5. Correlation of CPS-9 Radar Echo Intentisy	
	with Clock-liquity Precipitation	

TABLE OF CONTENTS - Continued

- 6. Correlation of Radar Echo Intensity With Clock-hourly precipitation.
- 7. Effect of Elevation and Watershed Size on Seasonal Distribution of Annual Maximum Events.

LIST OF FIGURES

Figure

- Location of Study Area.
- 2. Preliminary Estimate of Q10 from Physiographic Parameters.
- 3. Distribution of Error Curves Showing Departure of Plotted Points for Recurrence Intervals Greater than 10 years from "Benson" and "Potter" type Curves on Gumbel Frequency Paper.
- 4. Relation Between Q_{10} and Q_{40} (Benson method) for Selected Stations Outside D-13 and D-20 problem areas.
- 5. Relation Between Q_{10L} and Q_{40U}, Q_{10U} and Q_{40U} (Potter method) for Selected Stations Outside D-13 and D-20 Problem Areas.
- 6. Distribution of Error Curves for the Relations Shown in Figures 4 and 5.
- 7. Comparison of Q₁₀ vs Q₂₅ Relation for Stations Inside and Outside of the D-13 and D-20 Problem Areas.
- 8. Comparison of Q₁₀ and Q_{10U} for Stations Inside and Outside
 The D-13 and D-20 Problem Areas.
- 9. Dimensionless Profiles of Main Stem of Watersheds Included in graphical Correlation.
- 10. Profiles of Main Stem of Watersheds included in Graphical Correlation.
- 11. Coaxial Graph for Estimate of Q10.
- 12. Frequency of Occurrence of Complete Areal Rainfall Coverage Associated With Annual Maximum Flood Events as a Function of Area for Nine Watersheds in Eastern Colorado.
- 13. Location Map.
- 14. Relations Among Area, Slope Factor, Q₁₀ and Q₄₀ for the D-13 and D-20 Problem Areas.
- 15. Distribution of Error Curves for Estimates of Q_{10} from Fig. 14 (dependent data)

LIST OF FIGURES (Continued)

Figures

- Recommended Maximum and Minimum Q₁₀ as a Function of Watershed Size.
- 17. Location Map for Relative Wetness Study.
- 18. Frequency Analysis of Precipitation Data at Greeley, Colo.
- 19. Relation Between R and d for Subarea IV (Seasonal 2-Year Values.)
- 20. Distribution of Error Curve for The Relation Shown in Fig. 19.
- 21. Relations Between R and d for Subareas I-IV Shown in Fig. 17. (Seasonal 2-Year Values.)
- 22. Accumulated Relative Frequency of Annual Maximum Flood Events For 20 Watersheds in Colorado Between 7800 and 11,000 feet Elevation.
- 23. Accululated Relative Frequency of Annual Maximum Flood Events For 18 Watersheds in Colorado Between 6091 and 7683 Feet Elevation.
- 24. Accumulated Relative Frequency of Annual Maximum Flood Events For 23 Watersheds in Colorado Between 2798 and 6080 Feet Elevation.

ABSTRACT

This is a report of an investigation into meteorological and hydrologic records which established that watershed area and a slope parameter could be utilized for prediction of the peak rates of runoff having a ten-year recurrence interval in a part of the High Plains in eastern Colorado, western Kansas and Nebraska, and southeastern Wyoming.

Tentative relations were developed for peak rates of runoff having a ten-year and forty-year recurrence interval. Peak rates of runoff can be predicted from ungaged watersheds within the study area by using these relations.

Because limited data were available for this study from watersheds having a contributing area less than 100 square miles, results from this study must be considered as tentative and subject to revision as more data become available.

Results from related studies are presented.

The investigation is being continued to refine the techniques developed and to extend the study to adjacent areas.

I. INTRODUCTION

OBJECTIVES

The objectives of the investigation were:

- 1. To develop techniques for predicting magnitude and frequency of floods in semi-arid regions on watersheds having a contributing area less than 1000 square miles and lying within a region having similar lithologic and physiographic characteristics.
- To evaluate the influence of certain physiographic parameters on peak rates of runoff.
- 3. To investigate the possibility of utilizing weather radar data to provide more adequate areal coverage of precipitation events for use in making estimates of runoff.

HISTORY

This investigation, initiated in July, 1958, is a research project sponsored by the United States Bureau of Public Roads.

II. ESTIMATES OF RUNOFF FROM PHYSIOGRAPHIC PARAMETERS

PRELIMINARY STUDIES

Collection of Data - Preliminary studies were confined to parts of eastern Colorado, western Nebraska and Kansas, southeastern Wyoming, and a part of southern South Dakota, as shown in Fig. 1. The "D-13" and "D-20" problem areas shown in Fig. 1 were established by the Soil Conservation Service on the basis of having similar physiographic features and similar problems in soil conservation. The D-13 area is called the "Northern Brown Plains." The D-20 area is called the "Plains of the Upper Arkansas and Purgatorie Rivers." Brief descriptions of the area in an unpublished manuscript of the Soil Conservation Service follow:

D-13 Northern Brown Plains

"The Northern Brown Plains occupy a total area of 48,938,000 acres located in northeastern Colorado, northwestern Kansas, southeastern Wyoming, western Nebraska, and a small area extends into south central South Dakota... It has a relief that is characterized by nearly level to gently rolling tableland areas that break off into steeply rolling valley slopes. In the eastern part of Colorado and southwestern Nebraska there are several relatively large areas of sandhills...

"The average annual precipitation is about 14 to 18 inches ... Rainfall is quite variable .. (with) the greater portion of the precipitation falling (at high rates) with high runoff and erosion rates..

"In the area as a whole 42 per cent of the land is in cultivation and 54 per cent is in range....

"Soils of the area are.. of four types and all of them can be found in each of the four states. They are: (1) deep medium textured soils on nearly level tableland areas; (2) medium depth, medium textured soils on upland ...; (3) shallow medium textured soils and gravel; and (4) sandy soils on aeolian sand deposits."

D-20 Plains of Upper Arkansas and Purgatorie Rivers

"This area is located in southeastern Colorado and covers an area of 6,795,000 acres. The relief is undulating to rolling, 4,000 to 5,000 feet elevation above sea level. Rainfall variable, 11 to 14 inches (annually). Shallow to moderately deep, medium to moderately heavy textured soils on range land.

"...Erosion - - slight sheet erosion on much of the area. Severe in local areas having poor cover...

"Seven per cent cultivated ... 90 per cent grassland classed as semi-arid grazing land ..., 3 per cent miscellaneous, no forest."

Records of maximum annual rates of runoff from stations within these areas were collected if the following criteria were satisfied:

- 1. The length of record was equal to or greater than 7 years.
- 2. No significant artificial flow control existed for high flows.
- The watershed contributing area was not more than 1500 square miles.
- 4. More than 50 per cent of the contributing area of each water-shed must lie within the boundary of the D-13 or D-20 problem areas.

Analysis - All available records of peak rates of runoff from stations within the D-13 and D-20 areas that met these criteria were plotted on Gumbel's Extreme-Value probability paper and analyzed on the basis of techniques developed by Potter $(1)^*$ and by Benson (2). Discharges having a recurrence interval of 10 years (Q_{10}) were determined by Benson's technique for further study for seventeen watersheds in the D-13 and D-20 problem areas. Values of Q_2 , Q_5 , and Q_{15} were determined for the same stations by the same method. An estimate of Q_{25} was made by extension of the curve drawn by Benson's technique.

Relations Between Short-Term and Long-Term Floods - Attempts were made to relate short-term discharge values (\mathbb{Q}_2 , \mathbb{Q}_5 , \mathbb{Q}_{10}) to longer-term discharge values (\mathbb{Q}_{15} , \mathbb{Q}_{25}) for stations with records of suitable length. Logarithmic plots were made of \mathbb{Q}_2 vs \mathbb{Q}_5 , \mathbb{Q}_5 vs \mathbb{Q}_{10} , \mathbb{Q}_5 vs \mathbb{Q}_{15} , \mathbb{Q}_5 vs \mathbb{Q}_{25} , and \mathbb{Q}_{10} vs \mathbb{Q}_{25} . Of these combinations, \mathbb{Q}_5 vs \mathbb{Q}_{25} and \mathbb{Q}_{10} vs \mathbb{Q}_{25} were considered to have the greatest potential usefulness. The departures from the fitted regressions of \mathbb{Q}_{25} on \mathbb{Q}_5 and \mathbb{Q}_{25} on \mathbb{Q}_{10} was such that more than 67 per cent of the sample had an error of less than ± 25 per cent, the criterion of

^{*} Numbers refer to appended references.

suitable accuracy followed in this study.

In order to make estimates of peak rates of discharge for recurrence intervals greater than 25 years, it was considered necessary to utilize records having longer records than those which were available in the D-13 and D-20 areas. The success in relating Q₁₀ to Q₂₅ as described above suggested that a sample of longer records from outside the study area could yield usable relations between floods of short and long-term frequencies that would be applicable to the study area. This approach is described in Chapter II.

Parameters Used in Multiple Correlation - Graphical multiple correlation techniques ⁽³⁾ were used in evaluating the relationship of Q₁₀ with the following parameters:

- The contributing area of the watershed, as listed in the
 U. S. Geological Survey Water Supply Papers.
- 2. A location factor, defined as the difference in degrees between the mean longitude and the mean latitude at the centroid of the watershed as determined by eye.
- 3. A drainage density factor, defined as the total length of channels in miles as indicated by the blue lines on 1:250,000 scale maps of the area prepared by the U. S. Geological Survey, divided by the contributing area in square miles as defined in item 1.
- 4. An orientation factor with respect to an east-northeast axis.
- 5. The mean elevation of the watershed, an average of the highest and lowest elevations.
- 6. The mean longitude in degrees at the centroid of the watershed as determined by eye.
- 7. The mean latitude in degrees at the centroid of the water-shed as determined by eye.

- 8. The ratio of width of the watershed divided by its length.

 The length of the watershed was the distance from the gaging station to the furthest point. The width was defined as the contributing area divided by this length.
- 9. A compactness ratio, defined as the circumference of the circle having the same area as the watershed, divided by the total perimeter of the watershed.
- 10. A precipitation parameter which was the 2-year, 1-day point rainfall in inches, at the station nearest the centroid of the watershed.
- 11. A precipitation parameter which was the 5-year, 1-day point rainfall in inches, at the station nearest the centroid of the watershed.
- 12. A precipitation parameter which was a 5-year, 1-day point rainfall in inches, expressed as an average of stations in and near the watershed.
- 13. A precipitation parameter which was a 5-year, 1-day point rainfall in inches, expressed as area rainfall with an appropriate reduction from point-rainfall.
- 14. The overall slope of the watershed in feet per mile, determined by dividing the elevation difference between gaging station and headwater (in feet) by the distance (in miles) between these two points.
- 15. A slope parameter for the upper and lower halves of the watershed, determined as for item 14.

Of these parameters, the first three listed gave errors of estimate within acceptable limits of accuracy, using the dependent sample. The graphical relationship that was derived is shown in Fig. 2. It was not possible to develop a graphical correlation that gave acceptable accuracy using combinations of the other parameters.

REFINED ESTIMATES

Collection of Basic Data - The general requirements considered in the collection of basic data for refined estimates of rates of runoff included the following criteria in addition to those described previously:

- Records of annual maximum stream flow had to be derived from recording gages only; those records derived from staff gage readings were discarded.
- 2. Records were not used if there were more than four years break in records.
- No record was utilized where there had been a change in location of site greater than two miles up or down stream.

Records of rates of runoff from stations outside of the D-13 and D-20 problem areas were compiled in order to relate short-term to long-term floods. For these stations, the additional requirement of a minimum length of record of 23 years was established and a watershed size of not more than 2000 square miles.

Records of rates of runoff were also collected from stations inside the D-13 and D-20 problem areas. For these stations, the contributing watershed areas were all less than 1500 square miles with most stations having a contributing area less than 1000 square miles. The minimum acceptable length of records for stations in these areas was limited to seven years.

Relations Between Floods of Short-Term and Long-Term Frequency A comparison was made between two techniques currently used in the
analysis of data plotted on Gumbel paper. Potter's method (1) approximates an array of points on Gumbel paper by two straight lines, giving
a "dog-leg". Benson's method (2) consists of drawing a curved line
that best fits the array of plotted points. A frequency plot of the annual
maximum runoff from each of the stations was made on Gumbel plotting

paper. Each of the stations used in the study is identified in Appendix No. 1. Gumbel plots from the individual stations are included in Appendix No. 2. Stations included in the study that were located outside of problem areas D-13 and D-20 fell into four general geographic locations: northwest, east, southeast and southwest of the problem areas D-13 and D-20, defined respectively by the following locations:

Northwest: 45 to 49 degrees north by 106 to 113 degrees west.

East: 37 to 43 degrees north by 94 to 100 degrees west.

Southeast: 29 to 35 degrees north by 94 to 101 degrees west.

Southwest: 34 to 38 degrees north by 102 to 107 degrees west.

Curves were drawn on the Gumbel plots by the two methods described previously. Using the Potter approximation, two straight lines were drawn to best fit the data, and using Benson's method a curved line was drawn that best fit the plotted points.

Approximately 58 separate records were available in the geographic locations mentioned previously outside the problem areas D-13 and D-20. A sample was selected, using the following criteria:

- 1. Equal numbers of stations were desired from each location.
- Equal numbers of stations were desired from watersheds less than 500 square miles from watersheds larger than 500 square miles.
- 3. Equal numbers of stations were desired from different lithologic areas having the following classifications:
 - a. Sandstone and shale.
 - b. Glacial drift and loess.
 - c. Unclassified.

Using these criteria, a total sample of 22 stations was selected from the 58 records available. Of these 22 stations, 19 were suitable for the Benson method of analysis; three stations being discarded because of extreme irregularities in the plotted curve on the Gumbel paper. The same 3 stations were discarded in utilizing the Potter method because the upper and lower frequency curves were nearly parallel. This gave a discontinuous curve utilizing the Potter method. In addition, two other stations were discarded for utilization by the Potter method because of an excess error in approximating the plotted points with the two straight lines by the "dog-leg" method. For these two stations, the accumulated percentage error in representing the data with the "dog-leg" was greater than $\frac{1}{2}$ 25 per cent for 2/3 of all plotted points having a recurrence interval of 10 years or more.

The next procedure was to compare the errors resulting from each of the two methods of curve fitting. Using the 17 stations that remained, the accumulated error curve was plotted for both methods. The distribution of error curve is shown in Fig. 3. It will be noted that both methods gave a good representation of the plotted points having a recurrence interval greater than 10 years. Approximately 95 per cent of the sample was within ± 17 per cent error for both methods.

An attempt was made to group the data from the regions outside of D-13 and D-20 problem areas by geographic areas and by geological parent material classifications. Variations in the relation between Q_{10} vs Q_{40} (Benson's method) and Q_{10U} vs Q_{40U} (Determined from the upper frequency curve by Potter's method) were considered to be sufficiently small to permit grouping together the data from northwest, southwest, and east of the problem area. Data from these locations were grouped together. A plot of Q_{10} vs Q_{40} (Benson method) is given in Fig. 4. Plots of Q_{10L} vs Q_{40U} and Q_{10U} vs Q_{40U} are given in Fig. 5. (Q_{10L} was determined from Potter's lower frequency curve.) Fig. 6 shows the distribution of error curves for both methods. Examination of Fig. 6 shows that a smaller error results from use of the Benson method, which gives 94 per cent of the sample having $\frac{1}{2}$ 25 per cent error.

The relations shown if Fig. 4 and Fig. 5 were derived from geographic locations outside of the D-13 and D-20 problem areas. The problem remained to compare this type of relation from outside the D-13 and D-20 areas with that inside the same area. Fig. 7 shows the relation between Q_{10} vs Q_{25} for points inside and outside the D-13 and D-20 problem areas. Since the points from inside the study area appear to be consistant with those northwest, east, and southwest of the study area, the assumption was made that the relation between Q_{10} and Q_{40} as shown in Fig. 4 also applied inside the D-13 and D-20 areas. This is in agreement with the tentative conclusion reached in the preliminary studies.

It should be noted that the apparent better fit for the "Benson" curve on Fig. 6 is not adequate justification for acceptance and use of the method described herein. Drawing a curved line on Gumbel paper departs significantly from the straight line that theoretically should represent extreme values. Acceptance of a curved line on Gumbel paper implies the existence of a limiting discharge for a curve that is concave downward or of a limiting recurrence interval for a curve that is concave upward. While a limitation on the maximum possible discharge may be possible on physical reasoning, a more common occurrence in the area studied was a curve that was concave upward.

The method of Potter in fitting two straight lines to the plotted points on Gumbel paper does not suffer these limitations, although for some records difficulty was experienced in obtaining a suitable fit for the data with two straight lines. Fig. 6 shows that this method approaches acceptable accuracy of "two-thirds of the sample having less than 25 per cent error."

A comparison of values of Q_{10} and Q_{10U} for stations inside and outside of the D-13 and D-20 problem areas is shown in Fig. 8. Based on this comparison it was concluded that differences between Q_{10} and Q_{10U} as used in this study were not significant.

Comparison of Figs. 4 and 5b show that for a given estimate of Q_{10} (or Q_{10U}), the difference between the resulting estimate of Q_{40} and Q_{40U} is less than 25 per cent for nearly all the range of values shown on Fig. 8.

For these reasons the estimate of Q_{10} and Q_{40} , obtained as described above, are considered to be consistent with estimates of Q_{10U} and Q_{40U} . In view of the length of records used to derive the relationship of Fig. 11, it was not considered feasible to extend the relationship shown in Fig. 14, beyond a frequency of 40 years. The relation shown in Fig. 4 was used to develop the relation between Q_{10} and Q_{40} as shown in Fig. 14.

Graphical Correlation - The preliminary studies described previously indicated that watershed contributing area, geographic location, and stream density were factors that could be used for making estimates of peak rates of runoff for ungaged watersheds. While it was believed possible to utilize these physiographic parameters in the final graphical correlation, several disadvantages in use of these parameters were evident. Computation of the drainage density factor was time consuming and laborious. The location parameter (longitude minus latitude), while probably directly related to the frequency of thunderstorm occurrence, is not a parameter that is directly involved in the runoff process. Consequently other suitable parameters were sought.

Work by Benson (5) indicated that channel slope showed considerable promise as a factor for explaining variations in peak rates of runoff from New England watersheds. Watersheds from inside the study area were examined and dimensional and dimensionless plots were made of the channel profiles. These profiles are shown in Figs. 9 and 10.

Each of three slope parameters were used in conjunction with contributing area in a graphical correlation process to derive a relation suitable for use in estimating Q_{10} . Following Benson's work (5) a slope parameter S_{01} was defined by

S.9L =
$$\frac{E_{0.9L} - E_{CS}}{D}$$

where E_{0.9L} = elevation in feet 9/10ths of the length of the watercourse upstream from the construction site.

E_{CS} = elevation in feet at the construction site.

D = distance in miles along the watercourse between these locations.

The second slope parameter utilized was the "T" factor, suggested in conversations with Mr. W. D. Potter. The "T" factor (indicating a measure of "time of travel,") is defined as follows:

$$T = T_1 + T_2 = \frac{0.3L}{\sqrt{S_1}} + \frac{0.7L}{\sqrt{S_2}}$$
where $\sqrt{S_1} = \sqrt{\frac{E_{HW} - E_{0.7}}{0.3L}}$
and $\sqrt{S_2} = \sqrt{\frac{E_{0.7L} - E_{CS}}{0.7L}}$

where the symbols have the following meanings:

E_{HW} = elevation (feet msl) at the headwaters of the watershed

elevation (feet msl) at a point 0.7 of the distance from the construction site to the headwaters, measured along the watercourse.

E_{CS} = elevation (feet msl) at the construction site.

E = distance (in miles) between construction site and headwaters.

A third slope parameter was defined by

$$S_{0.5L} = \underbrace{E_{0.5L} - E_{CS}}_{0.5L}$$

where E_{0.5L} and E_{CS} are the elevations in feet at the point 0.5 the length of the watercourse and at the construction site, respectively, and L has the same meaning as above.

These slope parameters were used in conjunction with contributing area "A" (in square miles) in a graphical correlation process to estimate Q_{10} . It was found that use of $A^{1/2}$ provided some improvement over $A^{1\cdot0}$ in some cases. The correlations were repeated, using $A^{0.75}$ and $A^{0.90}$. It was found that $A^{0.90}$ vs Q_{10} with the slope parameter $S_{0.9L}$ gave the best results. Fig. 11 shows the relation between these variables, and forms the basis of Fig. 14 which is presented in Chapter IV.

The parameters used in Fig. 11 and Fig. 14 are probably not the only ones suitable for use; in fact they may not be the best ones for the intended purpose. However, they are believed to be suitable for use because they satisfy the following criteria:

- 1. They are relatively simple to determine.
- 2. The accuracy of estimate of $\,Q_{10}\,$ is consistent with the accuracy of the basic data on contributing area and measured discharge that went into the study.

For these reasons, further refinements in Figs. 11 and 14 were not attempted. Additional records of peak rates of runoff having a high degree of accuracy are desirable for making further refinements in the relations shown in Figs. 11 and 14. It will be noted that five stations in Fig. 11 (Nos. 4, 20, 25, 33, and 34) have 12 years' record or less. For this reason their estimated Q_{10} in Fig. 11 may be subject to some revision. In view of this fact and because of the scarcity of data from drainage areas of less than 100 square miles, the correlation charts in Fig. 11 and Fig. 14 were modified from a consideration of unit discharges (cfs/mi⁻²) as a function of area. (Details are not presented in this report.)

A description of the use of Fig. 14 and some of the limitations and precautions to be observed, are given in Chapter IV, "SUMMARY OF RESULTS."

III. BACKGROUND STUDIES

OUTSIDE CONTACTS

A survey of all sources of information was first made with the following objectives:

- 1. Prevention of duplication in methodology and data analysis.
- 2. Obtaining only data essential to the proposed study.

Agencies contacted were (a) those making studies of runoff on small watersheds, (b) those making studies of severe storms in eastern Colorado, and (c) those responsible for flood and erosion control on small watersheds.

A list of persons and agencies contacted during the course of the study is given in Appendix 3.

CONFERENCES

A general description of the objectives of the project and some of the details of procedure being followed were presented at two separate conferences. The first presentation was given to meeting of the Committee on Surface Drainage of the Highway Research Board which met at Fort Collins on 15 September 1958.

The second presentation was given to a <u>Highway Drainage Conference</u> sponsored jointly by the U.S. Bureau of Public Roads and the Colorado Department of Highways on 3 March 1959 in Denver.

PRECIPITATION STUDIES

Because runoff in semi-arid regions is a direct function of such parameters as the amount of precipitation, and precipitation pattern on the watershed, the rainfall distribution associated with annual floods in eastern Colorado was studied. The objective of the study was to define the size of watershed that gives floods from "random" thunderstorm activity, as opposed to general precipitation over a wider area. Using climatological data, this study was confined to nine stations in eastern Colorado.

Annual maximum peak flows from contributing watersheds of not more than 1000 square miles were recorded for the period 1930-1950. For each flood event the amount of precipitation at raingage stations-recording or non-recording--on or near the basin was determined. The precipitation data were then given a weight, as follows: If 0.1 inch per day fell at a raingage station, a weight of 1 was given; if less than 0.1 inch were recorded, a weight of 0.5 was given; and zero rainfall was given a weight of zero.

The drainage basin was divided into sub-areas by the Thiessen method using the foregoing weighted values to compute the per cent

of basin area covered by precipitation for the given flood event. A weight of "one" was used when the entire sub-area received rainfall. The ratio of the number of the annual maximum floods associated with 100 per cent coverage of watershed to the total number of flood events was then expressed as a per cent. This value was then plotted against basin area, as shown in Fig. 12. Fig. 12 shows that for watersheds with contributing areas larger than about 900 square miles, two-thirds or more of the annual maximum flood events are associated with rains which cover the entire basin. For watersheds with contributing areas less than about 50 square miles, one-third or less of the annual maximum floods are caused by such rainfalls.

A factor related to flood runoff from small, intense storms is clock-hourly precipitation amounts. Because of the paucity of such data at non-recording rain gage stations in eastern Colorado, an investigation was made to determine if daily or monthly values of rainfall could be used to make estimates of clock-hourly precipitation values.

The investigation was confined to the analysis of precipitation amounts having a two-year recurrence interval. The data collected included daily, monthly, seasonal and annual precipitation amounts having a two-year recurrence interval. Through interrelationships termed "relative wetness," estimates were made of clock-hourly precipitation amounts for the two-year recurrence interval from precipitation amounts of longer duration. Preliminary results show that these estimates are acceptable as a substitute for recorded clock-hourly precipitation amounts. Details of this study are given in Appendix 4.

The use of weather radar echo data received from rainstorm patterns was also investigated as a possible means of obtaining more

adequate areal coverage of rainfall events. Two sources were considered (a) hand-drawn sketches of the PPI scope of the United Air Lines 5.5 cm radar set, and (b) reconstructed records of CPS-9 data (a 3.0 cm set) from Lowry Air Force Base. Both radar units were located in Denver, Colorado and have a range of approximately 200 miles.

To test the suitability of these data as a means of providing more adequate areal coverage of rainstorm events, attempts were made to correlate echo intensity with clock-hourly rainfall amounts concurrent with the time of echo occurrence.

No satisfactory correlation was obtained for either the United Air Lines data or the CPS-9 data.

Details of these studies are given in Appendix 5 and Appendix 6.

Attempts were made to correlate certain physiographic parameters with precipitation parameters, as had been done by Spreen (4) for western Colorado, where mean seasonal and annual precipitation was correlated to factors of elevation, exposure, and zone. Results indicated that a statistically significant correlation could be obtained between mean monthly rainfall (the month of May was used in the study) and simple parameters of location (latitude, longitude and elevation.) Detailed results of these correlation analysis for a number of stations in eastern Colorado are given in Chapter IV.

These precipitation studies were undertaken in an attempt to find a precipitation parameter suitable for reducing some of the observed variation in rate of runoff from watersheds of comparable size. As noted previously, several precipitation parameters (2 and 5-year) 24-hour precipitation amounts, both for point and areal rainfall) were tried unsuccessfully for this purpose. The reason for this failure is not clear. It is possible that the entire region under study may have precipation characteristics sufficiently homogeneous that explanations for variations in rate of runoff cannot be explained by precipitation parameters.

A further limitation in the precipitation parameter using clock-hourly data is the relative shortness of record for many stations. This may have caused difficulty, for example, in attempts to relate a precipitation parameter having a 2-year or a 5-year recurrence interval to a 10-year rate of runoff. The relative wetness study described previously indicates that there is a close relation between daily and clock-hourly precipitation amounts. Hence, use of daily rainfall amounts from stations having long records can help to overcome this difficulty.

RUNOFF STUDY

The Seasonal Distribution of Annual Maximum Flood Events - A study was made to determine the effect of elevation and contributing area on the seasonal distribution of annual maximum flood events for sixty-two stations in the North Platte, South Platte, Republican, Arkansas, and Colorado River watersheds covering all of Colorado, except the San Luis Valley.

The stations were first divided into three nearly equal groups according to elevation. These groups were then divided into three more groups according to watershed area, making a total of nine classes with varying numbers of cases in each class. Each class was then plotted using accumulated frequency of annual maximum flood events in percent vs month of occurrence of the maximum flood event.

Results of this study indicate that the average date of occurrence of 67 per cent of annual maximum floods advances with increase in water-shed size, and that for watersheds below 7683 feet elevation, the date of occurrence of 67 per cent of annual maximum floods advances with decreasing elevation.

Details of this study are given in Appendix 7.

IV. SUMMARY OF RESULTS

PRESENTATION OF RESULTS

Area of Application - A map showing the area of application of these results is shown in Fig. 13.

Estimates of Q_{10} and Q_{40} from Physiographic Parameters - Fig. 14 shows the relations among area (A), slope factor $(S_{0.9L})$, Q_{10} and Q_{40} .

USE OF FIG. 14 FOR ESTIMATES OF PEAK RATES OF RUNOFF

Geographical Limitations - Use of Fig. 14 should be limited to the
confines of problem areas D-13 and D-20, shown in Fig. 13.

<u>Limitation on Size of Contributing Watershed</u> - Fig 14 is applicable for watersheds having a contributing area less than 1000 square miles.

Example of Use - The following example illustrates the use of Fig. 14 for making estimates of magnitude and frequency of peak rates of runoff. Assume a watershed within the region shown in Fig. 13, having the following characteristics:

Contributing area: 400 square miles.

Length of watercourse (measured from appropriate map or aerial photograph): 45 miles.

Elevation at construction site: 5608 feet.

Elevation at 9/10ths of the distance from the construction site to the headwaters: 7320 feet.

From these values, the following parameters are determined:

A = 400 squarε miles.

$$S_{0.9L} = E_{0.9L} - E_{CS} = \frac{7320 - 5608}{40.5} = 42.3$$

Enter Fig. 14 with A = 400, $S_{0.9L} = 42.3$, and read

o₁₀ = 14,000 cfs

କ୍₄₀ = 20,500 cfs

These values of Q_{10} and Q_{40} are the desired estimates of peak rates of runoff having a recurrence interval of 10, and 40 years, respectively.

LIMITATIONS AND PRECAUTIONS

Limitations in Basic Data - Because of the limited amount of basic data that went into this study, particularly for watersheds having less than 100 square miles contributing area, the results presented in this study must be considered as tentative and subject to revision as new data become available.

Errors of Estimate - The distribution of error curve for estimates of C_{10} from Fig. 11, from which Fig. 14 was derived, are given in Fig. 15. This shows that approximately 76 per cent of the dependent sample has an error of estimate less than 25 per cent. This is considered as acceptable accuracy for field design purposes, since it complies with the basic accuracy requirement of at least 67 per cent of the sample having a departure of less than $\frac{1}{2}$ 25 percent from the fitted regression.

Recommended Maximum and Minimum Values of Q_{10} as a function of Area - Maximum recorded peak rates of flow as a function of watershed size are shown in Fig. 16 as an envelope curve. (Source: U.S. Bureau of Reclamation, compiled largely from U.S. Geological Survey records.) Observed values of Q_{10} at 15 stations in D-13 and D-20 areas are plotted on Fig. 16 with small circles. The curves giving maximum and minimum recommended values of Q_{10} were computed on the basis of the graphical correlation in Fig. 11. These two curves serve as envelopes for estimation of Q_{10} for the streams with $S_{.9L} = 10 - 50$ ft/mi in the study area. The upper curve corresponds to $S_{.9L} = 50$ ft/mi, and the lower curve to $S_{.9L} = 10$ ft/mile.

SUMMARY OF RESULTS OF RELATED STUDIES

Characteristics of precipitation associated with annual maximum flood events. - From a study of precipitation amounts associated with annual maximum flood events from nine watersheds in Colorado in the foothills of the Rocky Mountains, it was concluded that for watersheds equal to or greater than about 900 square miles, two-thirds or more of the annual maximum floods were probably caused by rains covering the entire watershed; while for watersheds smaller than about 50 square miles, one-third or less are produced by such rains.

Estimates of clock-hourly precipitation from precipitation amounts of longer duration - A study was made to determine the interrelations among precipitation amounts for various time periods for a given recurrence interval for precipitation records for stations located in eastern Colorado. Preliminary studies show that estimates of clock-hourly precipitation can be made with satisfactory accuracy from records of precipitation amounts of longer duration.

Utilization of weather radar data to provide increased areal coverage of rainfall events - Attempts were made to utilize two types of weather radar data to extend the areal coverage for individual rainfall events. Hand-drawn sketches of the Plan-Position Indicator (PPI) scope from a 5.5 cm set used by United Air Lines in Denver, and sketches of a PPI scope reconstructed from coded descriptions of radar echo data from a 3.0 cm set at Lowry Air Force Base in Denver were studied. It was concluded that the data in this form were not suitable for the intended purpose.

Correlation of precipitation factors with physiographic parameters— Mean monthly precipitation for May was correlated with position (latitude, longitude, and elevation) for 48 stations in eastern Colorado. Details of this correlation follow:

Independent variable (Y) = mean monthly precipitation for May,
inches

Independent variables

X, = latitude, less 30 degrees.

X2 = longitude, less 100 degrees.

 X_3 = elevation, in 10 thousands of feet.

Station groupings

Group 1: Nineteen (19) stations in Colorado in the Arkansas River drainage.

Group 2: Twenty-nine (29) stations in the Platte and Kansas drainage in Colorado.

Results included the regression equation, the correlation coefficient \overline{R} , the standard error of estimate \overline{S} , and the standard deviation (σ) of the individual coefficients.

Results

Group 1:
$$Y = 2.99 - 0.045X_1 - 0.55X_2 + 2.95X_3$$

$$\overline{R} = .72 \text{ (Significant at 99 per cent level)}$$

$$\overline{S} = .38$$

$$\sigma_1 = .17$$

$$\sigma_2 = .12$$

$$\sigma_3 = .73$$

Group 2:

Y = 3.33 + 0.03
$$X_1$$
 + 0.15 X_2 - 3.43 X_3
 \overline{R} = 0.67 (Significant at 99 per cent level)
 \overline{S} = 0.37
 σ_1 = .13
 σ_2 = .09
 σ_3 = .90

Seasonal distribution of annual maximum flood events - A study was made to determine the effect of watershed contributing area and elevation on the seasonal distribution of annual maximum flood events from sixty-two (62) stations drawn from all parts of Colorado except the San Luis Valley.

Results indicate that the average date of occurrence of 67 per cent of the annual maximum floods advances with increase in watershed size.

For watersheds having an elevation less than 7683 feet, the date of occurrence of 67 per cent of the annual maximum floods advances with decreasing elevation.

These results can be interpreted in terms of summertime rains as a cause of flood events on the plains, as compared to snow melt, or a combination of snow melt and rain as a cause of flood events in the mountain areas.

V. DISCUSSION

RESEARCH NEEDS

Need for additional records from small watersheds - Throughout the course of this study it was evident that there was an acute shortage of suitable records of runoff from small watersheds in the study area. This scarcity of adequate records of runoff was probably the most severe limitation in the statistical sampling procedure for determining relations suitable for prediction of magnitude and frequency of runoff from small watersheds.

One of the most valuable contributions to knowledge in the field of small watershed hydrology in this region would be the establishment of additional records of runoff for watersheds having a contributing area less than 100 square miles.

Such re-examination should have the objective of determination of a distribution that would produce the characteristic of linearity when applied to peak flood flows from small watersheds in semi-arid regions.

PLANS FOR FUTURE STUDIES

The similarity in relations between Q_{10} and Q_{40} for stations over a broad region (See Fig. 7) suggests that an analysis of the type presented in this study could also produce usable results in the areas adjacent to the D-13 and D-20 problem areas. Adjacent areas wherein records of peak rates of runoff were compiled in this study include the following (See Fig. 1 for locations.)

South: D21, D22, D29.

East: D14, D16, D18, D19.

North: D6, D7, D8, D9.

In addition to these areas, the Rocky Mountain Foothills region (E-5) contains a considerable number of records of runoff suitable for analysis.

Plans for future studies include the analysis of records collected from these areas, with the objective of developing suitable relations for estimates of peak rates of runoff from ungaged watersheds.

Priority of endeavor will be determined in consultation with the sponsor of the work.

In addition, work is underway on certain of the items mentioned previously in the section "RESEARCH NEEDS." Specific items include the following:

Study of the effects of diversions for irrigation on peak rates of runoff.

A better delimitation of precipitation characteristics as they affect runoff from small watersheds.

Study of methods of representing frequency distribution of peak rates of runoff with linearity in semi-arid regions.

ACKNOWLEDGMENTS

Throughout this study, a large number of persons have contributed suggestions and comments from their experience which have been of considerable help in the study.

Particular thanks are due Mr. W. T. Miller of the Denver office of the U. S. Geological Survey for his assistance in collection of the basic data included herein. Thanks are also due the U. S. Geological Survey offices in Lincoln, Nebraska and Topeka, Kansas for their assistance in collection of basic data.

The assistance of H. T. Harrison and W. B. Beckwith of United Air Lines is gratefully acknowledged.

The authors wish to thank Mr. Carl Izzard and Mr. W. D. Potter for their helpful suggestions made during the study, and to thank the U. S. Bureau of Public Roads for sponsoring the study.

REFERENCES

- (1) Potter, W. D., Upper and Lower Frequency Curves for Peak Rates of Runoff, American Geophysical Union Transactions, 39(1): 100-105, February 1958, (With discussions 39(3): 497-500 and 39(6): 1165-1170.
- (2) Benson, M. A., Discussion of "Upper and Lower Frequency Curves for Peak Rates of Runoff," American Geophysical Union Transactions, 39(6): 1167-1169, December 1958.
- (3) Linsley, R. K., M. A. Kohler, and J. Paulhus, Hydrology for Engineers, McGraw-Hill, 340p, 1958.
- (4) Spreen, Wm. C., Topographically adjusted normal isohyetal maps for western Colorado, U.S. Weather Bureau Technical Paper, No. 4, 1947.
- (5) Benson, M. A., Channel-Slope Factor in Flood-Frequency Analysis, Journal of the Hydraulics Division, A.S.C.E., p. 1-9, April 1959.
- (6) Hourly precipitation data for Colorado, Annual Summary, 1957.
- (7) Rainfall intensities for local drainage design in western United States, U.S. Weather Bureau Technical Paper, No. 28, 1956.
- (8) Alexander, G. N., Letter to the editor, Journal of Geophysical Research, 64(1): 132, January 1959.
- (9) Rowe, R. R., et al. Flood frequency by regional synthesis.

 Transactions of American Geophysical Union 38: 879-84,
 1957.

APPENDIX NO. 1.

IDENTIFICATION OF STATIONS USED IN THE STUDY

Appendix No. 1, Part s-1

Identification of Caging Stations Used In The Study (Includes every station for which frequency plot was made)

Stations Inside D-13 & D-20 Problem Areas, Including Fringes

Number	Name .	Refer to U.S.G.S.
	Described to Control of the Control	Water Supply Paper
	Fountain Creek at Pueblo, Colo.	159-1311
	Apishapa River near Fowler, Colo.	184-1311
	Timpas Creek near Rocky Ford, Colo.	186-1311
	Horse Creek near Sugar City, Colo.	191-1311
	Rawhide Creek near Lingle, Wyo.	126-1310
	Blue Creek near Lewellen, Nebr.	155-1310
	Birdwood Creek near Hershey, Hebr.	165-1310
	Cherry Creek near Franktown, Colo.	201-1310
	Cherry Creek near Melvin, Colo.	202-1310
	Cherry Creek below Cherry Creek Dam, Colo.	203-1310
	Cherry Creek at Denver, Colorado	204-1310
	Lodgepole Creek at Bushnell, Nebr.	288-1310
	North Fork Republican River at Colorado-Nebraska State Line	387-1310
	Buffalo Creek near Haigler, Hebr.	388-1310
	Rock Creek near Parks, Hebr.	389-1310
	Frenchman Creek below Champion, Hebr.	400-1310
	Sappa Creek near Oberlin, Kansas	420-1310
	White River at Crawford, Nebr.	332-1439
	Niobrara River above Box Butte Reservoir, Hebr.	343-1439
	Pumpkin Creek near Bridgeport, Nebr.	151-1310
	Landsman Creek near Hale, Colo.	392-1310
	South Fork Republican River near Idalia, Colo.	391-1310
	Cottonwood Creek at Wendover, Wyo.	99-1310
	Frenchman Creek near Hamlet, Nebr.	403-1310
37	Purgatoire River at Trimidad, Colo.	193-1311
38	Vermejo River near Dawson, New Mexico	323-1311
	Sheep Creek near Morrill, Nebr.	137-1310
41	Dry Spotted Tail Creek at Mitchell, Nebr.	139-1310
42 !	Tub Spring near Scottsbluff, Nebr.	141-1310
43	Winter Creek near Scottsbluff, Nebr.	143-1310
44	Minemile Drain near Mcgrev, Nebr.	146-1310
45	Bayard Sugar Factory Drain near Bayard, Mebr.	147-1310
46 1	Red Willow Creek near Bayard, Nebr.	149-1310
47 1	Bijou Creek near Wiggins, Colo.	282-1310
	Buffalo Creek near Darr, Nebr.	300-1310
• • •	Buffalo Creek near Overton, Nebr.	301-1310
	Elm Creek near Overton, Nebr.	302-1310
	Wood River near Riverdale, Nebr.	305-1310
	Wood River near Cibbon, Nebr.	306-1310
	Middle Loup River at Dunning, Nebr.	312-1310
	Middle Loup River at Arcadia, Nebr.	318-1310 318-1310

			7 3			
40.00						
COCE	tion	Drainage A	res in Sq. Mile	Period of	. Problem	
Longi tude			Contributing		Area	* **
104-35-40		926		700		
					D-50	
103-59		1125	2000	2	D-20	
103-13-20					D-20	
103-37-16					D-50	
104-19-20	42-01-30				D-13	
102-10	41-20	561		24	D-13	4
101-04-3	41-13	286		23	D-13	
104-45-5	39-21-30	172		19	D-13	
104-10-13	39-30-20	369		19	D-13	
104-21-40	39-39-10	38 6 420	24		D-13.	
105-00-05	139-44-20°	420		10	D-13	A
		1090			D-13	
102-03-03	40-04-10	320 18 0		20 24 18	D-13	1
101-54-15 101-13-10	140-02-45	180			D-13	W 107
101-4	NO-28-00	E SALA			D-13	
100 23 0	33-18-13	940 1040			D-13	
1.00 Januar	12-41				J-1	
103-25	10 OF 76	8 313 1400	Can Marie		D-13	
103-10-13	1228	1080			1	C 10.5
103-04-04	31-38. · · · · · · · · · · · · · · · · · · ·	450			D-13	
102-1	20 27 00	1306	700		D-13	6.4
104-5	39-37-00 12-19-37	# 15 9 8		2 2 100	D-13	
	MD-22-30	- 148G	ocar i		P D IF	
104	37-10-15	795		77.	D-26	
101-1-1	14-10-50	or			D-2	301
103	\$1-58 ·	44.19				200
1034	11.57		***			
	4T-55	18 1 E		1		146
103	1.52			200		
10				200		
	A1-44			20		1.20
103-3	47.43			20		
103-1 104-0	MG_1A_53	TASIO				100
99-5484	40-54-00	1000		10		
99-78-20	40-44	7.76				
	40-50-40	21/2		7 Resemble 1265	Anna a Maria	l
	40-47-50	175 - 31 - 373		12		P. (4)
	40-46-10	KIL		1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1	Loss land	
100-06-20	41 -49-50	572	A Part No. 10 and	10		
	11-25 22	1760	O.	Proposition State National State		100
79-00-0	41-25-20	4730	020	19 19 19 19 11 11 11 11 11 11 11 11 11 1	1	l i
	S				÷ ;	İ
726		7				Ī

		mel Slope Fac	
L in Miles	Eo.qL in Pt.	EGS in Ft.	50.9L
71.4	6920	4663	35.2
100.3	7530	4317	35-5
48.0	5270	4220	24.3
1000	7-1 -	(55	
			Ì
22.6	2700	2200	12.77
31.6	3700	3309	13.7
29.0	3200	2920	10.7
25.7	7380	6150	53.3
44.9	7320	5608	42.3
1	•		
	•	•	
93.9	7120	4812	27.3
25.5	3760	3336	18.5
10.2	3460	3204	27.9
9.2	3250	3093	19.0
80.6		3240	
۵.6	4190	3240	13.1
	i		
	0-	1	
122.0	5180	4015	10.6
55.9	4330	3636	13.8
56.1	4615	3720	17.7
88.1	5210	3680	19.3
1	·	_	
I			,
•		·	
1	1		
(Power's)			

(Remark)

- 1. L Main Channel Length from Caging Station to Headwater
- 2. E_{0.9L} Elevation of Channel Bed at 90% of L upstream from Gaging Station
- 3. EGS Elevation of Channel Bed at Gaging Station
- 4. $S_{0.9L} = \frac{E_{0.9L} E_{GS}}{0.9L}$

Appendix No. 1, Fart a-1 (continued)

Identification of Gaging Stations Used In The Study (Includes every station for which frequency plot was made)

Stations Inside D-13 & D-20 Problem Areas, Including Fringes

Serial	Name	Refer to U.S.G.S.		
Number	D. LANGER	Water Supply Paper		
	South Loup River at Ravenna, Nebr.	322-1310		
	Mud Creek near Sweetwater, Nebr.	324-1310		
57	Oak Creek near Dannebrog, Nebr.	326-1310		
	Arikaree River at Haigler, Nebr.	385-1310		
	South Fork Republican River near Hale, Colo.	393-1310		
60	Frenchman Creek near Imperial, Nebr.	401-1310		
61	Frenchman Creek near Enders, Mebr.	402-1310		
	Frenchman Creek at Palisade, Nebr.	404-1310		
	Stinking Water Creek near Wauneta, Nebr.	405-1310		
	Stinking Water Creek near Palisade, Nebr.	406-1310		
65	Blackwood Creek near Culbertson, Nebr.	408-1310		
	Driftwood Creek near McCook, Nebr.	409-1310		
67	Red Willow Creek near McCook, Nebr.	41 1- 1310		
	Red Willow Creek near Red Willow, Nebr.	412-1310		
69	Medicine Creek at Maywood, Nebr.	290-1440		
	Fox Creek at Curtis, Nebr.	292-1440		
	Dry Creek near Curtis, Nebr.	293-1440		
	Medicine Creek above Harry Strunk Lake, Hebr.	413-1310		
	Mitchell Creek above Harry Strunk Lake, Mebr.	41#-1310		
	Medicine Creek at Cambridge, Nebr.	417-1310		
	Muddy Creek at Arapahoe, Nebr.	300-1440		
	Prairie Dog Creek at Norton, Kans.	426-1310		
77	Cottonwood Creek near Bloomington, Nebr.	311-1440		
	Rose Creek near Wallace, Kans.	441-1310		
79	North Fork Smoky Hill River near McAllaster, Kans.	442-1310		
	Big Creek near Hays, Kans.	448-1310		
	Bow Creek near Stockton, Kans.	351-1440		
	North Fork Solomon River at Kirwin, Kans.	461-1310		
_ 1	Fountain Creek near Fountain, Colo.	158-1311		
	St. Charles River near Rueblo, Colo.	163-1311		
	Apishapa River near Aguilar, Colo.	181-1311		
	Purgatoire River near Alfalfa, Colo.	194-1311		
	Cimarron River near Guy, New Mexico	240-1311		
	Canadian River near Hebron, New Mexico	316-1311		
_	White River below Cottonwood Creek near Whitney, Nebr.	333-1439		

(continued)

is Used In The Study 'equency plot was made)

reas, Including Fringes

Loca	tion	Drainage Ar	rea in Sq. Mile	Period of	Problem
Longi tude	Latitude	Nominal	Contributing	Record in Years	Area
98-54-45	41-00-35	1660	890	14	ļ 1
98- 59- 45		678		11	!
98-38-30		122		8	
101-57-25		1460		19]
102-09-45		?		9	
101-37-30		1220	760	17	[]
101-30-35		1300	820	12	<u> </u>
101-07-40		1500	980	8	[
101-19-50		1260	3 4 0	10	}
101-06-50		1390	430.	9	
100-48-25		290		12	1
100-39 - 40	40-08-50	360		13	
100 - 39	40-21	600	300	7	1
100-30-00	40-14-10	710	400	18	}
100-36-40		207		8	
100-29-20		77		8 8 8	
100-26-40		20			i
100-19-20		?		9 9	1
100-15-25		53			1
100-10-35	40-17-55	1070	6 80	20	}
99-54-40	40-18-20	243		8	}
99-53	39-50	721		15	Ī
99-03-55	40-05-10	17		7 7	•
101-38	38-53	28		7	{
101-22	39-01	670		7	<u> </u>
99-19	38-51	594		13 8	
99-17	39-34	337			
99-07	39-40	1360		11	1
104-40-13		676	A-A-#-	16	
104-31-40		468		12	
104-39-50	37-22-50	126		11	
104-07-30		1320		7	1
103-25-25	36-59-15	5 ¹ 45		17	}
104-27-45		229		12	
103-10-05	42-46-35	6 76		8	į <u> </u>

Appendix No. 1, Part a-2

Identification of Gaging Stations Used In The Study (Includes every station for which frequency plot was made)

Stations Outside D-13 & D-20 Problem Areas

Serial	Name	Refer to U.S.G.S.
Number	Neme	Water Supply Paper
		2 222
101	Floyd River at James, Iowa	3-1310
102	Elkhorn River at Neligh, Nebr.	352-1310
103	Tarkic River at Fairfax, Mo.	371-1310
104	Nodaway River near Burlington Junction, Mo.	377-1310
105	Little Blue River near Endicott, Nebr.	478-1310
106	Soldier Creek near Topeka, Kans.	484-1310
107	Delaware River at Valley Falls, Kans.	485-1310
108	Wakarusa River near Lavrence, Kans.	487 - 13 1 0
109	Stranger Creek near Tonganoxia, Kans.	488-13 1 0
110	Marais des Cygnes River near Ottaws, Kans.	524-1310
111	Pawnee River near Larned, Kans.	2 19- 1311
112	Little Arkansas River at Valley Center, Kansas	224-1311
113	Walnut River at Winfield, Kans.	229-1311
114	Spring River near Waco, Mo.	296-1311
115	Rayado Creek at Sauble Ranch, near Cimarron, New Mexico	333-1331
116	Cimarron River at Springer, New Mexico	338-1311
117	Mora River near Golondrinas. New Mexico	346-1311
118	Coyote Creek near Golondrinas, New Mexico	348-1311
119	Mora River near Shoemaker, New Mexico	351-1311
120	Mountain Fork River near Eagletown, Okla.	533 -1311
121	Kiamichi River near Belzoni, Okla.	528-1311
122	Judith River near Utica, Mont.	67-1439
123	Musselshell River at Harlowton, Mont.	75-1439
124	Flatvillow Creek near Flatvillow, Mont.	81-1439
125	South Fork Milk River near International Boundary	86-1439
126	North Fork Milk River above St. Mary Canal near Browning, Mont.	87-1439
127	North Fork Milk River near International Boundary	88-1439
128	Battle Creek at International Boundary	99-14 39
129	Woodpile Coulee near International Boundary	100-1439
13ó	East Fork Battle Creek near International Boundary	101-1439
131	Whitewater Creek near International Boundary	104-1439
132	Clarks Fork at Chance, Mont.	143-1439
133	Bull Lake Creek near Lenore, Wyo.	154-1439
134	Greybull River at Meeteetse. Wyo.	131-1439
135	Goose Creek near Sheridan, Wyo.	205-1439
/رد-	coons or their print runtle share	207-1409

s Used In The Study equency plot was made)

Problem Areas

	tion		rea in Sq. Mile	Period of	Problem
Longi tude	Latitude	Acminal	Contributing	Records in Years	Area
96-18-45	la al as	900			
		88 2	*******	22 .	
98-01-40		2200	1800	25	•
95-24-20		508		34	
95-05-20	40-26-40	1240		34	
97-08-10		2340		28	
95-43	39-06	268		25	
95-27	39-21	922		35	
95-16	38-55	458		28	
95-01-08		406		28	
95-15	38-37	1250		38	
99-20	38-11	2148	2010	32	
97-23	37-50	1327	1250	32 34	
97-00	37-14	1327 1840		35	
94-33-55		11 <i>6</i> 4		26	
104-58	36-22	6 5		33	
D4-35-50		1.032		. 27	
05-09-30		273		26	
05-09-50	35-54-40	257		28	
04-47	35-48	2304	1033	39	
94-37	3403	787		27	
95-29	34-32	1423		31	-
10-14	46-54	331		36	
09-51	46-26	1130		36 42	
08-37	46-47	195		<u>र्</u> या	
12-32-20	49-00	433		3 ¹ 4 43	
13-03	48-59	62	****	36 l	
12-58	49-02	101		36 42	'
09-25-20	49-00-10	726	}	39	
09-31-50	48-59-00	70		27	
09-08	48-58	95		26	
07-51	48-57	300			
09-05	45-00	1140		29 22	
	43-14-33	222	*****	30	
08-52-35	44-09-20	690		39 36	
07-11	44-42	120		.27	

Appendix No. 1, Part a-2 (continued)

Identification of Gaging Stations Used In The Study (Includes every station for which frequency plot was made)

Stations Outside D-13 & D-20 Problem Areas

Serial Number	Name	Refer to U.S.G.S. Water Supply Paper
1	Clear Fork Trinity River at Fort Worth, Texas Middle Concho River near Tankersly, Texas North Concho River near Carlsbad, Texas Pecan Bayou At Brownwood, Texas North Llano River near Junction, Texas Llano River near Junction, Texas Guadalupe River near Spring Branch, Tex. Guadalupe River above Comal River at New Braunfels, Tex. Blanco River at Wimberley, Tex. Plum Creek near Luling, Tex. Cibolo Creek near Falls City, Tex. Nueces River at Laguna, Tex. Frio River at Concan, Tex. Rio Grande near Del Norte, Colo.	
150 151 152 153 154 155 156 157 158	Cunejos River near Mogote, Colo. Red River near Questa, New Mexico Santa Fe River near Santa Fe, New Mexico Blue Water Creek near Bluewater, New Mexico Pecos River near Pecos, New Mexico Pecos River near Anton Chico, New Mexico Gallinas River near Montezuma, New Mexico Gallinas River at Montezuma, New Mexico Mimbres River near Mimbres, New Mexico	278-1442 298-1442 318-1442 333-1442 352-1442 353-1442 354-1442 403-1442

(continued)

s Used In The Study equency plot was made)

Problem Areas

	tion		rea in Sq. Mile	Period of	Problem
Longitude	Latitude	Nominal	Contributing	Records in Years	Area
	i				
97-21	32-44	5 26		24	
100-36-50			1128	26	
100-39	31-36	1533	1410	26	
98-58-30	_	1614		26	
99-47	30-30	914		31	
99:44	30-30	1874		24	
98-23	29-51-40	1282		30	
98-06-40	29-42-55	1516		28	
98-04	29-59	364		27	
97-37	29-42	356		26	
97-56	29-01	831		24	1.
99-59-50	29-25-45	764		29	
99-42	29-29	405		23́	
106-27-30	37-41-20	1320		46	
106-11-20	37-03-20	282		44	
105-34	36-42-10	112		25	
105-50-35	35-41-10	20		25	
108-91-40	35-17 -5 0	235		25	
105-41:	35-42-25		189	25	
105-06-20			1050 -	23	
105-19-10	35-39	84		24	
105-16-30	35-39-15	87		23	
107-59	32-52-20	152		25	

APPENDIX No. 1. Part b. Gaging Stations in Alphabetical Order.

Name	Serial Number	Refer. to USOS Water Supply Paper
Apishapa River near Aguilar, Colo. Apishapa River near Fowler, Colo. Arikaree River at Haigler, Nebr. Battle Creek at International Boundary Boyard Sugar Factory Drain near Bayard, Nebr. Big Creek near Hays, Kans. Big Creek near Heys, Kans. Big Creek near Hershey, Nebr. Blackwood Creek near Gulbertson, Nebr. Blackwood Creek near Gulbertson, Nebr. Blanco River at Wimberley, Tex. Blue Creek near Lewellen, Nebr. Bluewater Creek near Bluewater, N. M. Bow Creek near Stockton, Kans. Buffalo Creek near Bluewater, Nebr. Buffalo Creek near Overton, Nebr. Buffalo Creek near Haigler, Nebr. Buffalo Creek near Henore, Wyo. Canadian River near Eabron, N. M. Cherry Creek below Cherry Creek Dam, Colo. Cherry Creek near Falls City, Tex. Cimarron River at Springer, N. M. Clarks Folk at Chance, Mont. Clear Fork Trinity River at Fort Worth, Tex. Cottonwood Creek near Rloomington, Nebr. Cottonwood Creek near Mogote, Colo. Deleware River at Valley Falls, Kans. Driftwood Creek near McCook, Nebr. Dry Creek near Gurtis, Nebr. Dry Creek near Curtis, Nebr. East Fork Battle Creek at Mitchell, Nebr. East Fork Battle Creek at Mitchell, Nebr. East Fork Battle Creek near International Boundary Elkhorn River at Nellajh, Nebr. Flatwillow Creek near Flatwillow, Mont. Floyd River at James, Iowa. Fountain Creek at Overton, Nebr. Fountain Creek at Pueblo, Colo. Fox Creek at Curtis, Nebr. Frenchman Creek below Champion, Nebr.	85 388 58 47 11 55 44 10 33 18 48 19 49 33 88 14 15 12 13 16 77 55 88 15 15 15 16 71 14 15 16 16 18 17 18 18 18 19 18 18 18 18 18 18 18 18 18 18 18 18 18	181-1311 184-1311 385-1310 99-1439 147-1310 282-1310 165-1310 282-1310 282-1310 333-1442 351-1440 301-1310 201-1310 201-1310 201-1310 201-1311 202-1311 338-1311 143-1439 351-1440 99-1310 293-1440 99-1310 293-1440 139-1310 139-1310 139-1310 159-1311 159-1311 159-1311 159-1311

	t /	,
Frenchman Creek near Enders, Nebr.	61	402-1310
Frenchman Creek near Hamlet, Nebr.	36	403-1310
Frenchman Creek near Imperial, Nebr.	60	401-1310
Frenchman Creek at Palisade, Nebr.	62	404-1310
Frio River at Concan, Tex.	148	239-1442
Gallinas River at Montezuma, N. M.	157	355-1442
Gallinas River near Montezuma, N. M.	156	354-1442
Goose Creek near Sheridan, Wyo.	135	205-1439
Greybull River at Meetectse, Wyo.	134	181-14 3 9
Guadalupe River above Comal River at New Braunfels, Tex.	143	214-1442
Guadalupe River near Spring Branch, Tex.	142	213-1442
Horse Creek near Sugar City, Colo.	5	191-1311
Judith River near Utica, Mont.	122	67-1439
Kiamichi River near Belzoni, Okla.	121	528 -1311
Landsman Creek near Hale, Colo.	33	392-1310
Little Arkansas River at Valley Center, Kans.	112	224-1311
Little Blue near Endicott, Nebr.	105	478-1310
Illano River near Junction, Tex.	141	196-1442
Lodgepole Creek at Bushnell, Nebr.	16	288-1310
Marais des Cygnes River near Ottowa, Kans.	110	524-1310
Medicine Creek at Cambridge, Nebr.	74	417-1310
Medicine Creek above Harry Strunk Lake, Nebr.	72	413-1310
Medicine Creek at Maywood, Nebr.	69	290-1440
Middle Concho River near Tankersly, Tex.	137	172-1442
Middle Loup River at Arcadia, Nebr.	54	318-1310
Middle Loup River at Dunning, Nebr.	53	312-1310
Mimbres River near Mimbres, N. M.	158	403-1442
Mitchell Creek above Harry Strunk Lake, Nebr.	73	414-1310
Mora River near Golondrinas, N. M.	117	346-1311
Mora River near Shoemaker, N. M.	119	351-1311
Mountain Fork River near Eagletown, Okla.	120	533-1311
Mud Creek near Sweetwater, Nebr.	56	324-1310
Muddy Creek at Arapahoe, Nebr.	75	300-1440
Musselshell River at Harlowton, Mont.	123	75 - 14 3 9
Ninemile Drain near Megrew, Nebr.	44 .	146-1310
Niobrara River above Box Butte Reservoir, Nebr.	25	343-1439
Nodaway River near Burlington Junction, Mo.	104	377-1310
North Concho River near Carlsbad, Tex.	138	175-1442
North Fork Milk River near International Boundary	127	88-1439
North Fork Milk River above St. Mary Canal near	,,	-0 2.37
Browning, Mont.	126	87-1439
N.F. Republican River at Colorado-Nebraska State Line	18	387-1310
North Fork Smoky Hill River near McAllester, Kans.	79	442-1310
North Fork Solomon River at Kirwin, Kans.	82	461-1310
North Liano River near Junction, Tex.	140	195-1442
Nueces River at Laguna, Tex.	147	234-1442
Oak Creek near Dannebrog, Nebr.		
Pasmee River near Larned, Kans.	57	326-1310
Pecan Bayou at Brownwood, Tex.	111	219-1311 188-1442
Tooks before the minimiser's Toy's	139	100-1442
	I	Į .

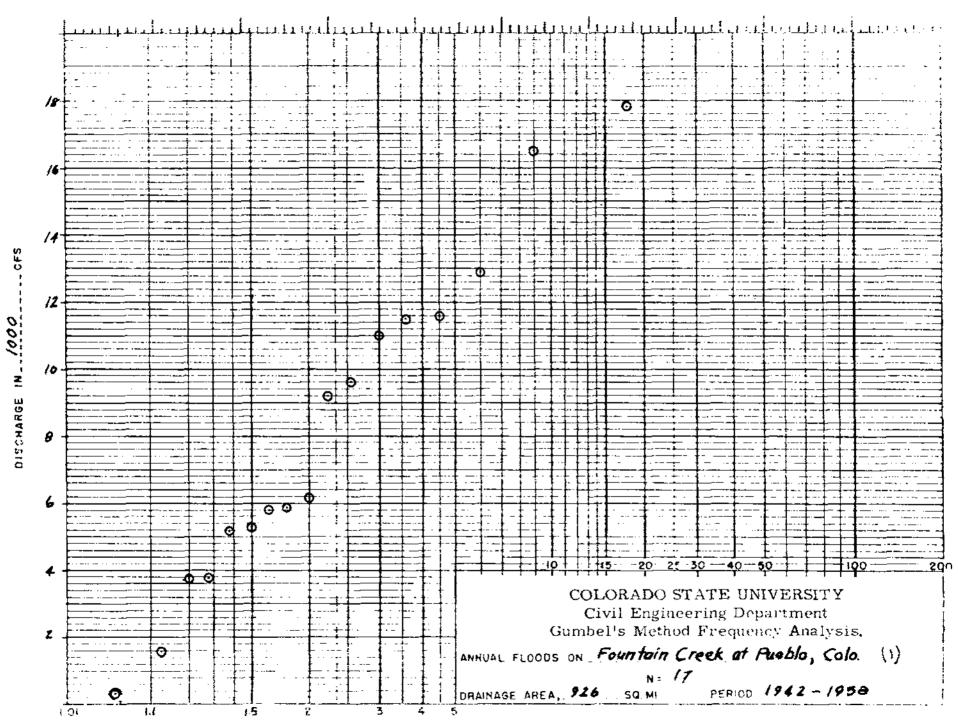
APPENDIX No. 1. Part b. Gaging Stations in Alphabetical Order (Cont'd)

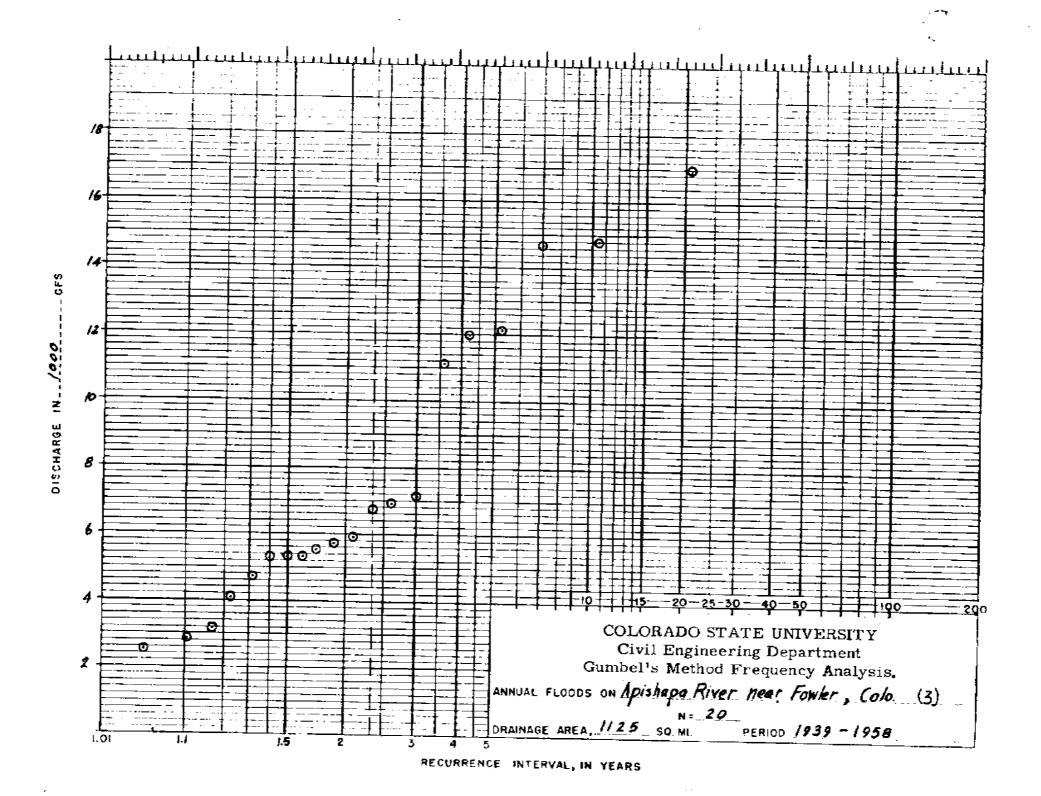
There Manager was the state of the state of	t -	1
Pecos River near Anton Chico, N. M.	155	353-1442
Pecos River near Pecos, N. M.	1.54	352-1442
Plum Creek near Luling, Tex.	145	219-1442
Prairie Dog Creek at Norton, Kans.	76	426-1310
Pumpkin Creek near Bridgeport, Nebr.	31 86	151-1310
Purgatoire River near Alfalfa, Colo.		194-1311
Purgatoire River at Trinidad, Colo.	37	193-1311
Rawhide Creek near Lingle, Wyo.	9	126-1310
Rayado Creek at Sauble Ranch, near Cimarron, N. M.	115	333-1311
Red River near Questa, N. M.	151	298-1442
Red Willow Creek near Bayard, Nebr.] 46	149-1310
Red Willow Creek near McCook, Nebr.	67	411~1310
Red Willow Creek near Red Willow, Nebr.	. 68	412-1310
Rio Grande near Del Norte, Colo.	149	259-1442
Rock Creek near Parks, Nebr.	20	389-1310
Rose Creek near Wallace, Kans.	78	441-1310
Santa Fe River near Santa Fe, N. M.	152	318-1442
Sappa Creek near Oberlin, Kans.	23	420-1310
Soldier Creek near Topeka, Kans.	106	484-1310
South Fork Milk River near International Boundary	125	86-1439
South Fork Republican River near Hale, Colo.		393-1310
South Fork Republican River near Idalia, Colo.	59 34	391-1310
South Loup River at Ravenna, Nebr.	55	322-1310
Sheep Creek near Morrill, Nebr.	40	137-1310
Spring River near Waco, Mo.	114	296-1311
St. Charles River near Pueblo, Colo.	84	163-1311
Stinking Water Creek near Palisade, Nebr.	64	406-1310
Stinking Water Creek near Wauneta, Nebr.	63	405-1310
Stranger Creek near Tonganoxia, Kans.	109	488-1310
Tarkio River at Fairfax, Mo.	103	371-1310
Timpas Creek near Rocky Ford, Colo.	4	186-1311
Tub Spring near Scottsbluff, Nebr.	42	141-1310
Vermejo River near Dawson, N. M.	38	323-1311
Wakarusa River near Lawrence, Kans.	108	487-1310
Walnut River at Winfield, Kans.	113	229-1311
White River below Cottonwood Creek near Whitney, Nebr.	89	333-1439
White River at Crawford, Nebr.	24	332-1439
Whitewater Creek near International Boundary	131	104-1439
Winter Creek near Scottsbluff, Nebr.	43	143-1310
Wood River near Gibbon, Nebr.		306-1310
Wood River near Riverdale, Nebr.	52 51	
Woodpile Coulee near International Boundary		305-1310
" " " A PARTY OF THE THE THE TOTAL PONTERTY	129	100-1439
	•	

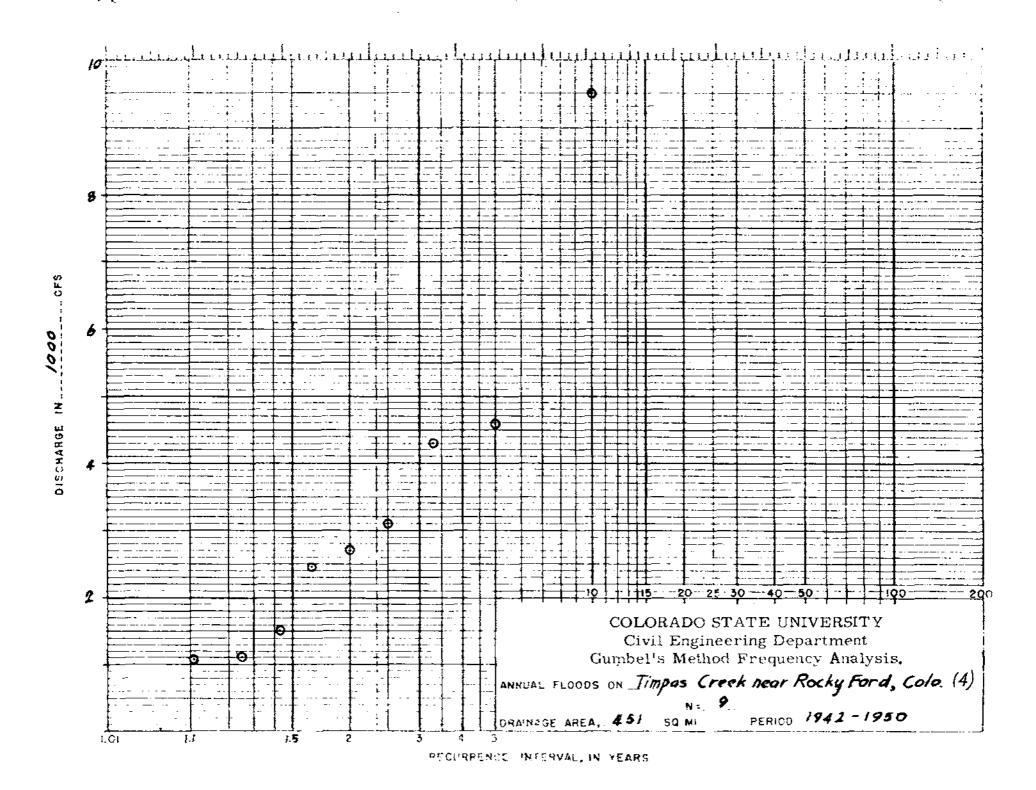
APPENDIX NO. 2

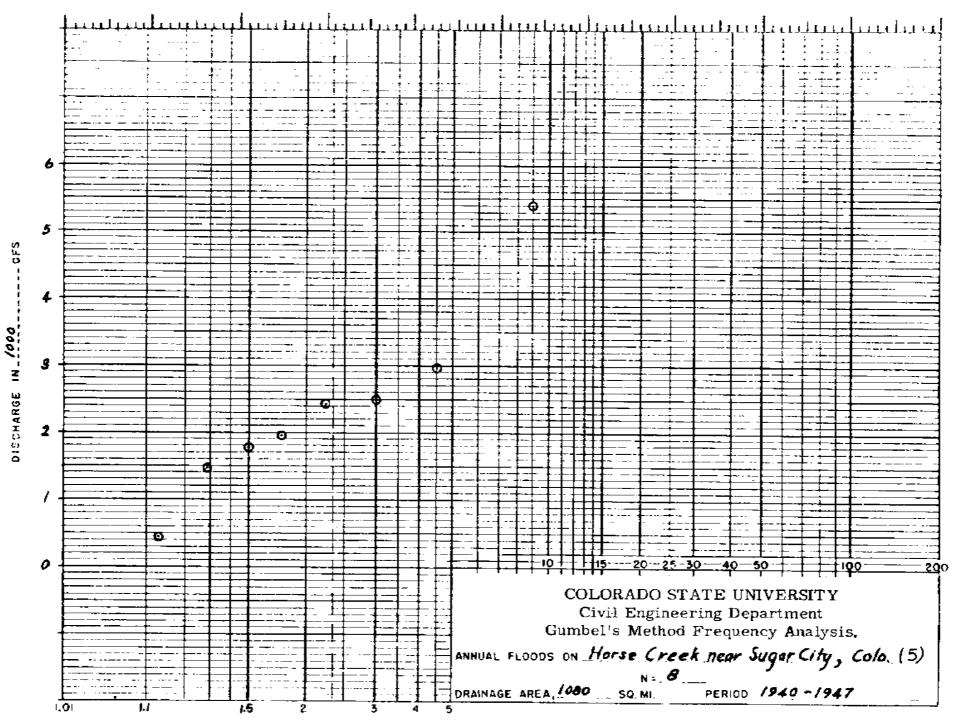
GUMBEL PLOTS FROM INDIVIDUAL STATIONS

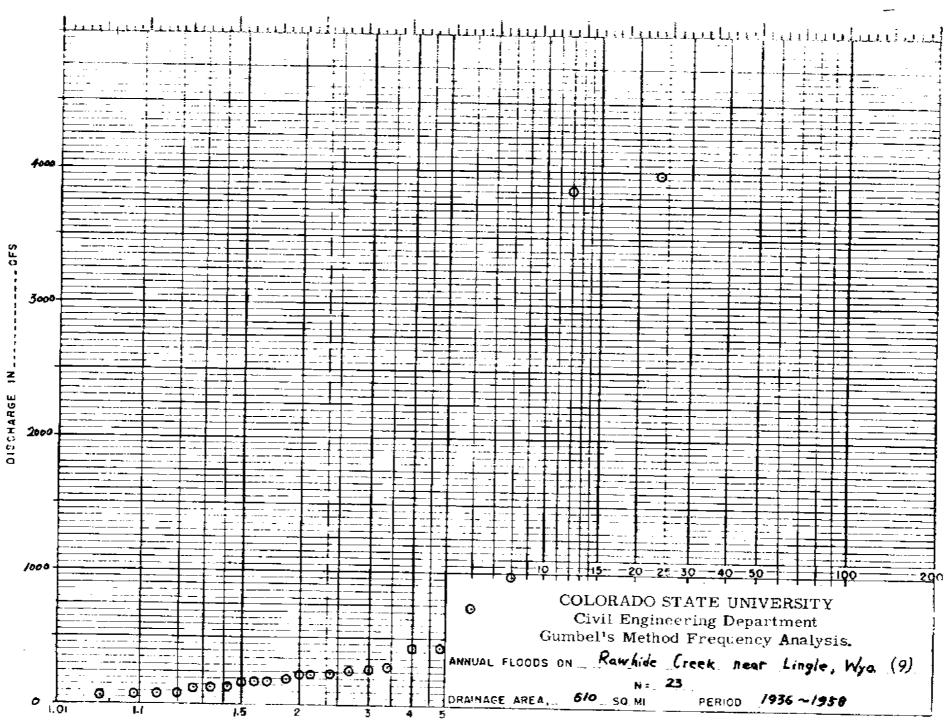
(Arranged in the Order
Given in Appendix 1-a)

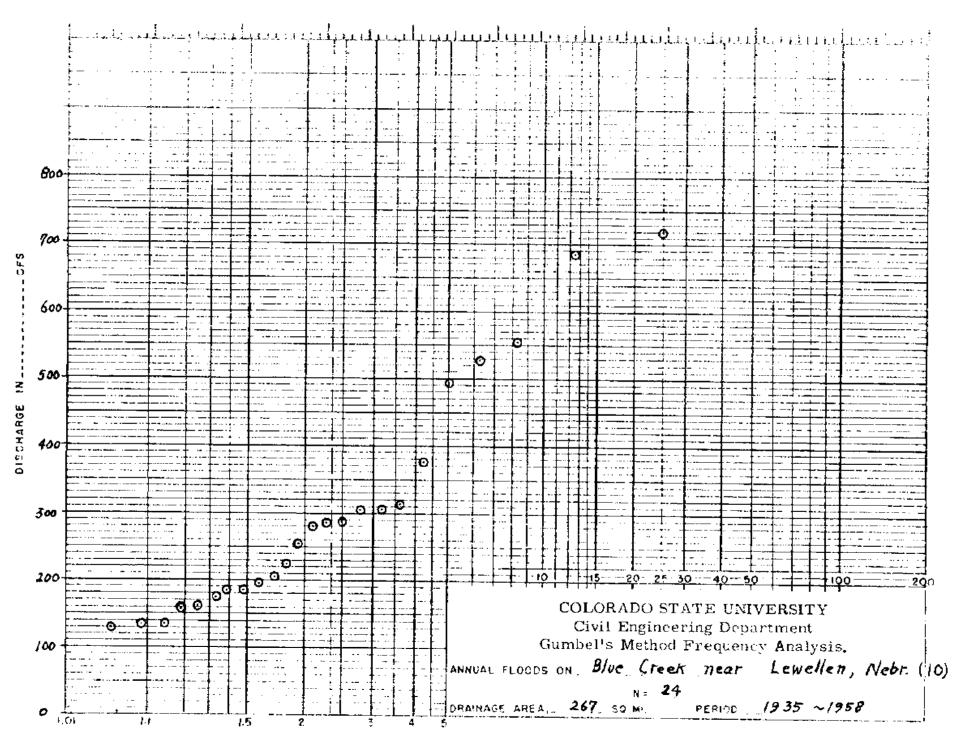


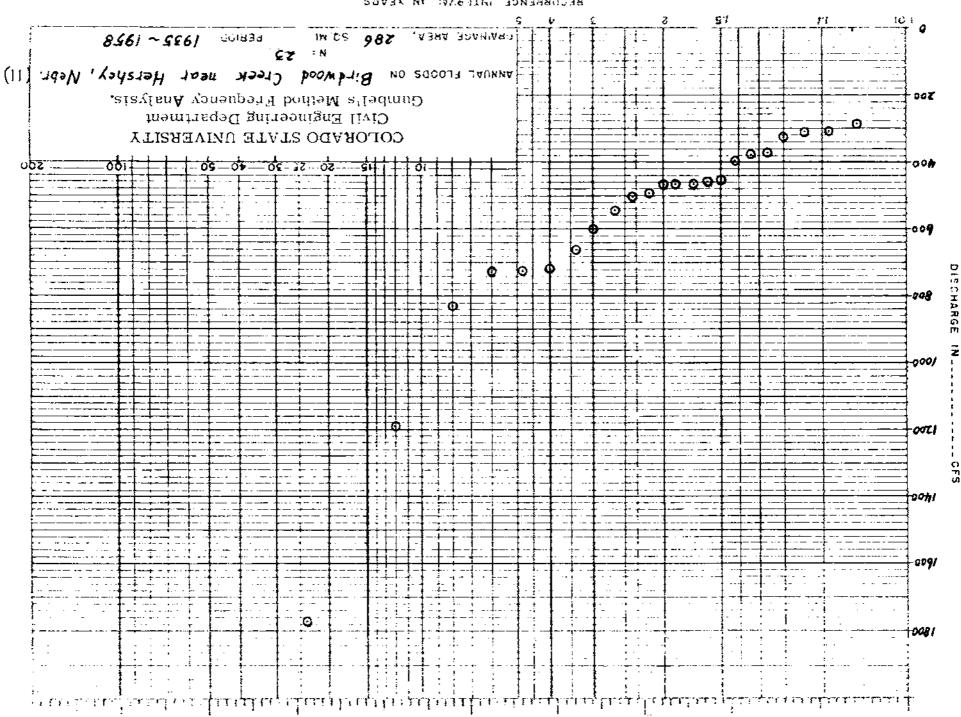


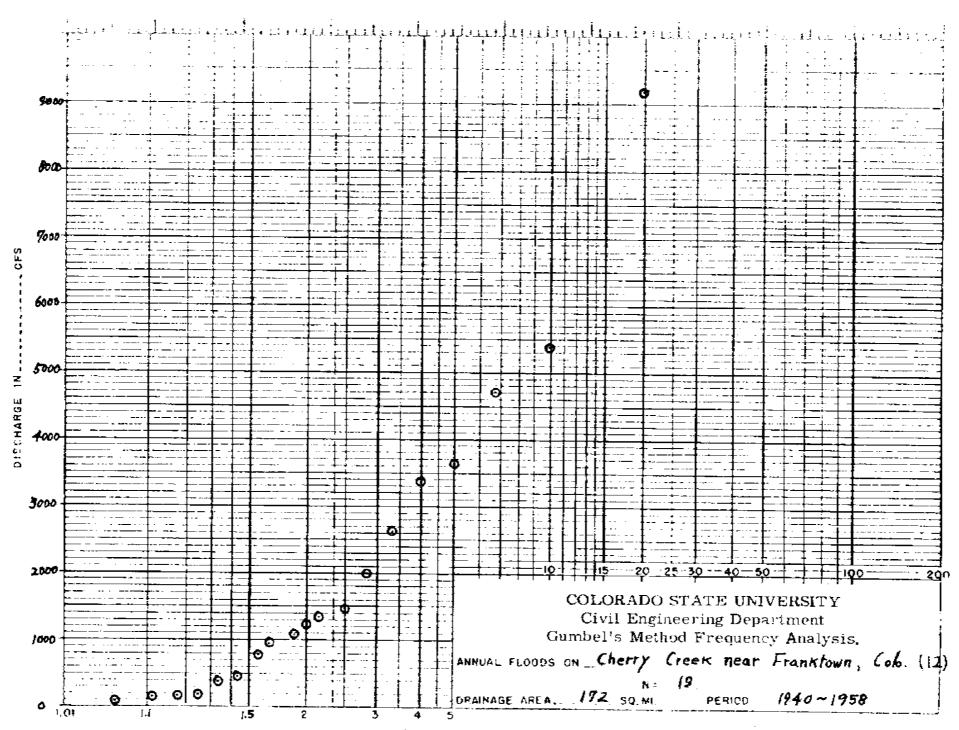


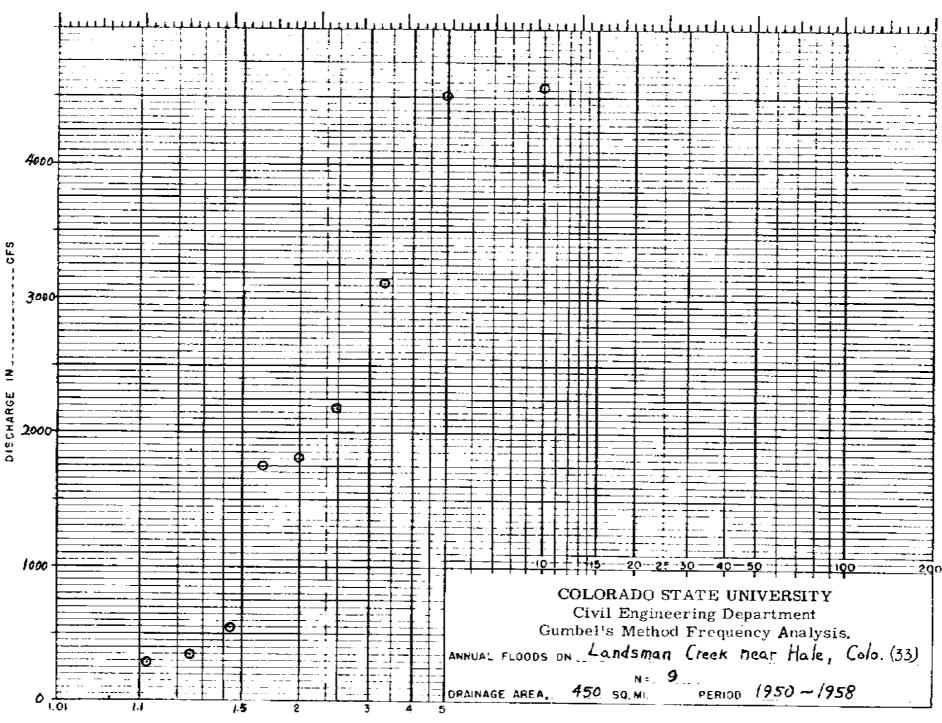


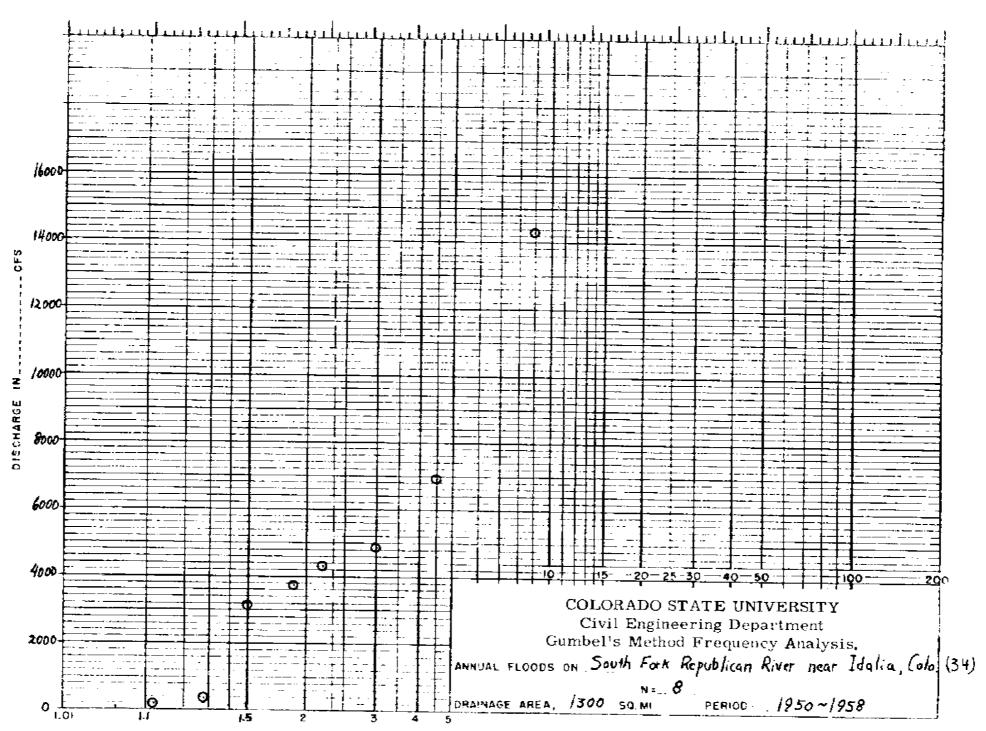


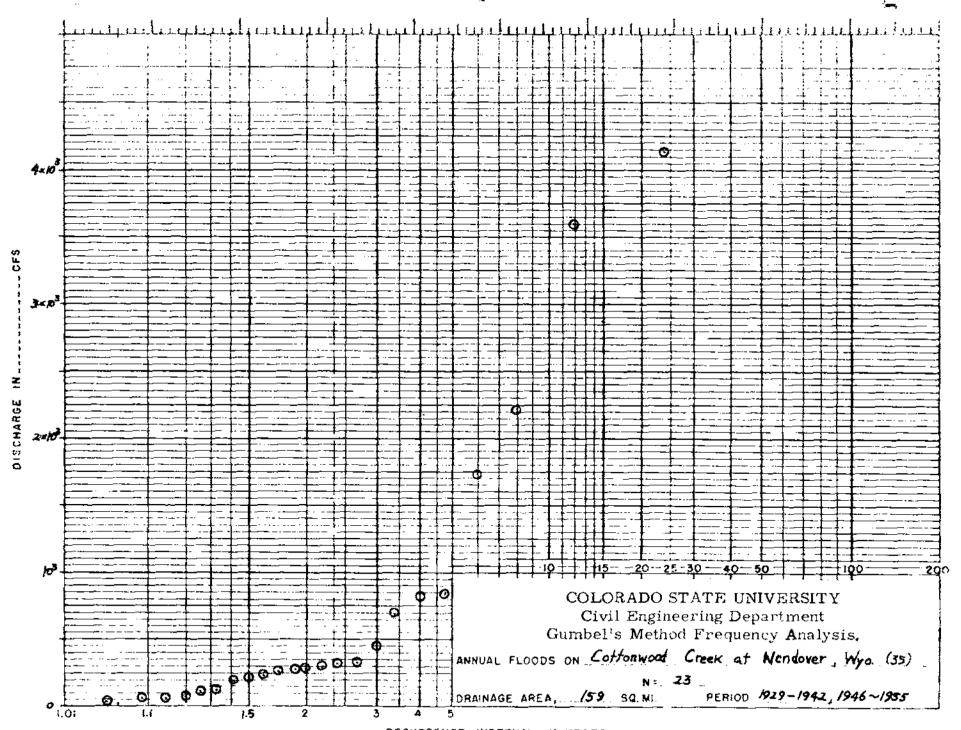


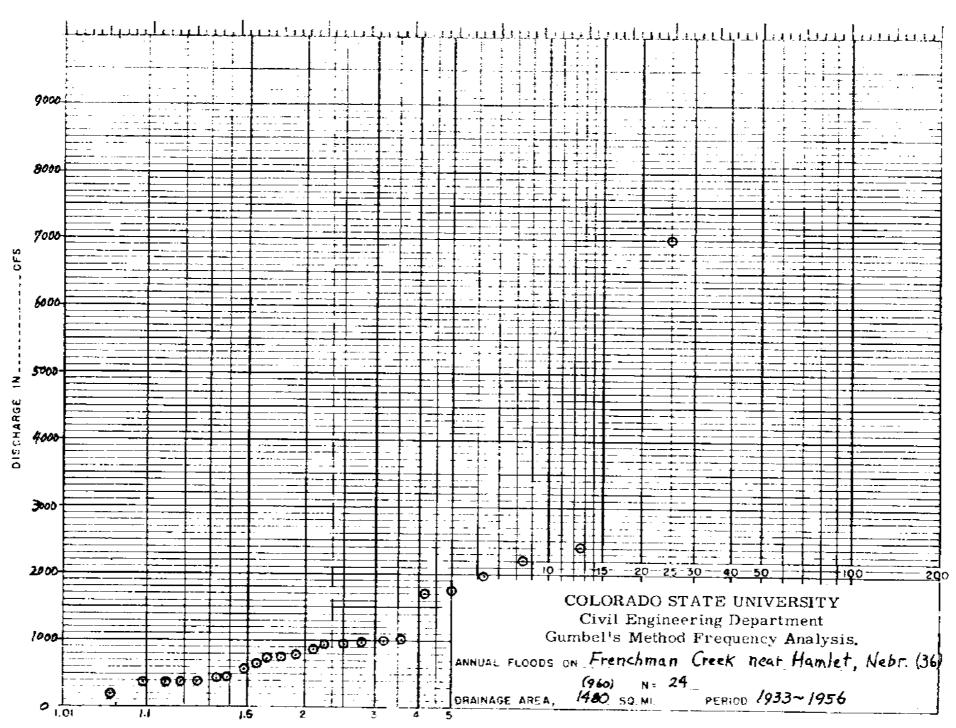


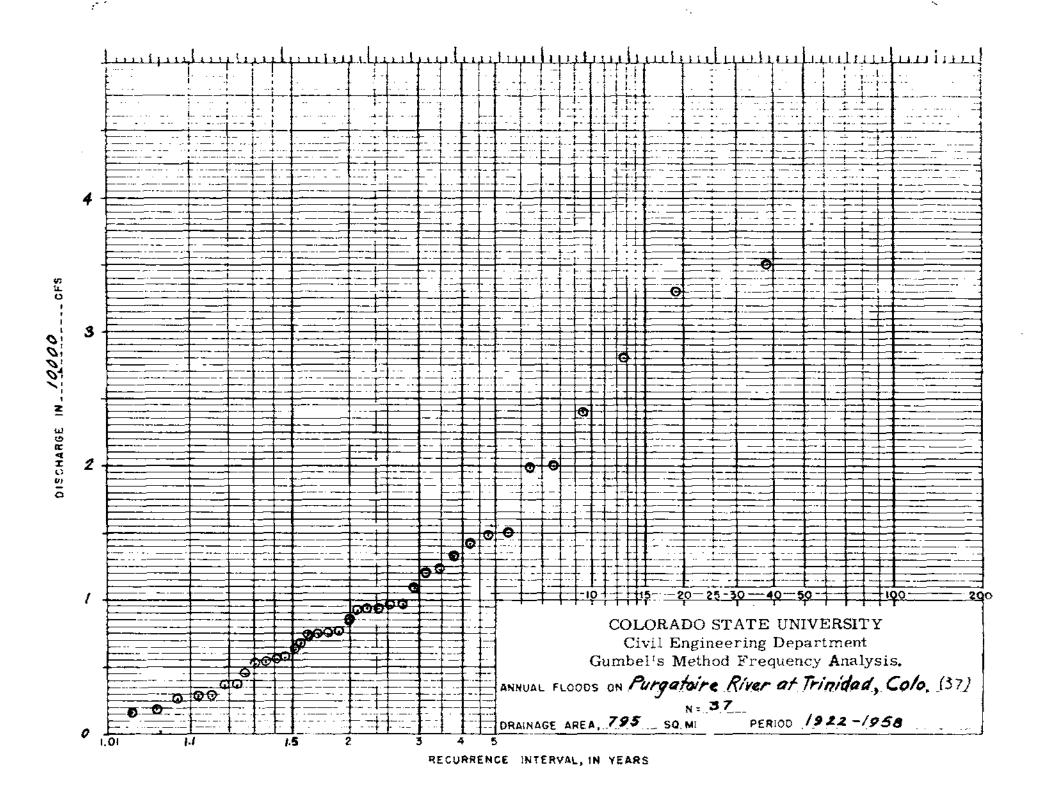


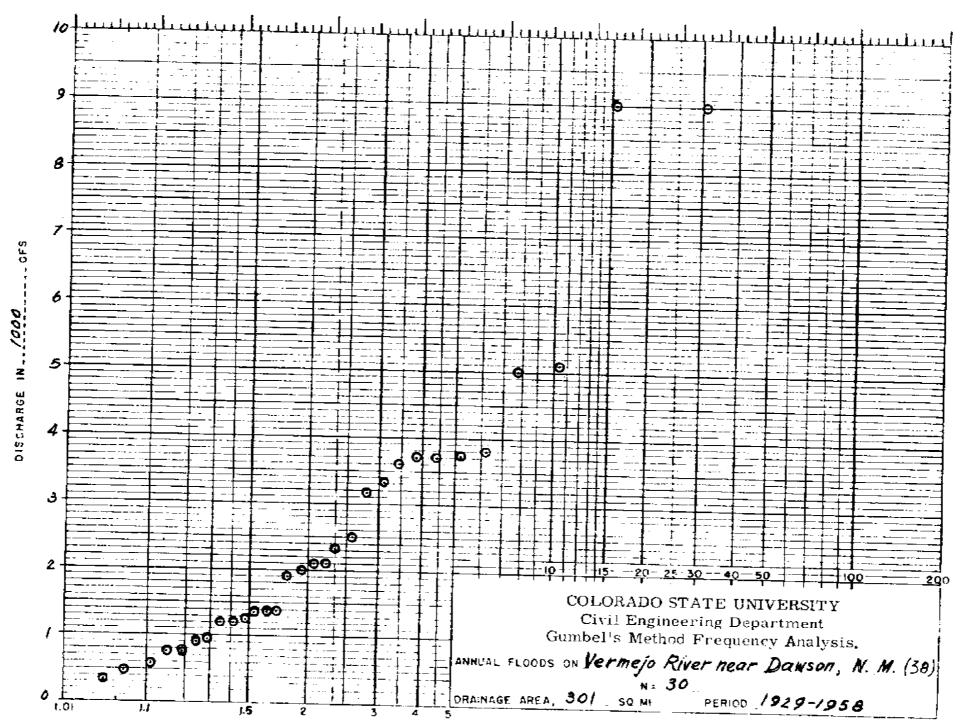


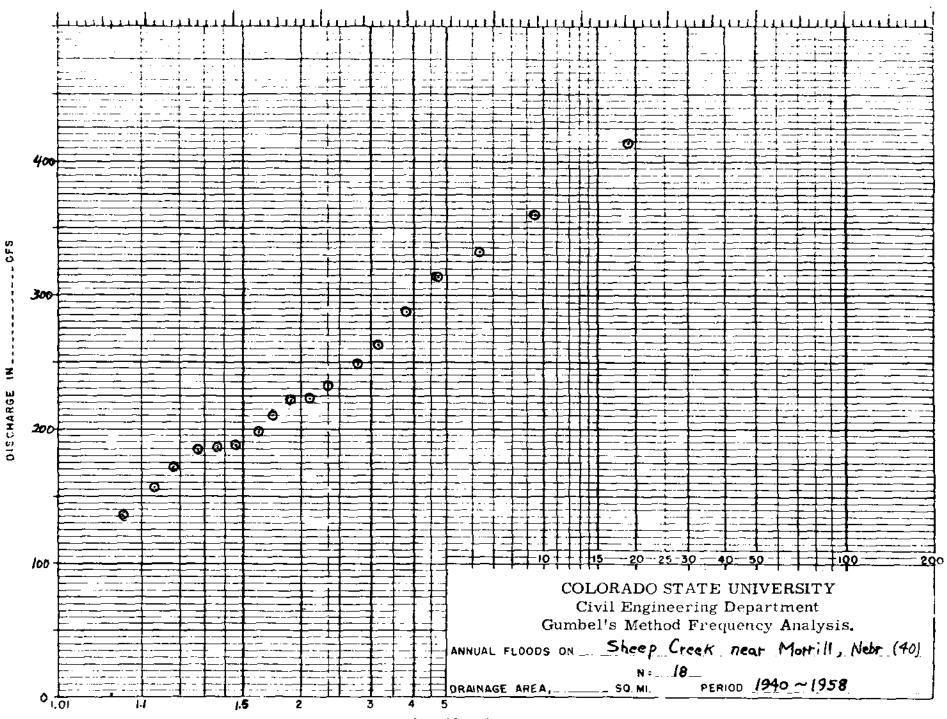


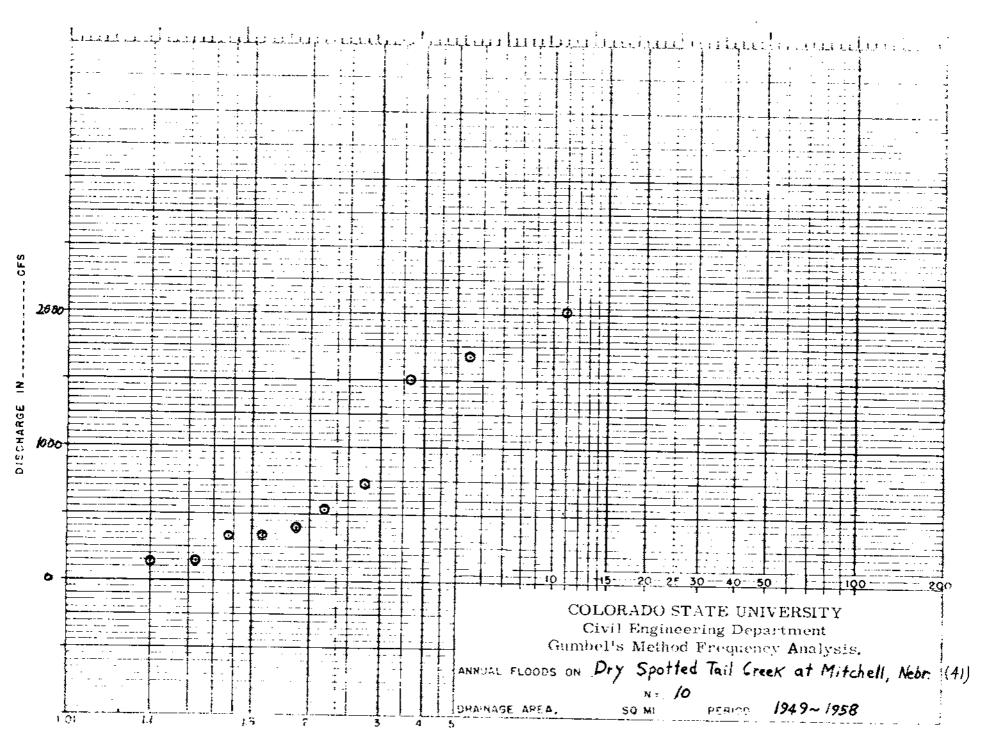


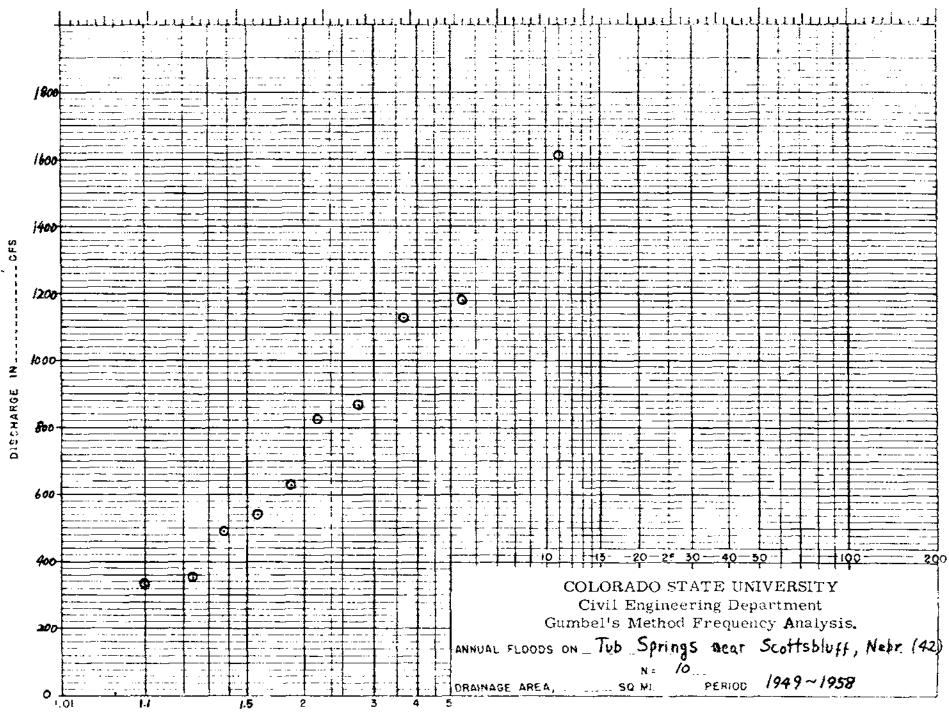


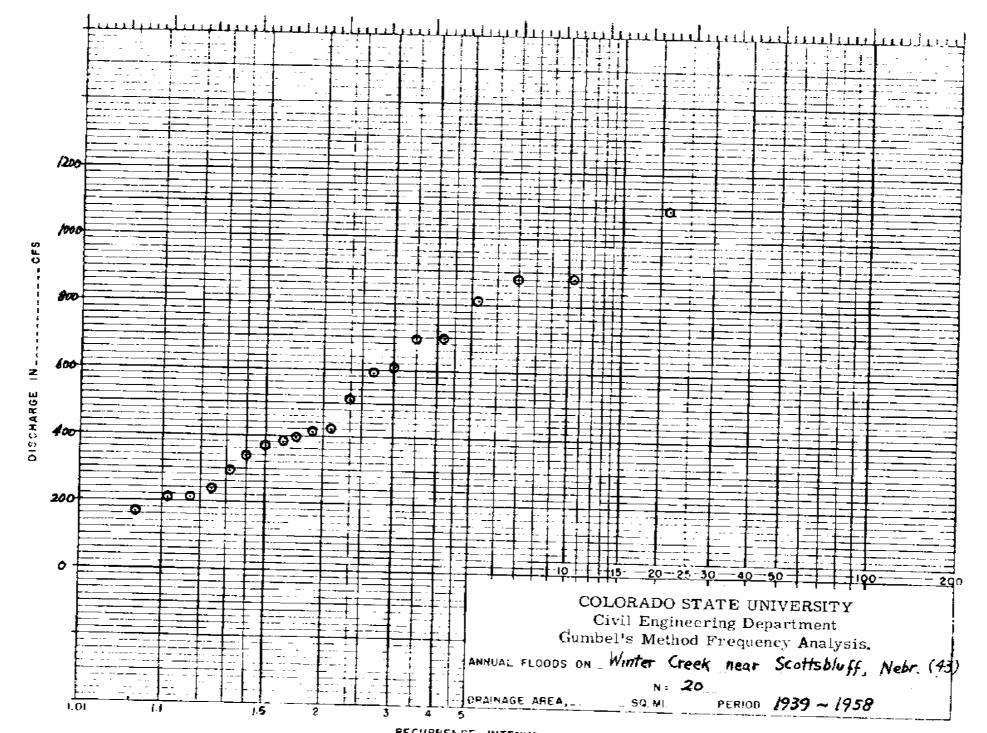


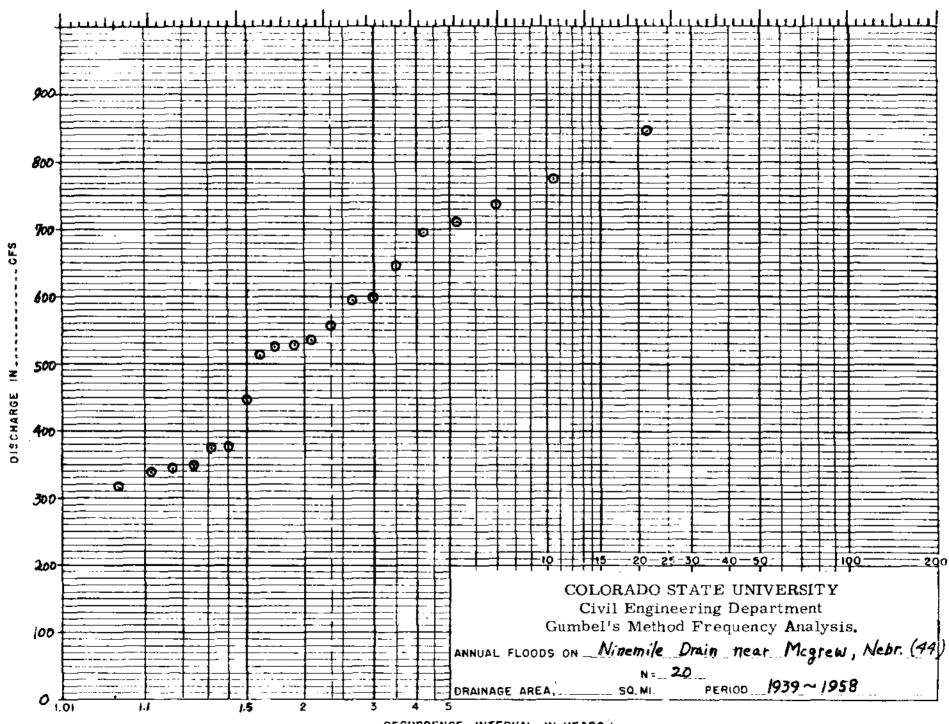


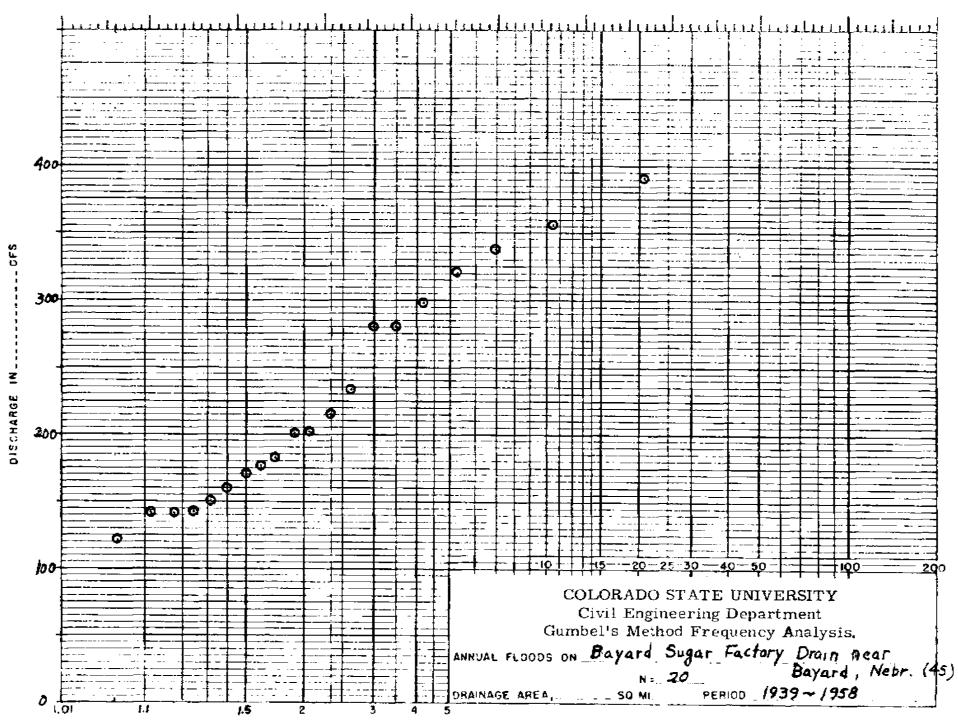


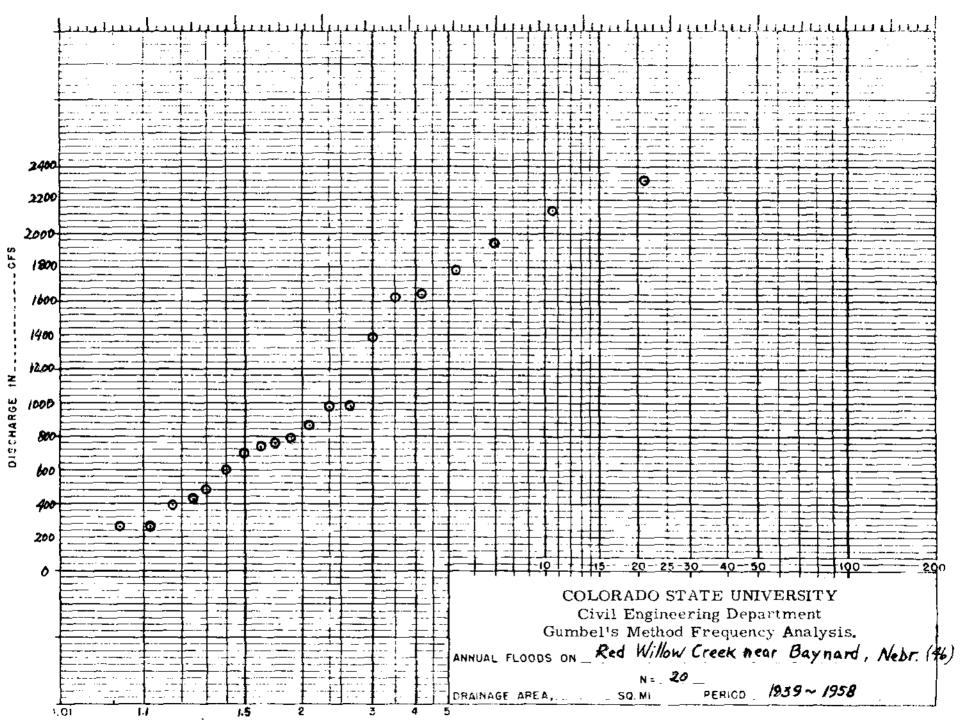


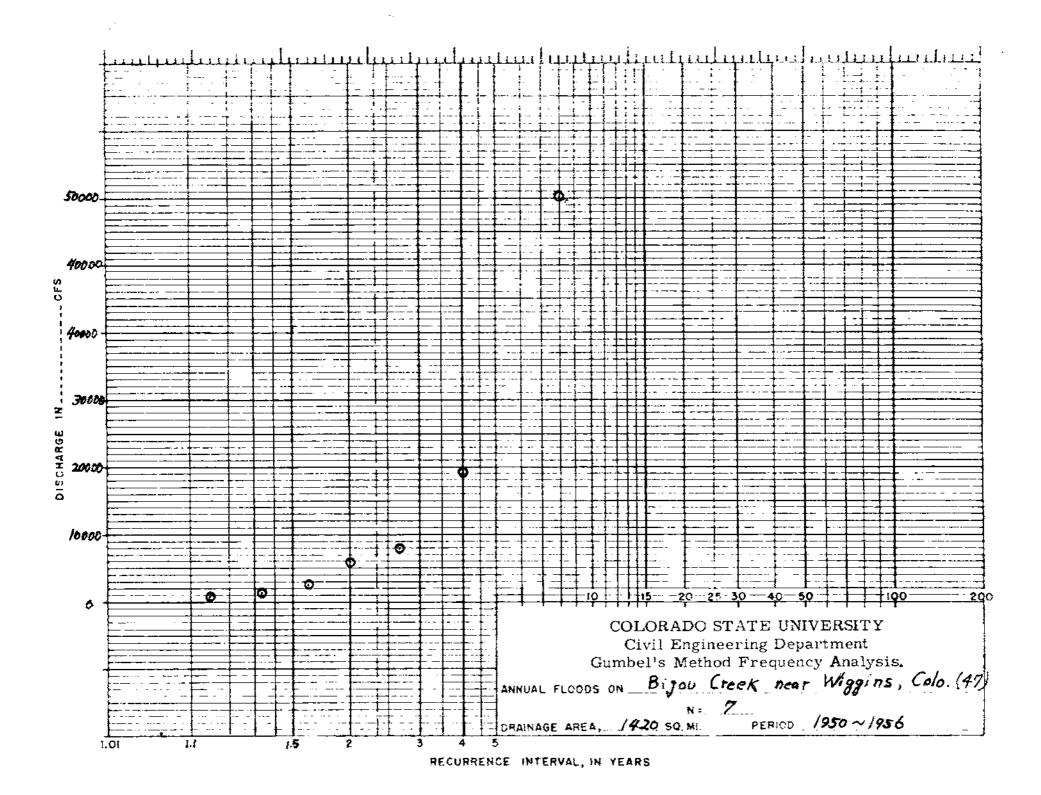


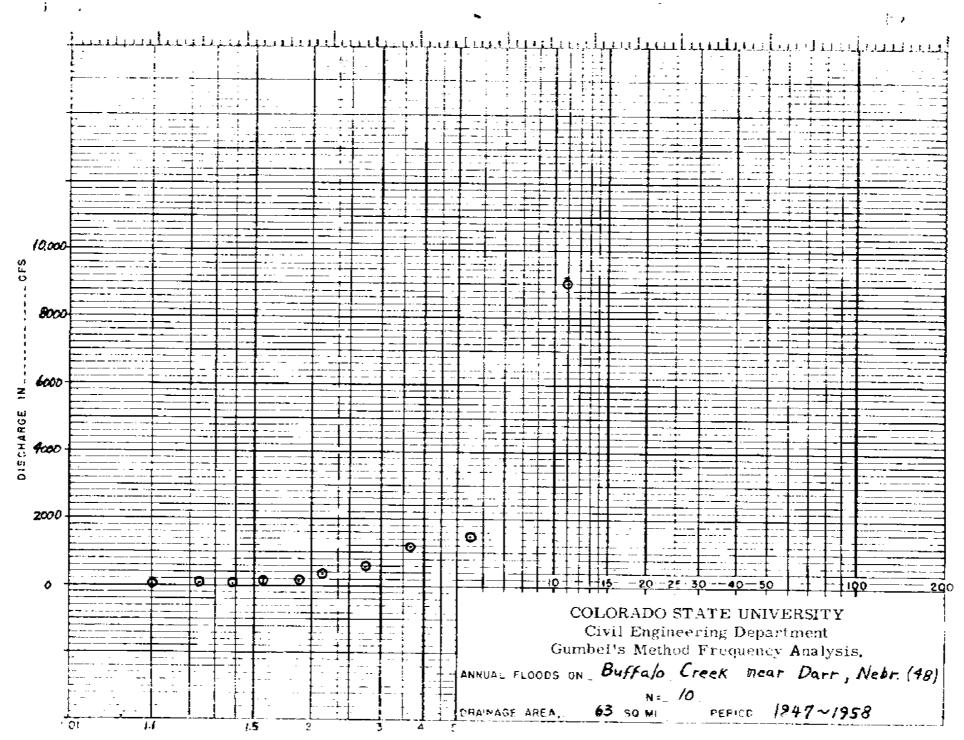


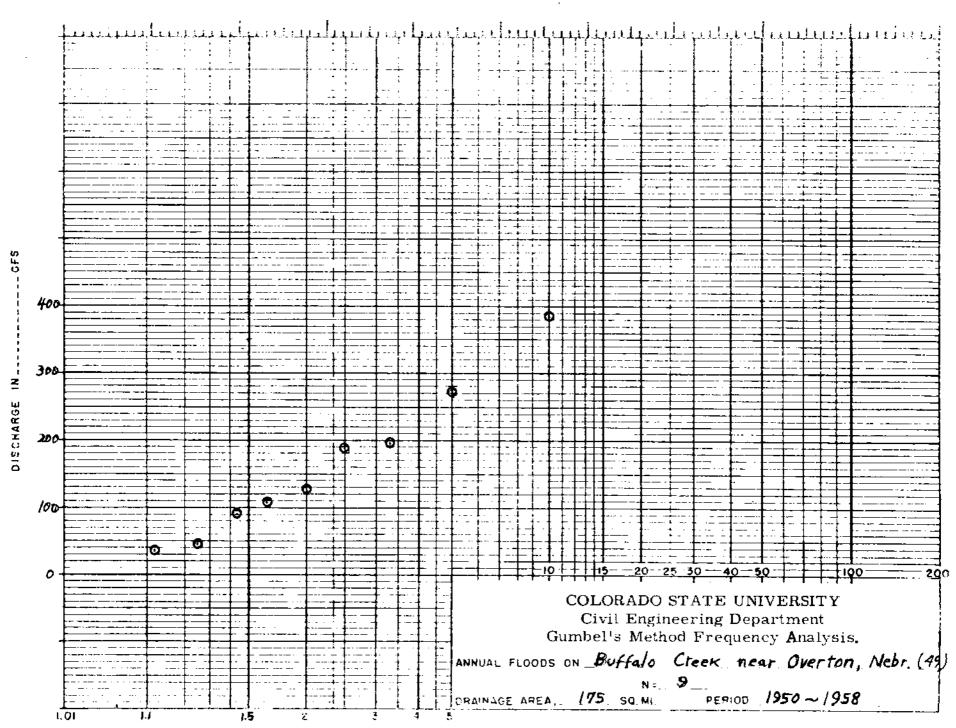


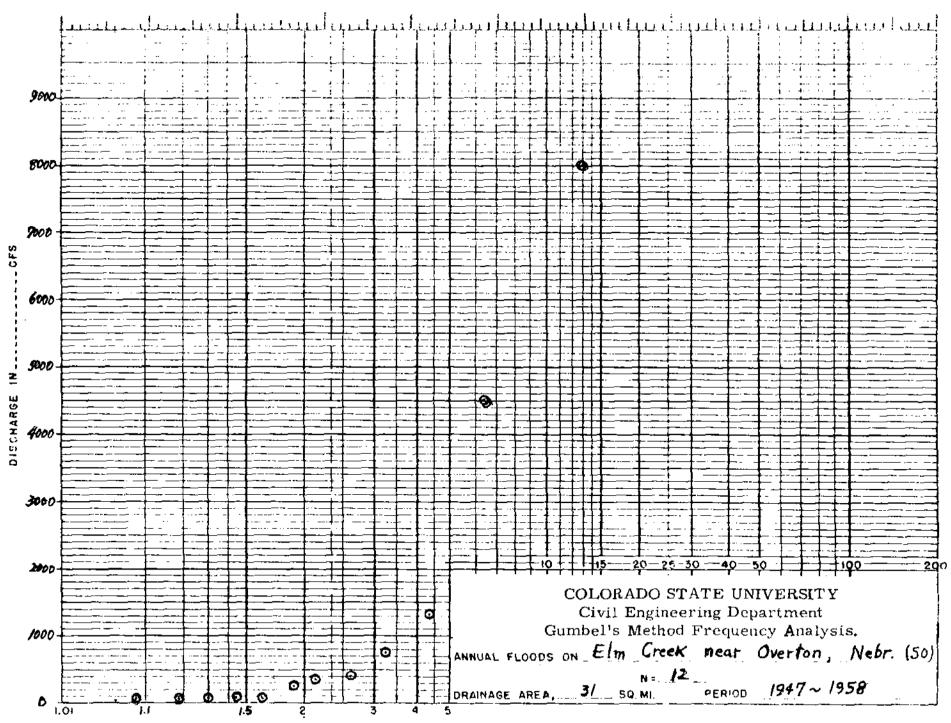




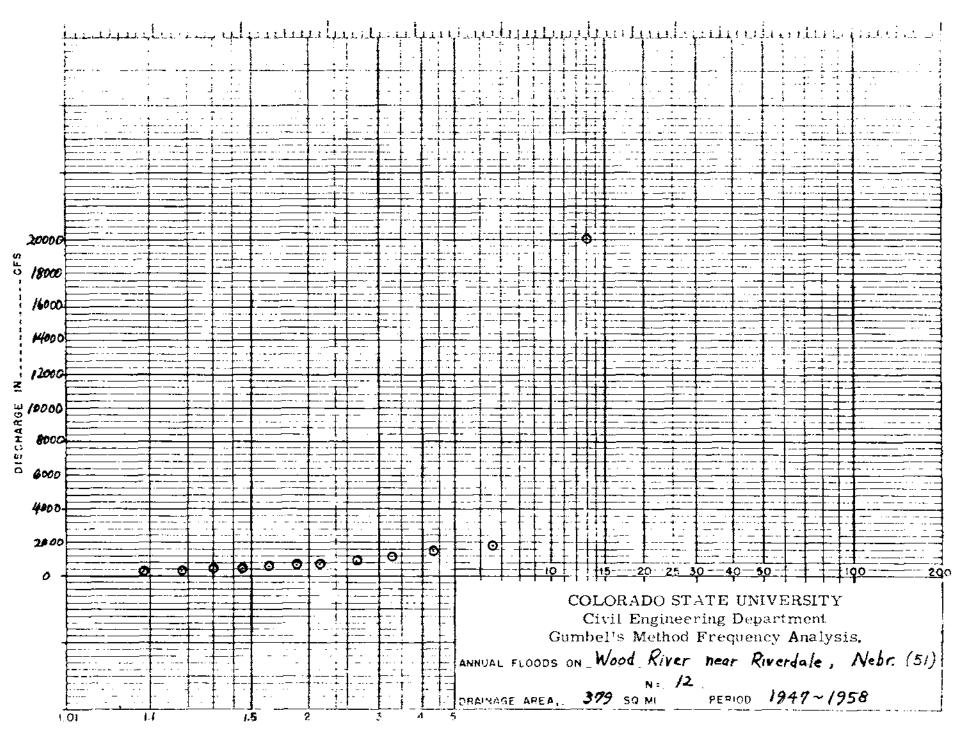


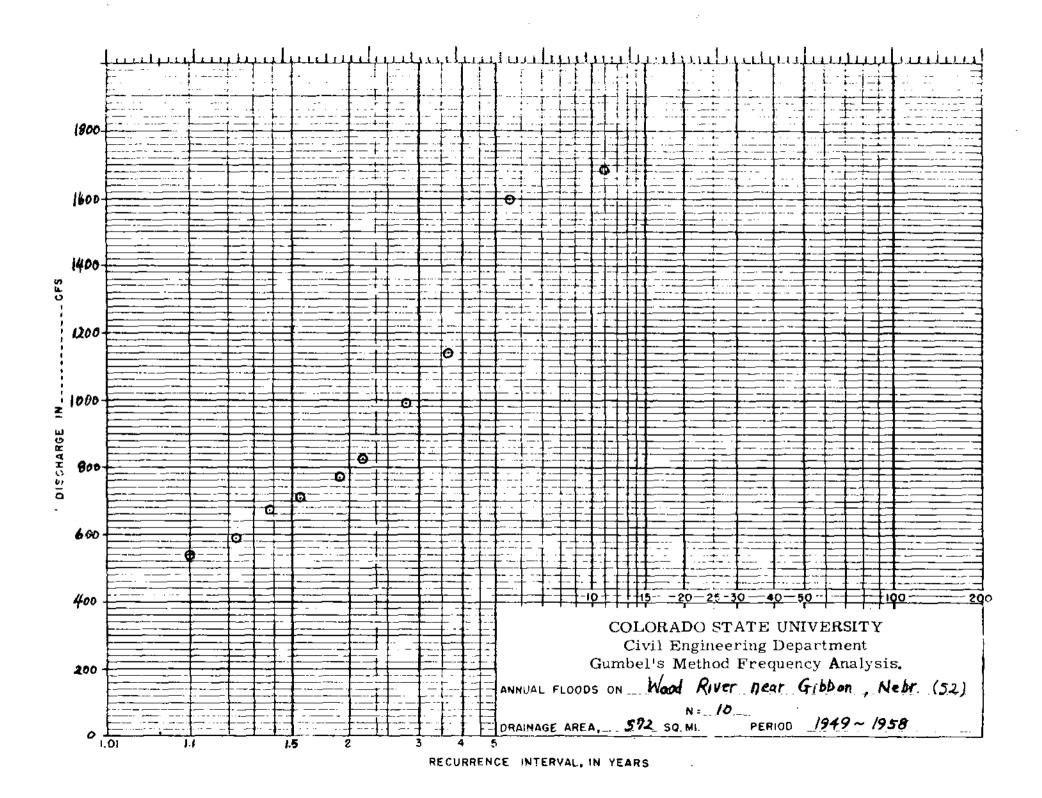


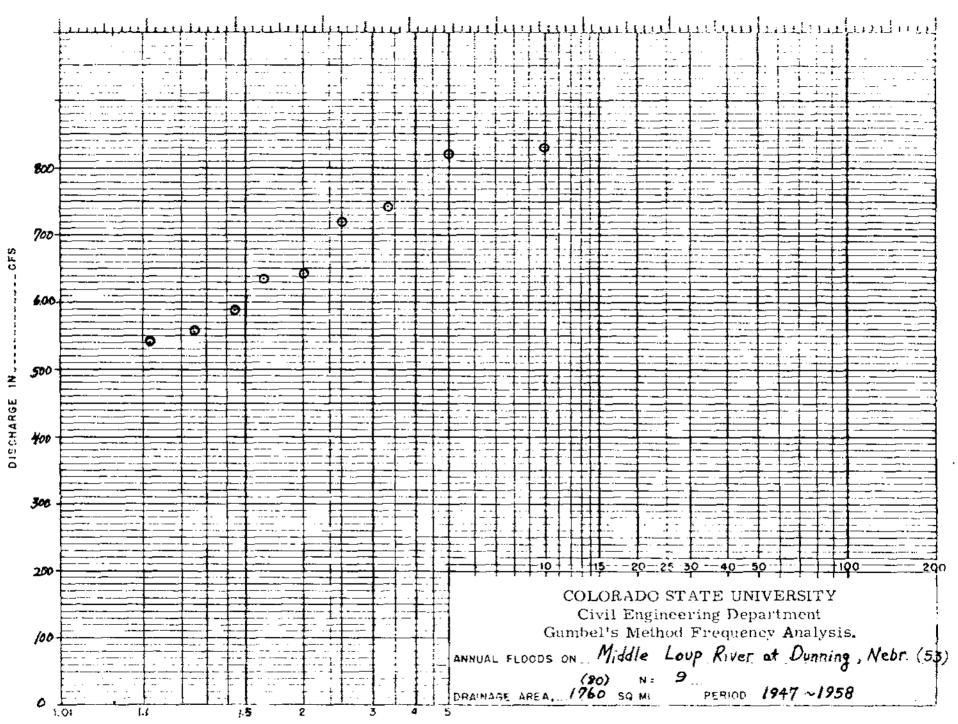


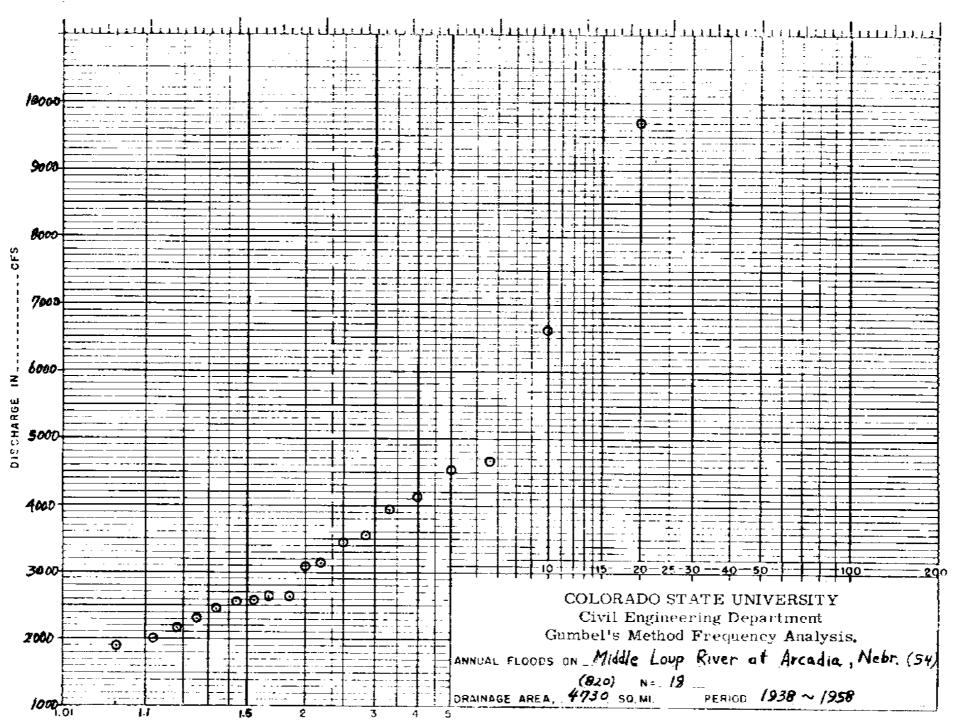


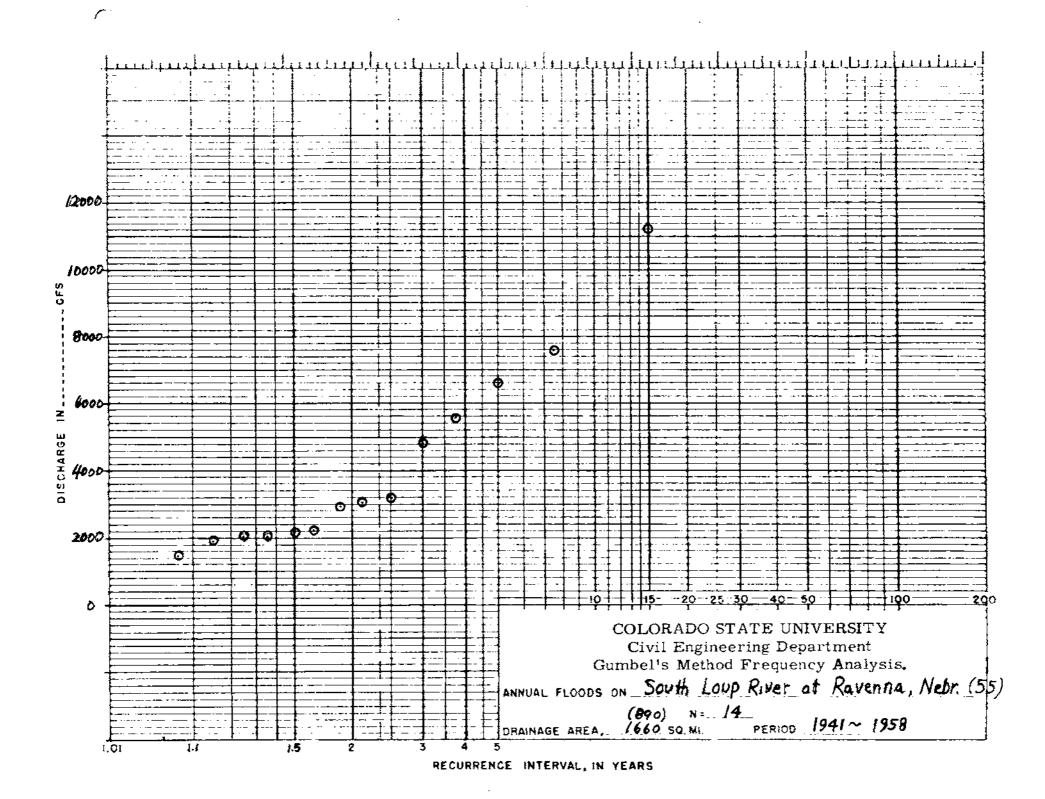
RECURRENCE INTERVAL, IN YEARS

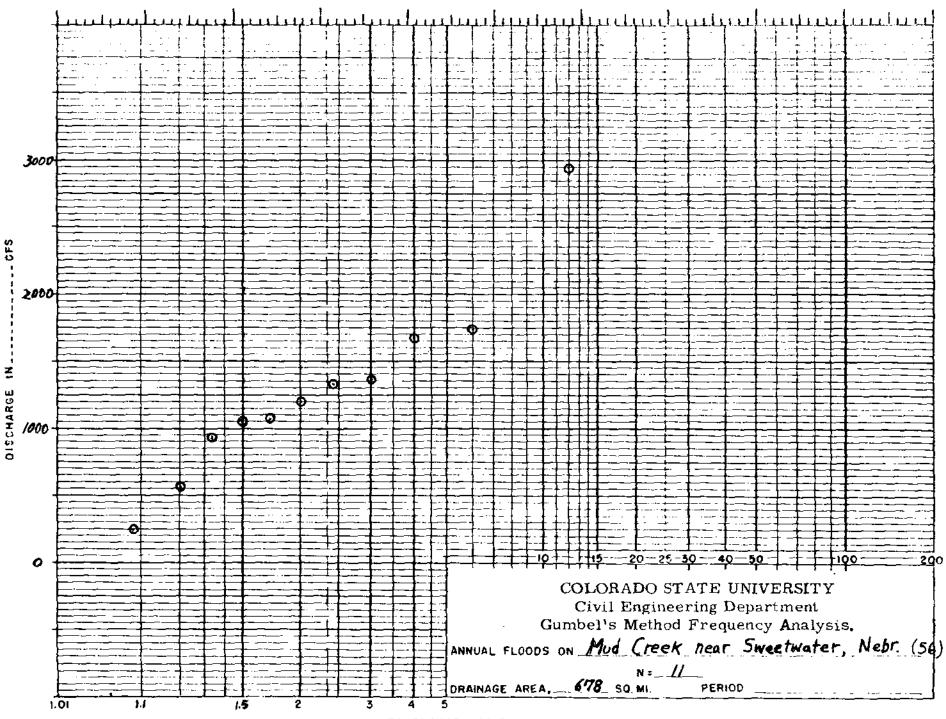


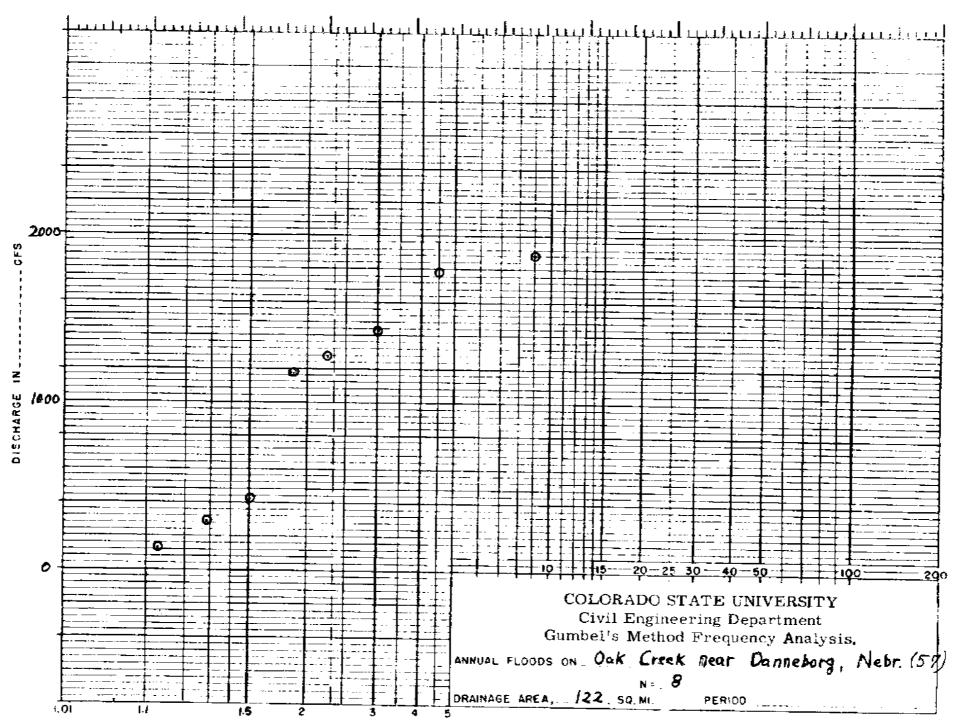


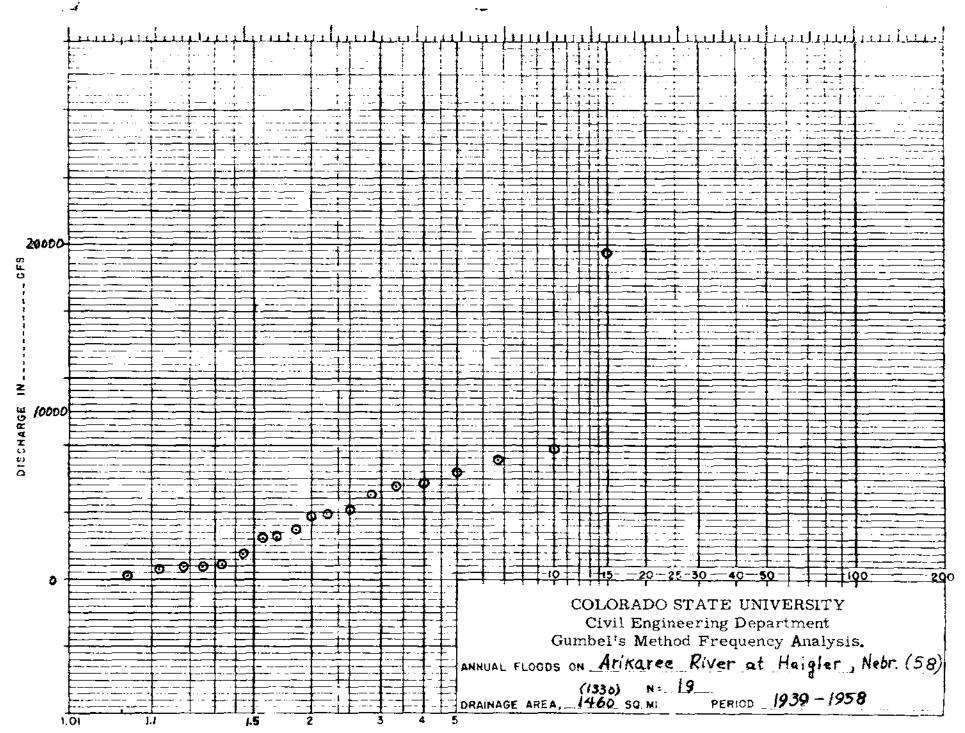


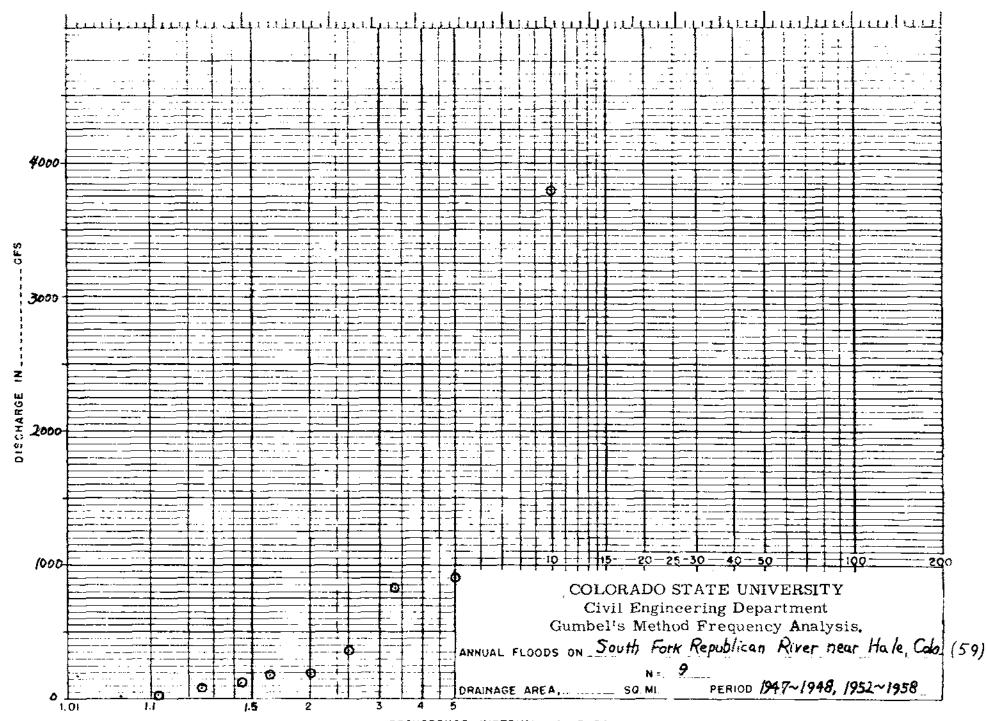


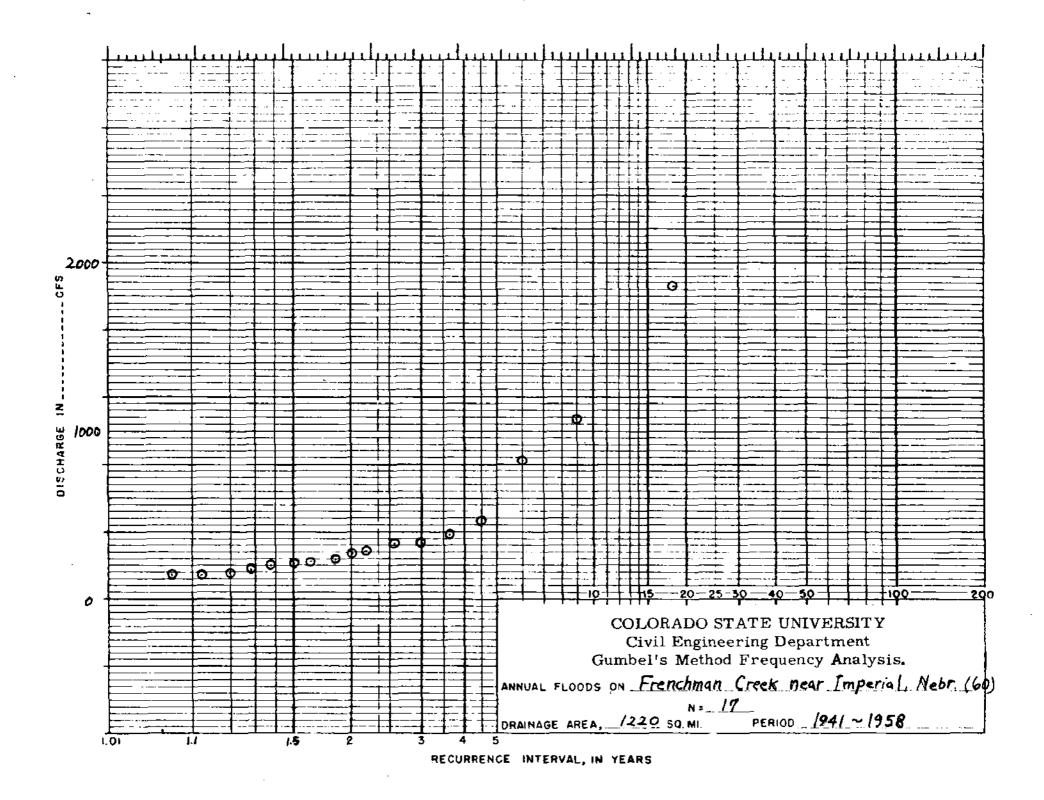


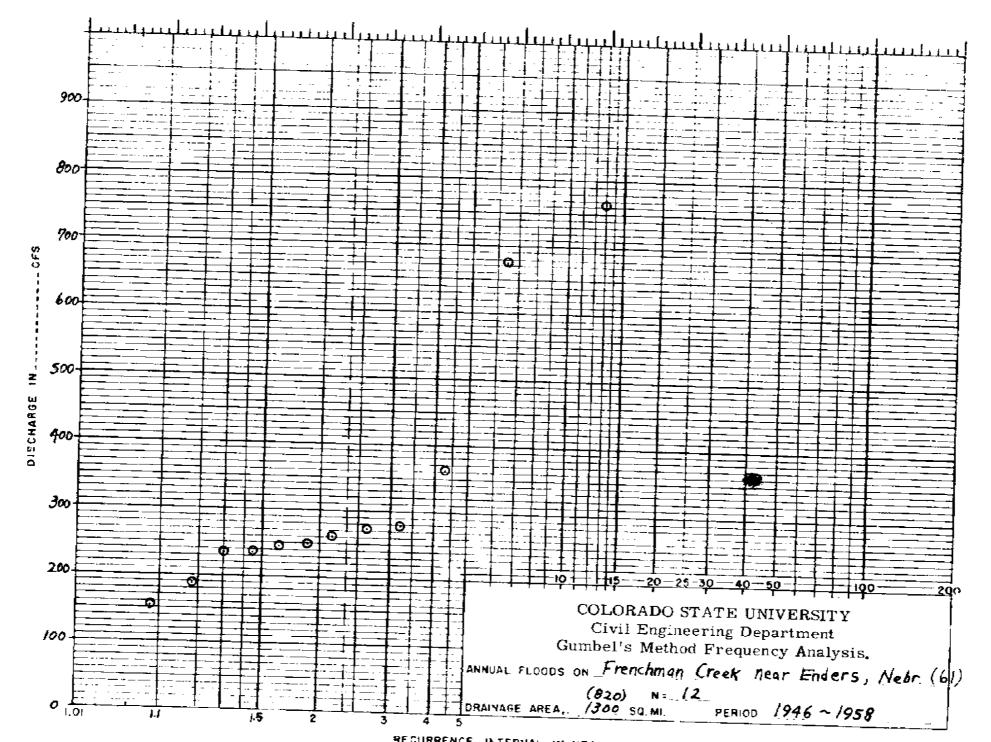


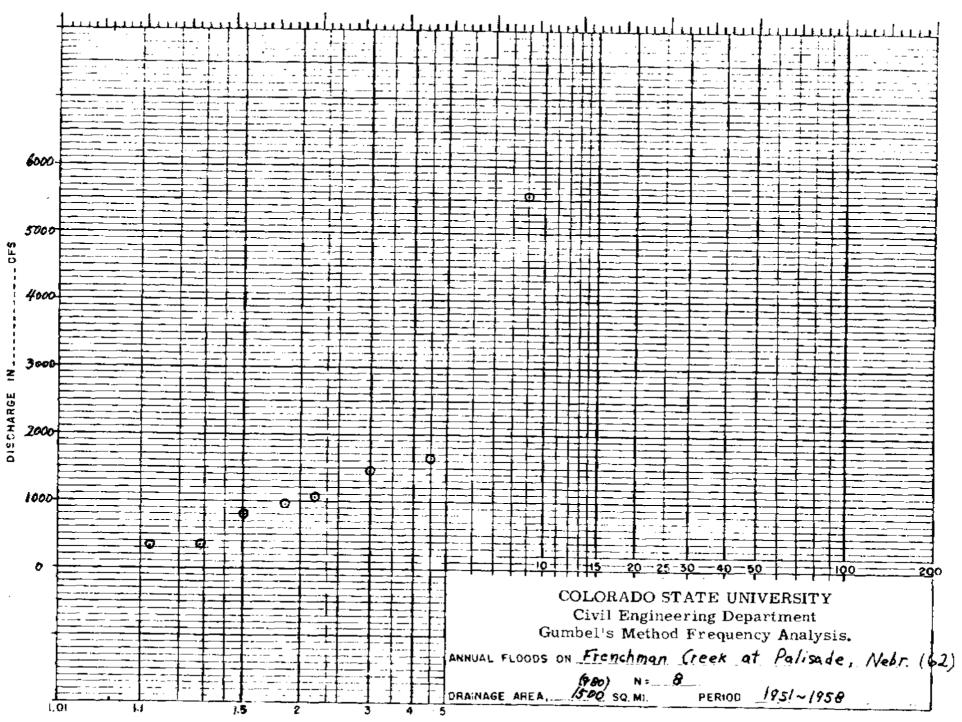


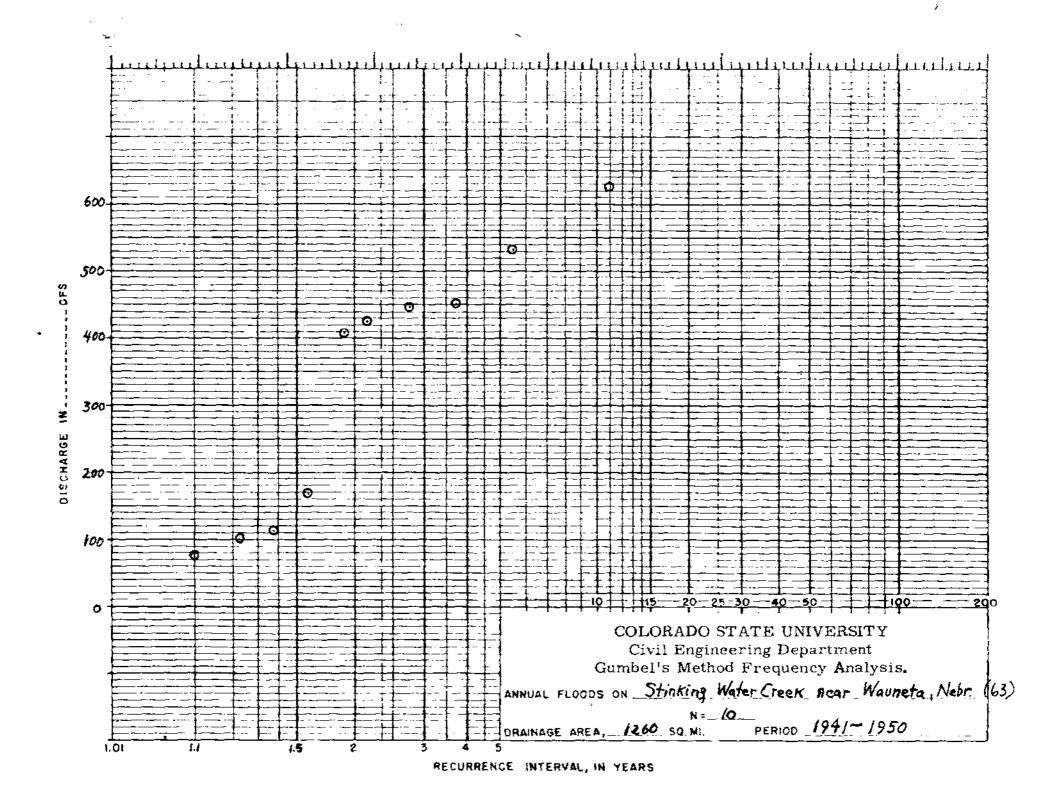


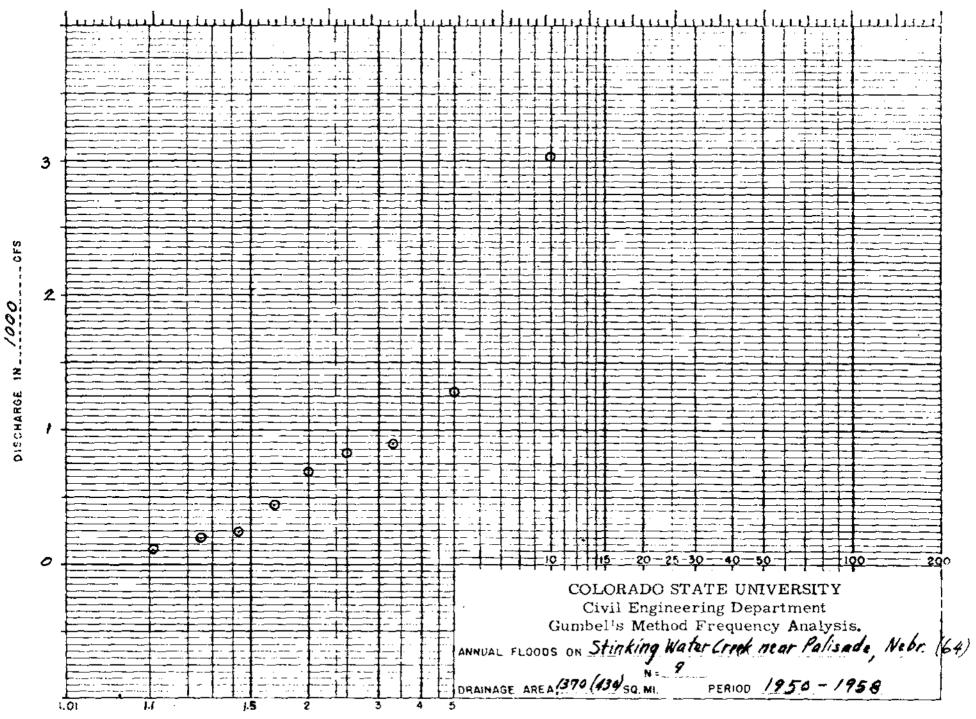


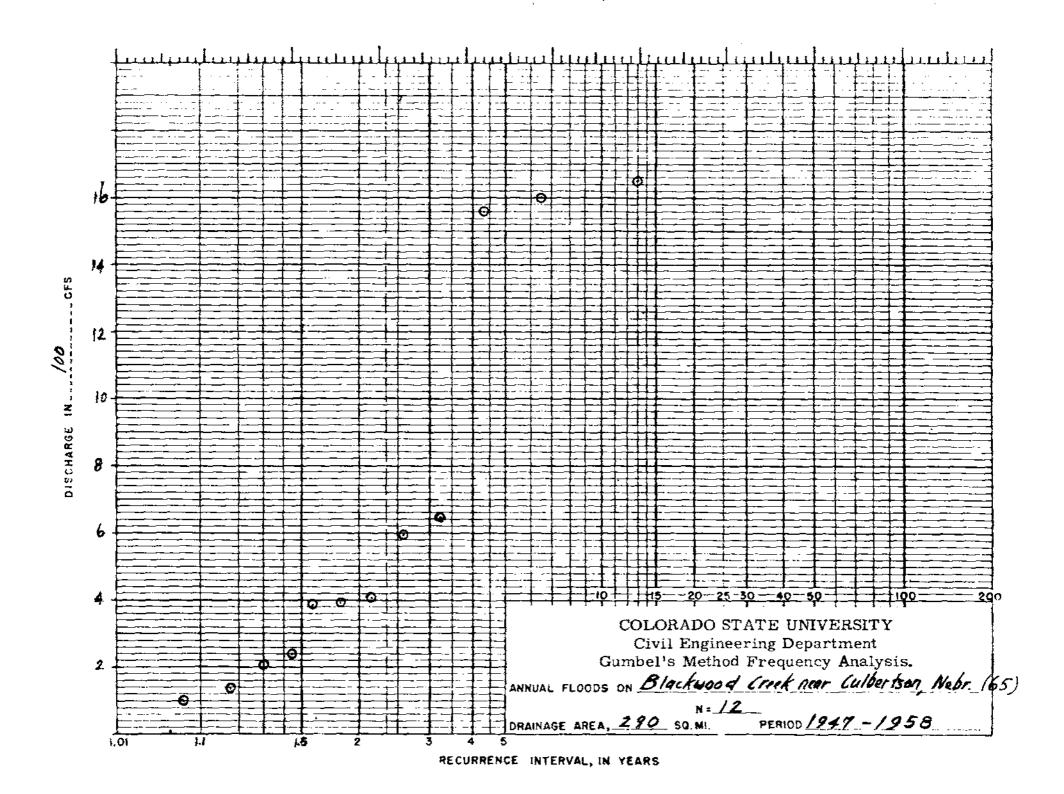


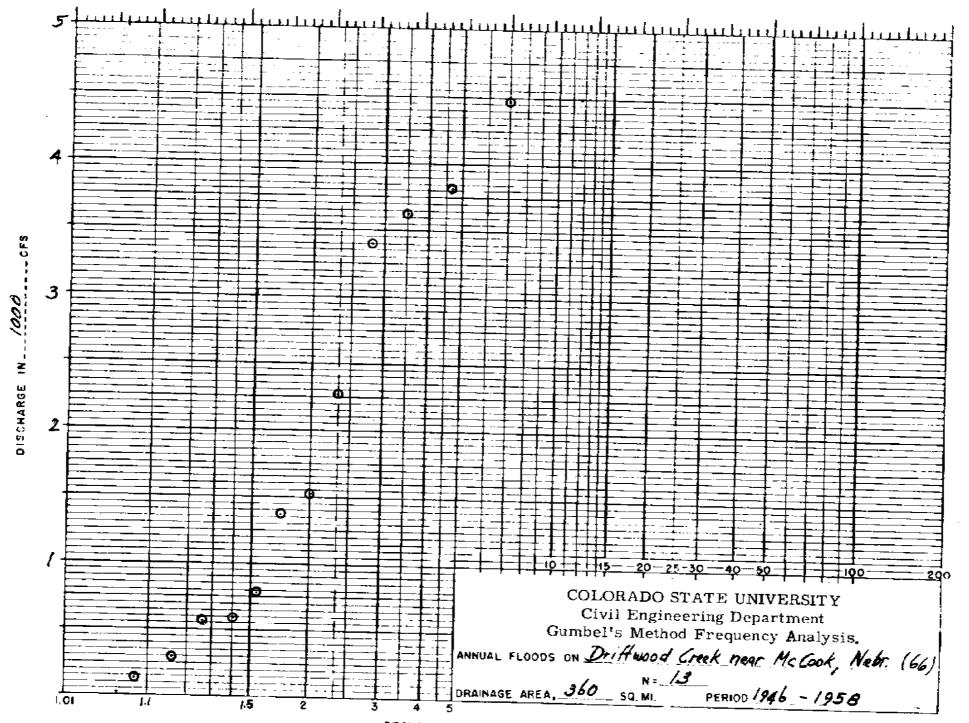


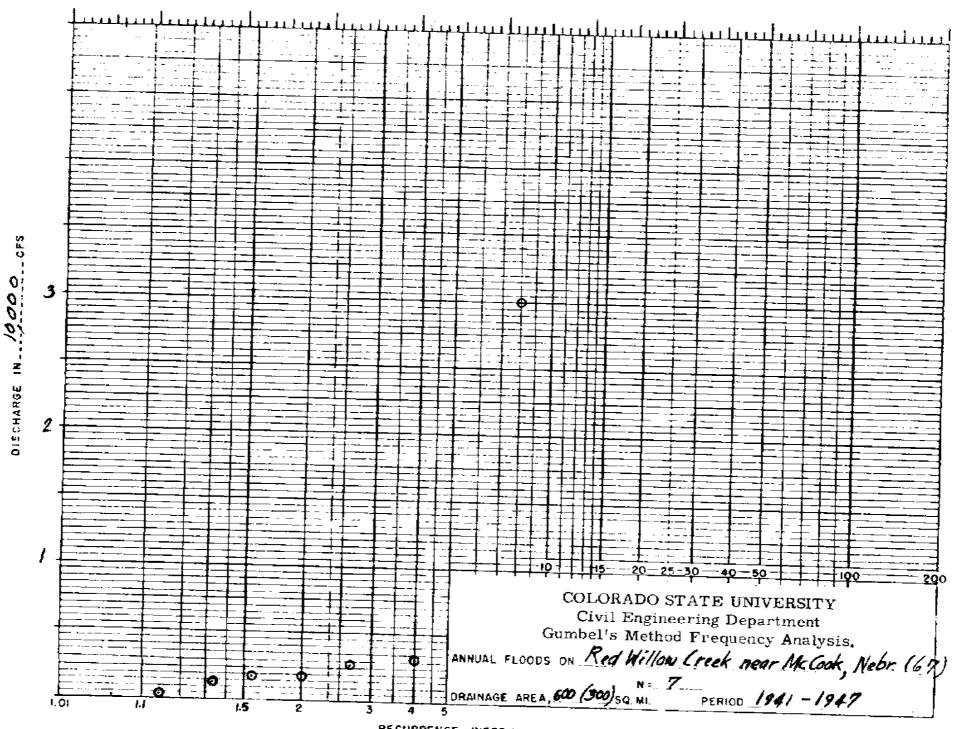


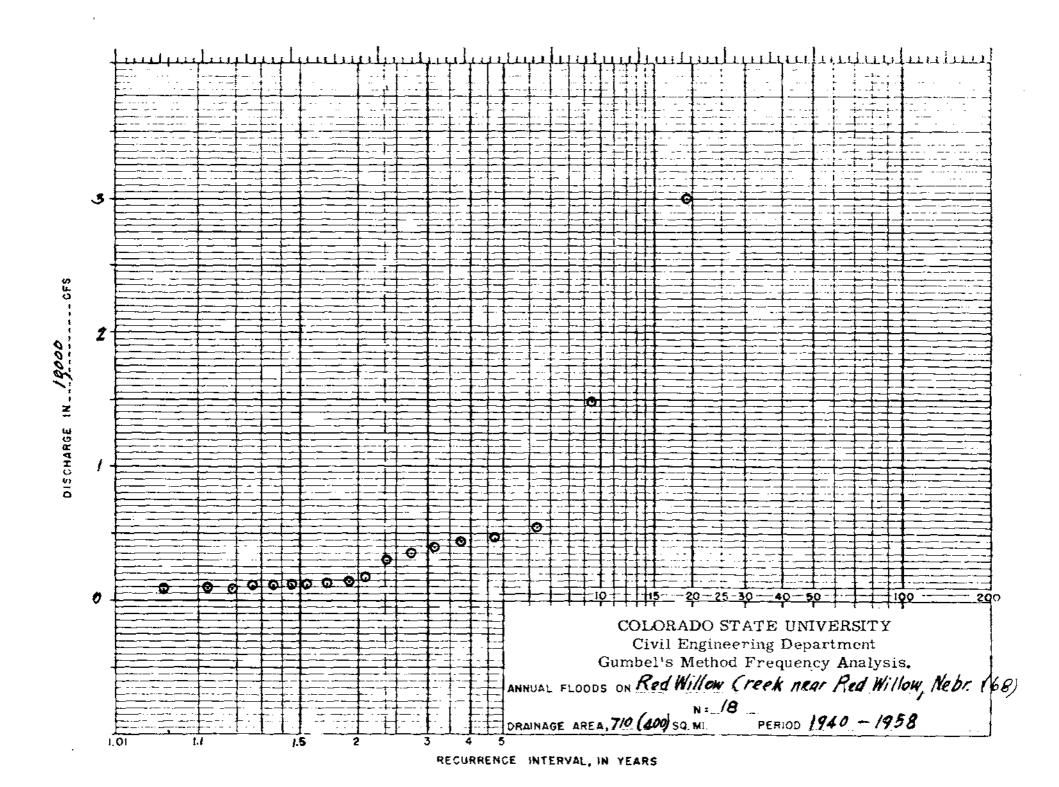


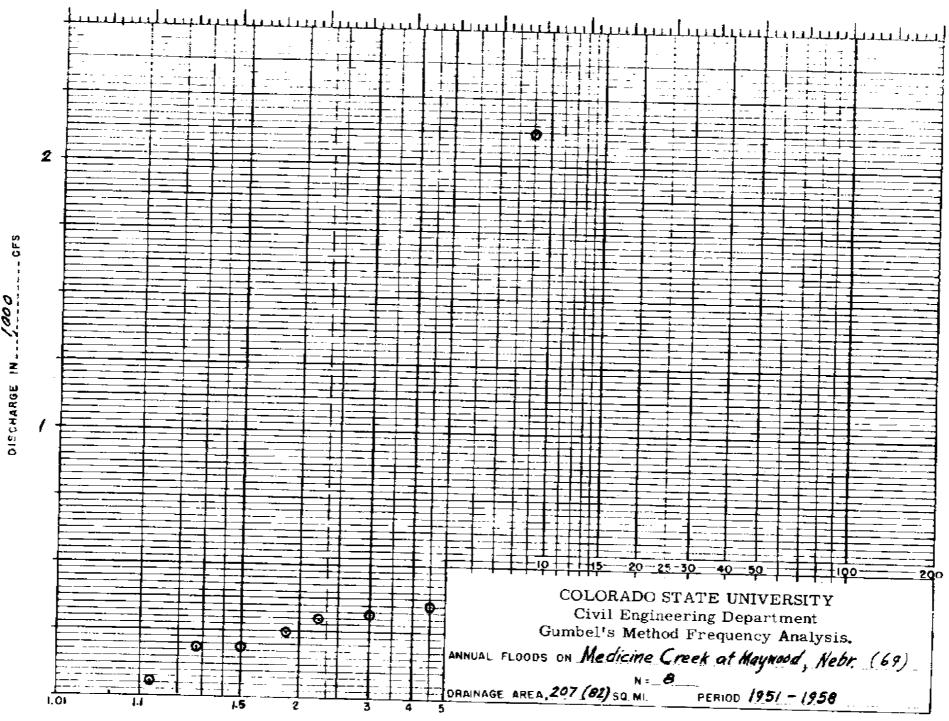


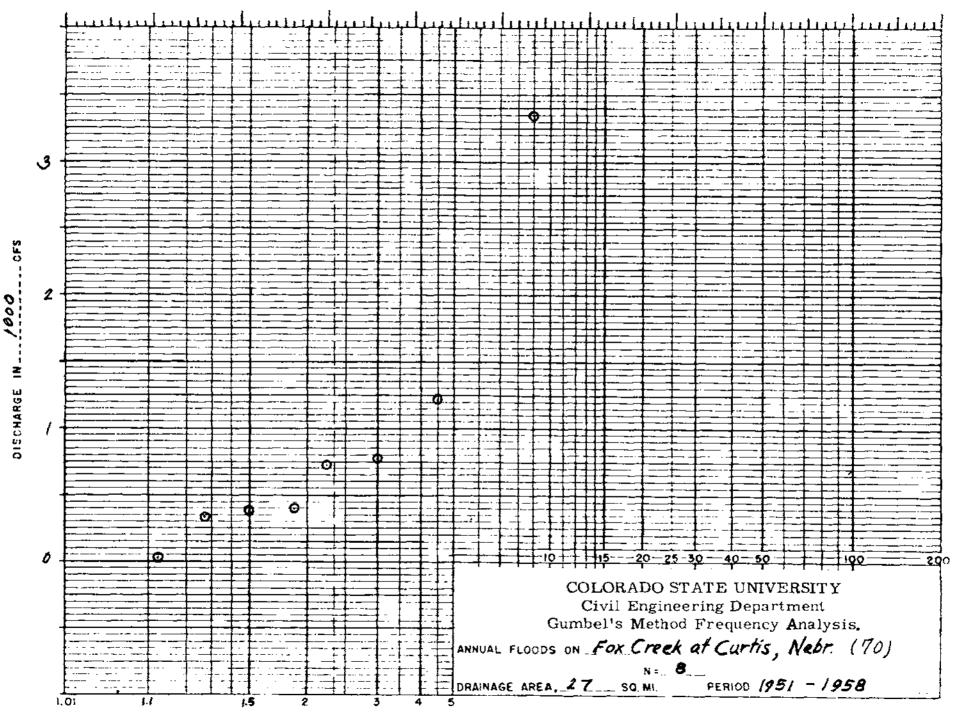


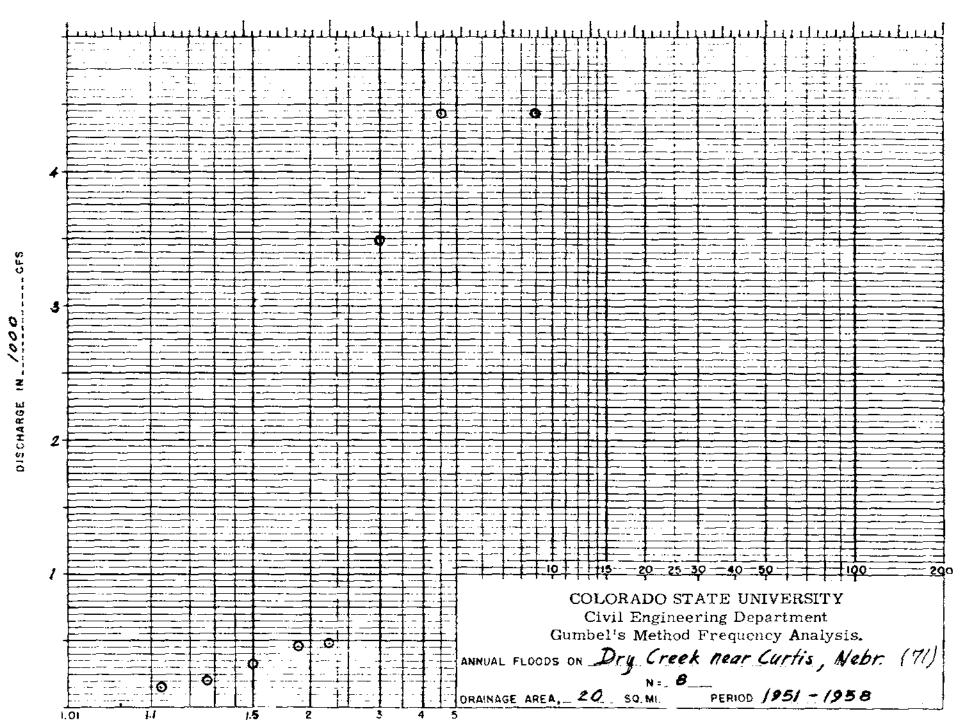


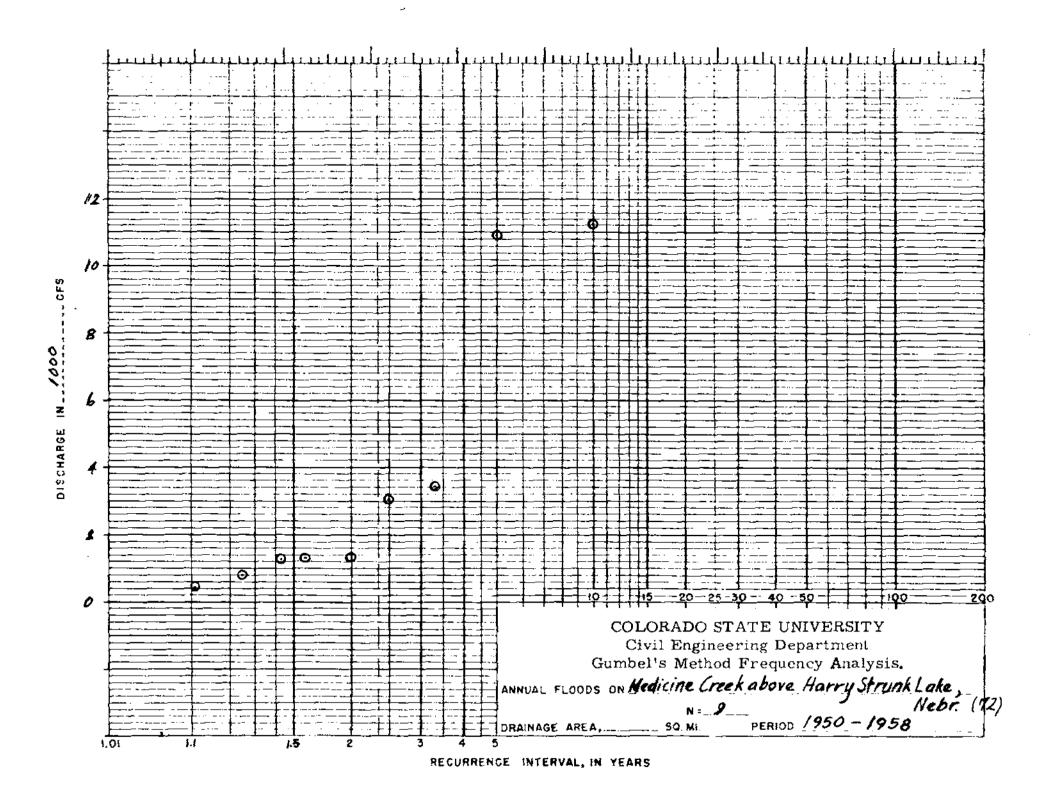


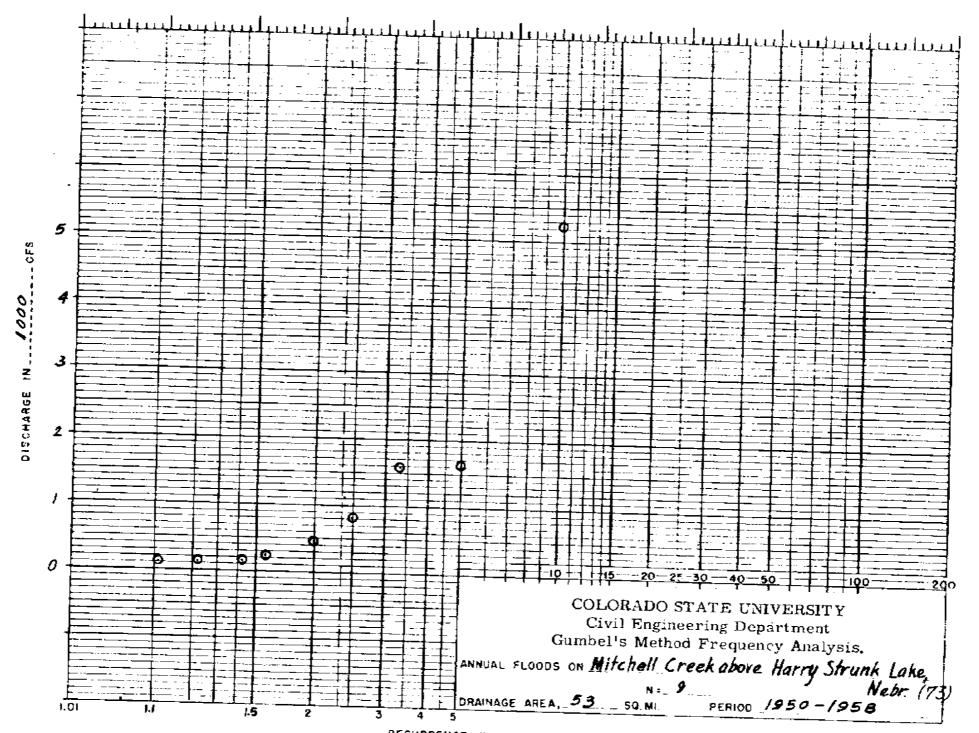


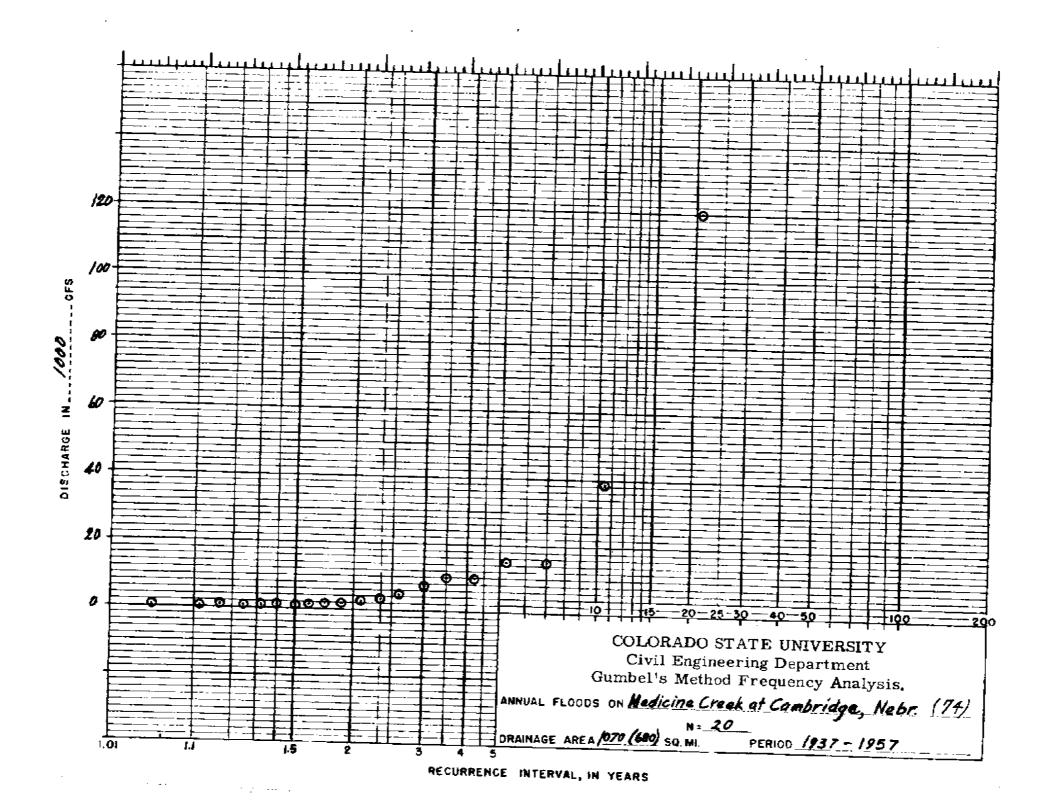


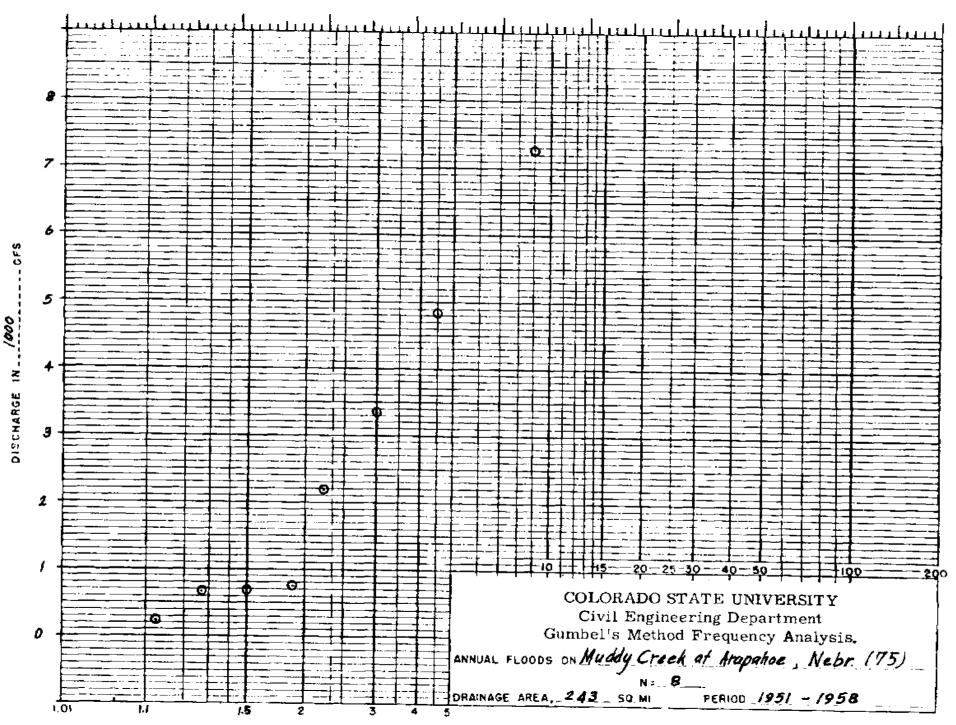


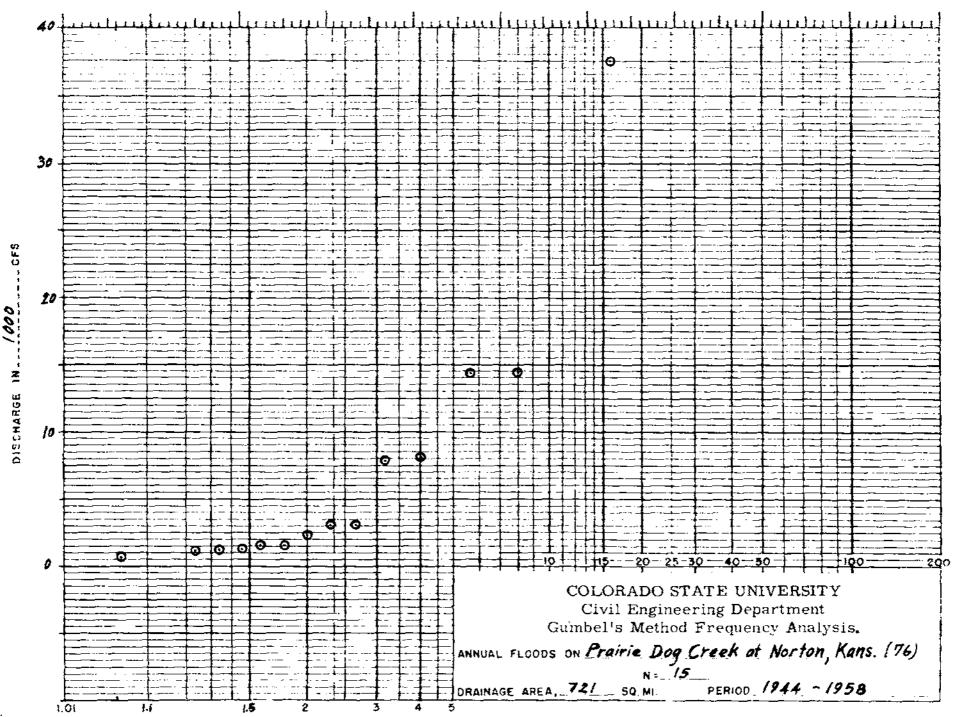


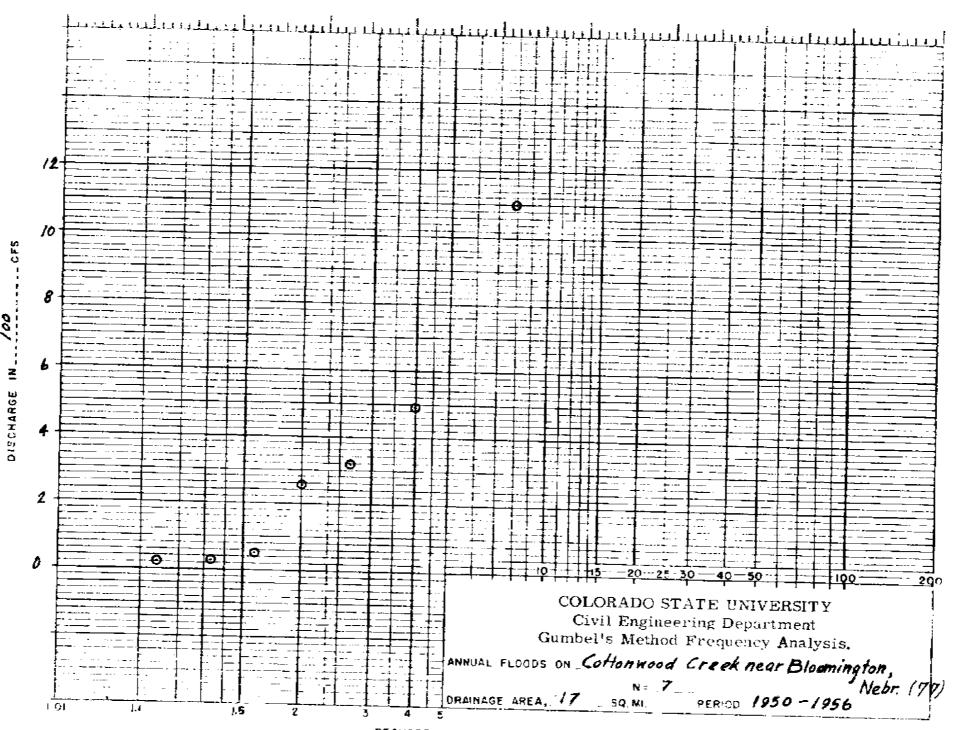


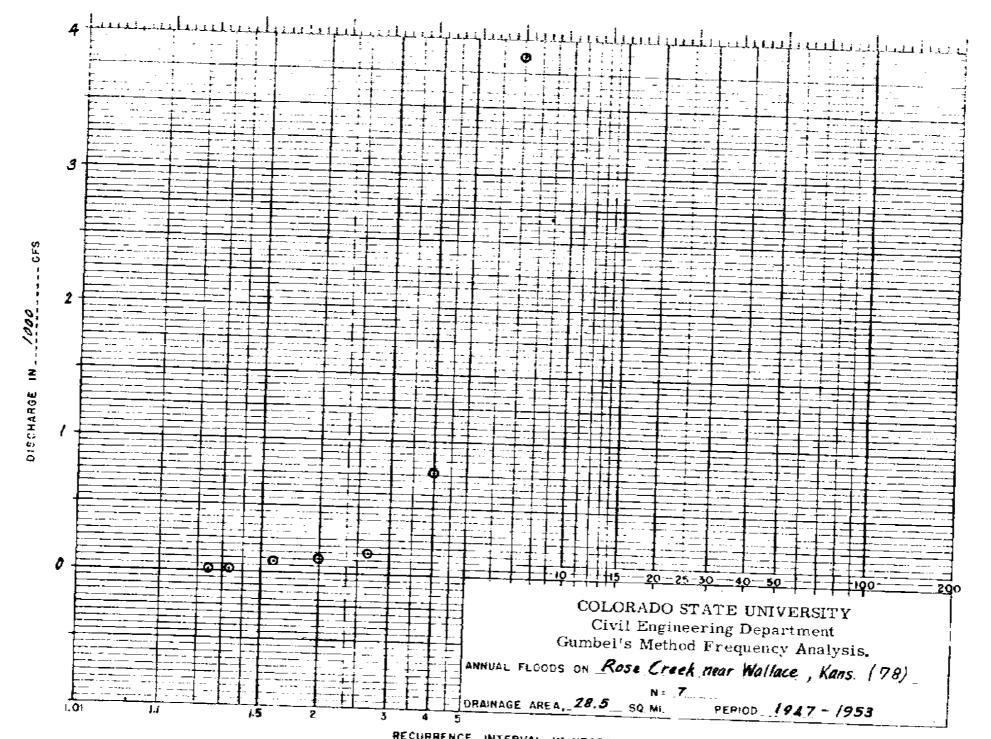


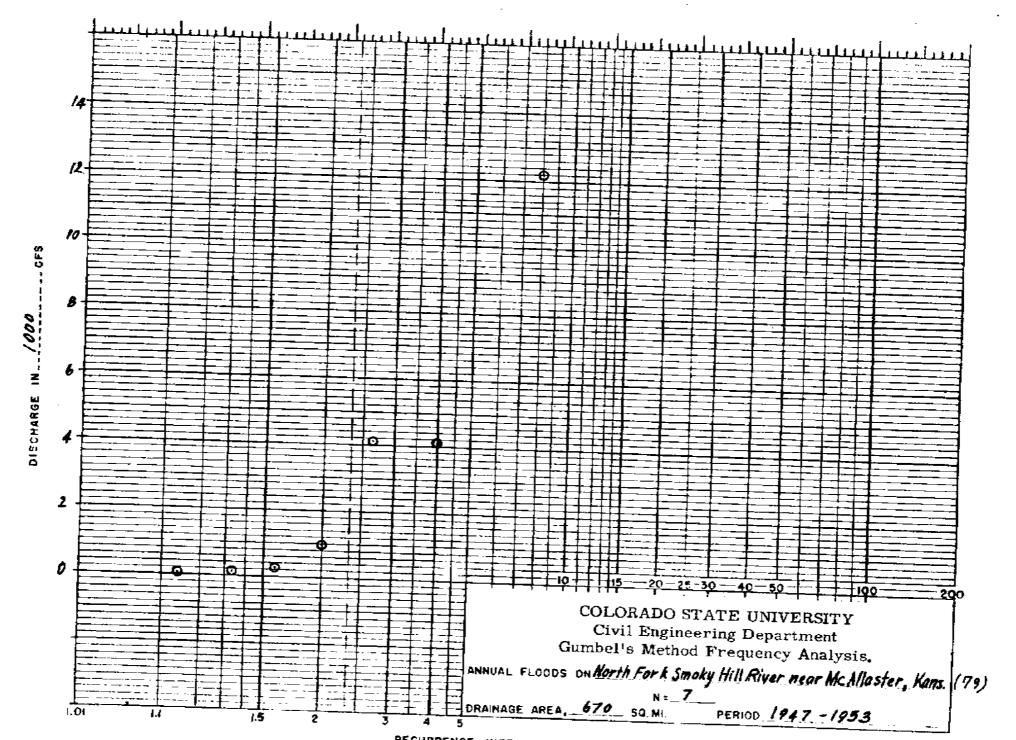


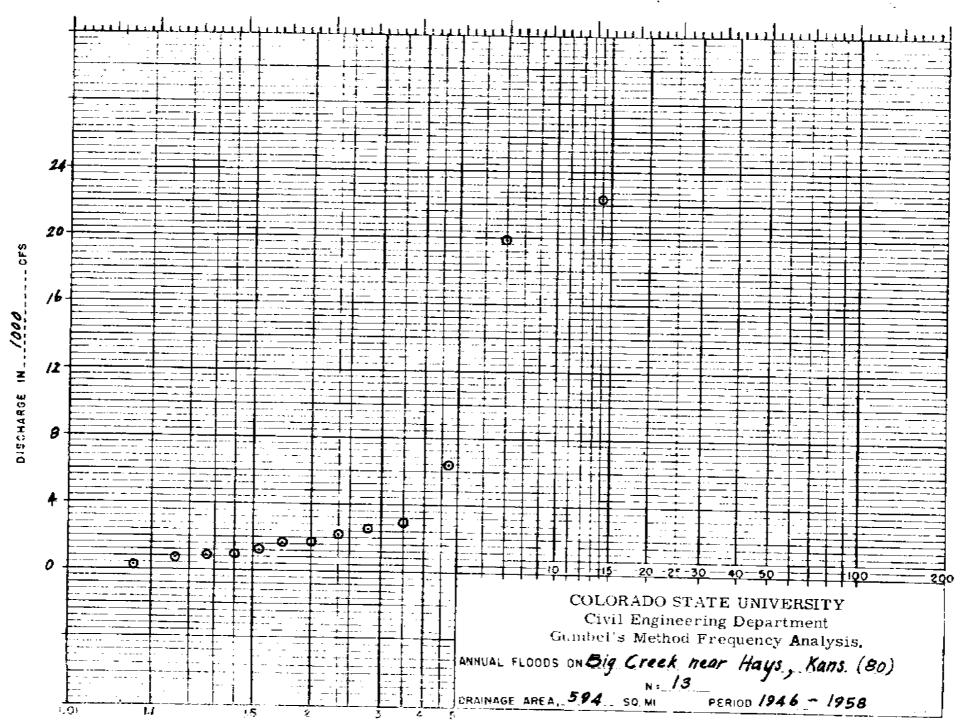


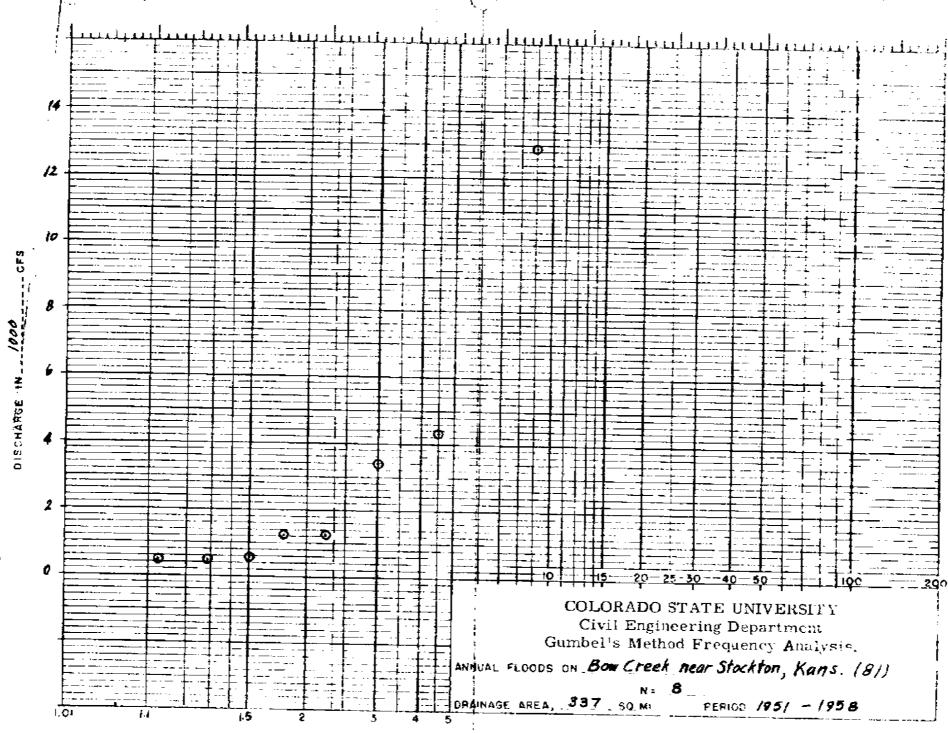




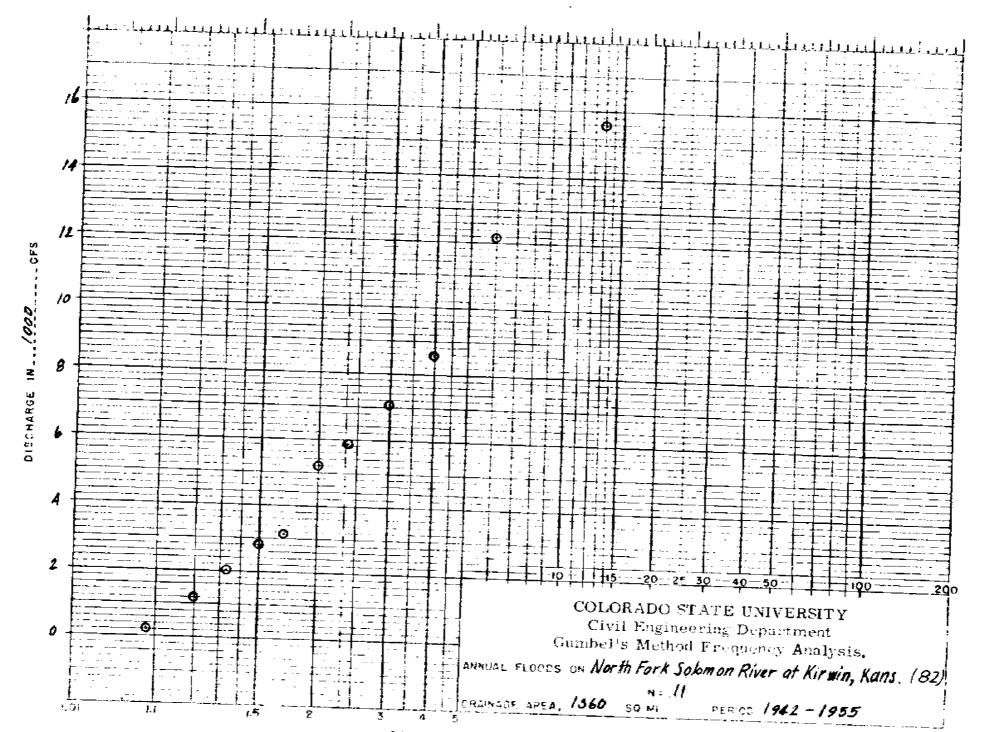


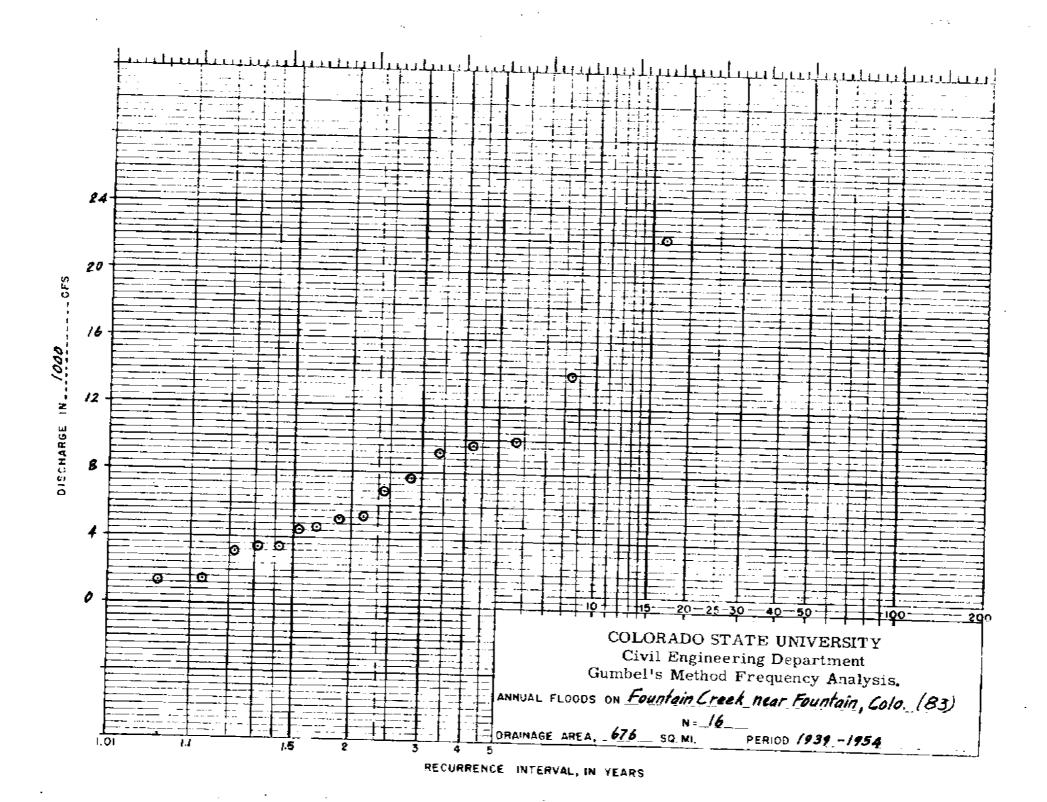


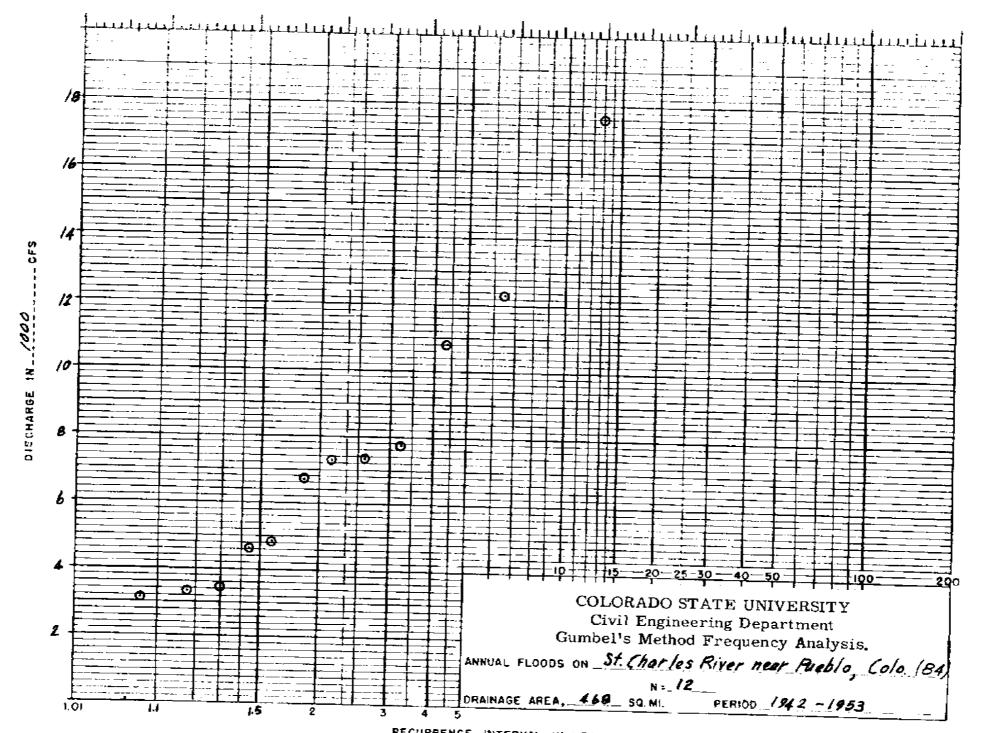


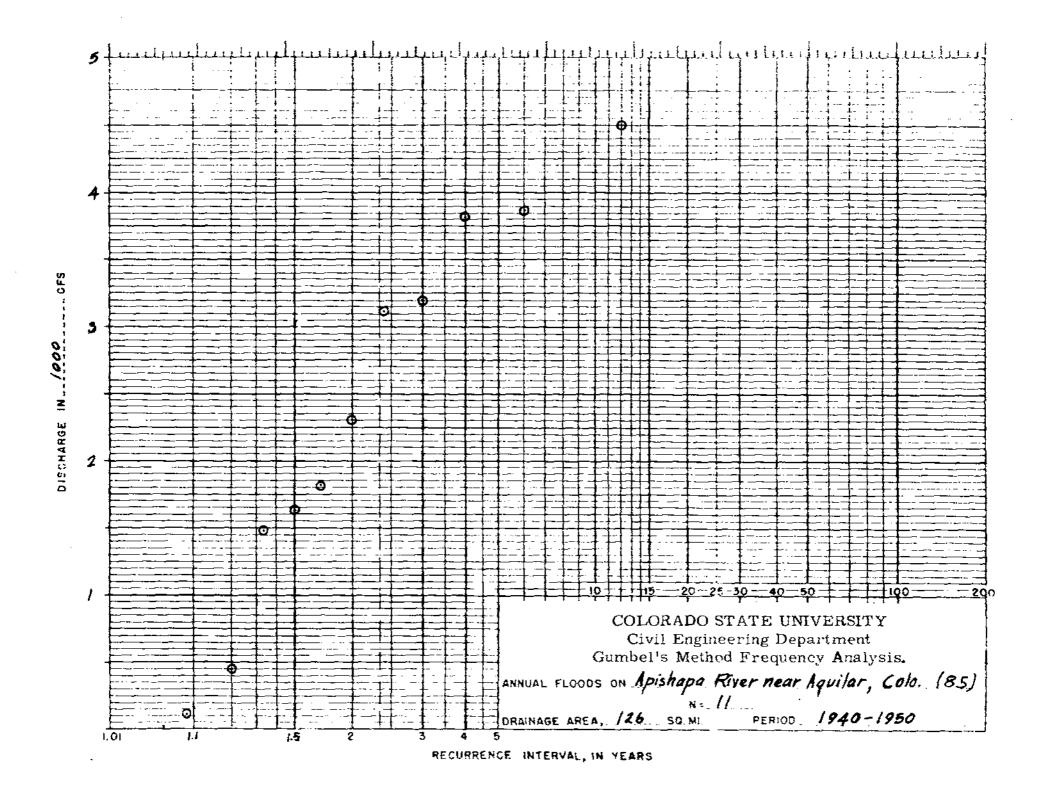


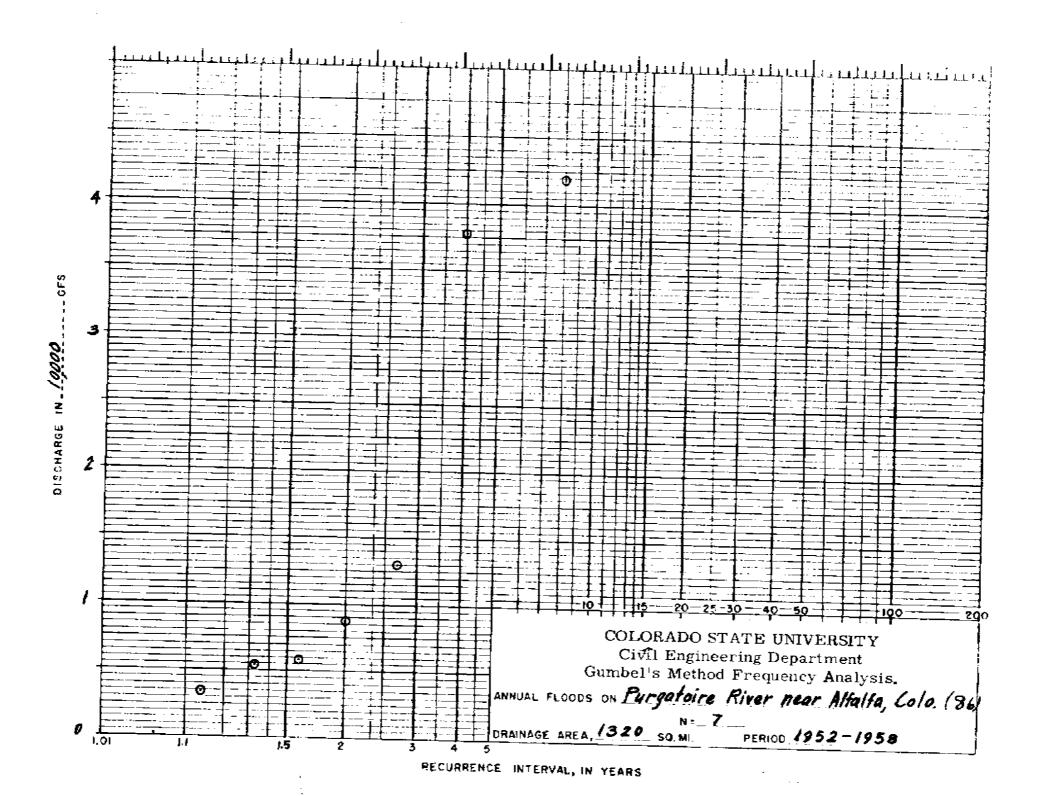
RELURRENCE INTERVAL, IN YEARS

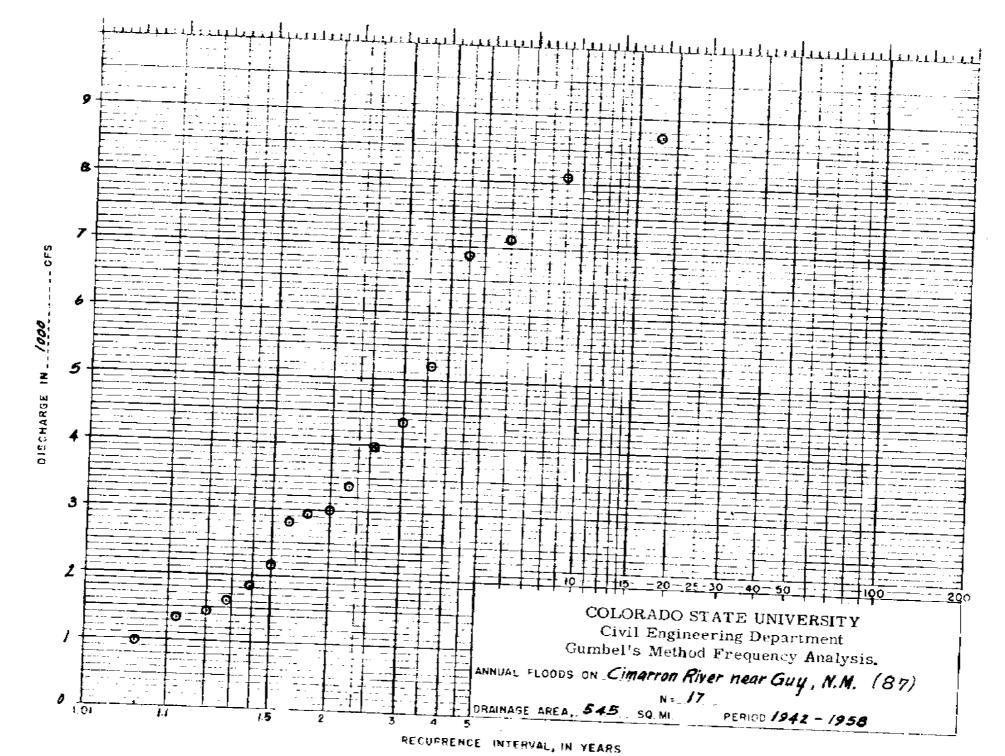


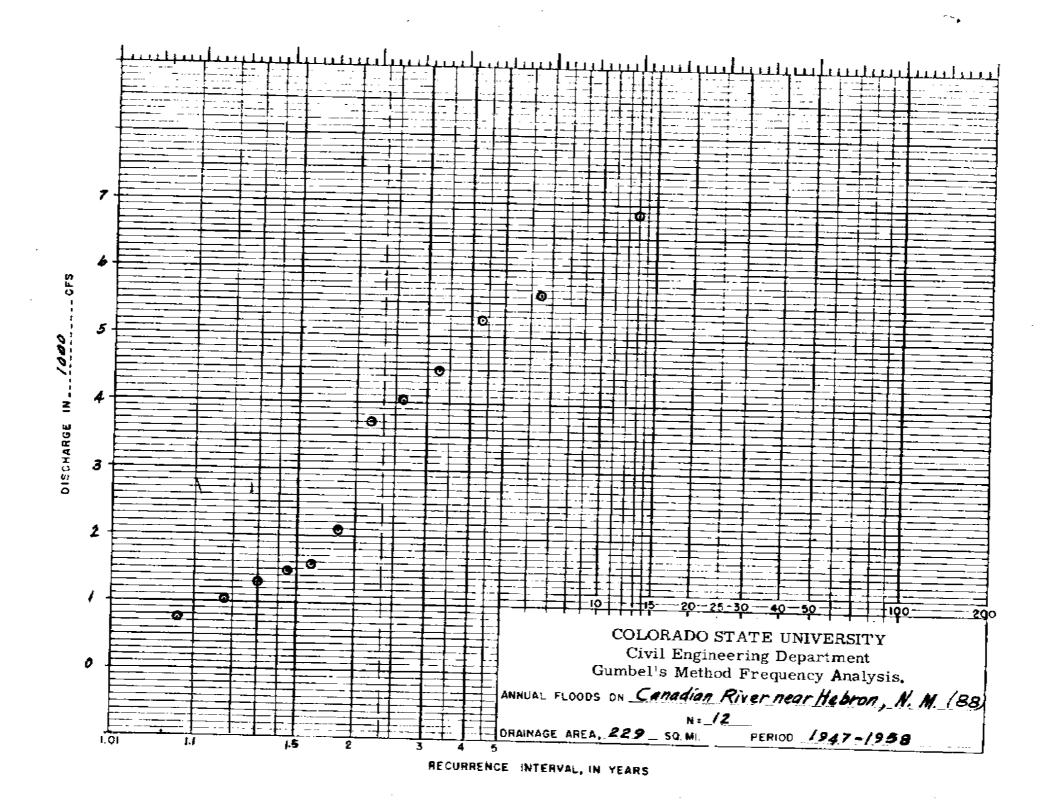


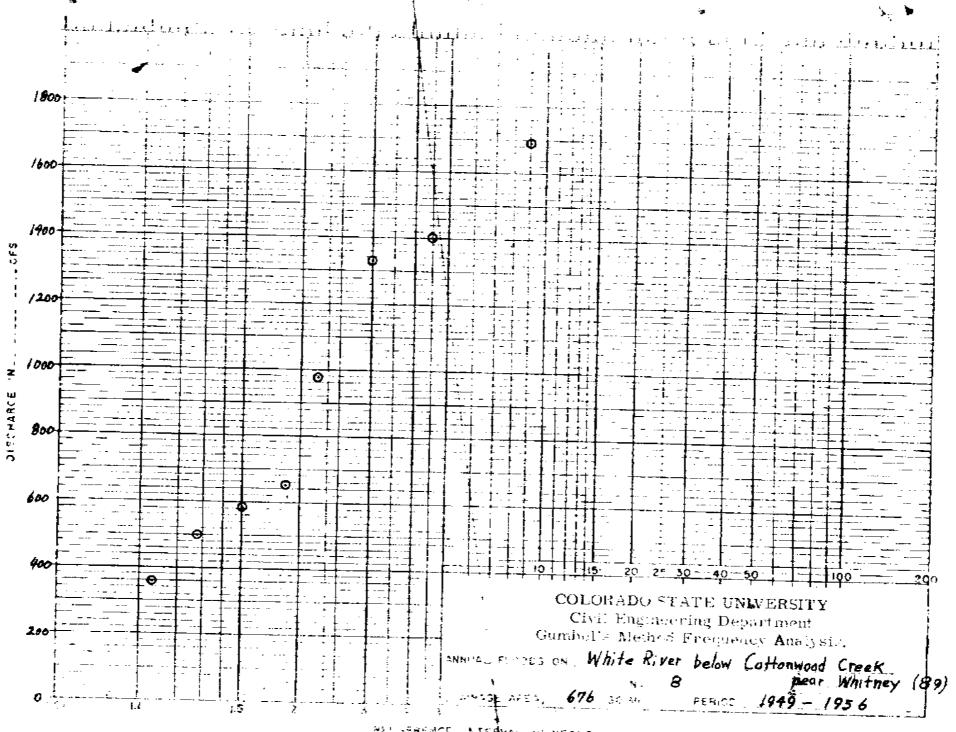




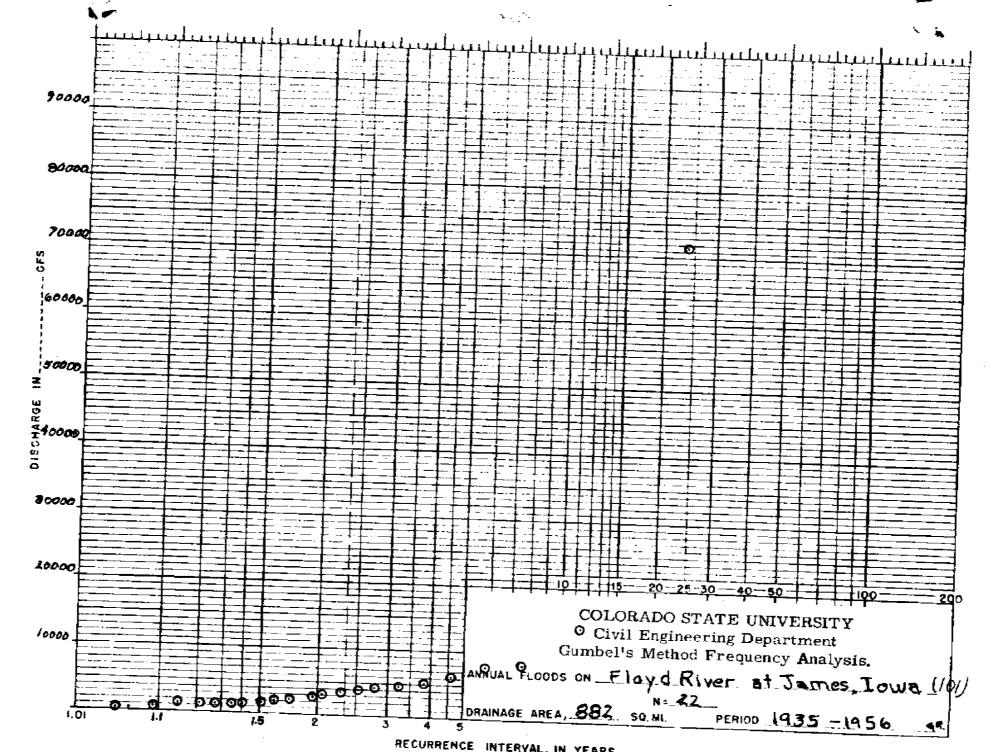


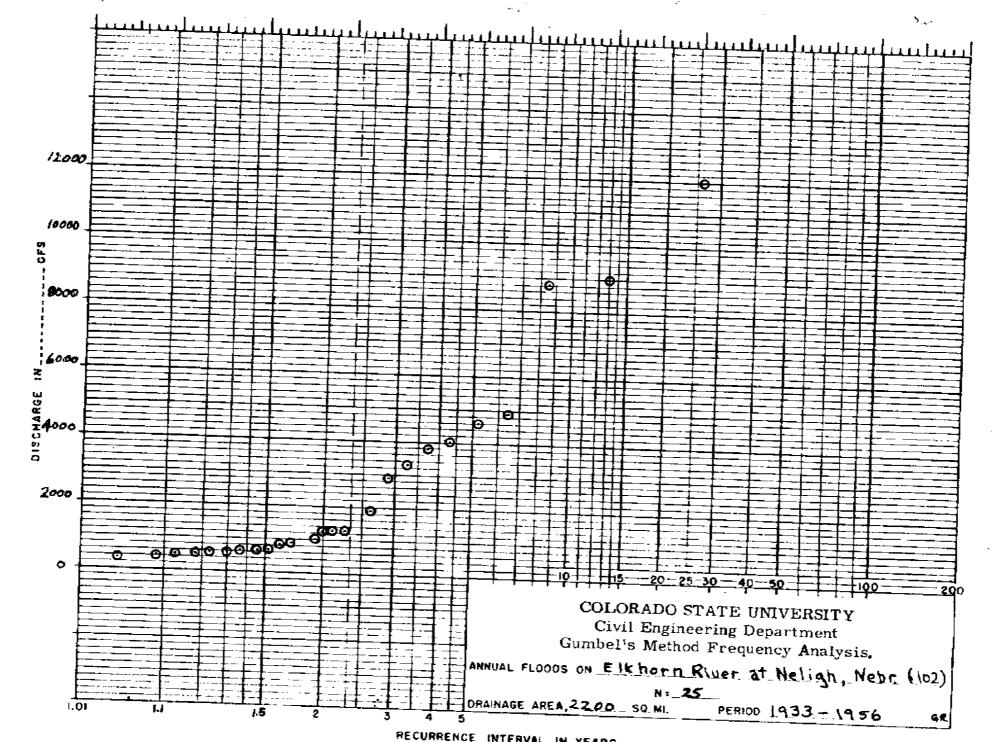


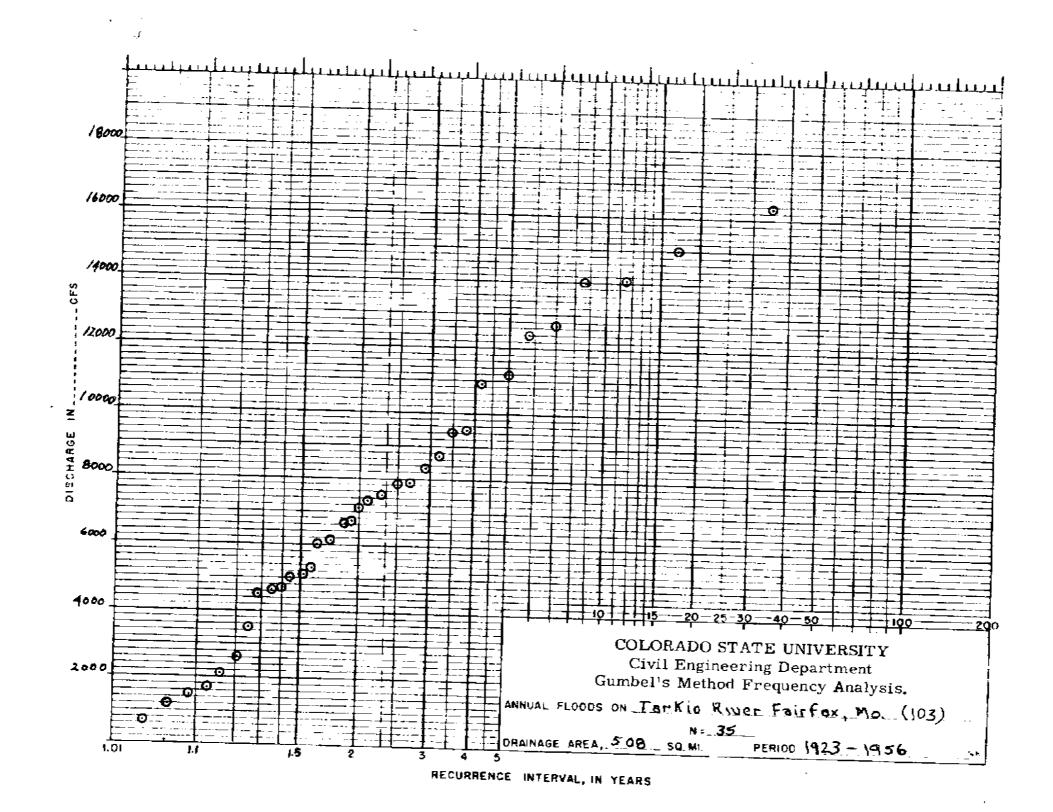


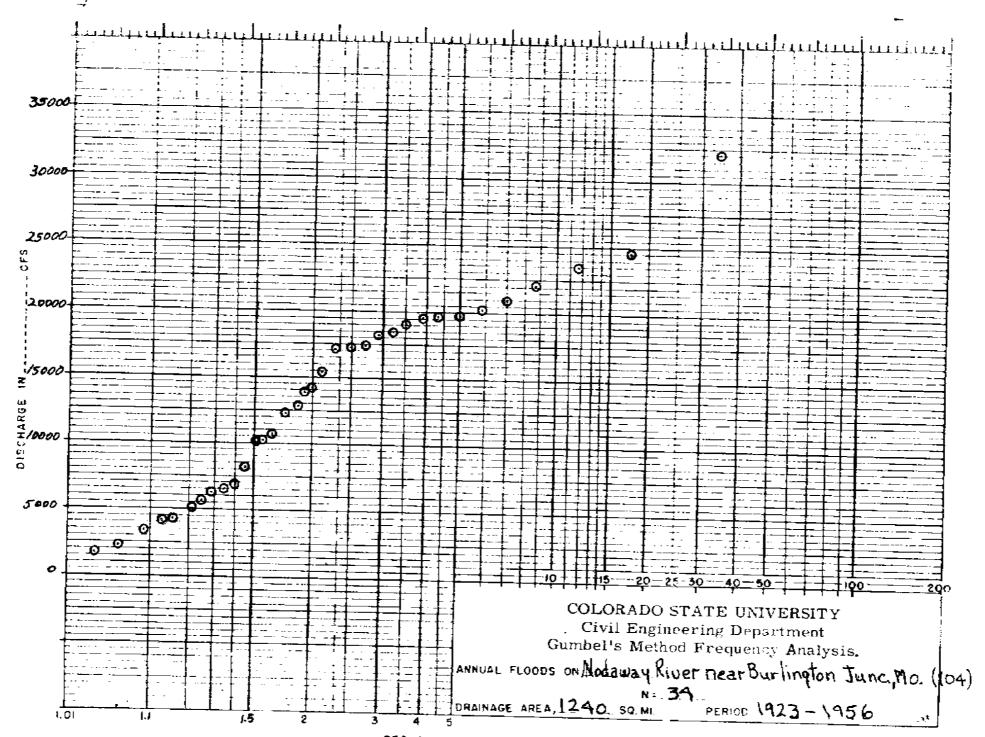


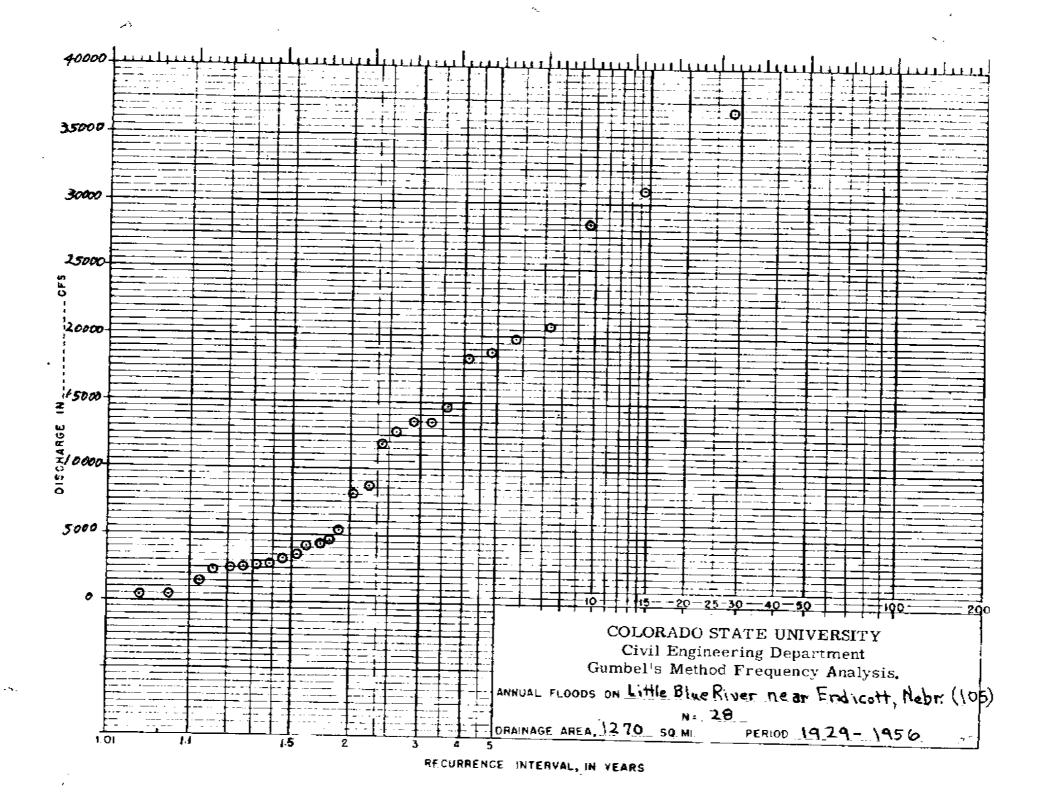
HET PREMOTE KTERVAL, IN YEARS

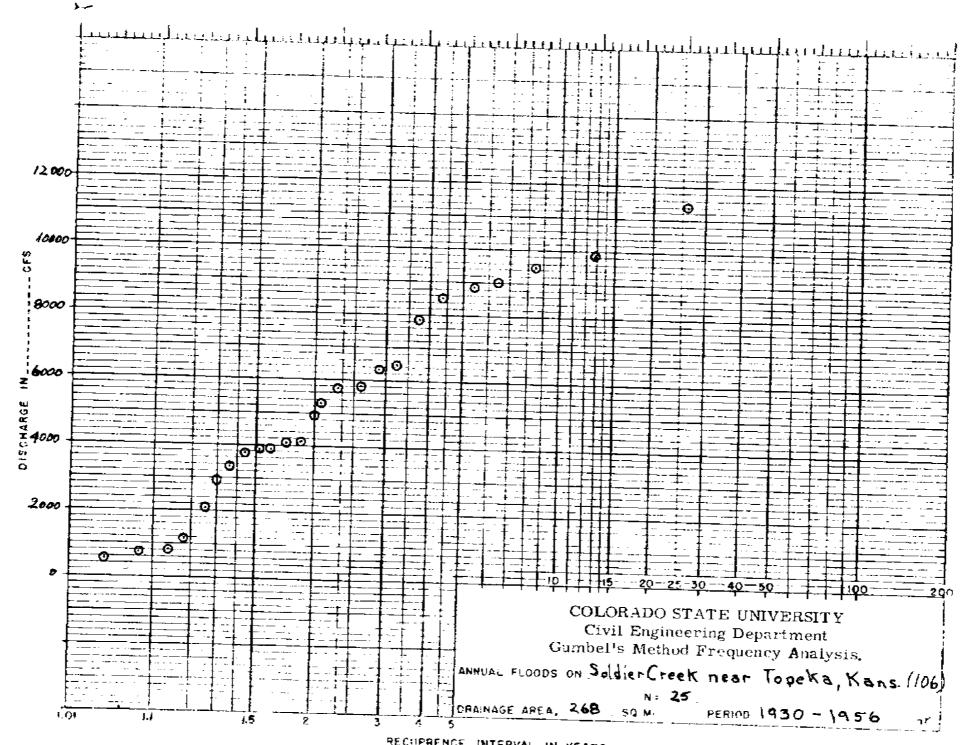


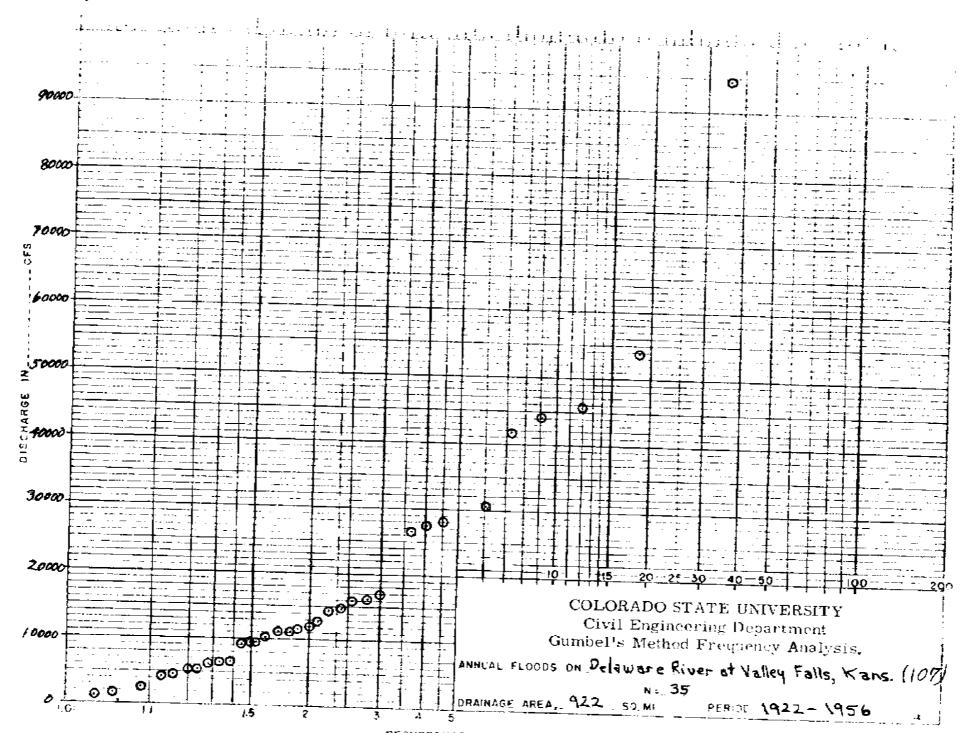


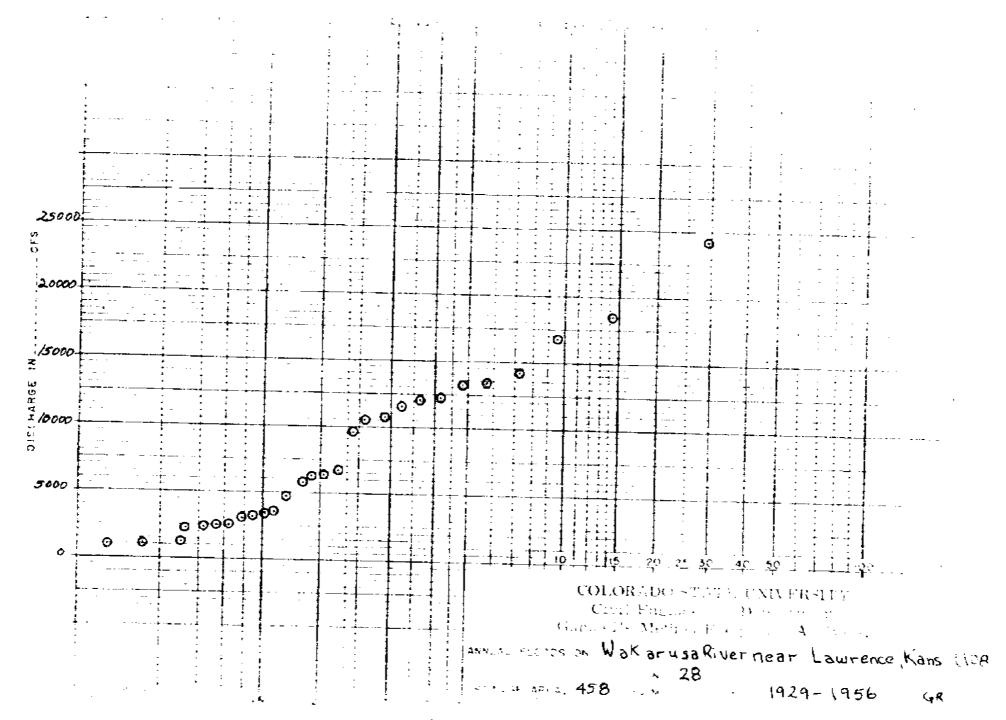


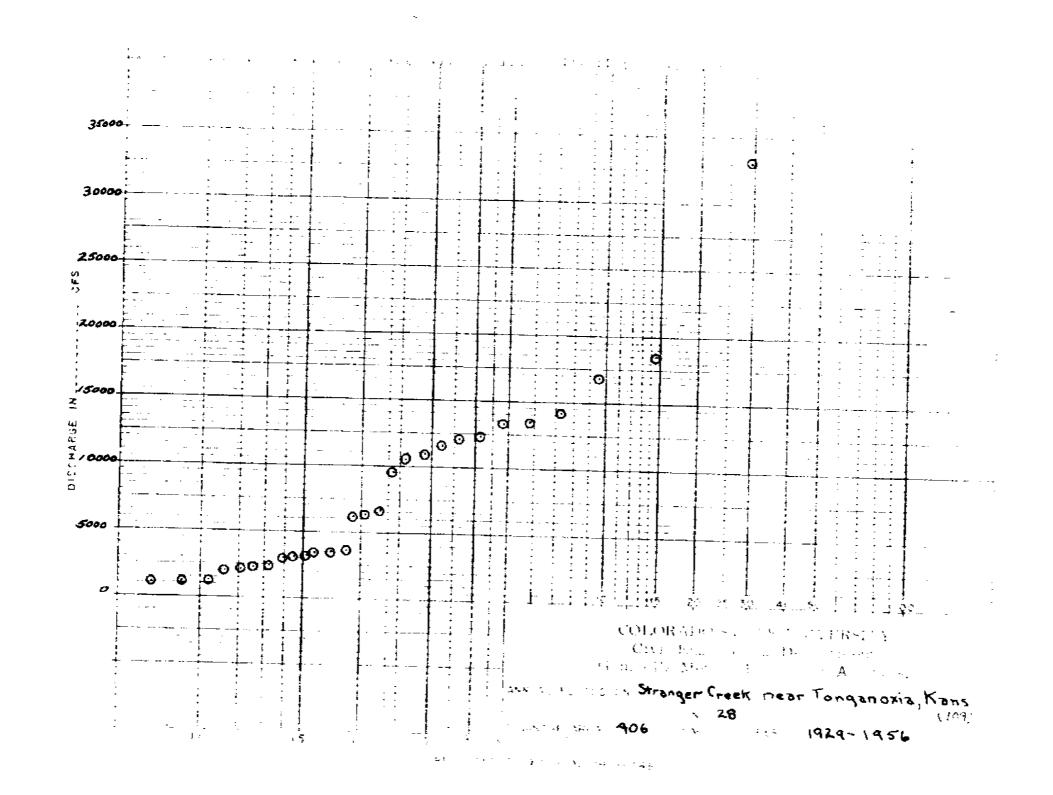


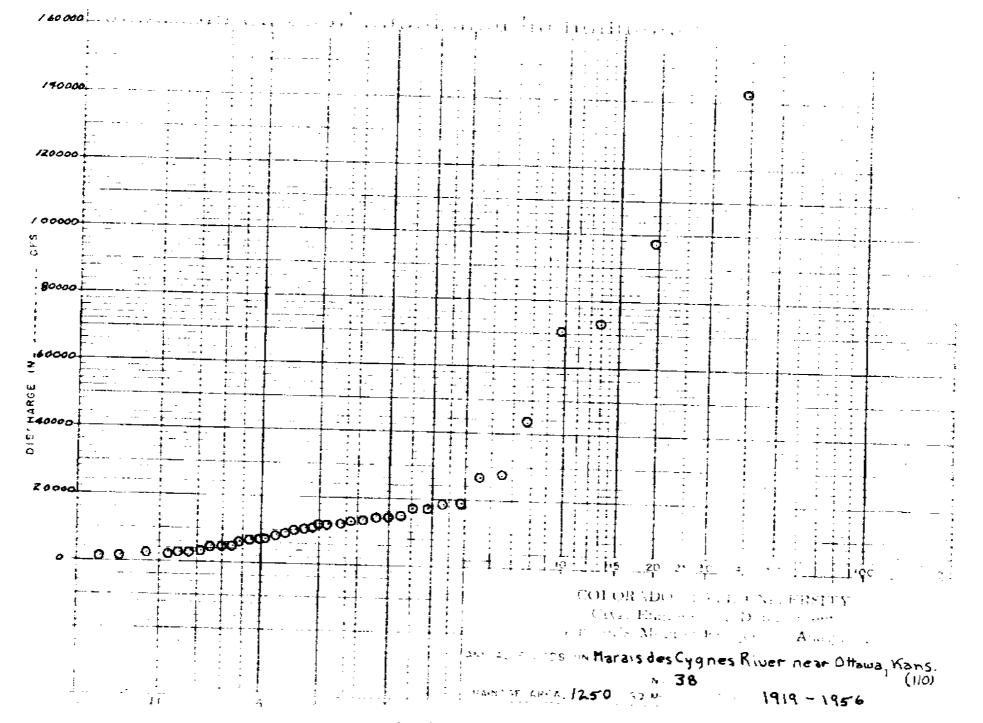


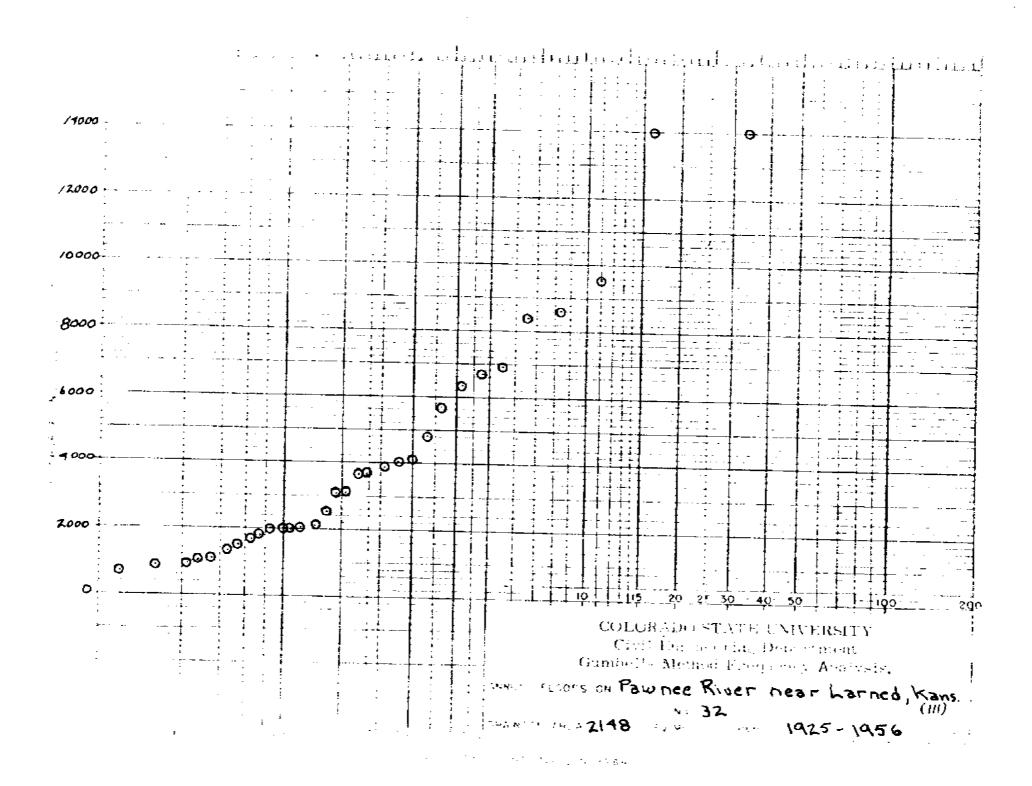


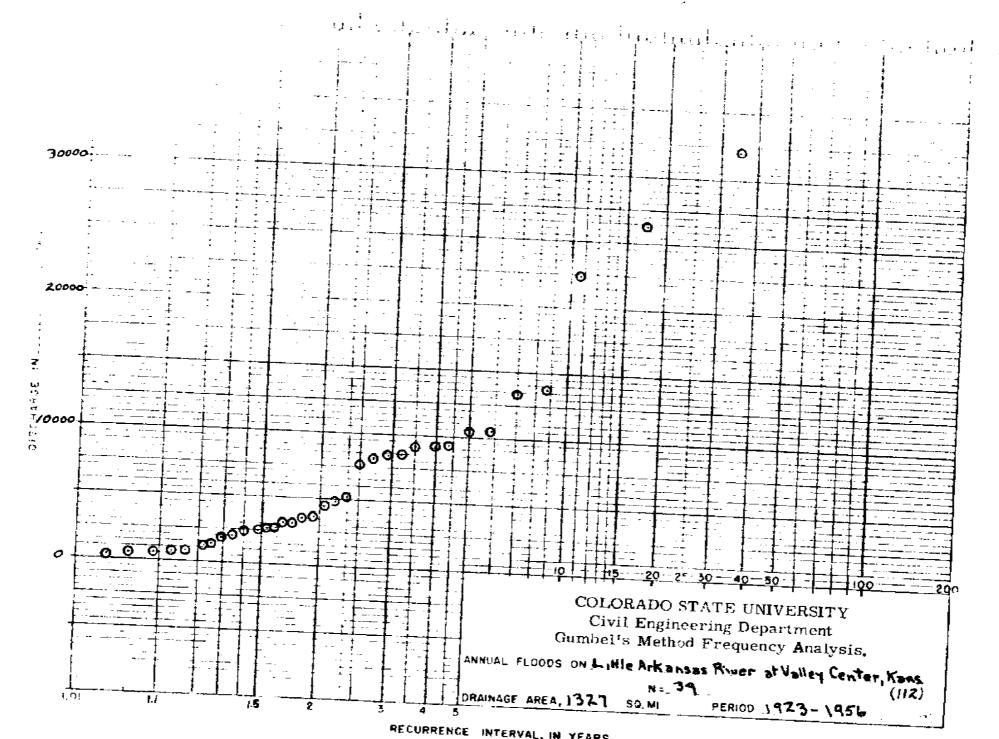


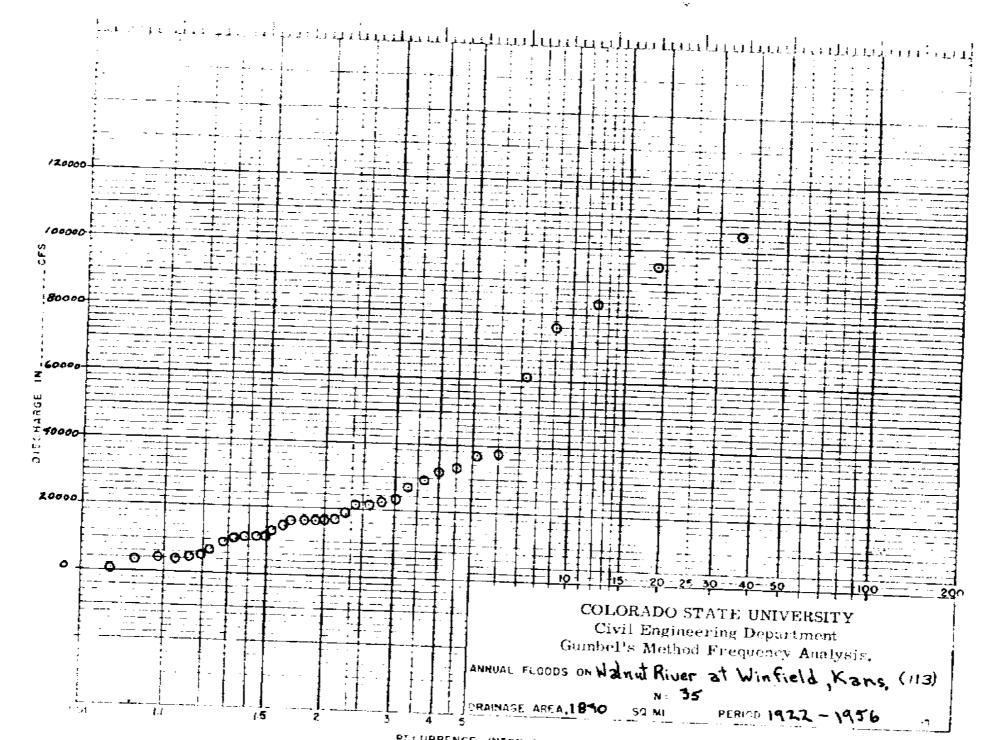


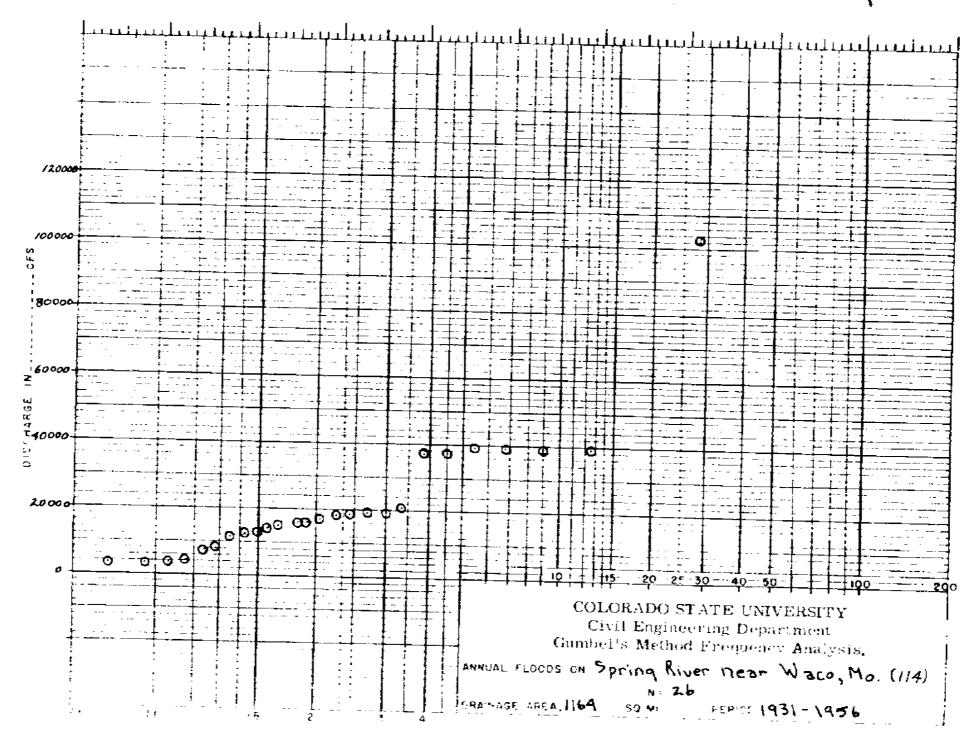


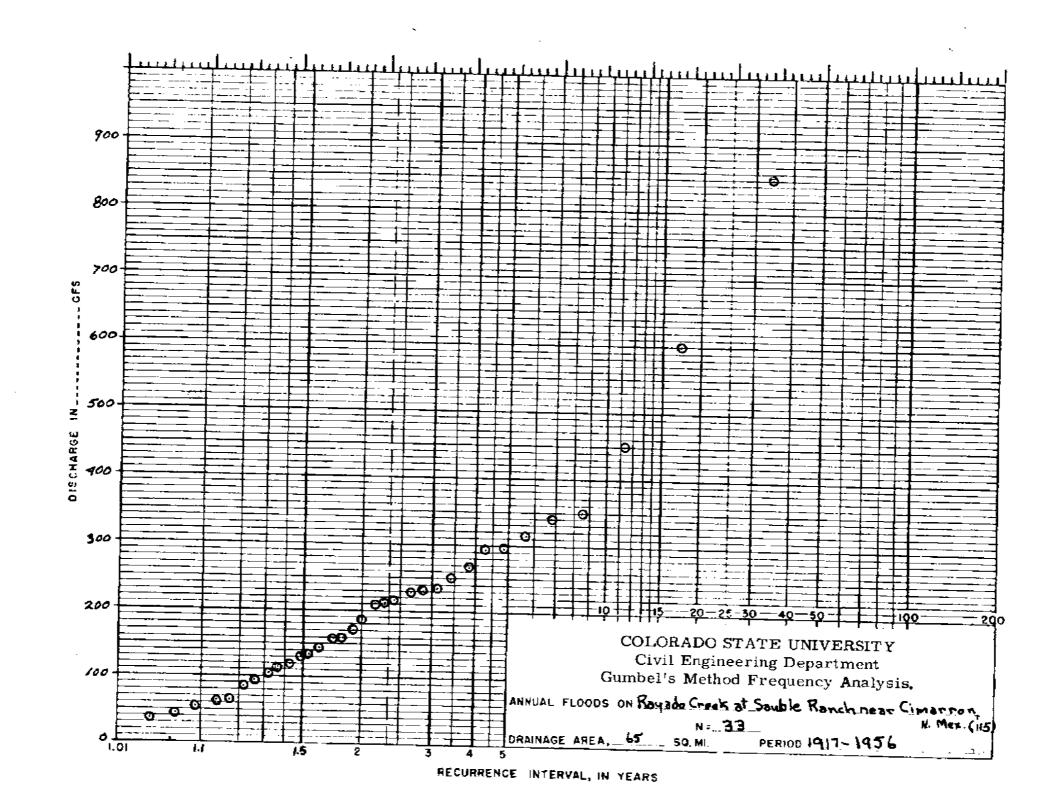


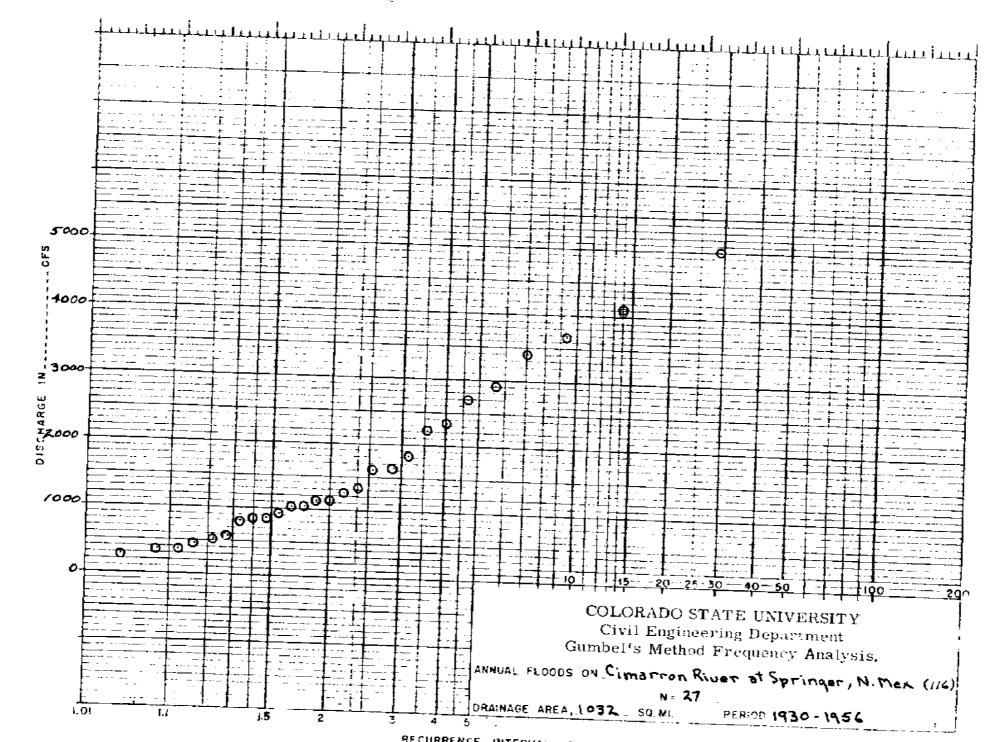


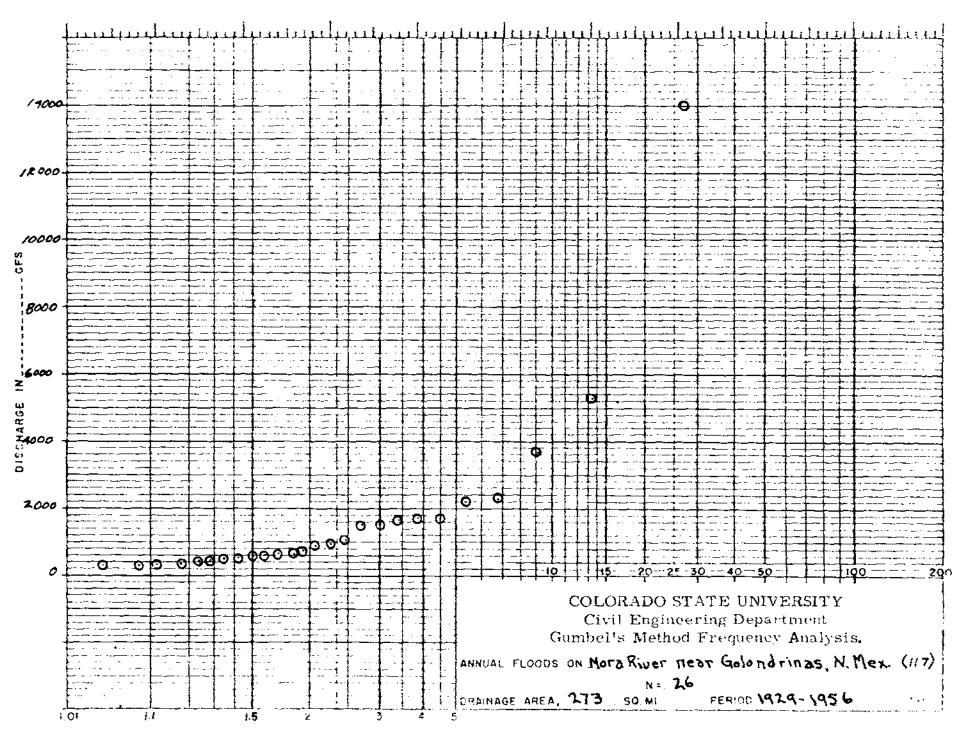


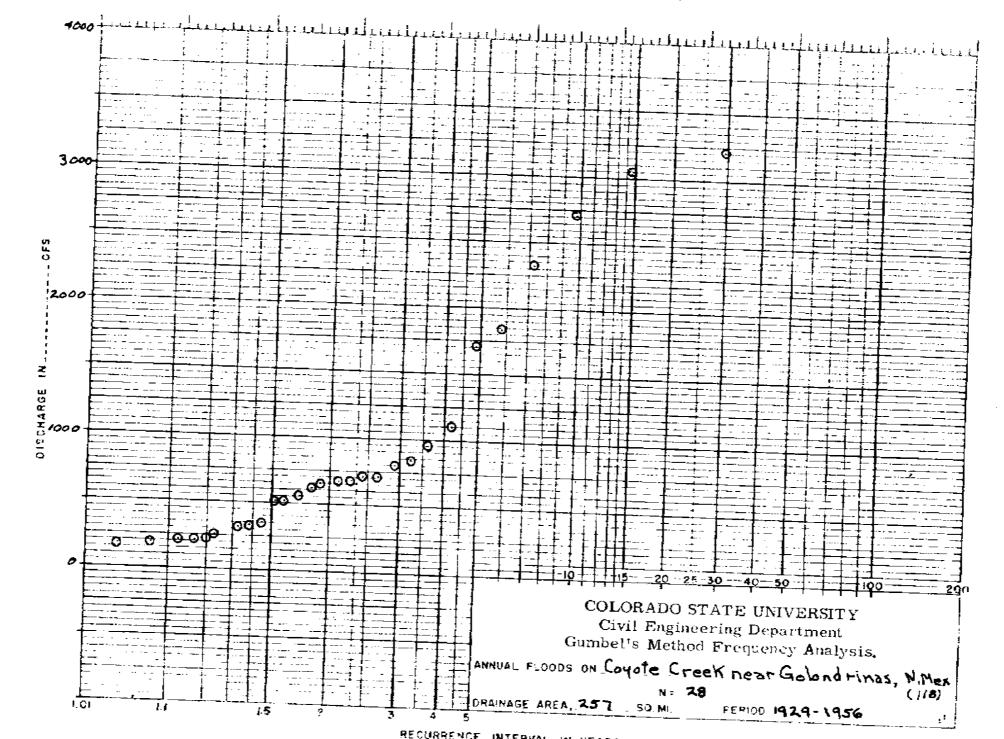


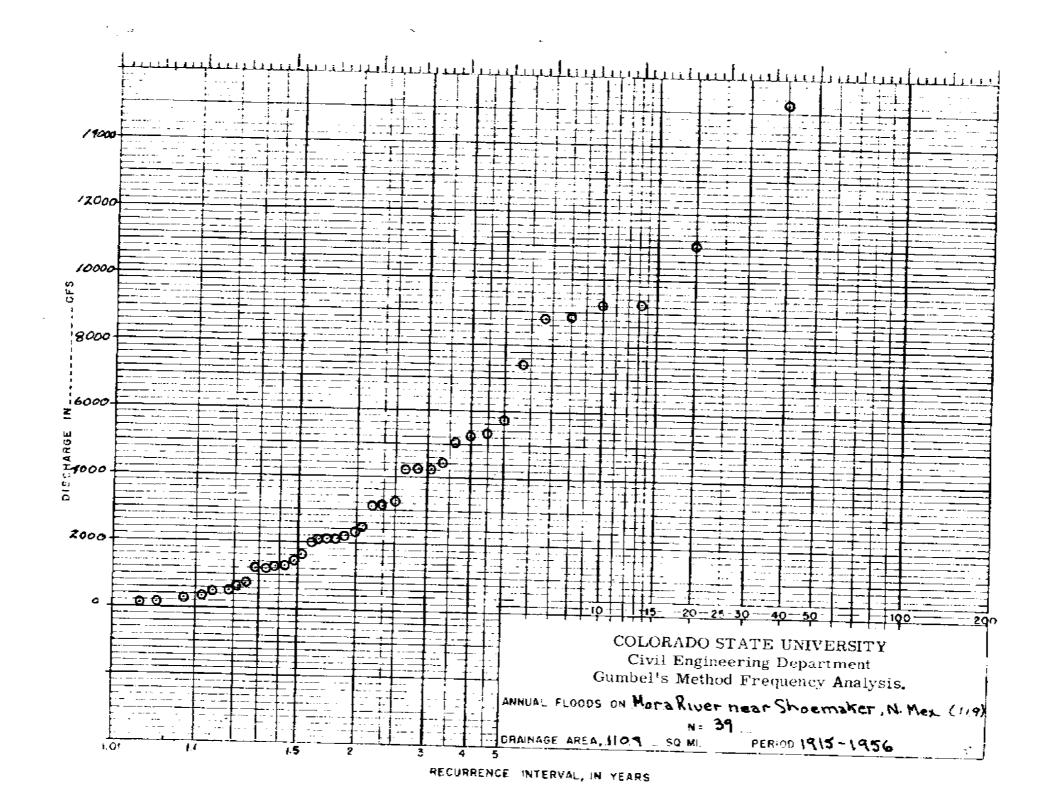


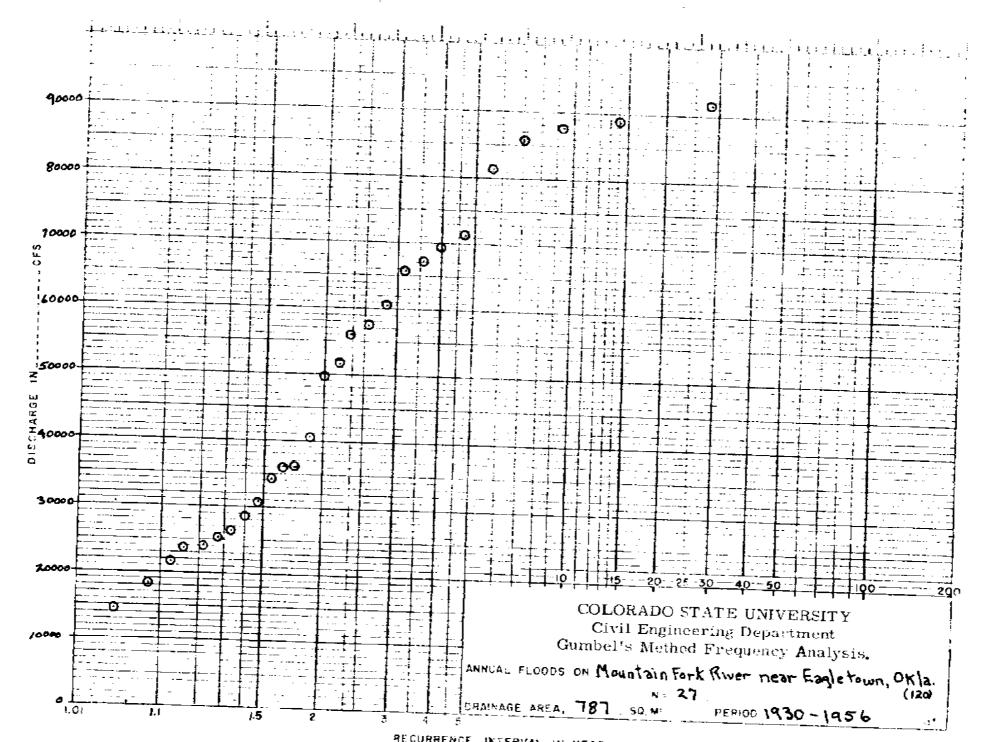


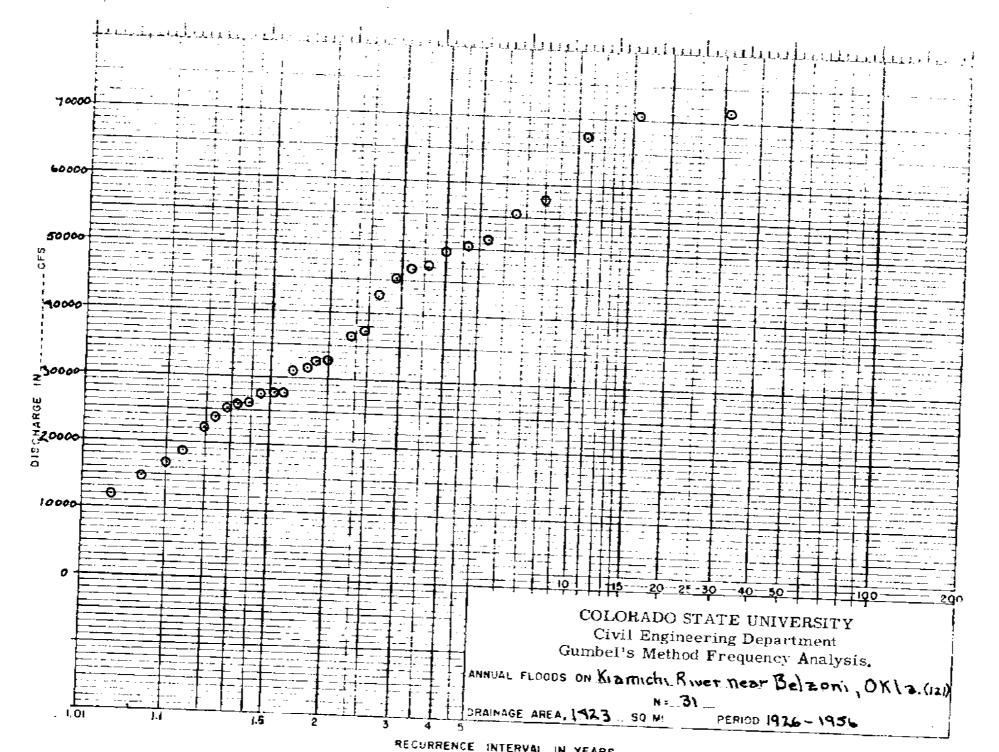


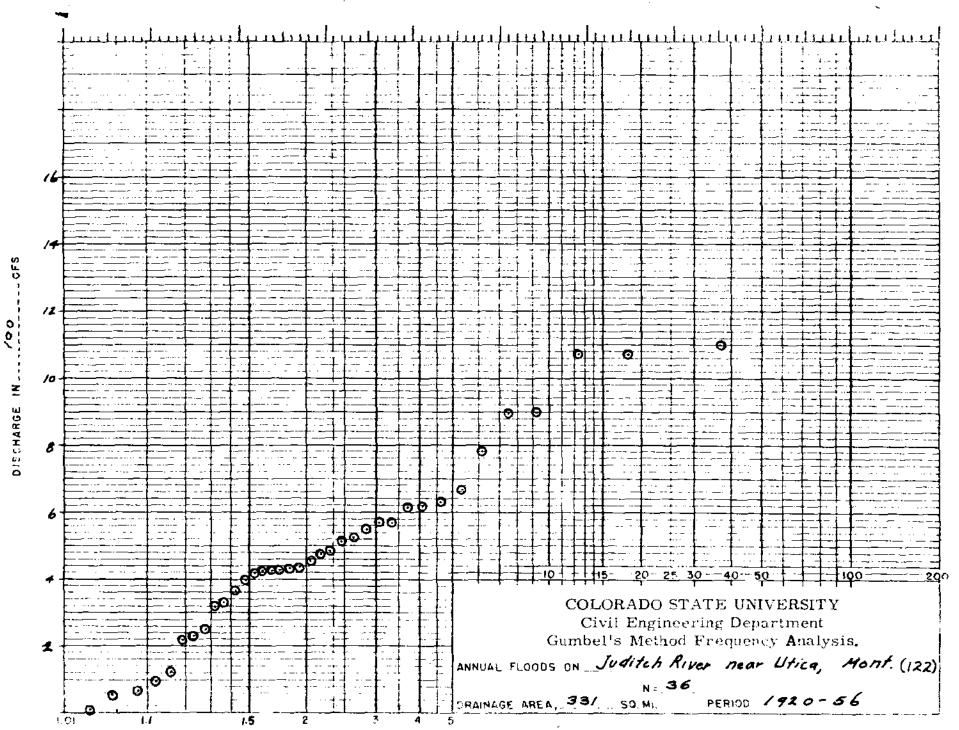


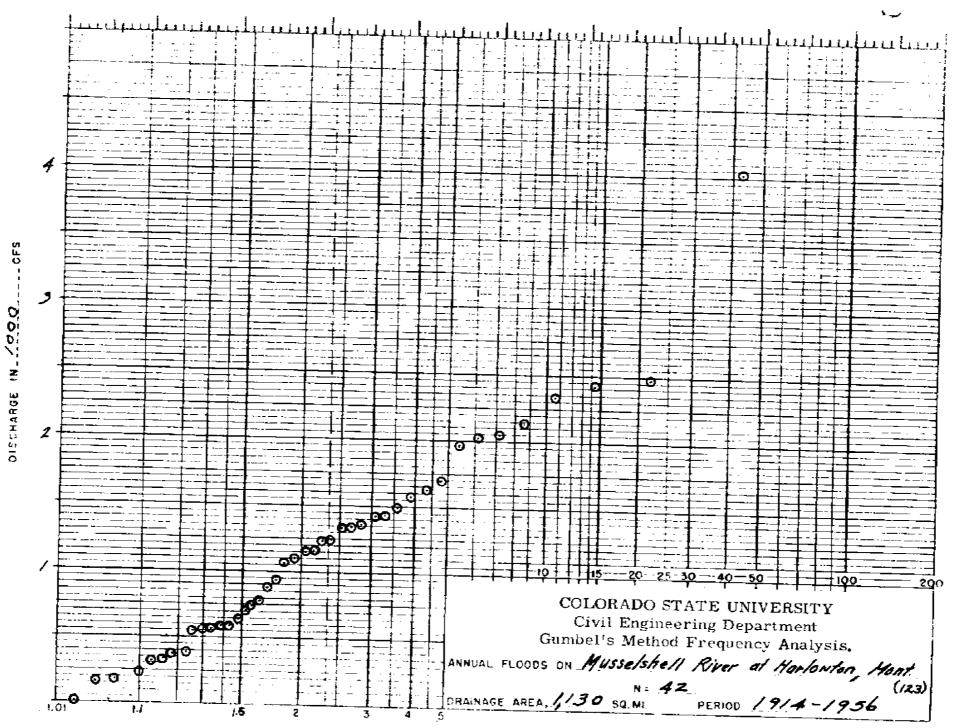


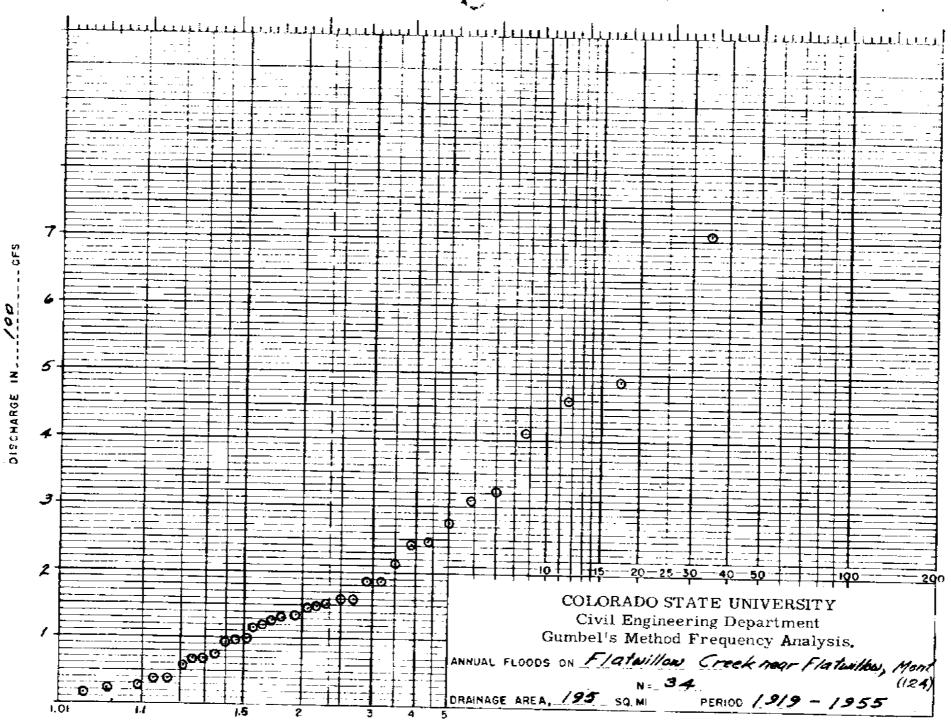


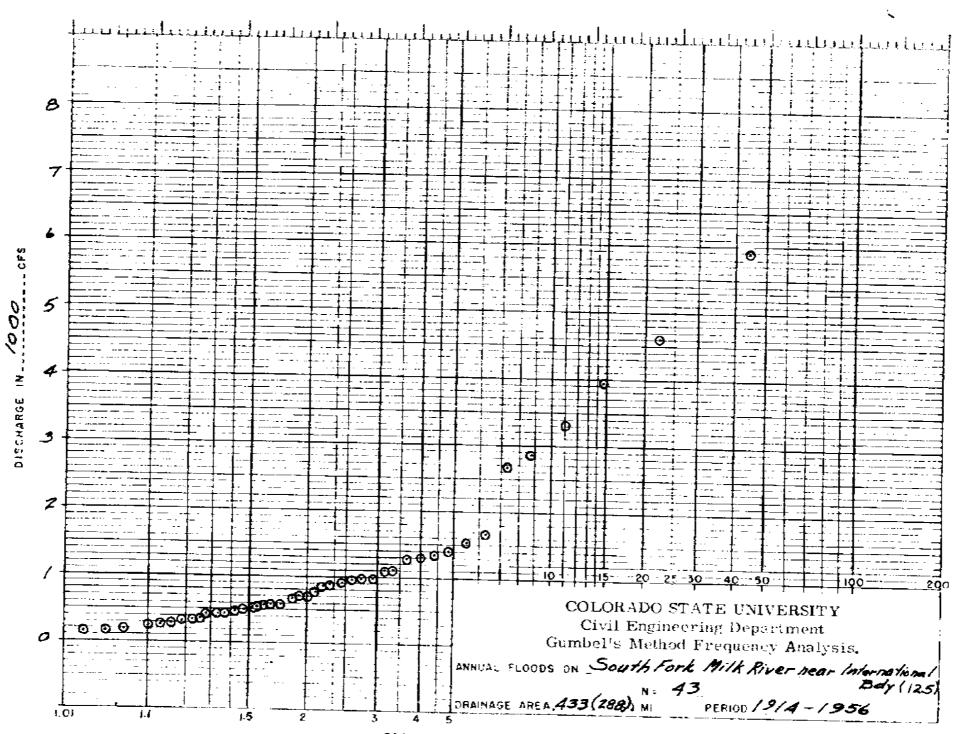


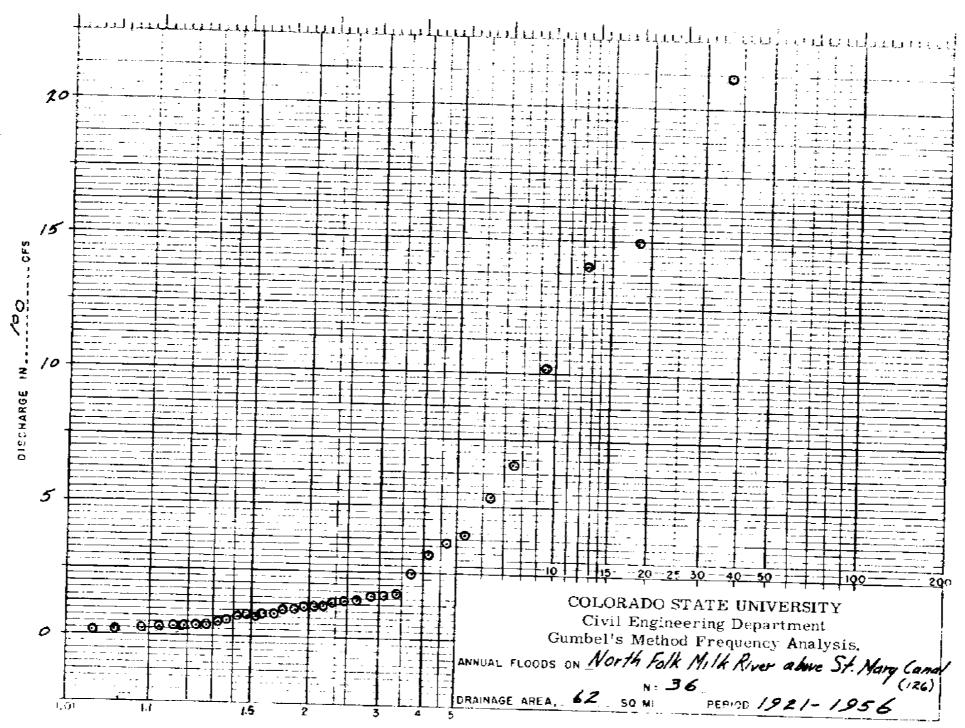


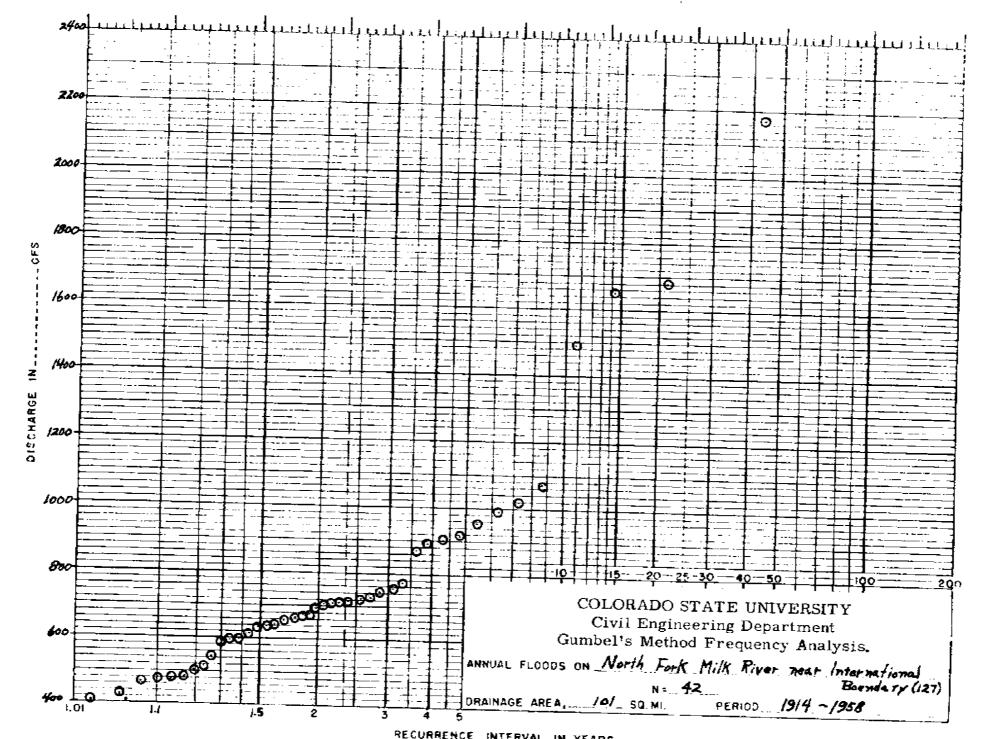


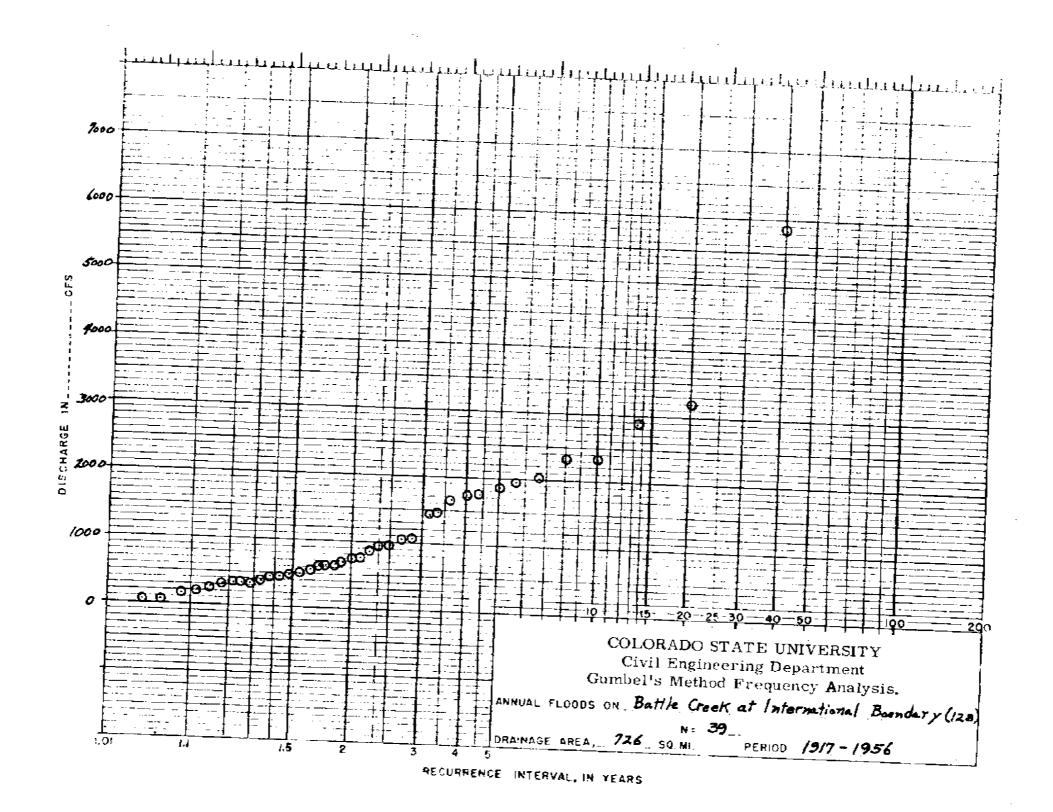


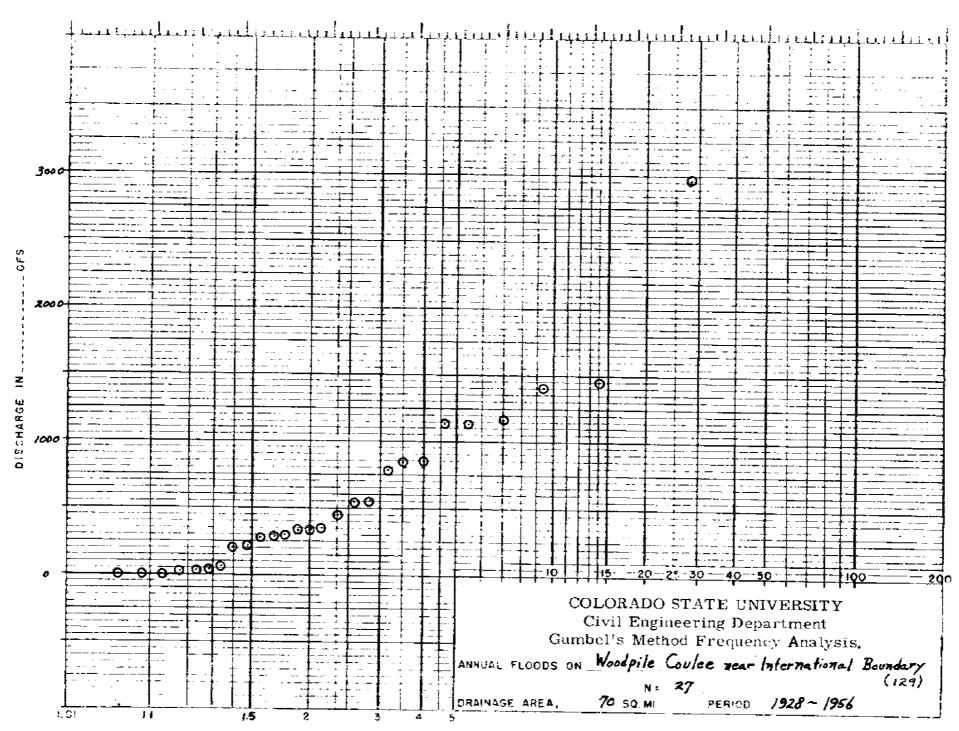


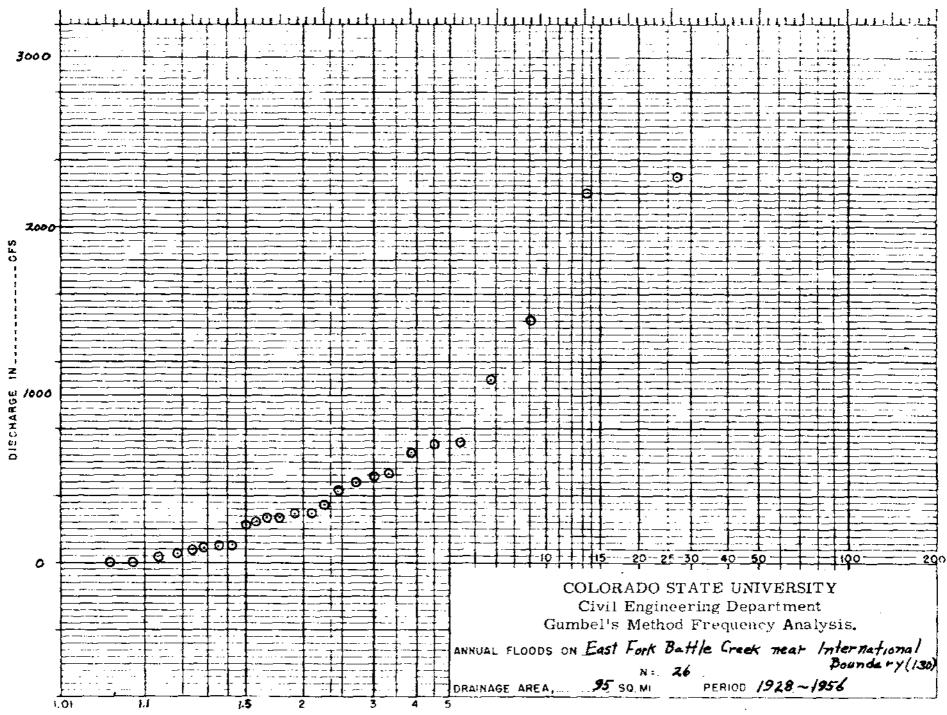


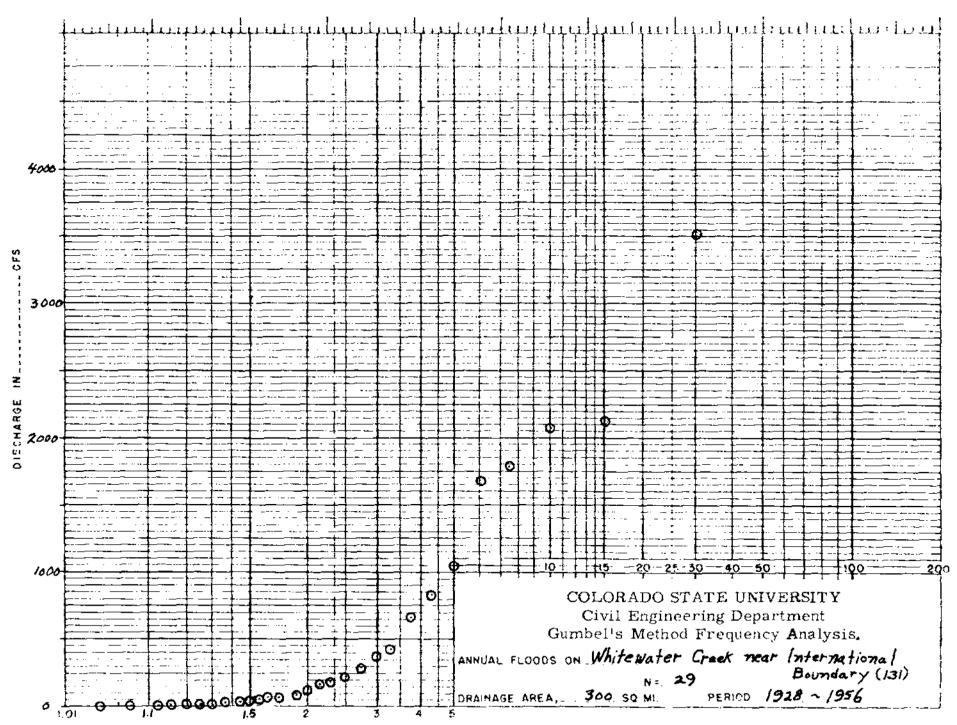


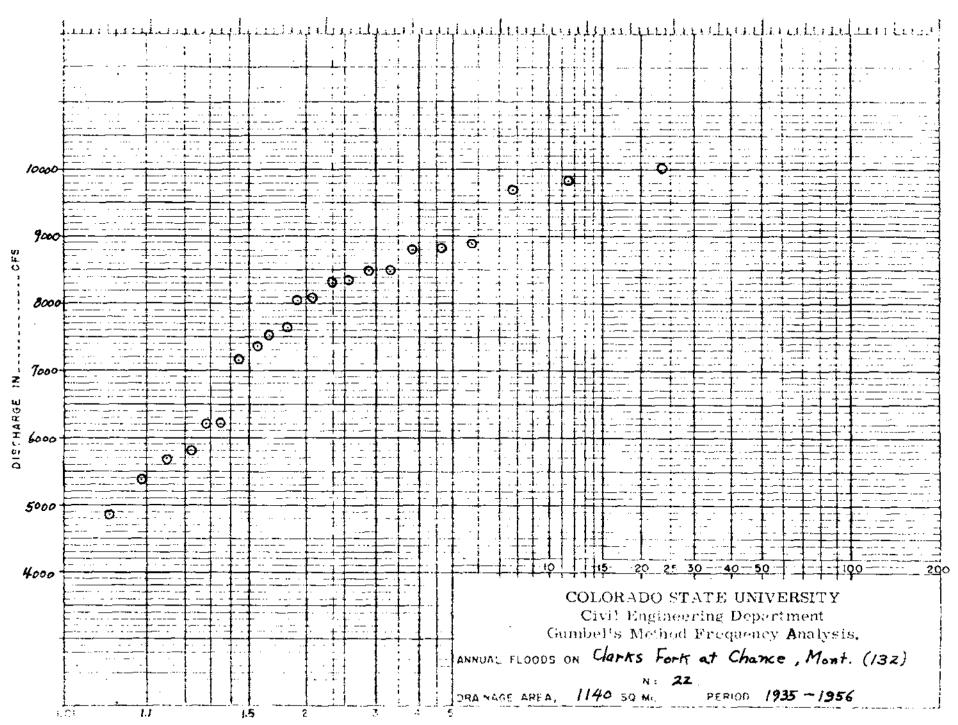


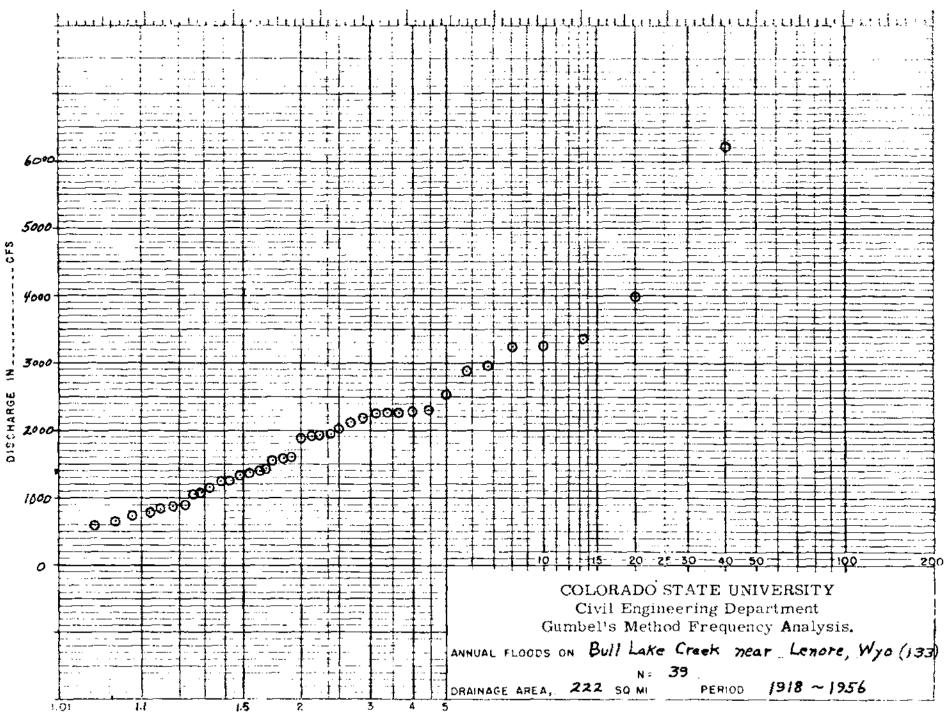


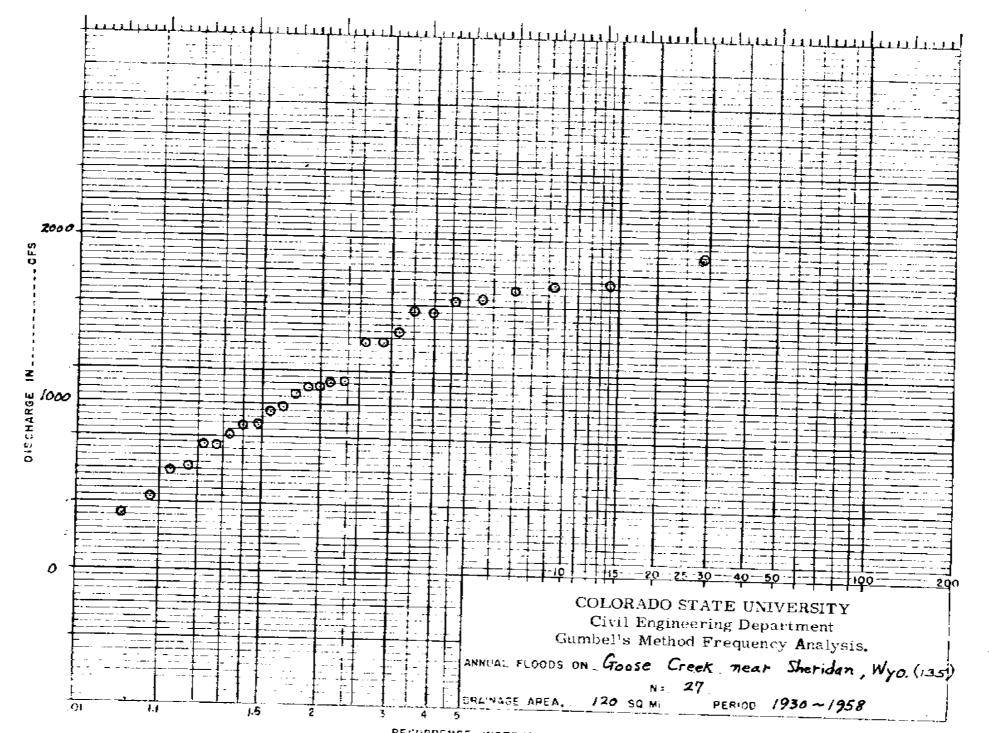


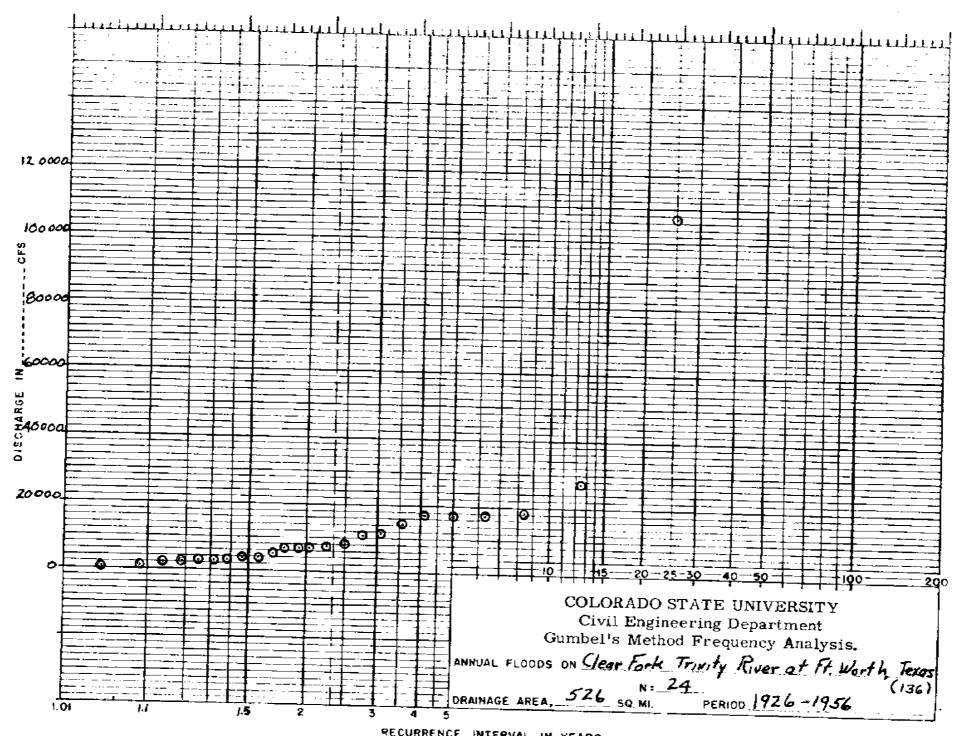


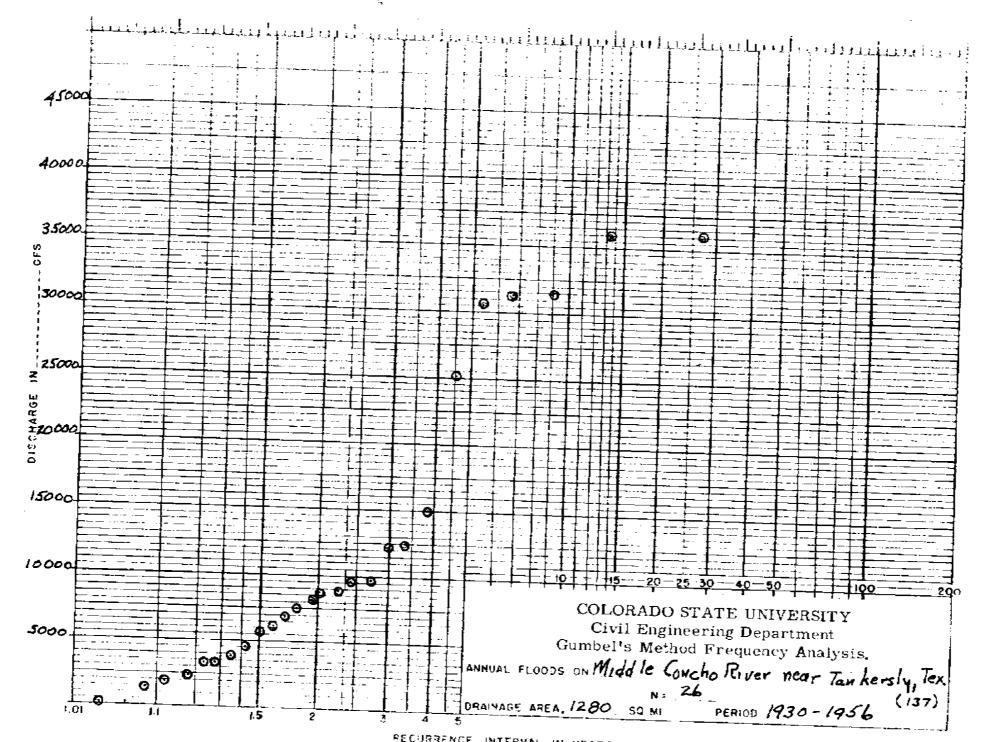


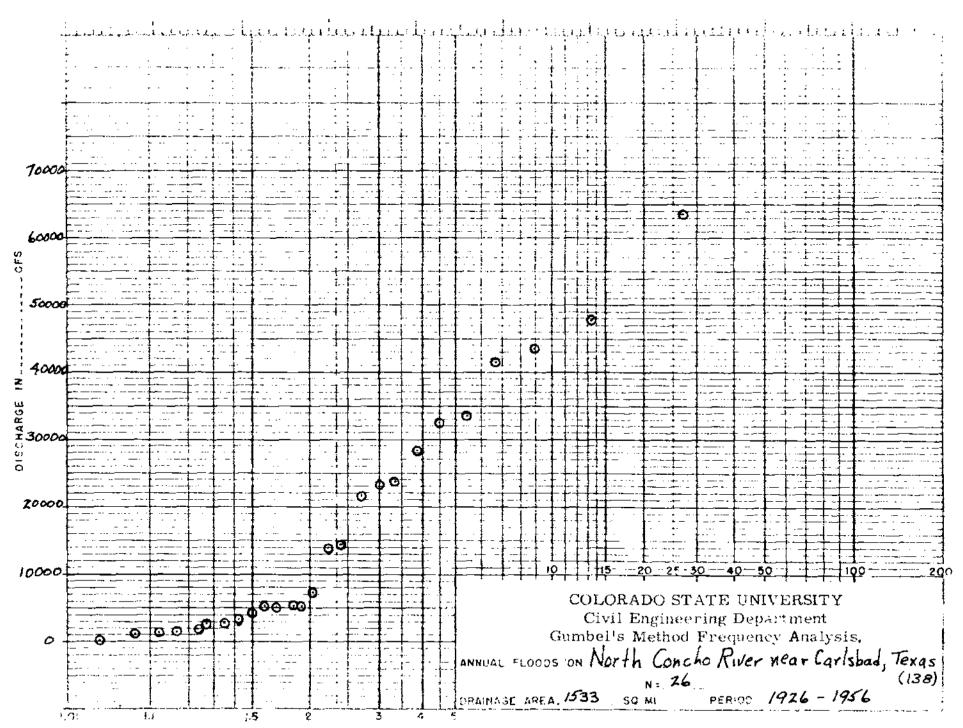


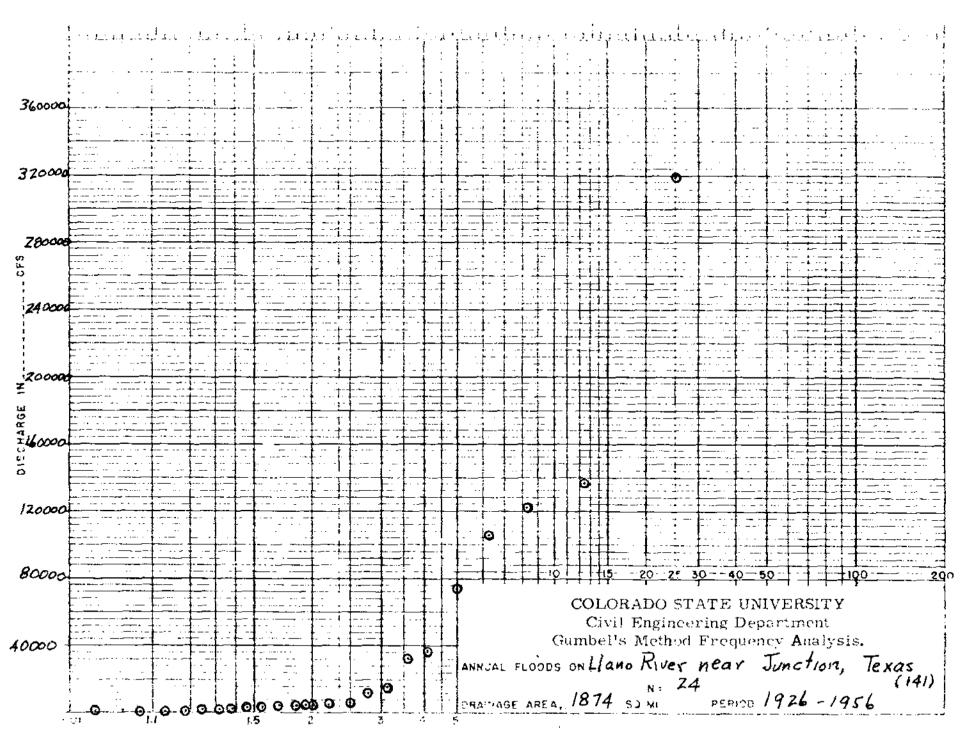


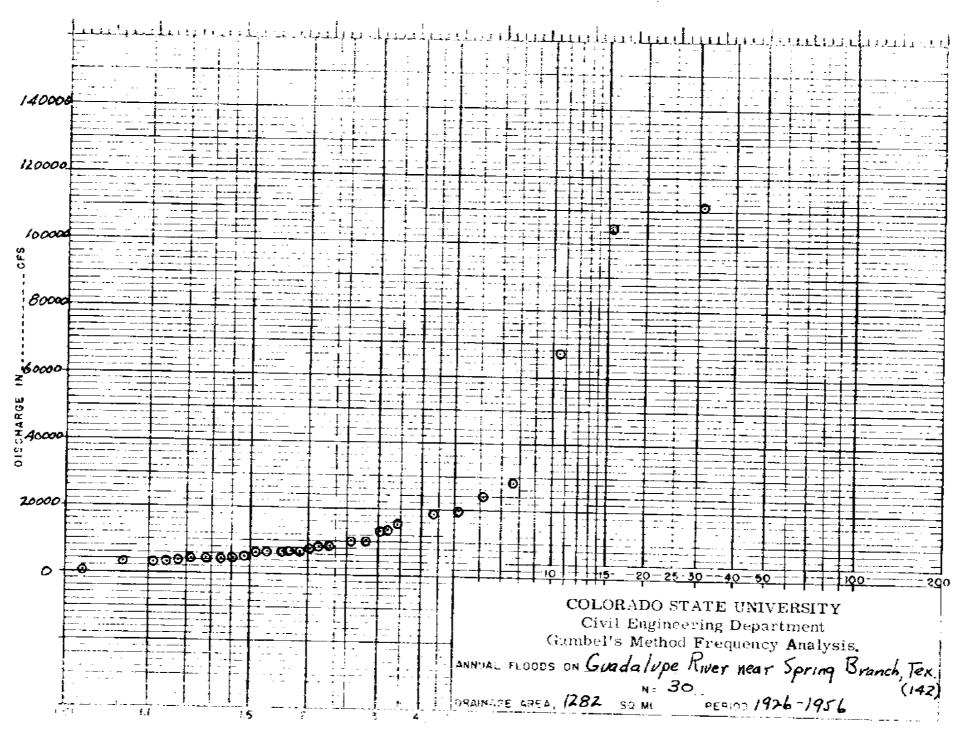


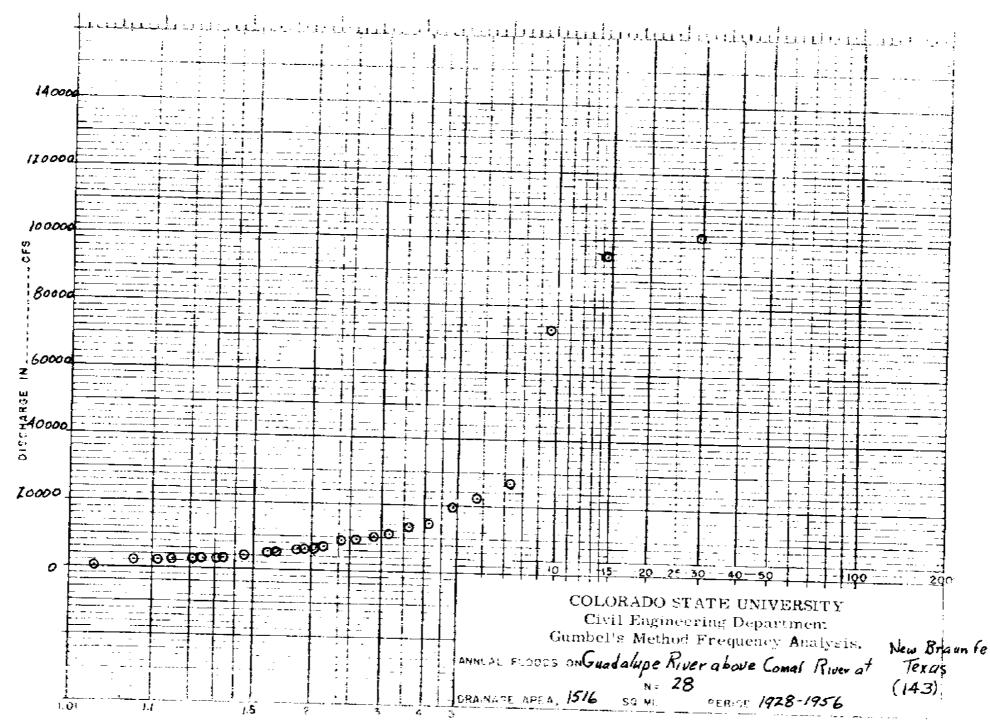


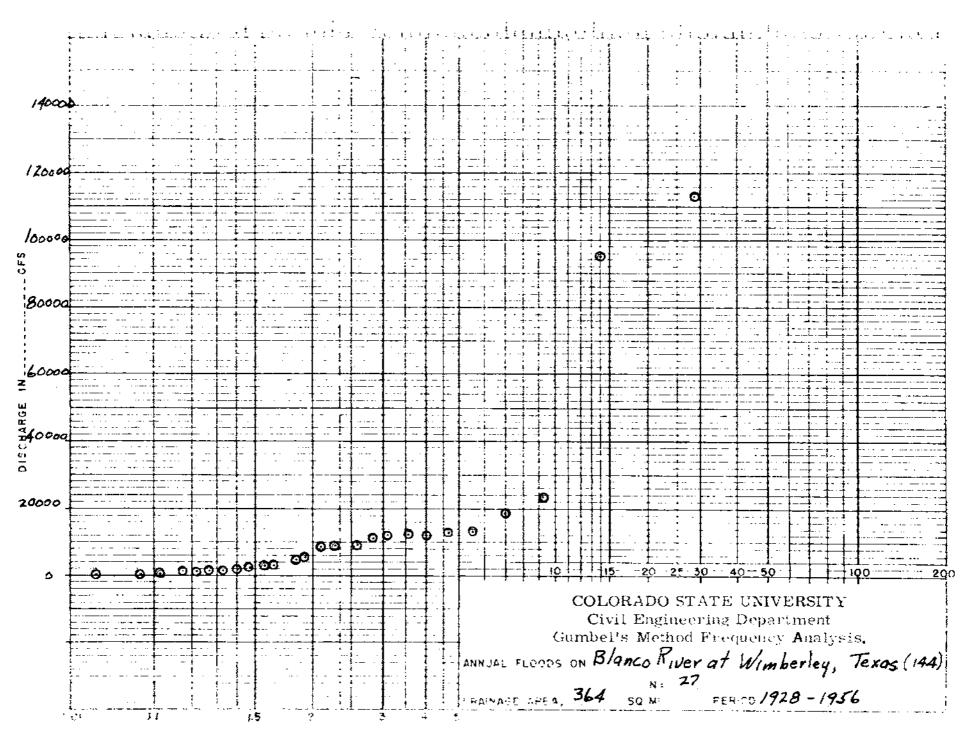


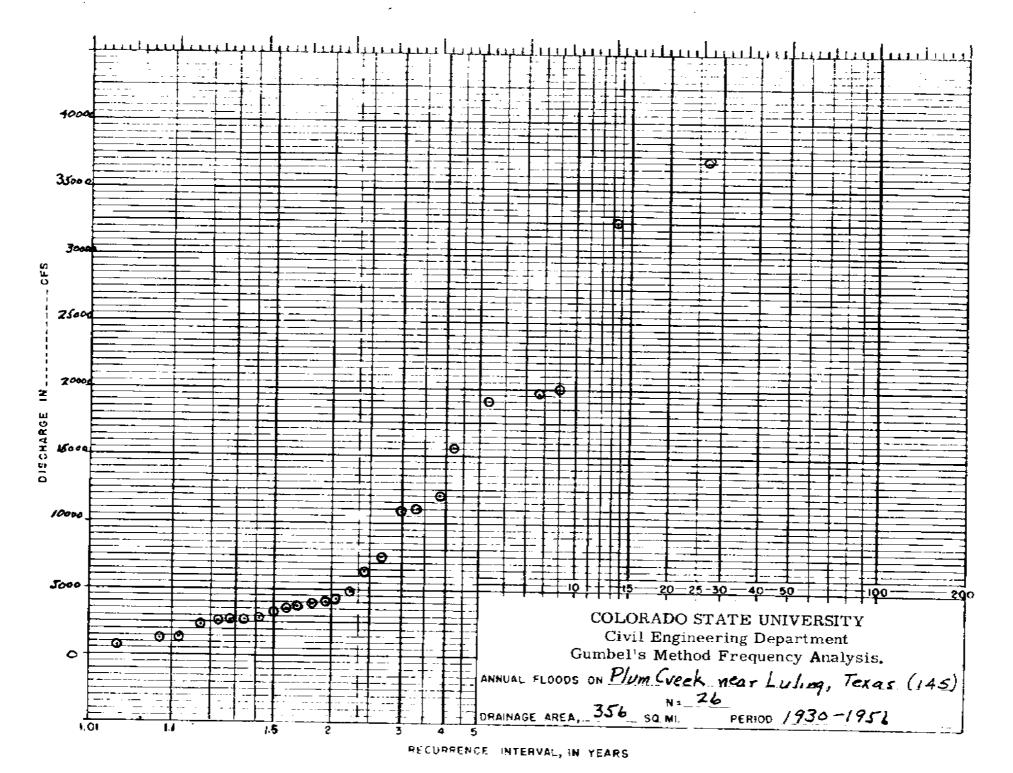


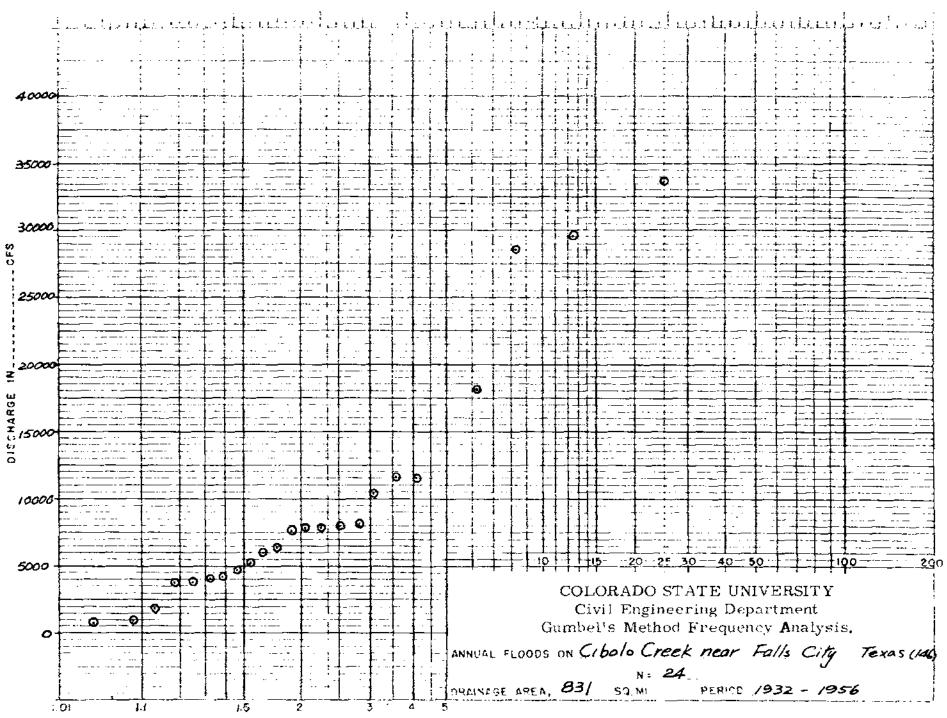


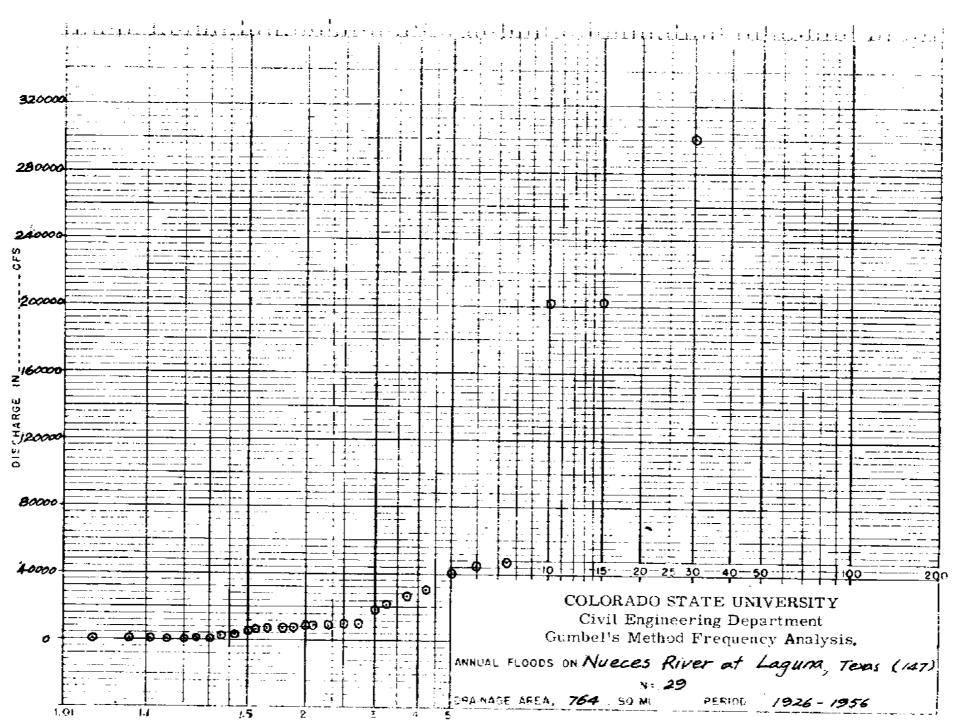


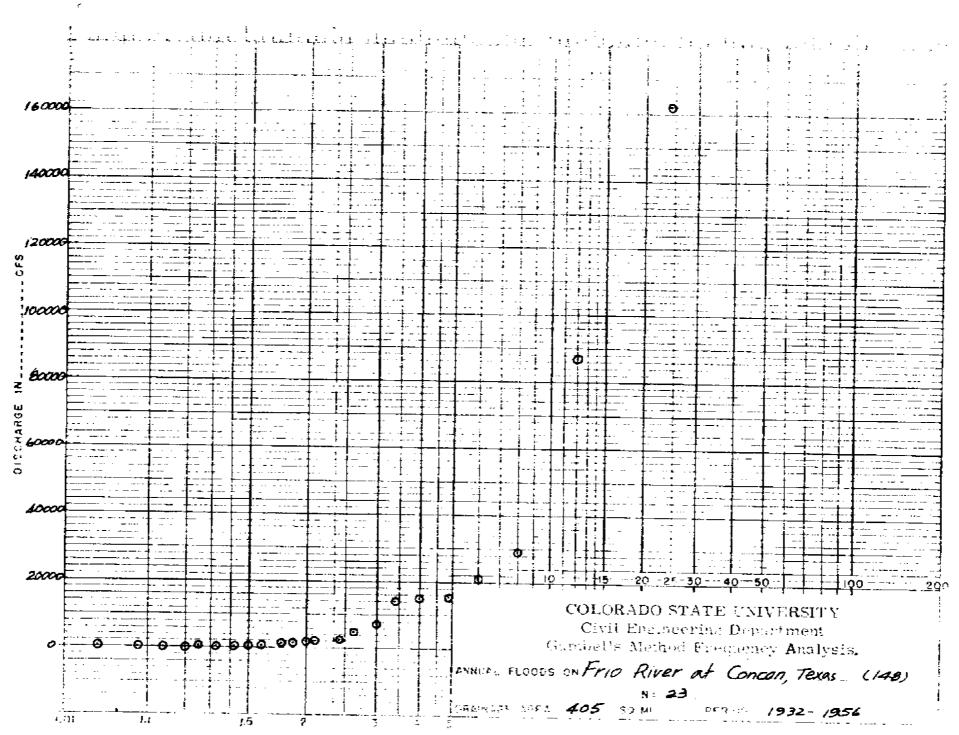


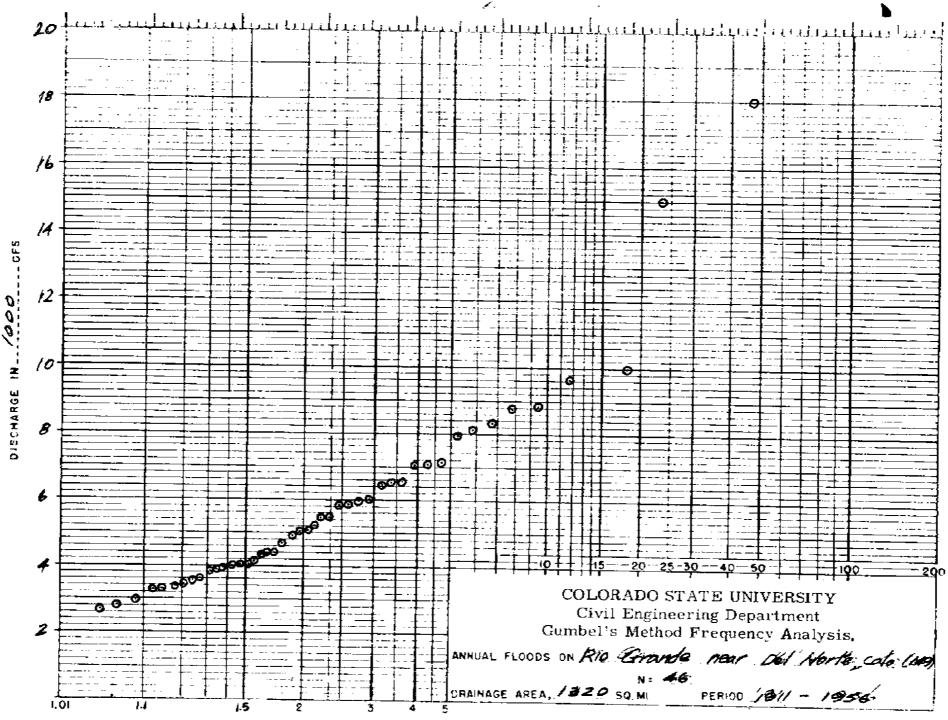


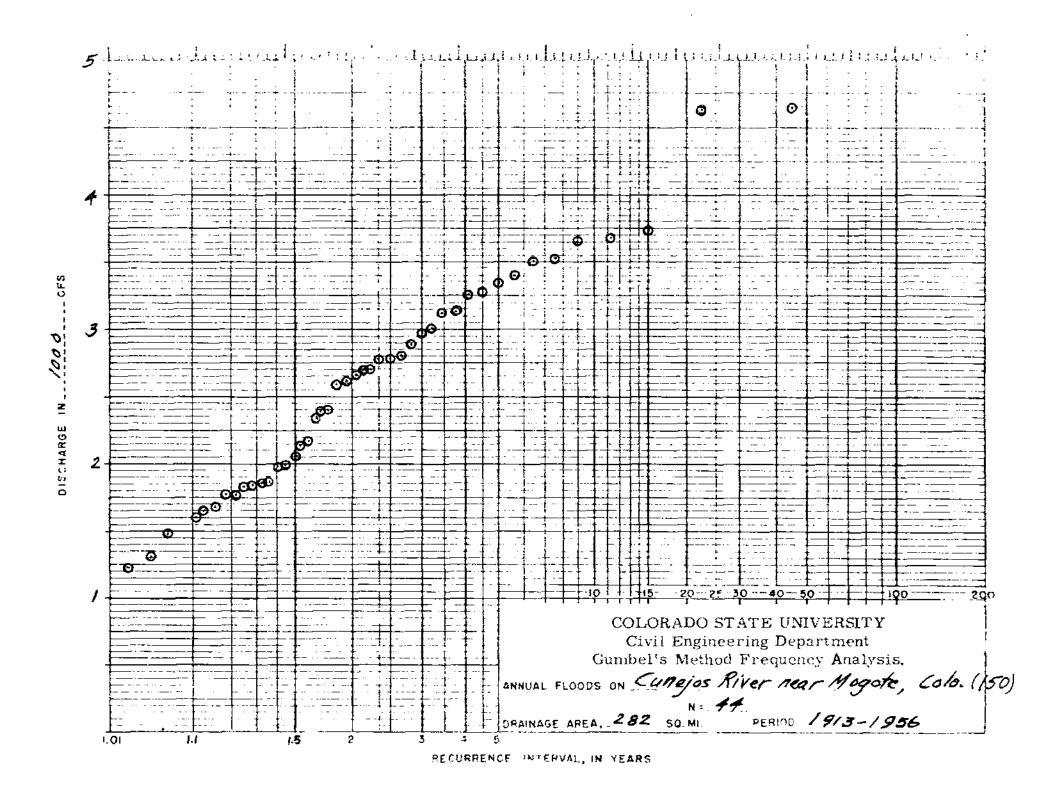


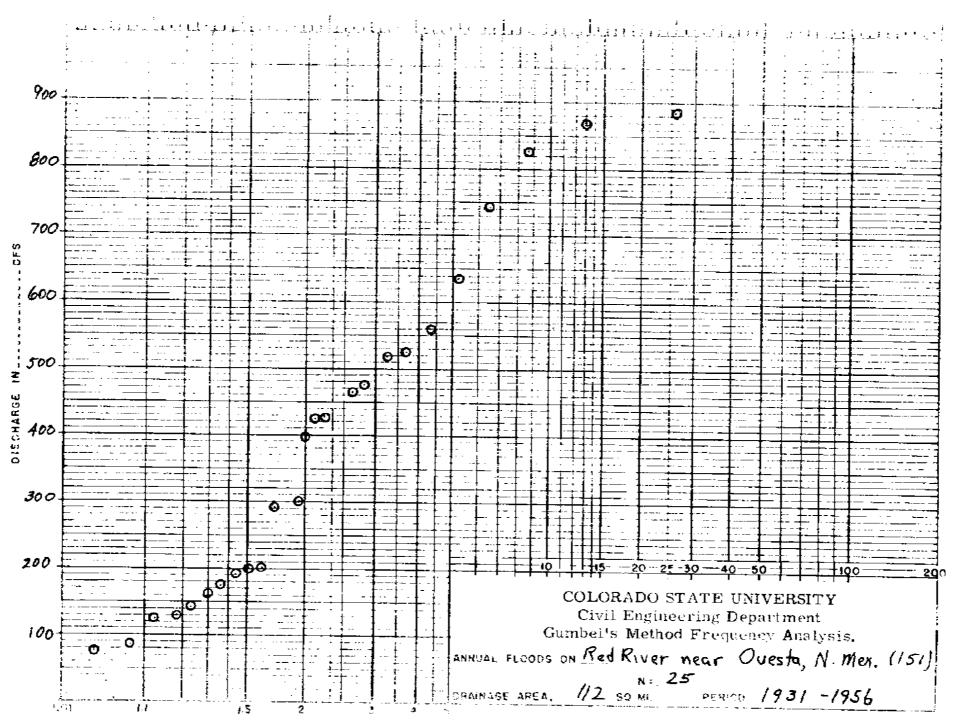


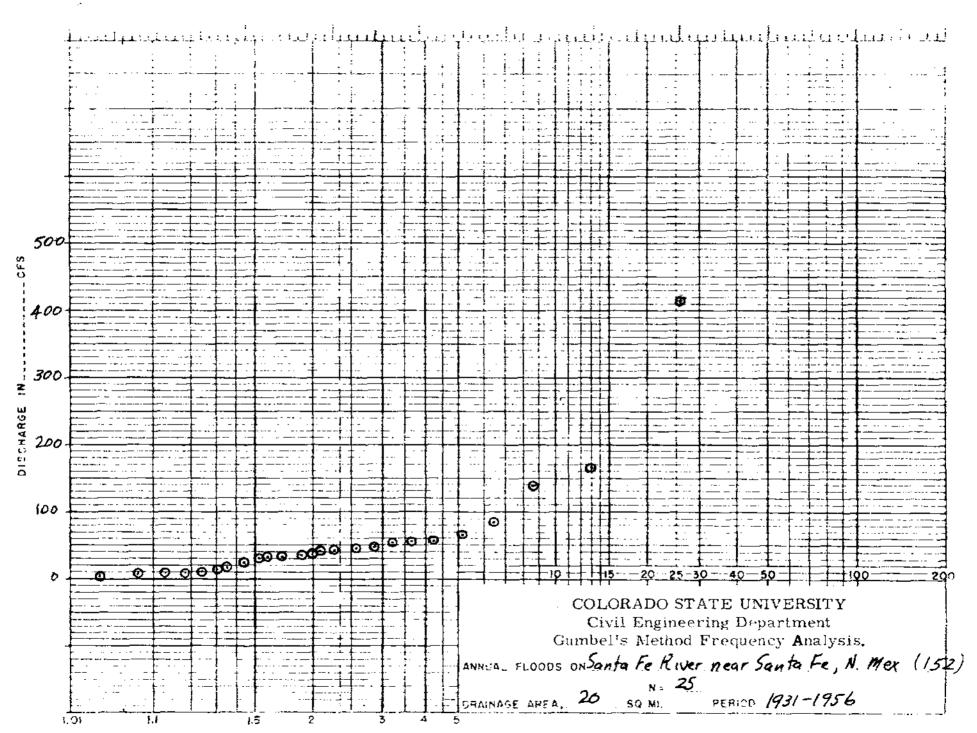


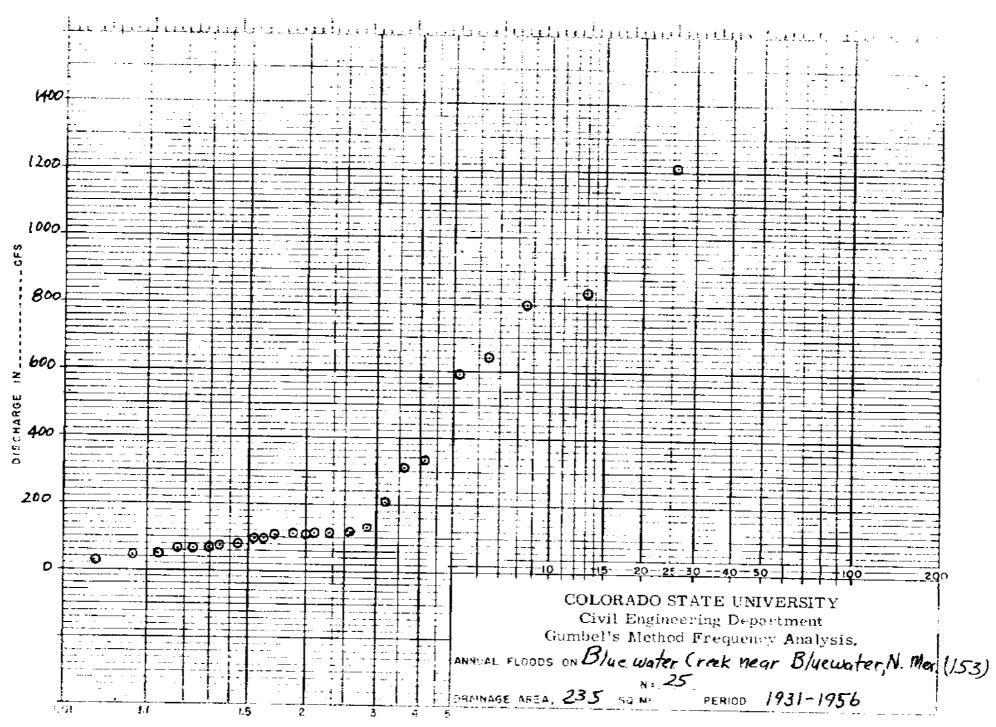


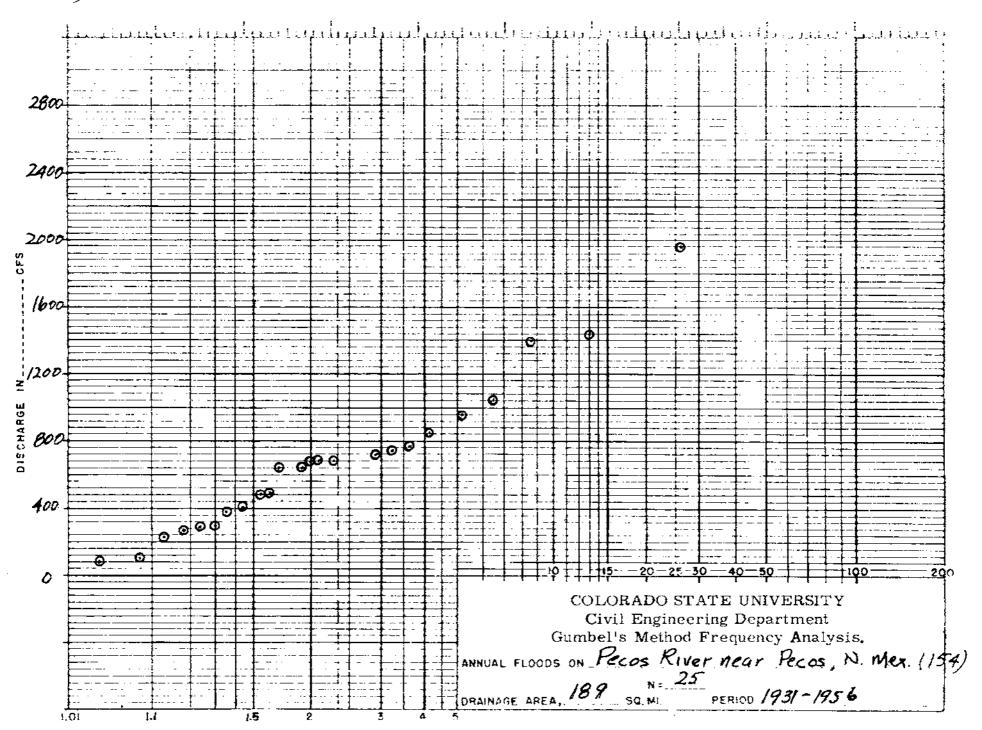


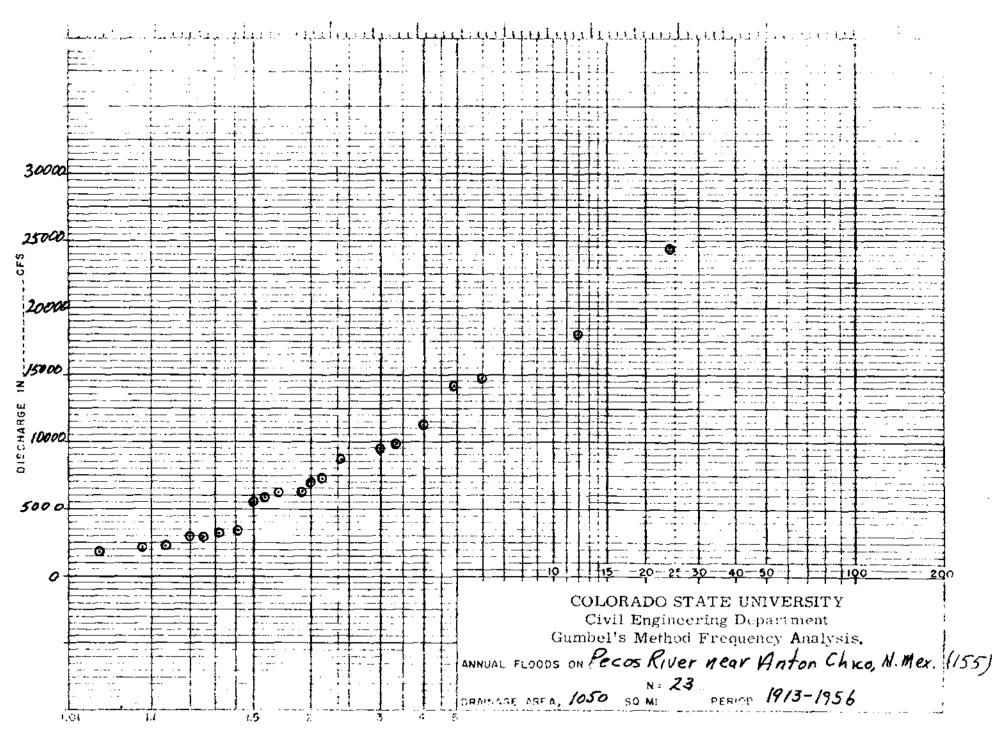


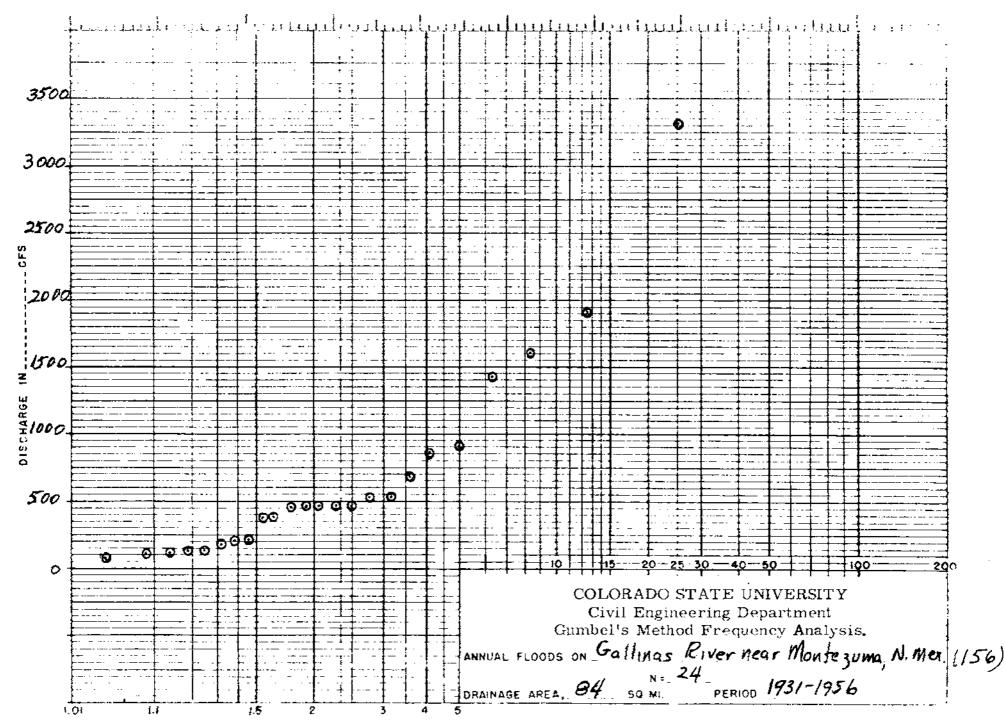


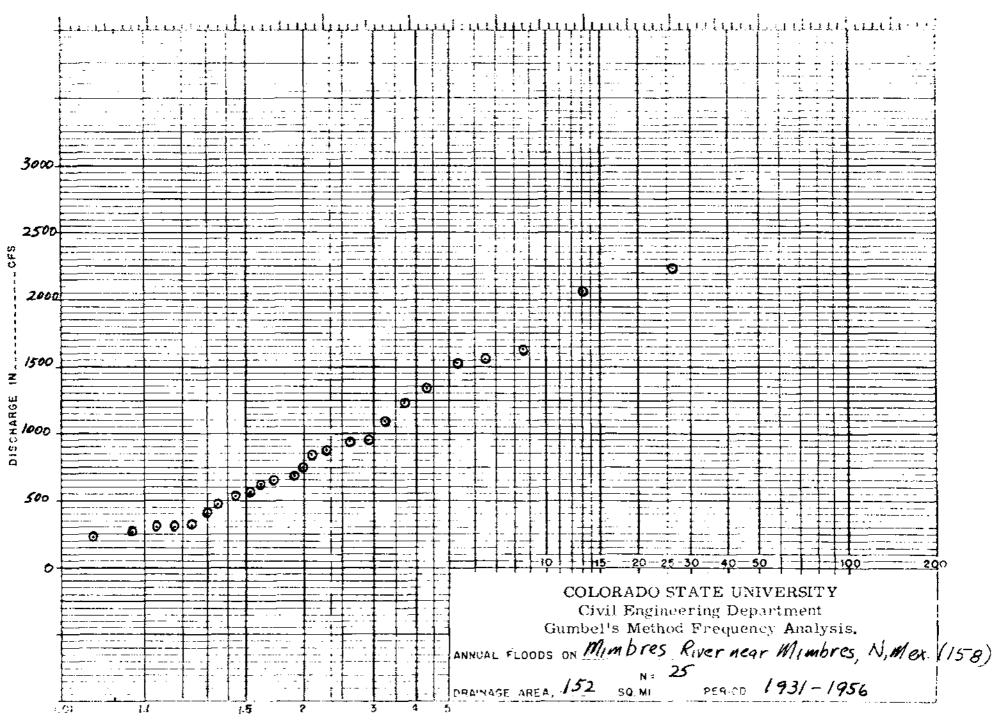












APPENDIX NO. 3.

LIST OF PERSONS CONTACTED

DURING THE STUDY.

APPENDIX NO. 4.

USE OF "RELATIVE WETNESS" PARAMETERS FOR

ESTIMATES OF CLOCK-HOURLY PRECIPITATION IN

EASTERN COLORADO

ABSTRACT

This is a report of an investigation made to determine the interrelations among precipitation amounts for various durations for a given recurrence interval for precipitation records from stations located in parts of eastern Colorado. The concept of "Relative Wetness" is introduced to provide estimates of clock-hourly precipitation amounts for the two-year recurrence interval from precipitation amounts of longer durations for the same recurrence interval. Comparisons are made of the accuracy of these estimates. Preliminary studies show that estimates of clock-hourly precipitation can be made with acceptable accuracy from the records of precipitation amounts of longer duration.

I. INTRODUCTION

The intensity and frequency of precipitation is one of the factors that affect runoff from small watersheds. For small watersheds, particularly in areas where thunderstorm precipitation predominates, short-duration precipitation is of paramount importance.

Records of short-duration rainfall are available in the published records of clock-hourly rainfall, derived from records of recording rain gages. These recording gages are fewer in number than the non-recording gages. In order to provide additional information on short duration rainfall, it would be desirable to utilize the more plentiful data from non-recording gages. In addition to being more plentiful, the non-recording gages offer the advantage that on the average, the records are of longer duration than the recording rain gages.

H. OBJECTIVE

The objective of the study was to evaluate methods for making estimates of clock-hourly precipitation for a given recurrence interval from parameters derived from non-recording rain gages for an area in eastern Colorado.

III. METHOD OF INVESTIGATION

STUDY AREA

The area covered in this study included part of Colorado east of the Continental Divide in the vicinity of Denver. The study area was divided into five subareas, each subarea containing five or more recording gage stations. A map of subareas is given in Fig. 17. In the course of the study, one additional subarea was selected. (Area VI, shown on Fig. 17.)

SOURCE OF PRECIPITATION INFORMATION

Precipitation records were obtained from climatological data published by the U.S. Weather Bureau.

DEFINITION AND COMPUTATION OF "RELATIVE WETNESS" Following are definitions of terms used in this study.

"Relative Wetness" refers to the ratio R_{ij}, d_{ij}, M_{ij}, A_{ij}, defined as follows:

- R_{ij} = 2 year freq. max. hourly precipitation in inches at station i Z year freq. max. hourly precipitation in inches at station j
- d_{ij} = 2 year freq. max. daily precipitation in inches at station i 2 year freq. max. daily precipitation in inches at station j
- M. = 2 year freq. max. monthly precipitation in inches at station i 2 year freq. max. monthly precipitation in inches at station j
- A ij = 2 year freq. max. yearly precipitation in inches at station i 2 year freq. max. yearly precipitation in inches at station j

The subscript is an index notation with $i = 1, 2, 3, \ldots, n-1$, and $j = i+1, i+2, \ldots, n$, where n is the total number of stations in the subarea. Ranking of stations was made in such a way that d_{ij} , M_{ij} , and A_{ij} were all less than unity.

From records from selected stations for the period 1948-57, annual maximum and summer seasonal (May - August) maximum precipitation values for each of the above durations were compiled for each station. Using Gumbel plotting positions, the precipitation amounts having a two-year recurrence interval was determined for hourly, daily, monthly and annual values. An example of the Gumbel plots is shown in Fig. 18.

Stations within each subarea were then ranked suitably to yield values of d_{ij} , M_{ij} , $A_{ij} \leq 1.0$, and values of R_{ij} , d_{ij} , M_{ij} , and A_{ij} , were then computed. Note that the subscript "s" refers to seasonal values (May - August) and the subscript "a" refers to annual values.

IV. RESULTS

٠.

RELATIONS AMONG
$$R_{ij}$$
 , d_{ij} , M_{ij} , and A_{ij} .

R_{ij} vs d_{ij}, R_{ij} vs M_{ij}, and R_{ij} vs A_{ij} were plotted both for annual and seasonal values for each subarea. Assuming an equation of the form y = mx, the best fit line was computed by the method of least squares. A distribution of error chart was prepared for each plot. These error charts were prepared to give a measure of the dispersion of the data from the best-fit curves. Results were presumed to be of acceptable accuracy if 67 per cent of the data fell within -25 per cent of the fitted regression line. Figures 19 and 20 illustrate these procedures.

Examination of 25 plots comparable to Fig. 20 revealed that only 3 had error greater than -25 per cent for 67 per cent of the plotted points. This indicates generally acceptable accuracy for the technique. To further delimit the dispersion of the data from the best-fit lines, the areas between the ordinate and the distribution-of-error curve for 0-67 per cent of the sample were determined. Using these areas as a measure of dispersion, the various combinations of relative wentess ratios were arranged in order of increasing error as shown in Table 1. Numbers shown in Table 1 indicate planimeter readings.

TABLE 1. RANKING OF RELATIVE WETNESS RATIO IN ORDER OF INCREASING ERROR OF ESTIMATE **OF R										
Area	l (lea	st error)	2		3		4		5 (Mos	t error)
I	Rs/ds	39.5	Ra/Ma	41.0	Rs/Ms	43.0	Ra/da*	43.0	Ra/A	57.0
п	Rs/Ms	18.0	Rs/ds	26.0	Ra/da	29.0	Ra/Ma	39.5	Ra/A	45.0
ш	Ra/A	5.0	Ra/Ma	10.0	Rs/Ms	23.0	Ra/da	28.0	Rs/ds	30.0
IV	Rs/Ms	10.0	Rs/ds	13.5	Ra/da	19.0	Ra/A	23.0	Ra/Ma	32.0
V	Rs/ds	30.0	Ra/A*	37.5	Rs/Ms	50.0	Ra/Ma	*56. 0	Ra/da	63.0

^{**}Based on areas between ordinate and distribution of error curve.

for 0-67 per cent of the sample.

It will also be noted from Table 1 that for four out of five cases, the best estimates resulted from use of seasonal data. Conversely, for four out of five cases, the worst estimates were associated with use of annual data.

Since R_{ij} vs d_{ij} (seasonal basis) showed least error for subareas I and V, and gave accuracy better than 67 per cent of the sample within ±25 per cent for each of the other three areas, R_{ij} vs d_{ij} was chosen as the relation for use throughout all five subareas, plus an additional subarea - No. VI. The relation between d_{ij} and R_{ij} for each of the subareas I through VI. is shown in Fig. 21.

15

^{*}Indicates greater error than 67 per cent of the sample within +25 per cent.

V. CONCLUSIONS

- 1. Summer season (May through August) and annual precipitation values can serve as suitable parameters for making estimates of clock-hourly precipitation amounts having a two-year recurrence interval.
- 2. Seasonal precipitation parameters give slightly more accurate estimates of clock-hourly precipitation than annual parameters in the areas studied.
- 3. Procedures for making estimates of clock-hourly precipitation described herein apply to two-year recurrence interval values only.
- 4. Techniques for obtaining a clock-hourly precipitation estimate from relative wetness data should be applicable for any point within the subareas.

VI. ILLUSTRATION OF THE USE OF RELATIVE WETNESS PARAMETERS FOR ESTIMATING CLOCK-HOURLY PRECIPITATION.

In this section a description is given of the procedure to be followed in estimating the clock-hourly precipitation amount (two-year recurrence interval) from the 24 hour precipitation amount having a two-year recurrence interval. The procedure illustrated is applicable to any of the areas shown in Fig. 17, using data for the appropriate areas as given in Fig. 21.

For this illustration the stations of Nunn and Greeley will be used. Both of these stations have clock-hourly precipitation records so a check can be made of the accuracy of estimating the two-year clock-hourly rainfall at Nunn from the 24-hour precipitation amounts at Nunn and the clock-hourly and 24-hour precipitation amounts at Greeley.

The procedure is described in the following steps:

1. Select station "j"; (in this case Nunn) a station for which daily rainfall records are available, and for which an estimate of the two-year

clock-hourly precipitation amount is desired.

- 2. From examination of the rainfall records in published climatological data, determine the maximum 24-hour precipitation amount that fell during each summer season (May through August) for each of the last ten years at Nunn.
 - 3. Plot these data on Gumbel frequency paper.
- 4. From such a plot (not illustrated here) the 24-hour amount having a two-year recurrence interval was determined to be 1.21 inches. The value is $d_{24i} = 1.21$ inches.
- 5. Select station "i"; (in this case Greeley) the nearest station having clock-hourly data available.
- 6. Repeat steps 2, 3, and 4 above for station "i", using both clock-hourly and 24-hour rainfall data. These data are plotted in Fig. 18, from which the values having a 2-year recurrence interval were determined to be
 - a. Clock-hourly amount, P_{60i} = .42 inch.
 - b. Daily (24-hour) amount, d_{24i} = .95 inch.
 - 7. Form the ratio d_{ij} , such that $d_{ij} \leq 1.0$

$$d_{ij} = \frac{d_{24i}}{d_{24i}} = \frac{.95}{1.21} = .785$$

- 8. Enter Fig. 21 with $d_{ij} = 0.785$ and estimate $R_{ij} = \frac{P_{60i}}{P_{60j}} = 0.725$, using the curve for area F.
- 9. Compute $P_{60j} = P_{60i}/0.725$, using the value of $P_{60i} = 0.42$ as determined in step six; $P_{60j} = \frac{0.42}{0.725}$ $P_{60j} = 0.58$ inch.

This is the estimate of the clock-hourly precipitation amount having a two-year recurrence interval at station j, Nunn, Colorado.

- 10. The actual value of P_{60j} as taken from a Gumbel plot (not reproduced here) was 0.69 inch.
- 11. Error of estimate = deviation x 100 = .69-.58 x 100 = 16 per estimate .69

VII. COMMENTS REGARDING USE OF "RELATIVE WETNESS" RELATIONS FOR ESTIMATES OF RUNOFF

The precipitation parameters that were used in attempts to explain differences in runoff characteristics (Page 5 of main report) consisted of 24-hour rainfall amounts having a 2-year and a 5-year recurrence interval. Sparcity of data from recording rain gages prevented use of comparable data for the 60-minute duration.

The "Relative Wetness" study described herein indicates that sufficient correlation exists between 60-minute and 24-hour rainfall amounts that the latter provides an acceptable estimate of the former. This principle is being used to develop a precipitation map of the study area for future use in runoff studies. Sufficient time was not available for completion of this study at the time of preparation of this report.

APPENDIX NO. 5.

CORRELATION OF CPS-9 RADAR ECHO INTENSITY

WITH CLOCK-HOURLY PRECIPITATION

CORRELATION OF CPS-9 RADAR ECHO INTENSITY WITH CLOCK-HOURLY PRECIPITATION

OBJECTIVE:

The objective of the study was to correlate clock-hourly precipitation amounts with CPS-9 radar echo intensity data reconstructed from original records of CPS-9 data from Lowry Air Force Base, Denver, Colorado, from 3-15 June 1958.

PROCEDURE:

The positions of radar echoes in relation to 45 recording rain gage. sites were determined by superimposing a plastic map of rain gage location over the reconstructed sketches of CPS-9 data. The presence or absence of an echo over the locations of the rain gaging stations were recorded, along with the type of echo. These data were then compared with concurrent clock-hourly precipitation data. The clock-hourly precipitation data were divided into four classes, and the radar echo data were divided into six classes, as shown in Table 1.

Table 1.	Clock-h	ourly prec	ipitation ar	id radar echo	classes				
Clock-ho	urly pr	ec ipit ation	classes, i	nch per hour:					
0	0, 0.01 - 0.05, 0.0610, and .11								
		1 / -							
Radar ec	ho clas	s code (Co	nvective e	choes only)					
Intensity:	Low	Medium	Strong						
	1	2	3	Scattered,	≤ .5 coverage				
	4	5	6	Broken	🤛 .6 coverage				
					_				

The number of occurrences of each of the precipitation classes was determined for each echo class. The results are given in Table II.

A similar procedure was followed to determine the number of occurrences of each echo class concurrent with each category of clock-hourly precipitation. The results are given in Table III.

RESULTS:

The results of the study are given in Tables II and III. In addition to the total number of occurrences, the fractions of the totals within each class are given as percentages.

DISCUSSION OF RESULTS:

- 1. Table II shows that 1348 out of 1545 radar observations, or 87 per cent of the total, were associated with zero precipitation.
- Nearly half (48 percent) of all echoes were in echo category
 i.e., scattered, strong convective echoes. The next largest group,
 per cent, were in category 6, i.e., broken, strong convective echoes.
- 3. This preponderance of echoes in categories 3 and 6 suggests an operator bias toward designation of echoes as strong, since the categories of low or medium intensity echoes were listed less than 15 per cent of all observations, despite the fact that low intensity precipitation (0.01 0.5 in/hr.) accounted for 71 per cent of all precipitation observations.

CONCLUSIONS:

- 1. The CPS-9 data do not provide a suitable means for determination of rainfall intensity.
- 2. An operator bias toward designation of echoes as strong is indicated.
- 3. The limited sample studied showed approximately 85 per cent of all clock-hourly precipitation having an intensity of 0.10 inch per hour or less.

TABLE II

Table II. Clock-hourly precipitation associated with various radar echo* categories. (Convective precipitation.)

RADAR ECHO CLASSES **

		1	2	3	4	5	6	No	Per Cent
83 83	No precip.	45	87	686	4	45	481	1348	87
lass ur	***	0	0	1	0	0	. 1	2	41
on Cla r hour	.0105	6	7	39	1	4	74	131	8
cípitatio Inch per	.0610	1	0	6	2	o	14	23	1
Precipitation Classes Inch per hour	> .10	1	0	14	0	0	26	41	3
	No:	53	94	746	7	49	596	1545	100
	Per cent:	3	6	48	(1	3	39		

^{*} Radar data from CPS-9 at Lowry Air Force Base Denver, Colorado 3-15 June 1959.

** Echo Glass Code:

Intensity	Low	Medium	Strong	
-	1	2	3 Scattered,	∠ .5 coverage
	4	5	6 Broken,	3.6 coverage

^{***} No record of precipitation distribution.

TABLE_III

Table III. Radar echo amounts and intensities associated with various measured clock-hourly precipitation amounts.

(Convective precipitation.)

CLOCK-HOURLY PRECIPITATION CLASSES Inches Per Hour

		* *	.0105	.0610	>.10	No.	Percent
	No 1	2	138	22	17	179	46
	1	0	6	1	1	8	2
ses*	2	0	7	0	0	7	. 2
Class	3	2	48	10	16	76	20
Echo Classes*	4	0	0	0	0	0	
	5	0	4	0	0	4	1
	6	0	74	17	24	115	30
	No: Perce	4 nt: 1	277 71	50 13	58 15	389	100

*Echo Class Code:

Intensity	Low	Medium	Strong	•	
_	1	2.	3	Scattered,	く.5 coverage
	4	5	6	Broken,	>.6 coverage

^{**} No record of precipitation distribution.

0 No echo.

APPENDIX NO. 6.

WITH CLOCK-HOURLY PRECIPITATION

CORRELATION OF RADAR ECHO INTENSITY WITH CLOCK-HOURLY PRECIPITATION

OBJECTIVE:

The objective of the study was to correlate clock-hourly precipitation amounts with radar echo intensity as determined from handdrawn sketches of the PPI scope of the United Air Lines 5.5 cm radar set at Stapleton Field, Denver, Colorado.

PROCEDURE:

The positions of the radar echoes in relation to recording rain gages were determined by superimposing a plastic map of rain gage location over the hand-drawn sketches of the PPI scope. The presence or absence of echoes and the intensities of echoes at the gage locations were then recorded. Two methods of analysis were used -- in the first method the position of the echo was interpolated between successive hourly (or half-hourly) observations. In the second method, the positions of the echoes were considered at hourly (or half-hourly) intervals only.

The frequency of occurrence of low, moderate, or high echo intensity was tabulated and compared with the concurrent clock-hourly precipitation amounts at each of 45 sites for each of the two methods of analysis described above. Precipitation class limits were established by dividing the maximum hourly precipitation rate by ten. This analysis was performed for ranges of 0-25, 25-50, 50-75, and 75-100 miles.

RESULTS:

- 1. No correlation was indicated between radar echo intensity and precipitation intensity for either of the two methods of analysis.
- 2. Most of the clock-hourly precipitation amounts fell into the lowest precipitation class interval.
- 3. The data indicate that the hand-drawn sketches of the PPI scope cannot be used successfully to indicate clock-hourly precipitation intensity.

DISCUSSION:

The lack of correlation between radar echo intensity and rainfall intensity could be caused by any combination of the following factors:

- 1. Error in drawing the sketches of the PPI scope.
- 2. Non-linearities in the scope presentation.
- Problems in relating the time of the echo to the time
 of the clock-hourly precipitation.
- 4. The problem of evaporation of raindrops between the cloud base and the ground, typical of the high-based clouds of this area.

No attempt was made to assess the relative importance of

each of these factors. Reports from other investigators* indicate that the lack of correlation between point rainfall rates and echo intensity may be a characteristic of present radar equipment.

^{*} Hiser, H. W., H. V. Senn, and L. F. Conover. Rainfall Measurement by Radar using Photographic Integration Techniques. American Geophysical Union Transactions, 39 (6) 1043 - 47, December, 1958.

APPENDIX NO. 7

WATERSHED SIZE ON SEASONAL DISTRIBUTION OF ANNUAL MAXIMUM FLOOD EVENTS.

maximum flood occurrence are later on the plains than in the mountains or the foothills.

A tabulation of the relative time of occurrence for annual flood events as a function of elevation and watershed size is given in Table 1.

TABLE 1.

TABULATION OF DATE OF OCCURRENCE OF
67 PER CENT OF ALL ANNUAL FLOOD
EVENTS AS A FUNCTION OF ELEVATION
AND WATERSHED SIZE.

Elevation Class ft, msl	Area Class sq. mi.	Approximate date of occurrence	Average date
н	Small	21 June	
M	ii.	30 May	11 June
L	11	12 June	
н	Medium	7 June	
M	11	12 June	14 June
L	n	24 July	
н	Large	l July	
M	11	9 June	23 June
L	11	29 June	

H = High elevation range:

7800-11,000 ft. msl

M = Medium elevation, range:

6091-7683 ft. msl

L = Low elevation, range:

2798-6080 ft. msl

Small: 1-127 sq. mi.

Medium: 139-448 sq. mi.

Large: 460-1766 sq. mi.

CONCLUSIONS:

- 1. Examination of Table 1 shows that the average date of 67 per cent of annual maximum floods advances with increase in watershed size.
- 2. For watersheds of 139-448 square miles, the date of 67 per cent of annual maximum flood advances with decrease in elevation.
- 3. For watersheds less than 139 square miles and between 460 and 1766 square miles, the date of occurrence of 67 per cent of annual maximum floods advances with decreasing elevation below 7683 feet msl.
- 4. The dates of flood occurrence are later on the plains than in the mountain areas and foothills. This can be interpreted in terms of summertime rains as a cause of flood events on the plains, as compared to snow melt, or a combination of snow melt and rain as a cause of flood events in the mountain areas.

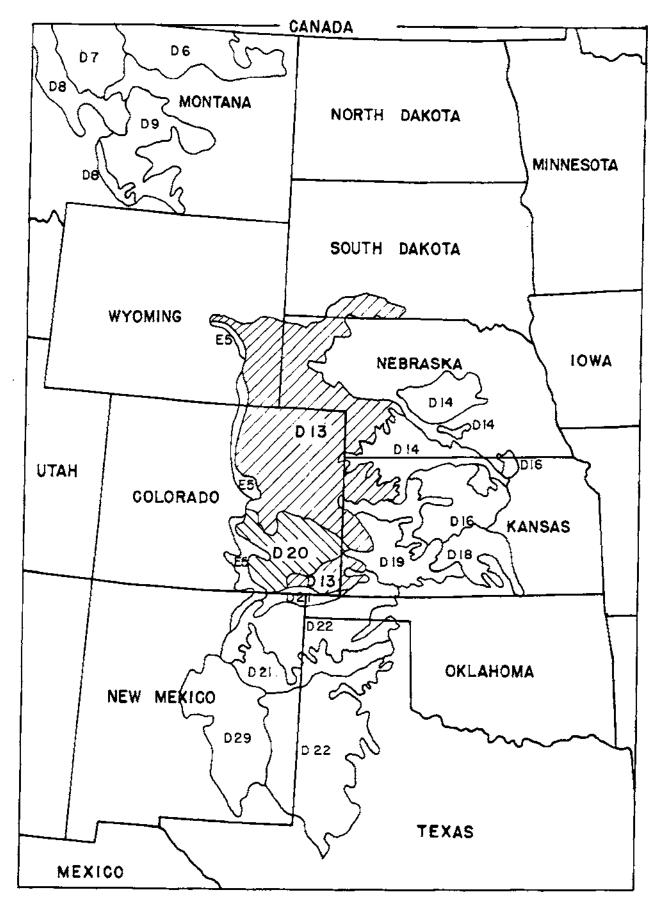


Fig. 1. Location of Study Area.

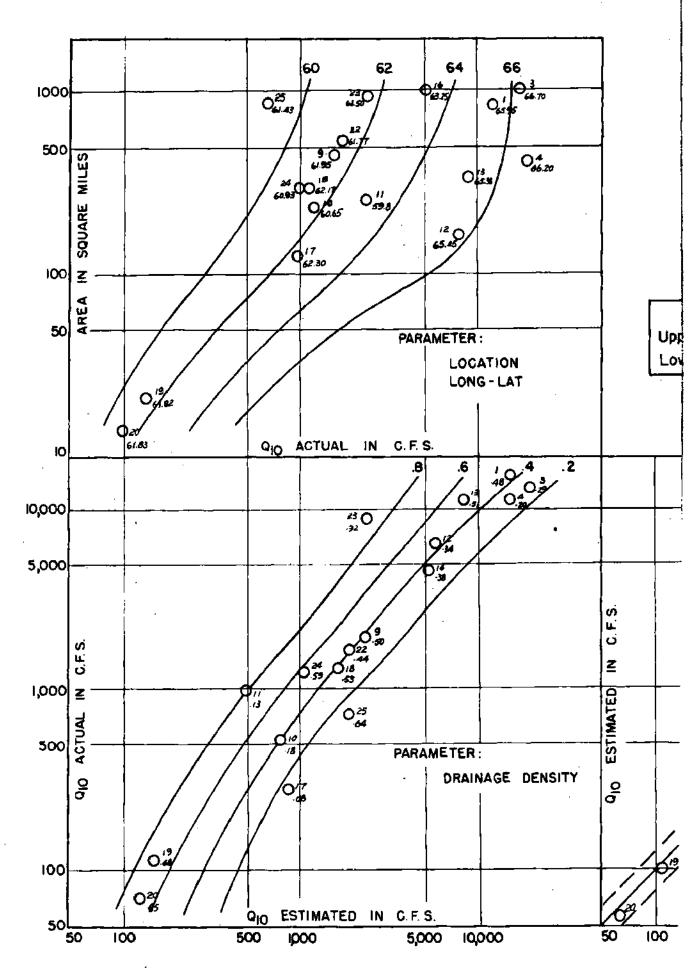
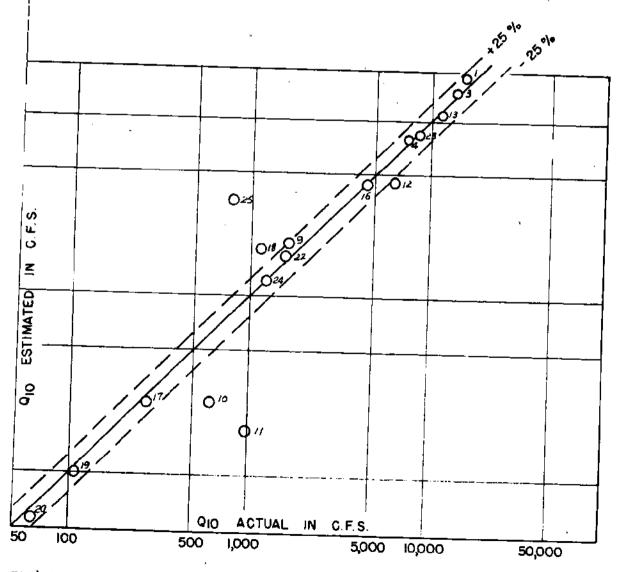


Fig. 2. Preliminary Estimate of Q₁₀ from Physiographic Parameters.

Note:

Upper Figure at Point Station Number. (see Appendix-I)
Lower Figure at Point Value of Parameter.



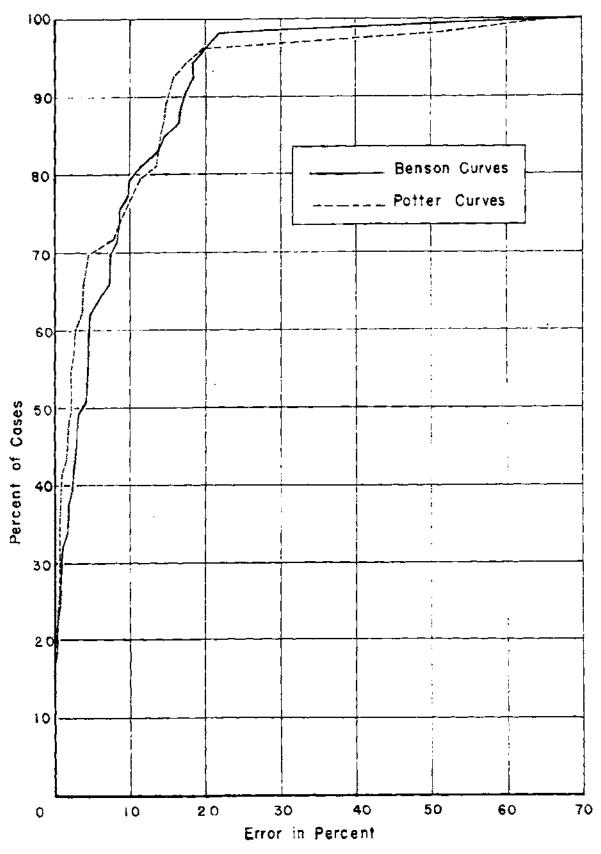


Fig. 3. Distribution of Error Curves Showing Departure of Plotted Points for Recurrence Intervals Greater Than 10 Years From "Benson" and "Potter" Type Curves on Gumbel Frequency Paper.

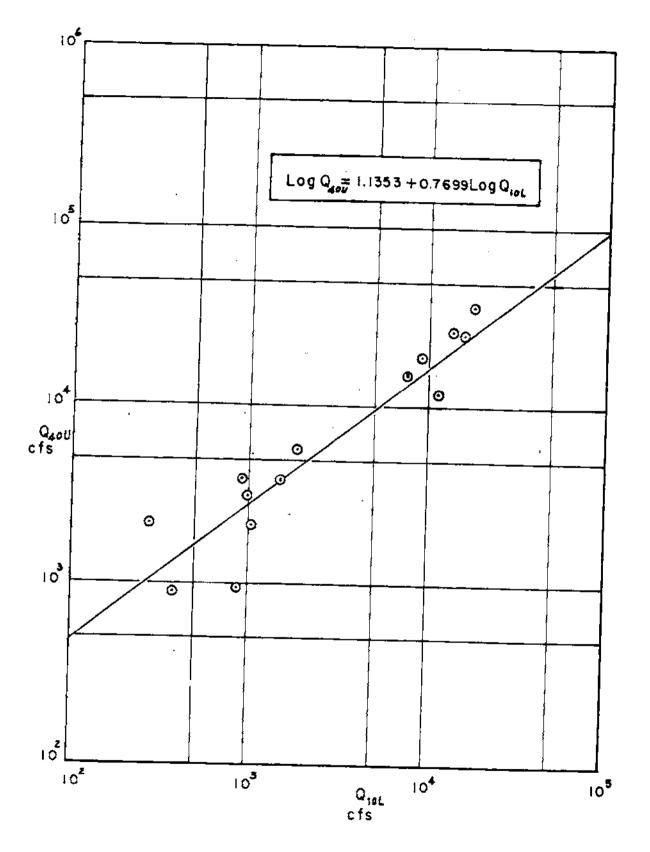
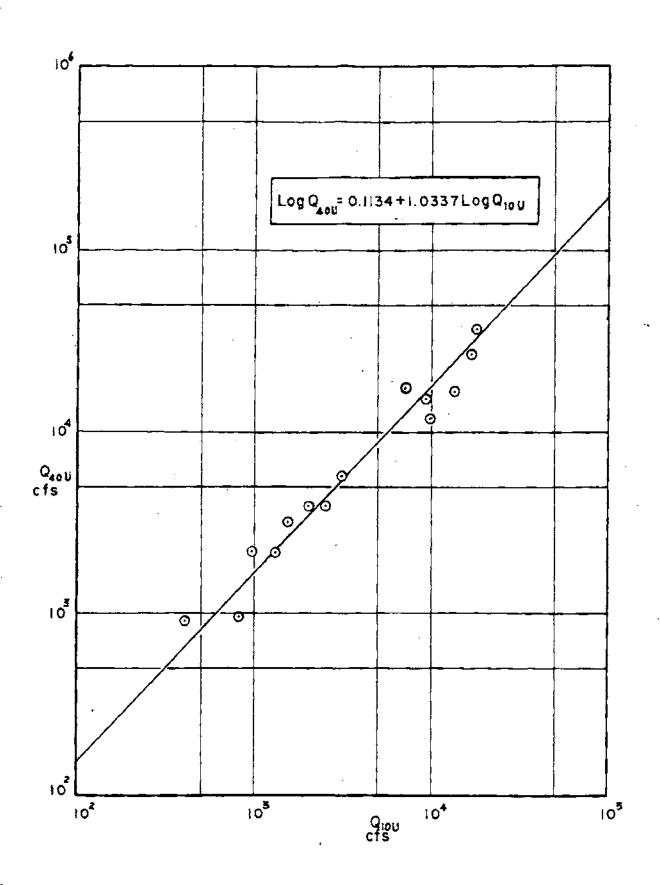


Fig. 5. Relation Between Q and Q , Q and Q (Potter method)



or Selected Stations Outside D-13 and D-20 Problem Areas.

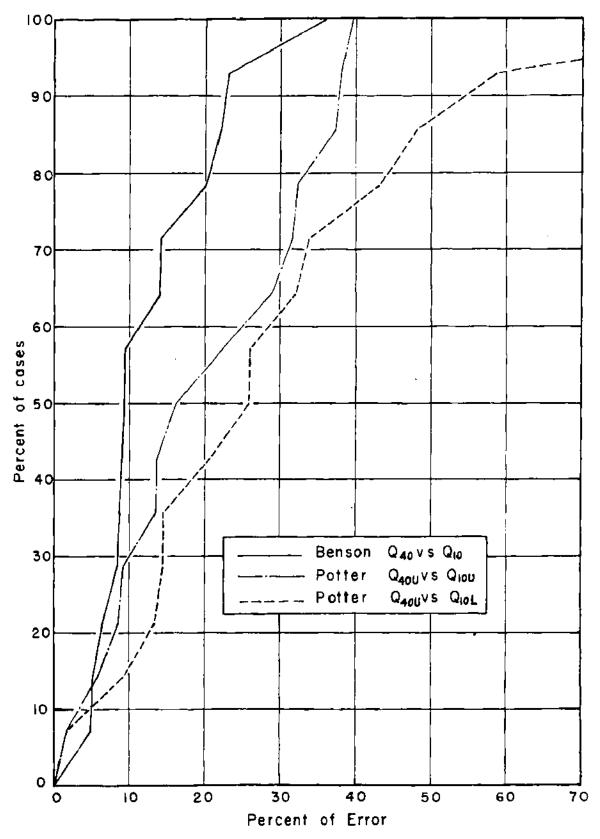


Fig. 6. Distribution of Error Curves for the Relations Shown in Figures 4 and 5.

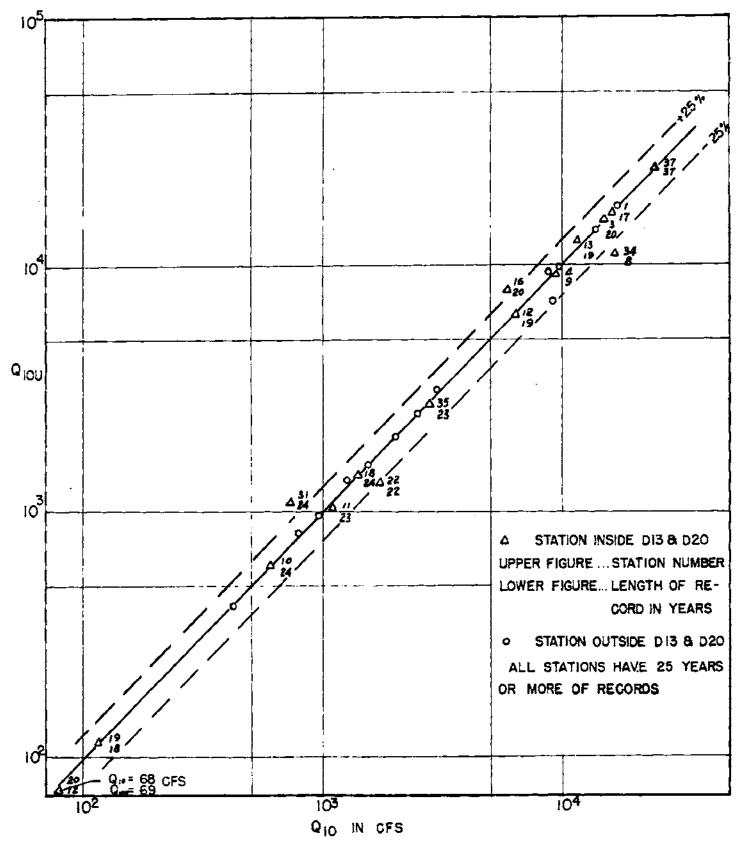


Fig. 8. Comparison of Q_{10} and Q_{10U} for Stations Inside and Outside the D-13 and D-20 Problem Areas.

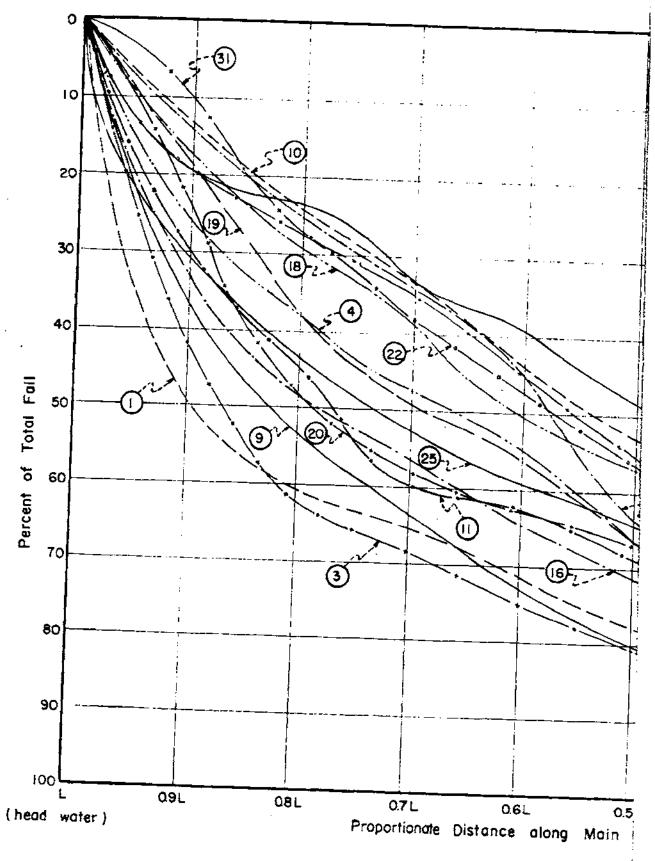
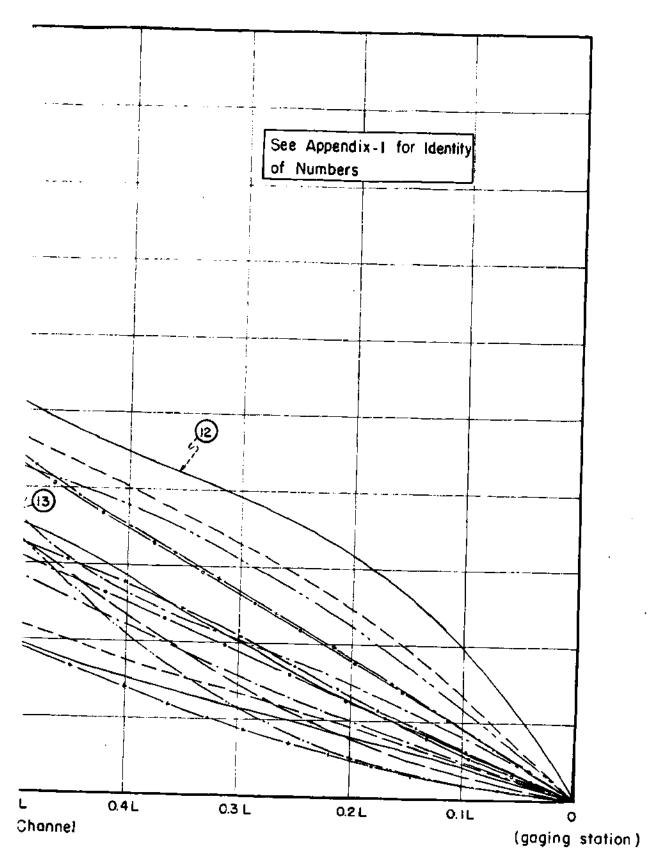
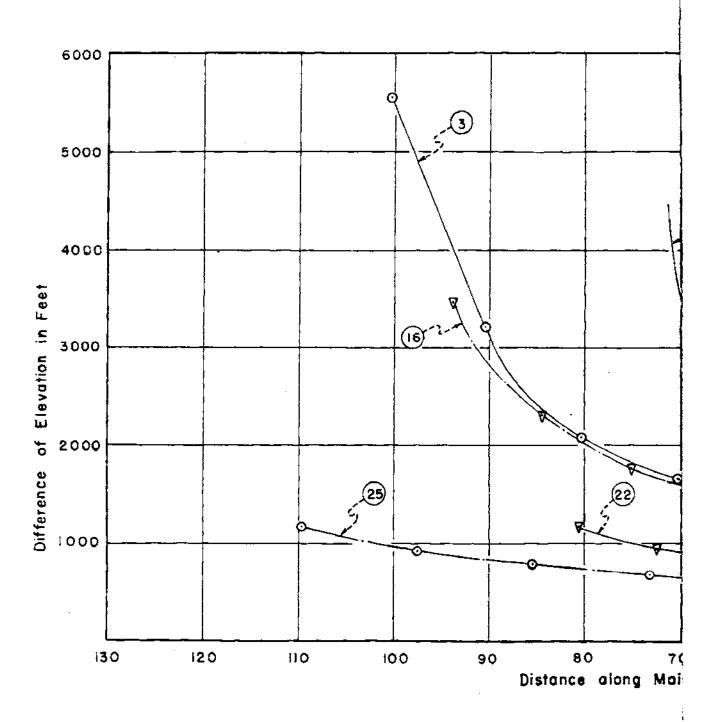


Fig. 9. Dimensionless Profiles of Main Stem of Matersheds Included

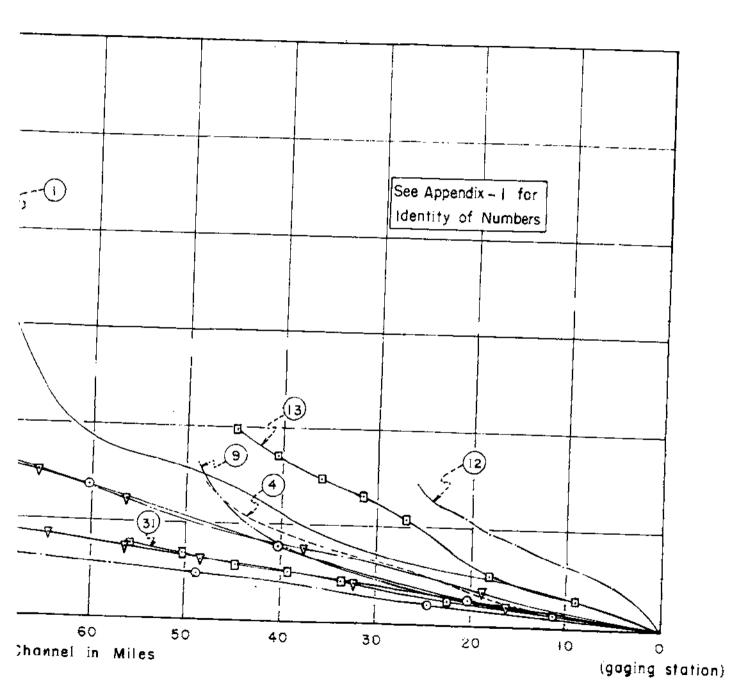


in Craphical Correlation.



٠, -

Fig. 10. Profiles of Main Stem of Watersheds Included in Graphical C



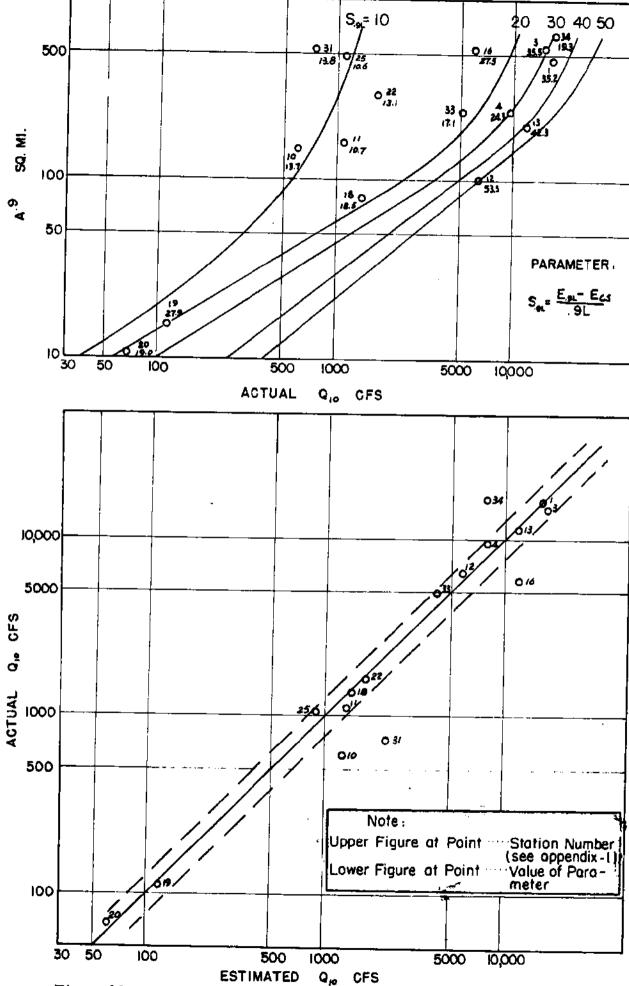


Fig. 11. Coaxial Graph for Estimate of Q₁₀.

Τ.

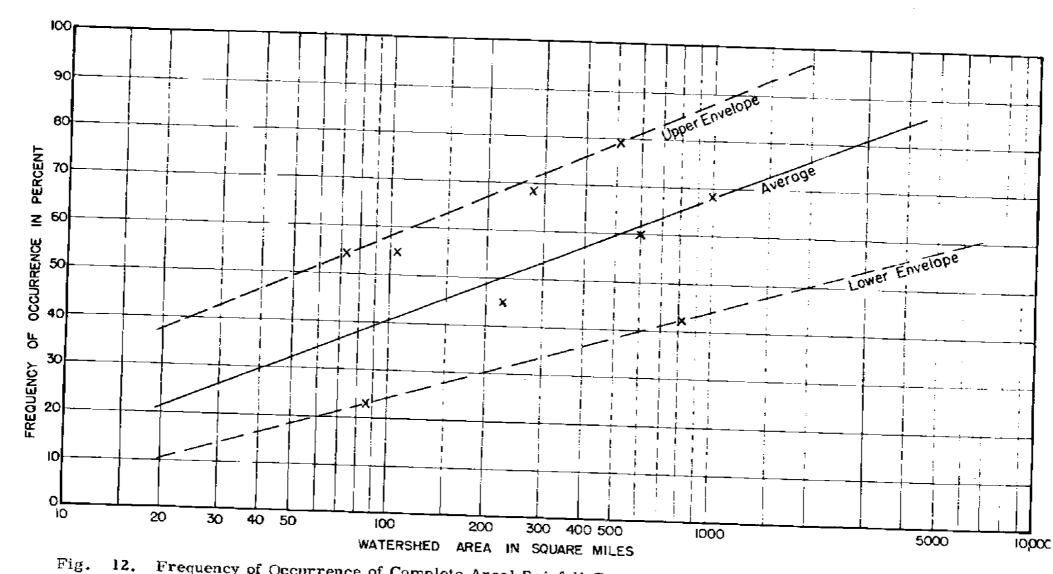


Fig. 12. Frequency of Occurrence of Complete Areal Rainfall Coverage Associated with Annual Maximum Flood Events as a Function of Area for Nine Watersheds in Eastern Colorado.

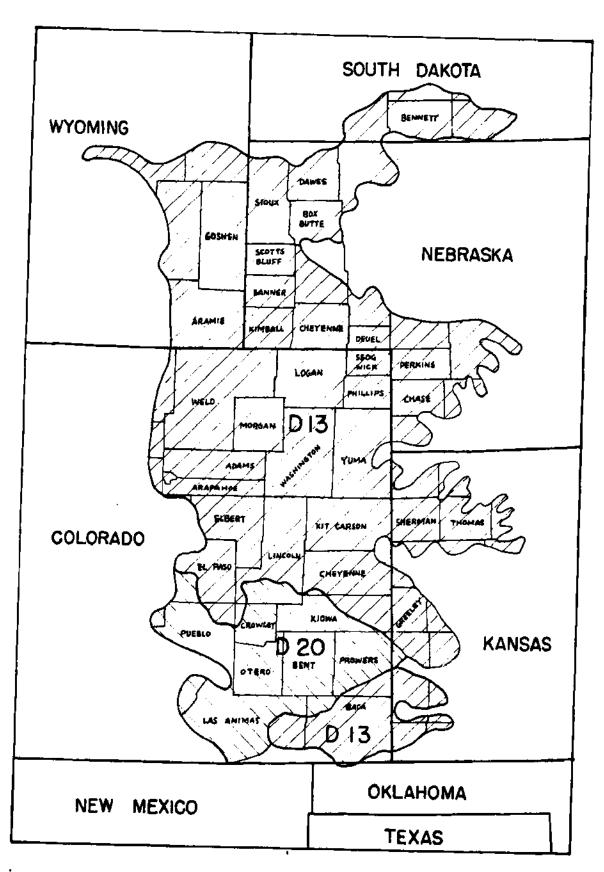


Fig. 13. Location Map.

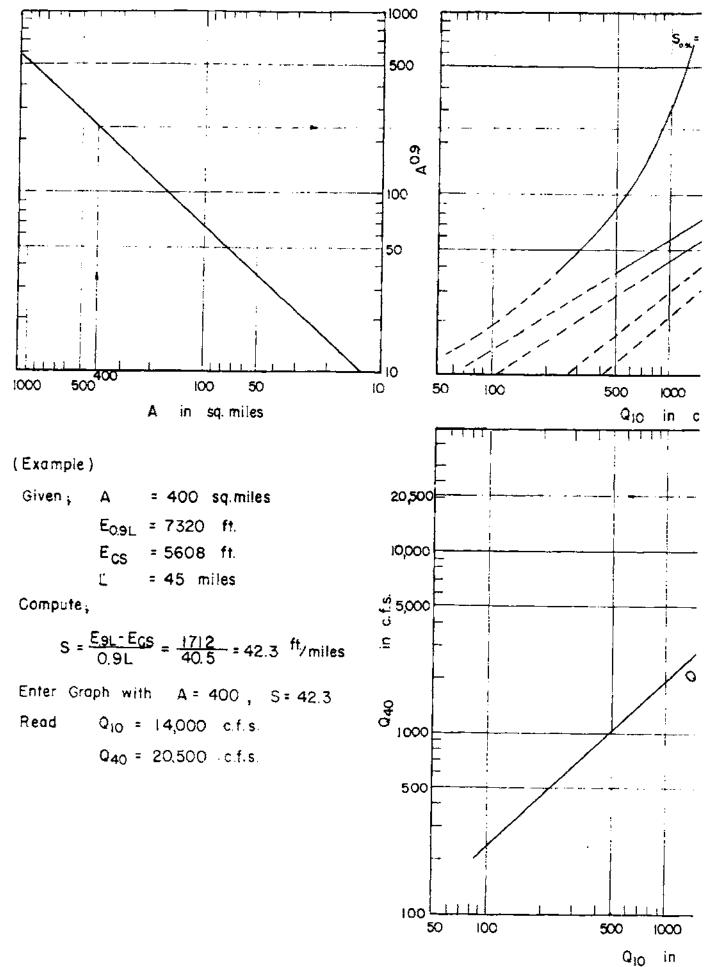
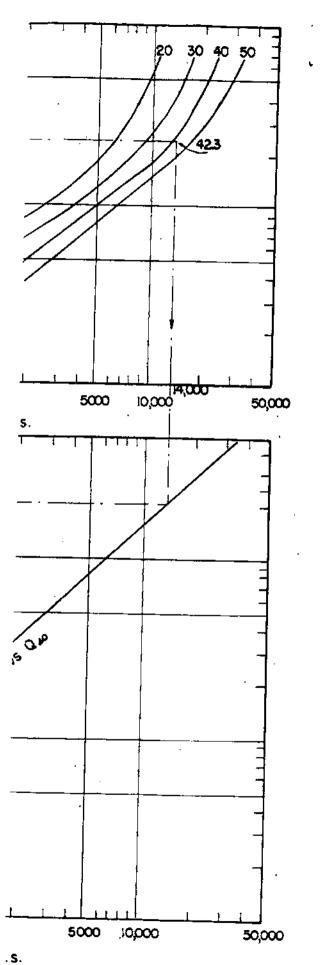


Fig. 14. Relations among Area, Slope Factor, O_{10} and O_{40} for the D-13 and D-2



Problem Areas.

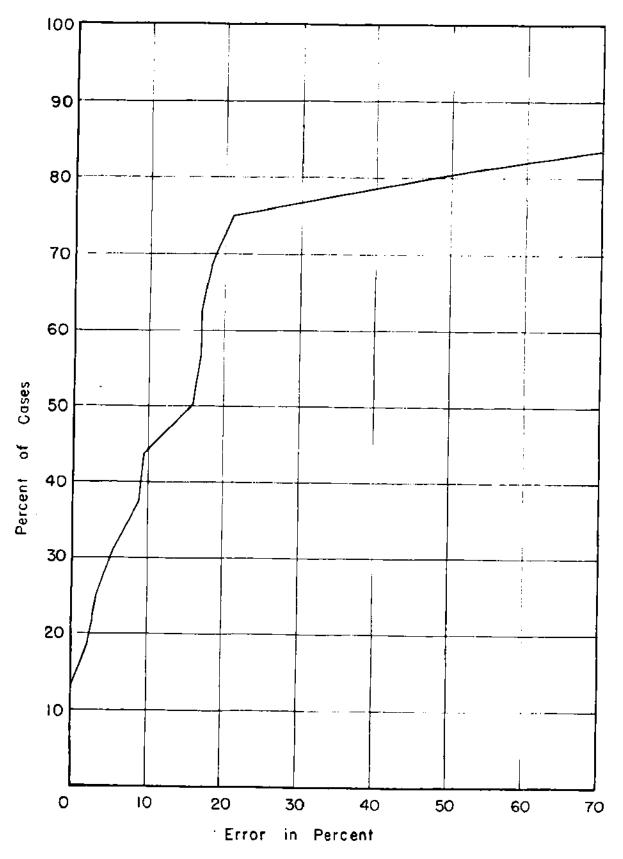


Fig. 15. Distribution of Error Curves for Estimates of Q₁₀ From Fig. 14. (Dependent data.)

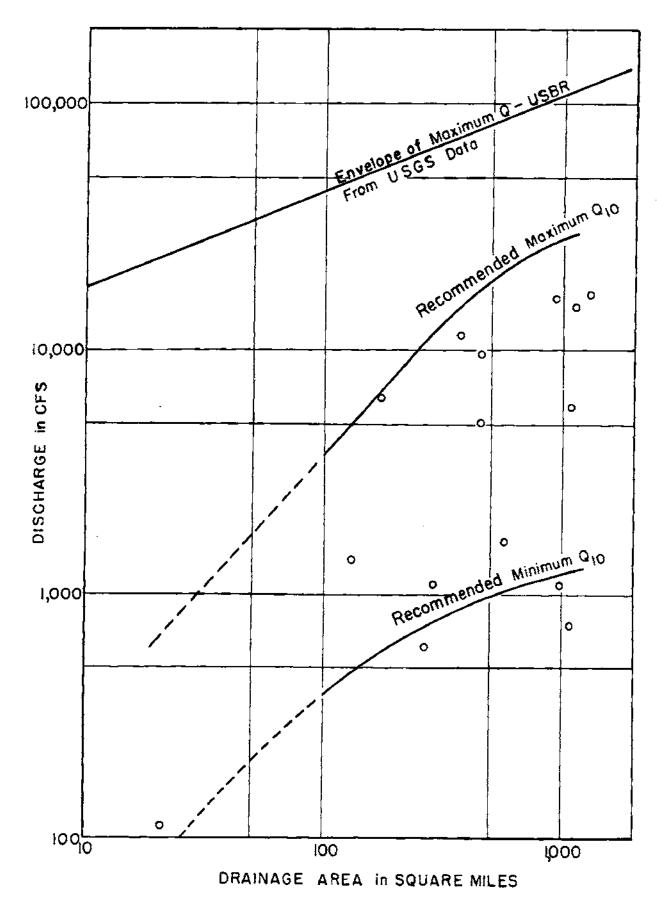


Fig. 16. Recommended Maximum and Minimum Q₁₀ as a Function of Watershed Size.

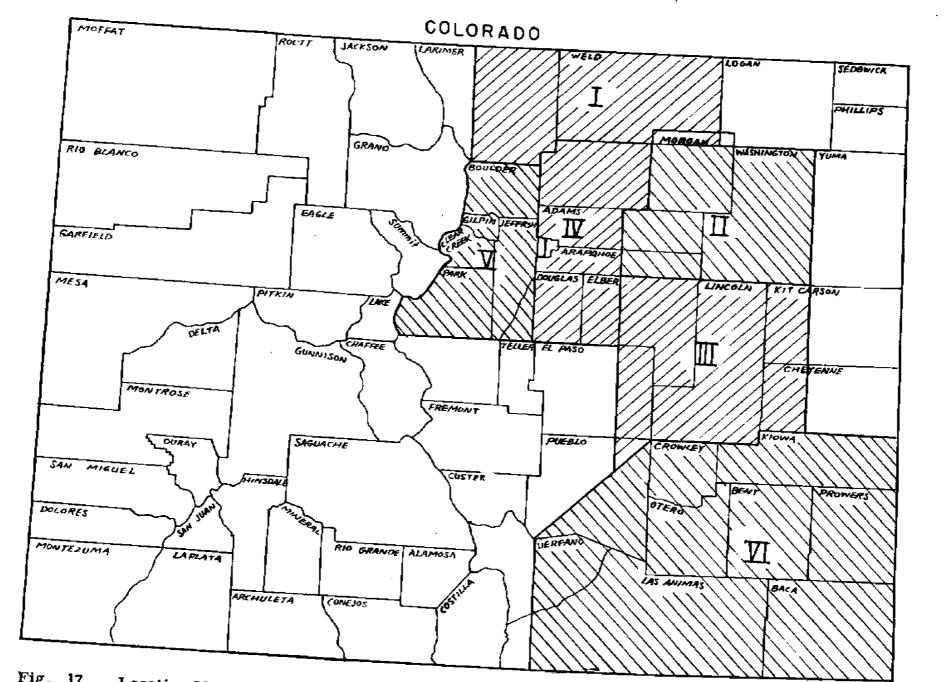


Fig. 17. Location Map for Relative Wetness Study.

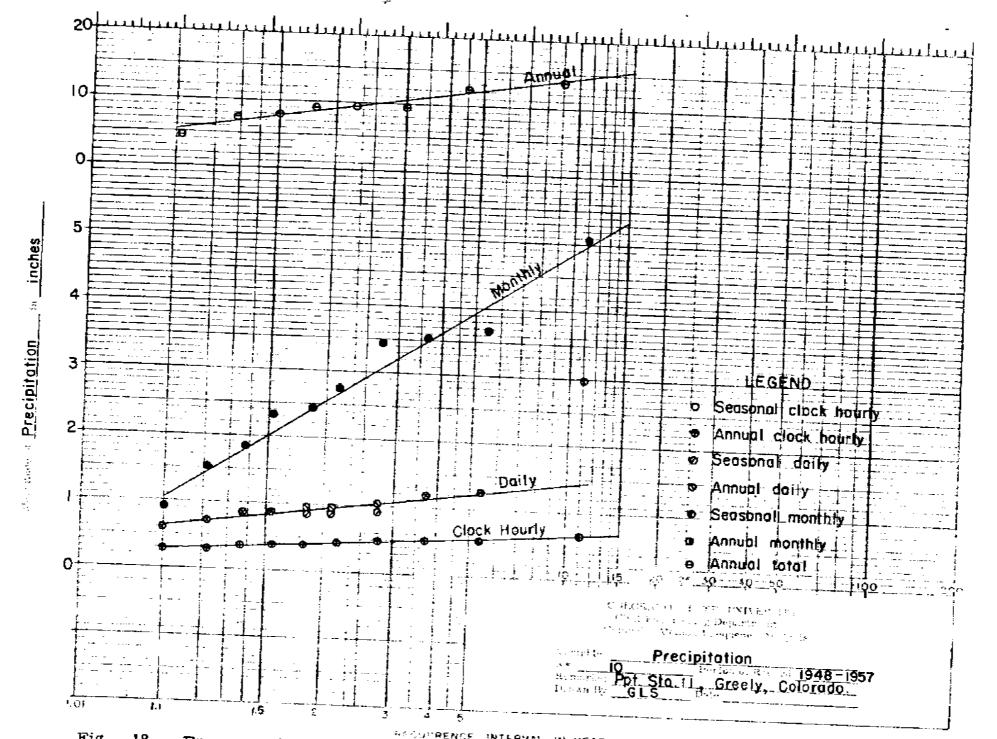


Fig. 18. Frequency Analysis of Precipitation Data at Greeley, Colorado.

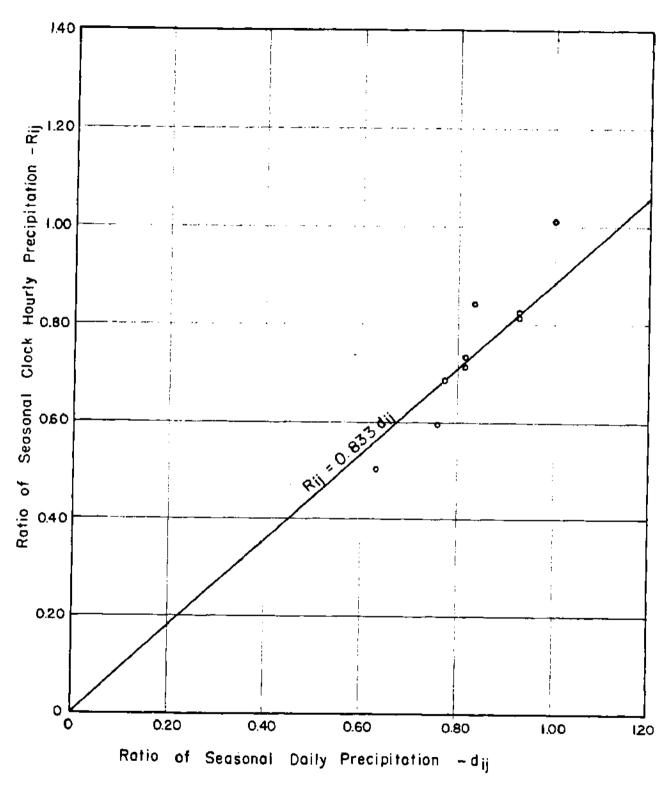
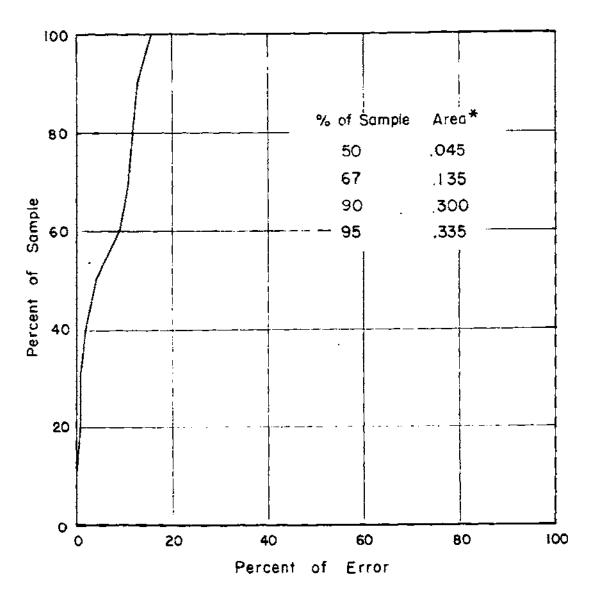


Fig. 19. Relation Between R $_{ij}$ and d $_{ij}$ for Subarea IV.(Seasonal 2 Year Values.)



* Area between Ordinate and Curve

D.

Fig. 20. Distribution of Error Curve for the Relation Shown in Fig. 19.

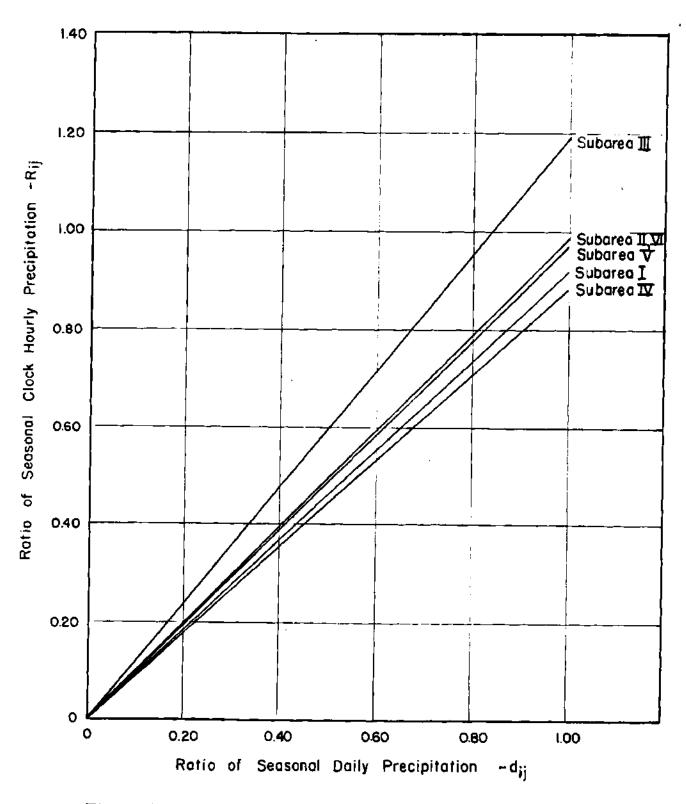


Fig. 21. Relations Between R_{ij} and d_{ij} for Subareas I-IV Shown in Fig. 17. (Seasonal 2 Year Values.)

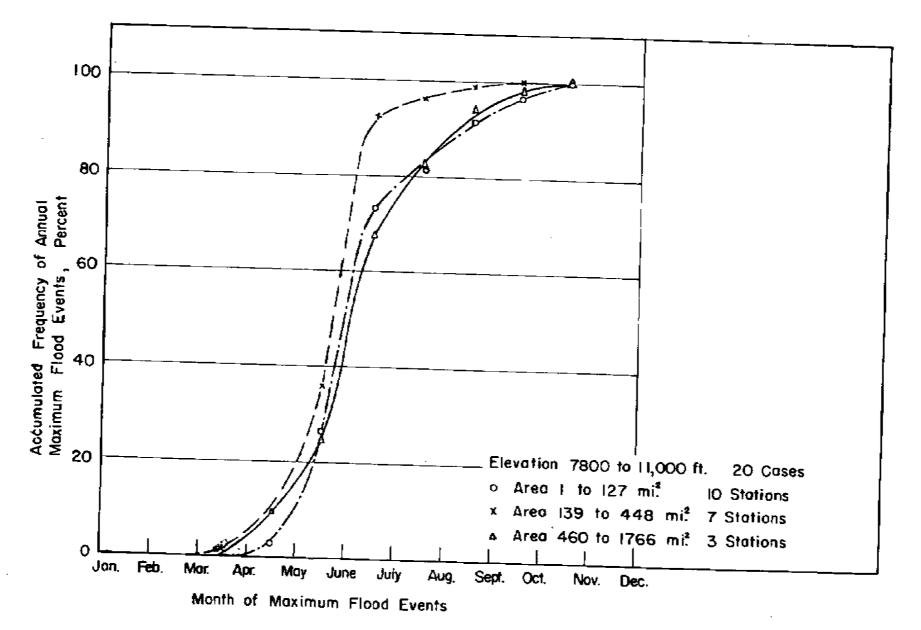


Fig. 22. Accumulated Relative Frequency of Annual Maximum Flood Events for 20 Watersheds in Colorado Between 7800 and 11,000 feet Elevation.

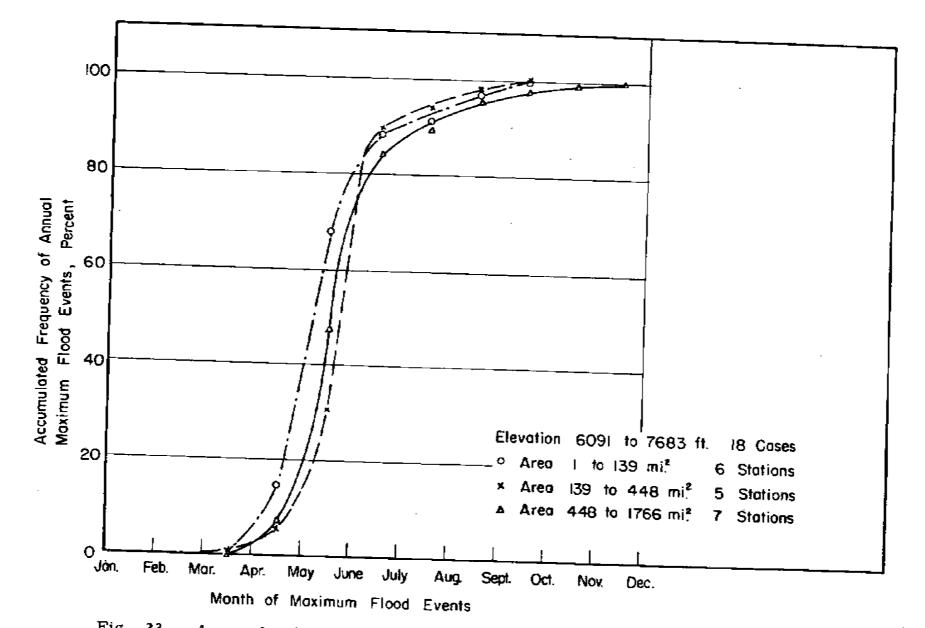
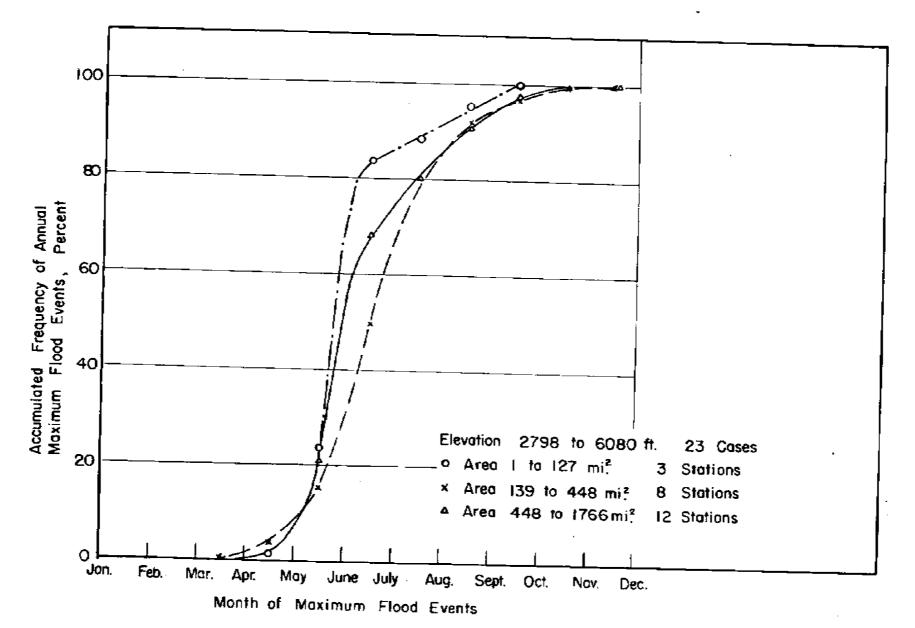


Fig. 23. Accumulated Relative Frequency of Annual Maximum Flood Events for 18 Watersheds In Colorado Between 6091 and 7683 Feet Elevation.



• 3

Fig. 24. Accumulated Relative Frequency of Annual Maximum Flood Events for 23 Watersheds In Colorado Between 2798 and 6080 Feet Elevation.