

# **RESISTANCE TO SHEET FLOW**

**by**

**Bahram Saghafian and Pierre Y. Julien**

**Center for Geosciences**

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**CER88-89BS-PYJ13**

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## RESISTANCE TO SHEET FLOW

### ABSTRACT

The results of a literature review on resistance to sheet flow are presented. The effects of surface roughness, rainfall, and vegetation are considered. At least in the case of laminar flow, it is found that the total flow resistance is the sum of the contributions of individual effects. The friction factor for the surface roughness effect in laminar flow is directly proportional to the relative roughness and varies inversely with the Reynolds number. A power function of rainfall intensity in laminar flow can represent the effect of rainfall on the product of friction factor and Reynolds number. For turbulent flow, however, the friction factor depends on the surface conditions which are partitioned into smooth, transition, and fully rough. The analysis of flow through vegetation is more complex and calls for further studies. For densely vegetated surfaces, the Darcy-Weisbach friction factor is shown to decrease significantly at Reynolds number well beyond the critical value of  $R_e = 2000$  for smooth surfaces. In some cases, the flow behaved as laminar flow at  $R_e = 100,000$ .

## 1. INTRODUCTION

Overland flow on natural watersheds and urban drainages due to excess rainfall is commonly referred to as thin sheet flow. When the rainfall intensity exceeds the infiltration rate of the surface, sheet flow begins; sheet flow is generally unsteady and non uniform. The discharge increases in the downstream direction during the rainstorm and surface runoff rushes down the slope of watersheds, paved roads, side walks, or parking lots in urban areas. After cessation of rainfall, runoff continues during the time in which base flow sources exist; thereafter the recession phase starts. Sheet flows can be dealt with as wide open channel flows except that if the flow is generated by rainfall, excess resistance will be induced by raindrop impact. Shallow flows are more sensitive to raindrop impact because of the reduced flow depth.

The mechanics of sheet flow is of interest for several practical purposes including evaluation of: (1) surface runoff from natural watersheds; (2) soil erosion from watersheds and farmlands; (3) design discharge for urban drainage systems; (4) hydraulic characteristics of shallow flows in border irrigation system; (5) the modeling of overland flow.

In one flow classification, the ratio of the inertia to viscous forces defines the Reynolds number,  $R_e$ . When viscous forces dominate the Reynolds number,  $R_e$  is small and usually thin flow depth exists. This kind of flow is called laminar sheet flow which classifies most of the cases of thin overland runoff. With large Reynolds numbers, the

inertia forces dominate the viscous forces and the flow is turbulent which corresponds to relatively large depths.

The primary parameter in mechanics of sheet flow is resistance to flow which determines other hydraulic variables such as velocity and shear stress. The focus of this paper is confined to the evaluation of the Darcy-Weisbach friction factor for steady laminar and turbulent sheet flows in wide channels under different surface roughness conditions, and with or without rainfall effect. The surface roughness conditions include smooth and rough boundaries in addition to roughness due to vegetation.

## 2. DIMENSIONAL ANALYSIS

The following analysis pertains to the general case of steady sheet flow in a wide channel over a rough boundary through vegetation with rainfall effect. The resistance coefficient, Darcy-Weisbach  $f$ , is then a function of all the relevant variables which describe the channel geometry, roughness, rainfall, flow and fluid characteristics. The variables fall into six categories: (1) channel variables such as bed slope  $S_0$ ; (2) roughness parameters such as boundary roughness height  $k$ , and roughness concentration  $C$ , defined as the ratio of the plan area of roughness elements to the total plane area of the base; (3) rainfall parameters such as rainfall size  $d$ , rainfall pattern  $\alpha$ , raindrop shape coefficient  $\lambda$ , rainfall intensity  $i$ , raindrop velocity entering main flow  $U$ ; (4) flow parameters such as average flow velocity  $V$ , average flow depth  $Y$ , head loss gradient  $S_f$ ; (5) fluid parameters such as fluid density  $\rho$ , specific weight of fluid  $\gamma$ , and dynamic viscosity  $\mu$ ; and (6) vegetation parameters classified into two categories: geometric and physical. Among the geometric characteristics are  $S_y$  = the average vegetation spacing at depth  $y$ ,  $d_y$  = the average diameter or width of the vegetation elements at  $y$ ,  $G_y$  = the average gap size at  $y$ , the pattern dimensionless quantity  $\psi$ , and the cross-sectional shape dimensionless quantity  $\theta$ . The physical characteristic of plants, as adopted by Kouwen and Unny (1973), is the flexural rigidity of the plants shown by  $EI$ . The deflected height of the vegetation,  $K$ , may be regarded as a parameter of the combination of geometric and physical characteristics.

The general form of functional relationship may be shown as follows:

$$\text{Func} (V, Y, S_f, S_o, k, C, d, \alpha, \lambda, i, U, S_y, d_y, G_y, K, \psi, \theta, EI, \rho, \gamma, \mu) = 0 \quad (1)$$

For flows over a rough surface without any effect of rainfall and vegetation, Eq. 1 takes the form:

$$f = \frac{8gYS_f}{V^2} = \text{func}(V, Y, S_o, k, C, \rho, g, \mu) \quad (2)$$

where  $f$ , instead of  $S_f$ , is the dependent variable. By selecting  $V, Y$ , and  $\rho$  as the independent variables and applying the  $\pi$  theorem for constant  $C$  (the maximum value similar to Nikuradse's experiments), one obtains:

$$f = \text{func} (S_o, k/Y, F, R_e) \quad (3)$$

in which  $F$  = Froude number and  $R_e$  = Reynolds number. The effect of Froude number can be dropped for laminar flow.

For boundary shear stress due to flow over a smooth surface with rainfall effect, Eq. 1 reduces to:

$$\tau = \text{func} (V, Y, S_o, d, \alpha, \lambda, U, i, \rho, g, \nu) \quad (4)$$

where  $\tau$  is the boundary shear stress equal to  $\gamma Y S_f$ . Yoon (1970) performed a dimensional analysis to present:

$$\frac{f}{8} = \frac{\tau}{\rho V^2} = \text{func} \left( \frac{VY}{\nu}, \frac{V}{\sqrt{gY}}, S_o, \frac{id}{\nu}, \alpha, \lambda, \frac{iY}{\nu}, \frac{U}{\sqrt{gY}} \right) \quad (5)$$

where  $V.Y/\nu$  and  $V/\sqrt{gY}$  are the conventional Reynolds number and Froude number respectively. Yoon experimentally found that: (1)  $iY/\nu$  and  $U/\sqrt{gY}$  showed a poor correlation with  $f$ ; (2) the effect of  $\alpha$  or rainfall spacing was negligible; (3)  $\lambda$  was kept constant and therefore dropped from the analysis; (4) Froude number appeared to be of secondary importance; and (5)  $id/\nu$  is proportional to  $i$  for constant  $\nu$ .

Therefore, Eq. 5 becomes:

$$f = \text{func} ( R_e, S_o, i ) \quad (6)$$

By applying the  $\pi$  theorem on Eq. 1 for the sheet flow through vegetation with rainfall effect and dropping unimportant terms of rainfall parameters based on the previous discussion, the following form is obtained:

$$\text{func} ( S_f, S_o, \frac{k}{Y}, \frac{id}{\nu}, \frac{S_y}{Y}, \frac{d_y}{Y}, \frac{G_y}{Y}, \psi, \frac{K}{Y}, \theta, \frac{EI}{\rho V^2 Y^4}, \frac{\gamma Y}{\rho V^2}, \frac{\nu}{VY} ) = 0 \quad (7)$$

Chen (1976) used the experimental results of Yoon (1970) and argues that the effect of rainfall would be maximum for flow on the horizontal smooth surface but would decrease with increasing  $k$  and  $S_o$ . He continues that since the roughness of turf surface is very high, the effect of rainfall intensity is believed to be insignificant. Also, the data by Chen (1976), Phelps (1970), and Hartley (1980) show that the flow resistance for flow through vegetation is much higher than that of flow only with rainfall.

After some modifications in Eq. 7 and using the relation  $V_{\max} \cdot G = V \cdot S$ , Hartley (1980) comes up with the following equation:

$$f = \text{func} \left( S_0, \frac{S_y}{Y}, \frac{d_y}{Y}, \frac{G_y}{Y}, \psi, \theta, \frac{K}{[EI/\rho V_*^2]^{1/4}}, \frac{V_{\max} \cdot d}{v}, \frac{V}{\sqrt{gY}} \right) \quad (8)$$

in which  $V_* = \sqrt{gYS_f}$ . The term  $k/y$  in Eq. 7 was dropped by assuming flow through vegetation having smooth boundary. However, the effect of roughness, if considerable compared to vegetation resistance, can be added to the vegetation resistance to yield total resistance.

In case of relatively sparse vegetation all of the terms in Eq. 8 should be considered. For grass with maximum density, however, the flow resistance is mainly due to drag on the roughness elements and concentration, shape, and pattern effects could be dropped from the analysis, as in Chen's study. In case of experiments with artificial cylinders, the restrictions and simplifications made by Hartley include: (1) the density of the system doesn't change with depth, so subscripts of the first three terms after  $S_0$  may be dropped; (2) the effect of pattern and shape will be represented by a constant in the final equations; and (3) flexibility effects can be dropped for the experiments with rigid cylinders. Also for rigid system,  $K = Y$ . Therefore:

$$f = \text{func} \left( S_0, \frac{S}{Y}, \frac{D}{Y}, \frac{G}{Y}, \frac{V_{\max} \cdot G}{v}, \frac{V}{\sqrt{gY}} \right) \quad (9)$$

In case of laminar sheet flow, usually with very shallow depth, the deflected height and flexural rigidity of the vegetation are not

important and Eq. 9 still applies. The Froude number contribution in laminar flow resistance equations has not been included so far. The experiments such as Chen's have been conducted with the attempt to eliminate surface instabilities. However, Hartley reported only small free surface effect even in turbulent flow. Hence, Eq. 9 takes the form of:

$$f = \text{func} ( S_o, S/Y, D/Y, G/Y, V_{\max} \cdot G/v ) \quad (10)$$

in which  $R_e = V_{\max} \cdot G/v = V \cdot S/v$  is the Reynolds number based on vegetation spacing.

### 3. GOVERNING EQUATIONS

One of the most common resistance factors is the Darcy-Weisbach friction factor,  $f$ . The Darcy-Weisbach formula was first developed for flow in pipes in the following form :

$$h_f = f \frac{L}{D} \frac{V^2}{2g} \quad (11)$$

where  $h_f$  = friction loss along length  $L$  of the pipe, given the pipe diameter,  $D$ , and the mean flow velocity,  $V$ . For open channel flow,  $h_f/L$  and  $D$  are substituted by  $S_f$  and  $4Y$  respectively :

$$f = \frac{8gYS_f}{V^2} \quad (12)$$

where  $S_f$  = friction gradient,  $V$  = velocity, and  $Y$  = flow depth equal to hydraulic radius in a wide channel. Eq. 12 may be applied to steady uniform flow in wide channels by substituting  $S_o$  for  $S_f$ . Other friction factors, such as Manning  $n$  and Chezy  $C$ , are mostly used for turbulent flow. The relationship between  $f$ ,  $n$ , and  $C$  in English units is as follows :

$$C = \frac{1.486Y^{1/6}}{n} = \left(\frac{8g}{f}\right)^{1/2} \quad (13)$$

The sheet flow with rainfall as lateral inflow is considered to be a shallow spatially varied flow which with constant rainfall intensity and constant base flow would be steady. The derivation of governing equations for steady spatially varied flow with rainfall has been studied by many investigators; among them, Chow (1959), Woo and Brater

(1962), and Yen and Wenzel (1970). Probably Yen and Wenzel (1970) derived the most comprehensive dynamic equation for this case by both momentum and energy approaches.

The continuity equation for the flow with rainfall in a wide channel can be written as :

$$q = q_0 + ix \quad (14)$$

where  $q$ , and  $q_0$  = total and base flow rates per unit width of the channel at  $x = 0$  . Under the following basic assumptions: (1) one dimensional steady flow; (2) hydrostatic pressure distribution; (3) constant channel slope; (4) constant momentum correction factor along the channel; (5) negligible air entrainment effect; and (6) impervious boundary, Yen and Wenzel (1970) using momentum approach came up with the equation of water surface profile for steady spatially varied flow as follows :

$$\frac{dY}{dx} \left( \cos \theta - \frac{\beta V^2}{gD} \right) = S_o - S_f + \frac{i}{gA} (U \cos \phi - 2\beta V) \quad (15)$$

where  $x$ =distance in the flow direction,  $D= A/T$ = hydraulic depth at  $x$ ,  $A$ = cross section area at  $x$ ,  $T$ = top width at the free surface,  $\theta$  = angle between  $x$  direction and horizontal direction,  $\beta$  = the momentum correction factor,  $S_f$  = friction slope defined as  $\tau/\gamma R$ ,  $R$  = hydraulic radius,  $\phi$  = angle between velocity  $U$  and  $x$  direction, and other variables have been already defined. For a wide channel,  $D$  and  $R$  are simply replaced by flow depth,  $Y$ .

#### 4. SURFACE ROUGHNESS EFFECT

##### 4.1. Laminar Flow

The study of laminar sheet flow over bare surface is the most simplified situation of interest in order to identify the variation of flow resistance coefficient due to surface roughness and Reynolds number. The following general formulation has been adopted by early investigators, such as Izzard (1944), and Woo and Brater (1961):

$$f = \frac{K}{R_e} \quad (16)$$

K value varies with the flow regime, surface roughness, rainfall effect, vegetation and probably slope. Theoretically speaking, K is equal to 24 for laminar flow over a smooth wide channel. This can be found by either applying Boussinesq equation, primarily developed for rectangular pipes having a width b and depth of 2Y, to a wide open channel with infinite width and depth of Y, or imposing equilibrium between the component of weight in the direction of flow and the shear resistance of the channel bottom. Horton, Leach, and Van Vliet (1934) experimentally confirmed the K value being 24 for laminar flow in a rectangular channel with a smooth surface, covered by white pine. Allen (1934) found the upper limit of  $R_e$  for true laminar flow regime being about 300 for smooth surfaces. The University of Illinois' data given by Landsford and Robertson (1958) and Chow (1959) determined the same K value as 24 for laminar flow when  $R_e < 500$ .

Woo and Brater (1961) tried to determine friction factor for different boundary surfaces. They partitioned the surfaces into smooth,

rough, and very rough. Woo and Brater evaluated the width effect for the flow in rectangular channels, estimating an error of less than 5 percent in K when the width-depth ratio was 25. Woo and Brater's data for flow over masonite surface representing a typical rough surface showed a value of 30.8 for K. The U.S. Waterways Experiment Station (1935) had already reported K being 31.6 for laminar flow over cement surface. The upper limit of  $R_e$  for laminar flow varied from 400 for a slope of 0.060 to 900 for a slope of 0.001.

Glued-sand with an average diameter of 1 mm on the masonite surface used by Woo and Brater (1961) as a very rough surface on which flow experiments were conducted. It was found that K increased with the slope (except for slopes less than 0.003), having a value of 39.2 for  $S_o = 0.001$  up to 100 for  $S_o = 0.060$ , Fig. 1. The upper limit of laminar flow range was confined between 400 to 800, varying inversely with the slope. Generally, the data in the laminar range seems inadequate to warrant the results.

If the  $f$  variation with slope is computed based on Woo and Brater's (1961) data, it will be found that for sand surface ( $k=1$  mm) when  $S_o > 0.003$ :

$$f = \frac{155.85 + 46 \log S_o}{R_e} \quad (17)$$

The application of the above equation is limited to slopes less than 0.020 after which the number of data points for each slope is lacking.

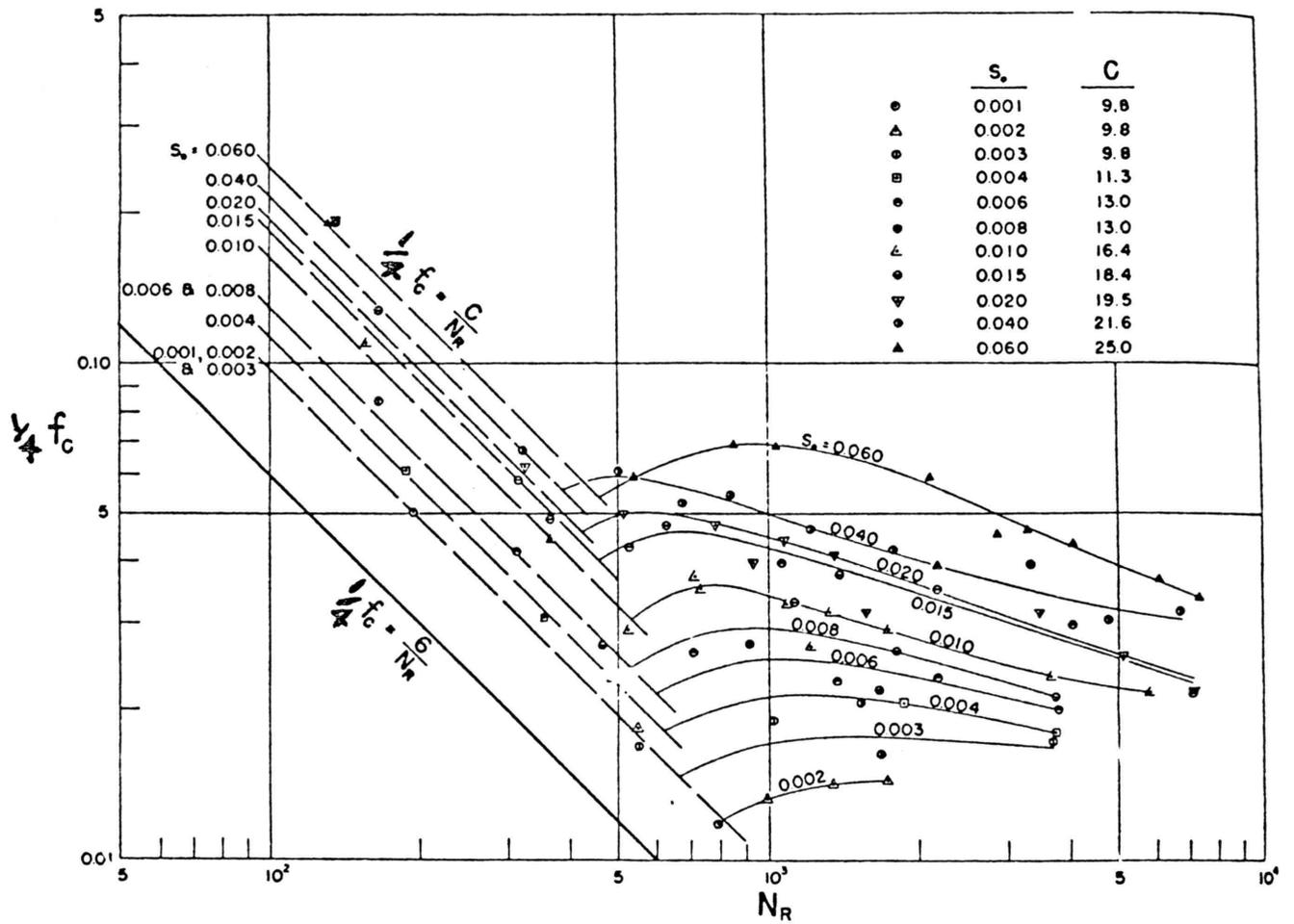


Fig. 1. The  $f$ - $R_e$  relationship for sand surface, after Woo and Brater (1961).

Through a different approach, Kruse et al (1965) attempted to define the friction factor for flow over rough surface in terms of roughness characteristics and channel slope. They came up with the following formula :

$$f = \frac{6000(\sigma/\lambda)S_0^{0.5}}{R_e} \quad (18)$$

where  $\sigma$  = soil roughness height, and  $\lambda$  = soil roughness spacing. The formula shows the correlation of friction factor with the ratio of roughness height to spacing and apparently the bed slope.

The idea of correlation of  $f$  with the relative roughness was investigated by Phelps (1975). Phelps tested the flow over spherical roughness elements with diameter of 1.17 mm (.046 in) and grain concentration of 0.1 in the slope range being 0.00048-0.0451. The data confirmed the variation of  $f$  with relative roughness not slope.

Having Phelps' data in Fig.2, the following power equation may be developed to confirm Eq.16 for constant  $k/Y$ :  $f = aR_e^b$ . Table 1 can be filled by using Fig.2 as the reference.

TABLE 1 - Values of  $a$  and  $b$  Based on Phelps' Data

Relative Roughness (1)	# of Data (2)	$a$ (3)	$b$ (4)	$K$ (5)
.23	4	35.889	-1.00195	35.498
.27-.28	5	43.584	-1.02503	38.161
.35	7	42.392	-1.00191	42.040
.52-.55	7	31.179	-0.88777	50.61

As it is seen, the exponent  $b$  is very close to  $-1.0$  except for the last series when  $k/Y = .52 - .55$ . As a result, the resistance equation may be written in this form:  $f = K/R_e$ , where  $K = \text{func}(k/y)$ . If a regression is to be performed, the result for  $K$  will yield:

$$K = 24 + 72.1 \left(\frac{k}{Y}\right)^{1.31}, \quad \frac{k}{Y} < .50 \quad (19)$$

The application of resistance equation in the form of  $f = K/R_e$  would be probably limited to  $k/Y$  values less than  $.50$ , according to Phelps' data. The result of the power model for  $k/y = .52 - .55$  is not satisfactory to verify the equation for that specific  $k/Y$ . It is possible that free surface instability effect for high  $k/Y$  cause the discrepancies such that the correlation of  $f$  with  $R_e$  decreases indicating the change in flow regime from laminar to transition and turbulent.

Phelps (1975) reported that Woo and Brater's (1961) data also validated Eq.16 as they were grouped based on relative roughness. Assuming so,  $K$  values deduced from Woo and Brater's data are higher than those of Phelps' as much as two times for a constant  $k/Y$ . One may reason that the roughness concentration used by Woo and Brater was the maximum possible similar to Nikurase's work, where Phelps' selected a concentration equal to  $0.1$  in his experiments.

Now, as it is clear, two different independent variables have been used for the evaluation of flow resistance, i.e. slope and relative roughness. Although Kruse et al. (1965) presented an equation in which slope was the independent variable besides the roughness size, they

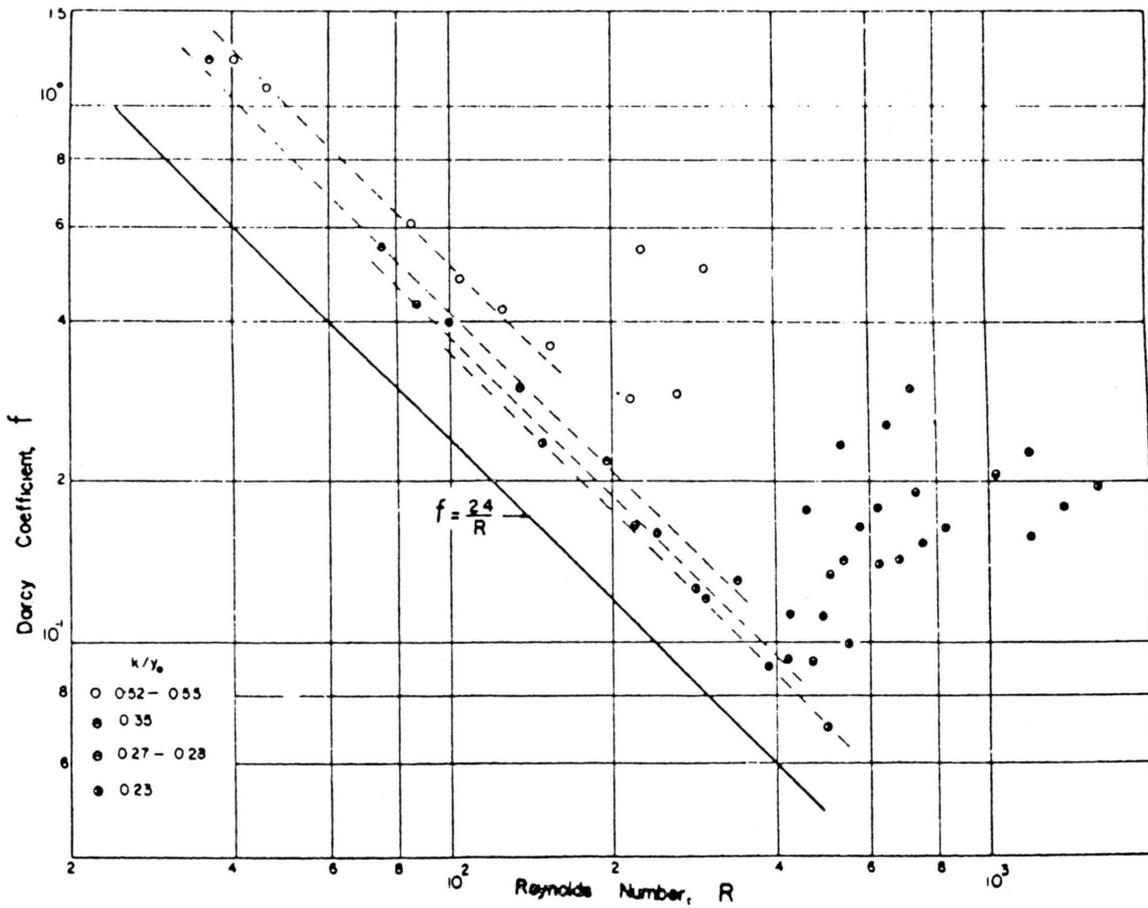


Fig. 2. The  $f$ - $R_e$  relationship for rough surfaces, after Phelps (1975).

$k/Y$  = Relative roughness

speculated that the apparent correlation between resistance and slope could be due to relative roughness and local turbulence at the tips of the roughness elements. When slope increased while discharge and hence Reynolds number were kept constant, depth would then decrease and more resistance would be induced due to larger portion of the flow being into contact with the roughness at a higher velocity. Therefore, the basic cause of resistance variation can be relative roughness rather than slope, which in turn is responsible for changes in relative roughness. In addition, working with slope as the primary variable requires a series of experiments for each roughness size whereas the  $k/Y$  ratio reflects both roughness size and depth which varies with bed slope in the case of constant discharge. Phelps' work successfully demonstrates the effectiveness of  $k/Y$  being independent variable and the validity of equation  $f = K/R_e$ .

Yet, some considerations must be taken into account when working with relative roughness. First of all, the roughness concentration has to be held constant for each diagram of  $f$  vs  $R_e$  and  $k/Y$ . Second, the  $k$  value, the height of the roughness, needs an accurate measurement. Third, for high  $k/Y$ , free surface instabilities may bring about additional energy dissipation whose effect on  $f$  in laminar flow region has not been quantitatively determined.

#### 4.2. Turbulent Flow

The flow over a bare surface becomes turbulent when  $R_e > 2000$ . There are three types of turbulent flow depending on size of the boundary roughness compared to laminar sublayer thickness. Smooth

conditions occur when the boundaries are hydraulically smooth such that the roughness elements are well covered under the laminar sublayer. On the contrary, turbulent flow over fully rough surface exists when the projections break through the laminar sublayer and dominate the flow behavior. Finally, transition region of turbulent flow is the region between smooth and fully rough conditions. It is noticeable that change from smooth to fully rough flow corresponds to increase in  $R_e$  and therefore in discharge, which shrinks the laminar sublayer thickness. The limits of these three kinds of turbulent flows are as follows :

1. Smooth condition :  $\delta > 3k$  or  $V_*k/\nu < 4$
2. Transition :  $k/5 < \delta < 3k$  or  $4 < V_*k/\nu < 70$
3. Fully rough :  $\delta < k/5$  or  $V_*k/\nu > 70$

where  $k$  = the median size of the boundary particles and  $\delta$  = the laminar sublayer thickness equal to  $11.6\nu/V_*$ .

The resistance equations were primarily developed for flow in pipes. The  $f$ - $R_e$  relationship for smooth pipes was derived by Blasius as the following :

$$f = \frac{0.223}{R_e^{0.25}} \quad (20)$$

in which hydraulic radius is used as the characteristic length in definition of  $R_e$ . The Blasius equation may be applied for turbulent flows over smooth boundary when  $R_e < 25000$ . Beyond that limit, the

Prandtl-von Karman equation based on logarithmic velocity profile is believed to hold :

$$\frac{1}{\sqrt{f}} = 2 \log (R_e \sqrt{f}) + 0.4 \quad (21)$$

The use of Eq. 20 and Eq. 21 for open channel flow has been investigated based on the data developed at the Univ. of Illinois given by Lansford and Robinson (1958) and also data of Univ. of Minnesota given by Straub et al. (1958). Fig.3 indicates that the equations for turbulent flows in smooth pipes may be representative of all smooth channels. In addition, the cross section shape of the channel in turbulent flow has little effect on friction factor whereas it is important in laminar flow. This means that for sheet flow assumed in a wide channel, Eq.20 and Eq.21 can approximate the friction factor when the boundary is smooth such as that of urban drainage systems.

Another alternative is to integrate the turbulent velocity profile over smooth boundary and then calculate the friction factor from average velocity. The final formula would be :

$$\frac{1}{\sqrt{f}} = a \log \left( R_e \frac{\sqrt{f}}{b} \right) \quad (22)$$

Basically, a is related to the von Karman's universal constant as 0.4, and b depends on the value of a as well as shape of the cross section of the channel. Keulegan's (1938) formula, which probably is the closest in result to Prandtl-von Karman equation, for a very wide, smooth

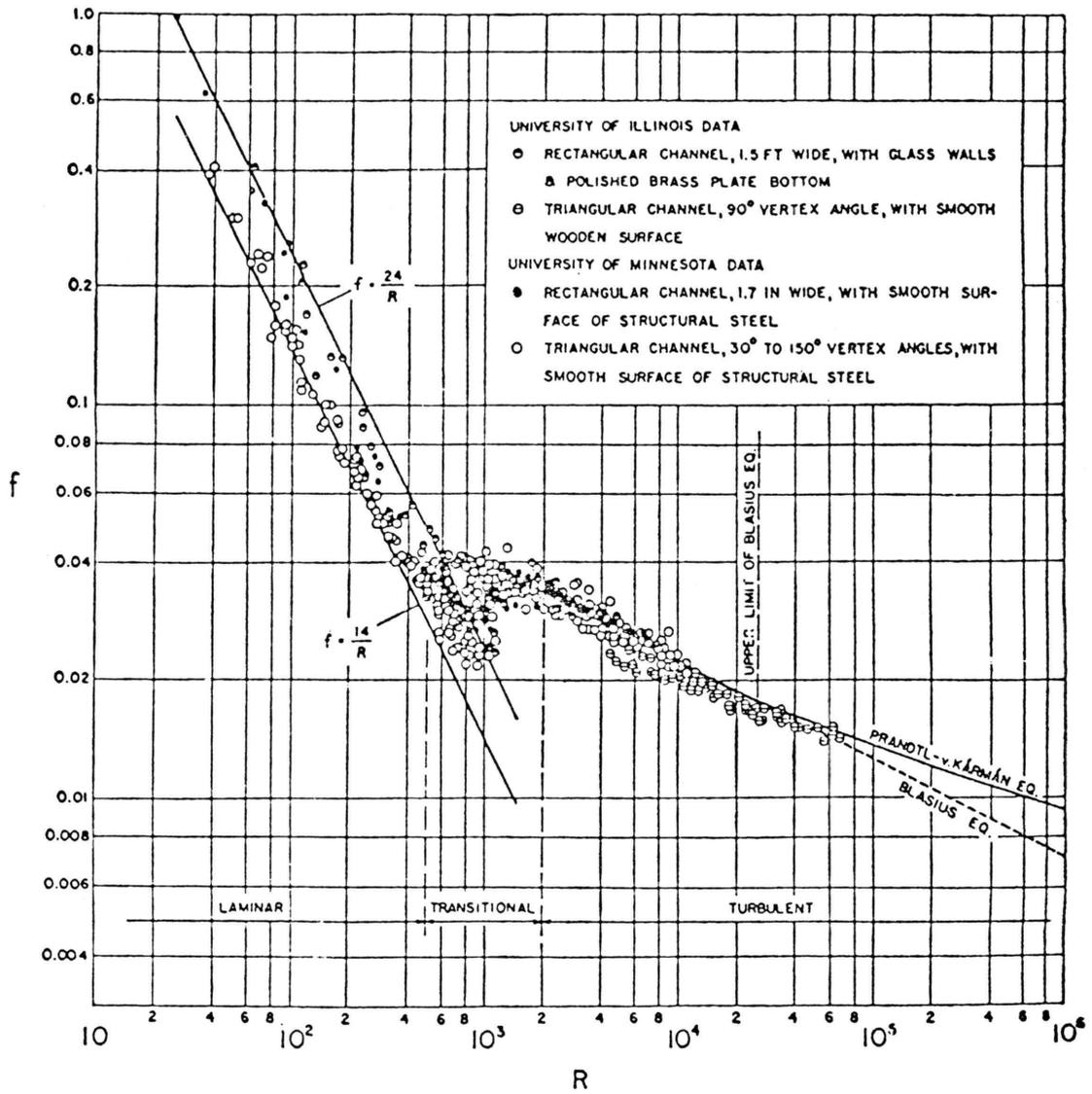


Fig. 3. The  $f$ - $R_e$  relationship for flow in smooth channels, after Chow (1959).

channel reduces to  $a=2.03$ , and  $b=0.853$ . In overland areas, however, the surface is mostly rough with fairly large relative roughness.

The flow resistance of turbulent flow in fully rough condition is entirely due to the ratio of hydraulic radius over the roughness size,  $R/k$ , and can be expressed as follows :

$$\frac{1}{\sqrt{f}} = \frac{C}{\sqrt{8g}} = a \log\left(b' \frac{R}{k}\right) \quad (23)$$

where  $R$ =hydraulic radius, and  $b'$  is a constant to be determined by experiments. The value of  $b'$  depends not only on the shape of the channel cross section but also on the spacing (roughness concentration) and form of the roughness elements. As a result, different investigators present different values based on the data they use. Keulegan (1938) found that  $a=2.03$  and  $b'=11.09$  for a very wide channel with sand-grain roughness in the fully rough regime. For a trapezoidal channel, however, Keulegan's formula gives similar  $a$  but  $b'=12.27$ . At the meeting of IAHR, Thijsee (1949) proposed a similar equation which after modifications results in  $a=2.03$  and  $b'=12.2$  for a very wide channel. In case of flow over commercial surfaces, such as concrete and wood, the  $k$  values have been presented by Ackers (1959).

If the variation of Chezy coefficient  $C$ , instead of Darcy  $f$ , is to be plotted versus  $R_e$  using Eq.20 and Eq.21 for smooth condition and Eq.23 for fully rough condition, a modified Moody diagram for open channel flow will show up. Fig.4, taken from Henderson's (1966) book, indicates that in case of turbulent flow over fully rough surfaces,  $C$  only depends on  $R/k$  ratio and independent of  $R_e$  effect. The  $R/k$  ratio

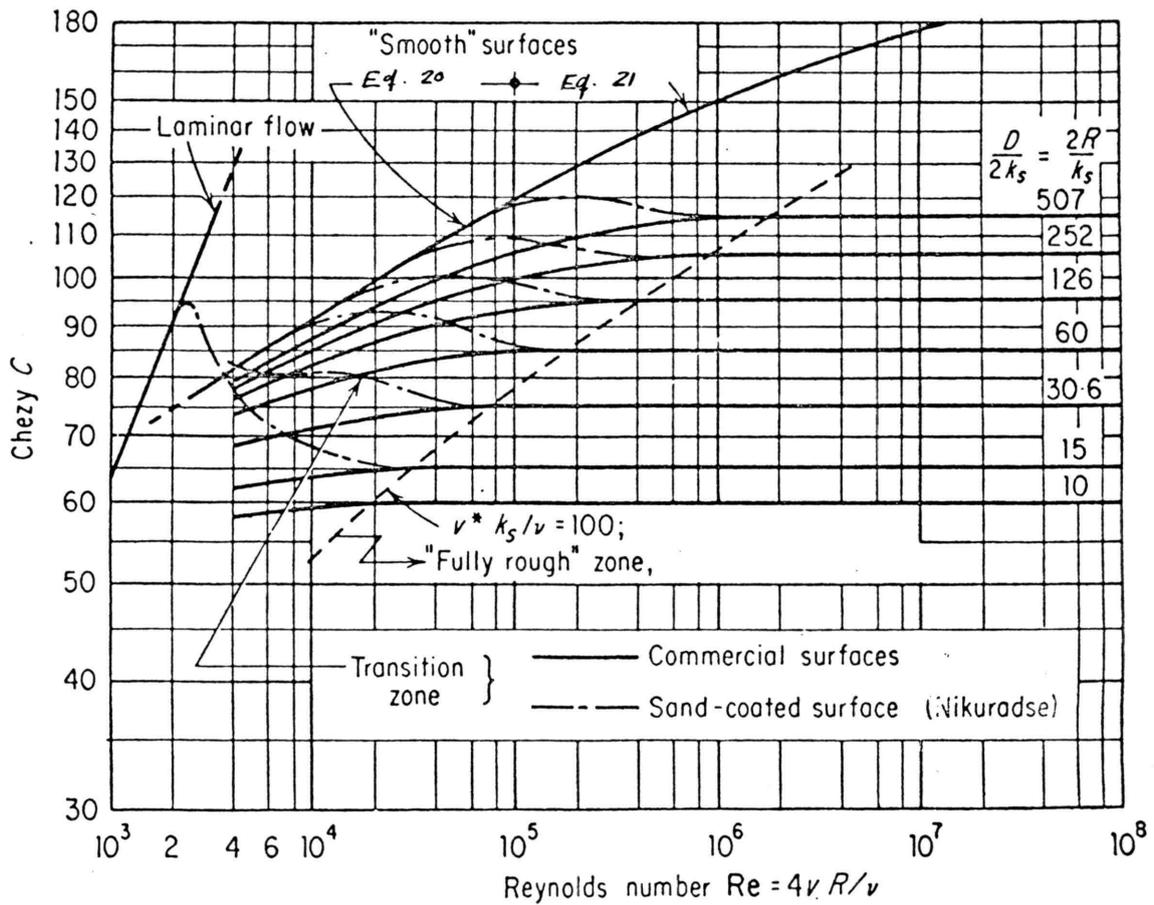


Fig. 4. Modified Moody diagram showing  $C-R_e$  relationship, after Henderson (1966).

covers from 5 to 235.5, probably based on range of available data. Although turbulent flow in fully rough condition usually occurs in relatively high R/k ratios, in overland regions with steep slope one may expect turbulent sheet flow with high relative roughness, or low ratios of R/k. In that case, the applicability of Eq. 23 needs more investigations in order to complete Fig. 4 for smaller R/k ratios.

A report by ASCE (1963) supports the use of Colebrook equation with slightly modified coefficients for flow in transition region to open channels. The equation is :

$$\frac{1}{\sqrt{f}} = \frac{C}{\sqrt{8g}} = -2 \log\left( \frac{k}{12R} + \frac{0.625}{R_e \sqrt{f}} \right) \quad (24)$$

However, the above equation is applicable to commercial surfaces. Therefore, for natural rough surfaces with k being the median particle size, Eq.24 has to be tested. In Fig.4 , the difference between the curves for pipe flow and open channel flow in transition region is shown.

Manning equation, as a flow resistance equation, is the most well known power relationship which has been developed for open channel turbulent flow over rough surfaces. For R/k ratios ranging from 10 to 10000, the Manning-Strickler relationship approximately gives equivalent resistance coefficients as the logarithmic equation by Keulegan:

$$n = \frac{1.486R^{1/6}}{\sqrt{8g/f}} = 0.0342 k^{1/6} \quad (25)$$

where  $k$ =median size of the roughness particles in feet. It should be noticed that Manning equation is suitable for all fully rough flows in which Manning's  $n$  is constant for a given particle size. For transition flows, however,  $f$  is the better resistance coefficient given by Eq. 24.

boundary shear stress,  $\tau$ , assuming  $\beta = 1$ . He found that the measured boundary shear stress, even with the difficulties in measuring flow depths with rainfall effect, was in excellent agreement with boundary shear stress computed using Eq.26. Therefore, the application of one dimensional dynamic equation of spatially varied flow appeared to be accurate enough for determination of water surface profile, provided a reasonable resistance law; i.e. an equation for  $f$ . It was also found that  $S_0$  overcame the other terms in magnitude while evaluating  $S_f$ . Each of  $S_1$  and  $S_2$  contributed nearly one tenth of  $S_0$  whereas  $S_3$  was negligible in magnitude.

### 5.1. Laminar Flow

Izzard (1944) first studied the resistance to laminar sheet flow with rainfall effect. He considered that the  $K$  value in general formula, Eq.16, could be the sum of a constant and a function of rainfall intensity. Therefore the following function was developed and then used by many other investigators:

$$f = \frac{K}{R_e} = \frac{K_0 + \phi(i)}{R_e} \quad (28)$$

where  $K_0$  is a function of surface roughness. Izzard used a paved rough surface in his experiments. As a result, he determined  $K_0$  being 27 for rough surface. The power function of rainfall intensity turned out to be  $5.67 i^{1.33}$ , where  $i$ (in/h). In addition, Izzard observed increase in  $f$  with increasing bottom slope. However, no slope parameter was included in friction factor equation.

Li (1972) conducted his tests to determine the independent variables of friction factor for laminar flow over smooth surface with rainfall through a dimensional analysis. He assumed the following power equation:

$$f = \beta_0 R_e^{\beta_1} i^{\beta_2} S_0^{\beta_3} \epsilon \quad (29)$$

where  $\beta_0, \beta_1, \beta_2, \beta_3$  are constants and  $\epsilon$  is the error in the regression equation. The data covered a range of  $R_e$  from 126 to 900 for laminar regime, 0 to 17.5 in/h for rainfall intensity, and slopes being .0108 and .0064. The result of multiple regression showed that :

$$f = 13.517 R_e^{-.958} i^{.413} S_0^{-.088} \epsilon \quad (30)$$

According to statistical tests made by Li (1972), bottom slope had an insignificant effect on the product of  $f.R_e$ . Furthermore, the exponent of  $R_e$  was approximated to -1.

Before Li (1972), Yoon (1970) had carried out several tests to identify the independent variables affecting friction factor. Yoon (1970) found that the effect of raindrop spacing and raindrop impact velocity were almost negligible on friction factor under his test conditions. However, friction factor increased with increasing rainfall intensity and relatively bottom slope.

Li (1972) performed a regression analysis using his data and Yoon's data to derive the following power function for  $\phi(i)$ :

$$\phi(i) = 27.162 i^{.407} , \quad \text{for } R_e < 900 \quad (31)$$

$i$  is in in/h. The agreement of the above equation with Yoon's data is shown in Fig.5 and with Li's data in Fig.6.

Fawkes (1972) approximated the flow with rainfall as a steady flow with a very flat water surface profile. As a result,  $S_f$  would be almost equal to  $S_0$ . Fawkes then presented  $\phi(i) = 9.982i$ .

Other data based on experiments on sheet flow over smooth and rough surfaces with rainfall given by Kisisel et al.(1973) indicated no significant change in  $f$  due to slope. The data seemed to obey the same general formulation for  $f$ , though no attempt was made to deduce a certain equation for  $f$ .

In order to define friction factor experimentally for sheet flow with rainfall, most of the investigators used the kinematic wave approximation as suggested by Woolhiser (1969). The approximation assumes that all the terms in the momentum equation are negligible except  $S_0$  and  $S_f$ , resulting in  $S_f = S_0$ . Then, depth and velocity in Eq.12 are measured for a cross section and the variation of  $f$  due to rainfall versus  $R_0$  will be defined. Izzard (1944), Kisisel et al.(1973), and Fawkes (1972) used the kinematic wave approximation to determine the  $f$  variation.

According to Yoon's study on Eq.26, the kinematic wave approximation may involve up to 20 percent error in  $S_f$  determination. Yoon (1970), and then Li (1972), directly measured the boundary shear stress by hot film sensors, in order to avoid any approximation in their analysis. Having shear stress and flow velocity, they computed friction factor,  $f = 8\tau/\rho V^2$ , for specific rainfall intensity and Reynolds number.

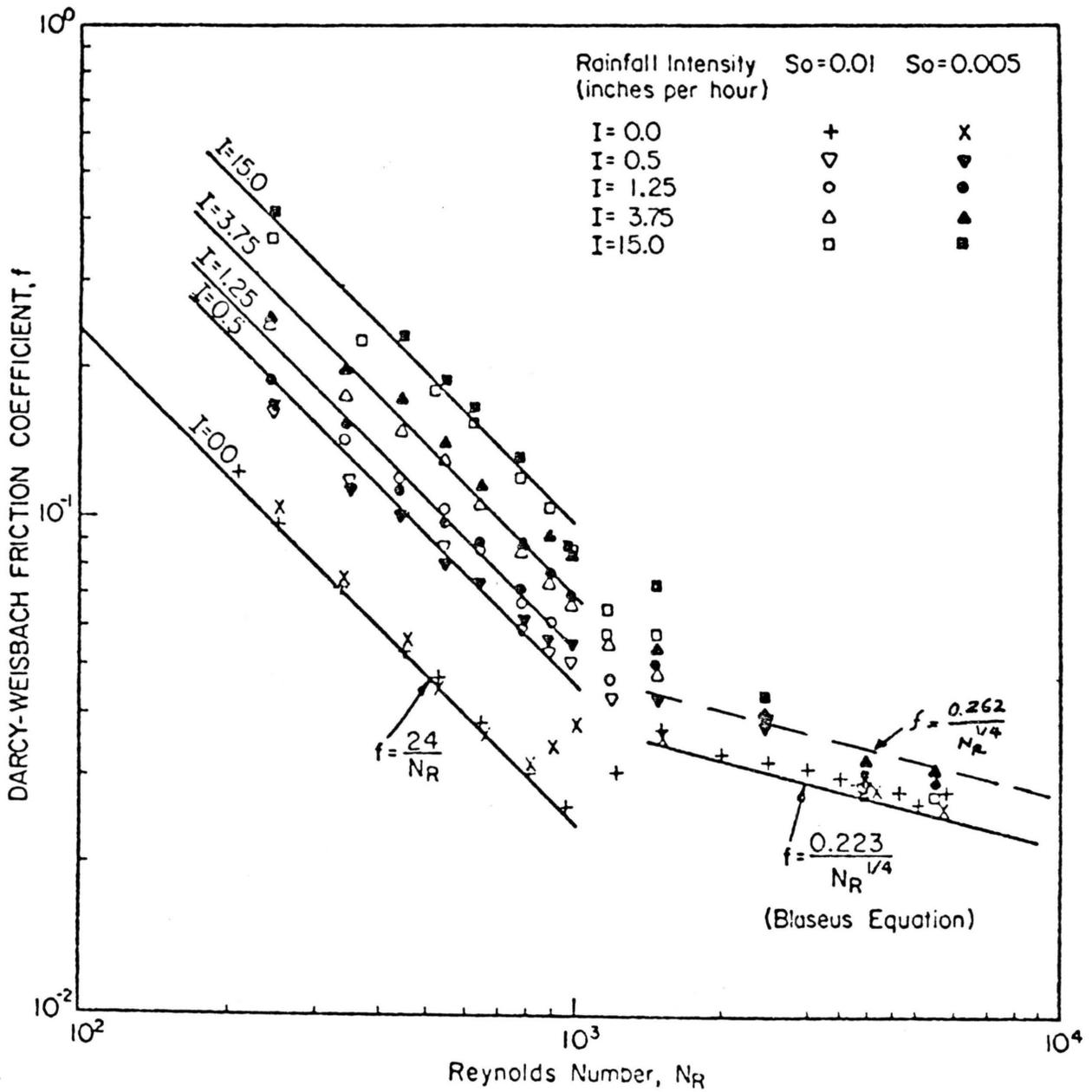


Fig. 5. The  $f$ - $R_e$  relationship for flow with rainfall, after Yoon (1970).

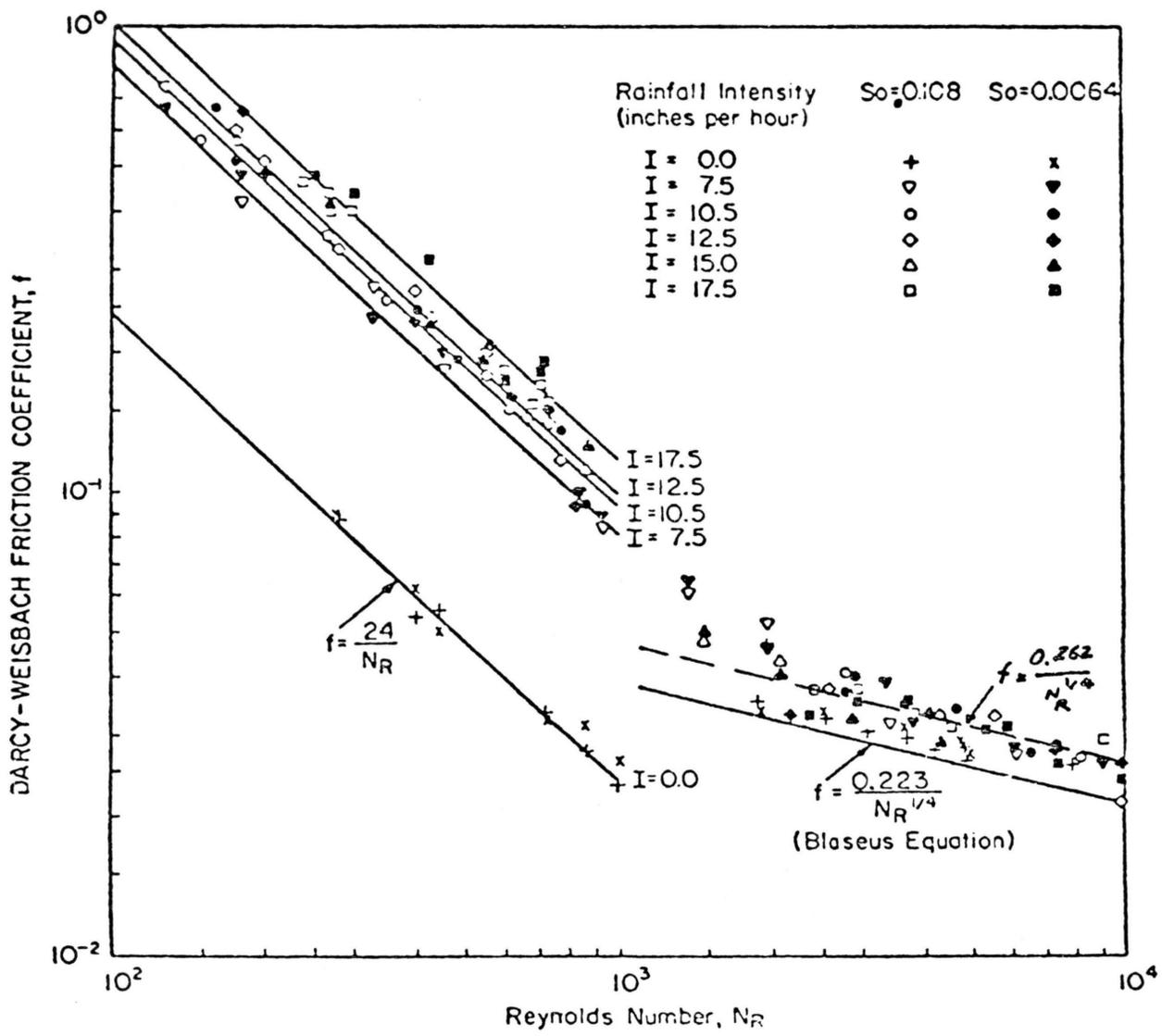


Fig. 6. The  $f-R_e$  relationship for flow with rainfall, after Li (1972).

Consequently, Eq. 28 substituted by  $\phi(i)$  from Eq. 31 is the most accurate equation for solving dynamic equation of spatially varied flow.

As already discussed,  $K$  in Eq.16 may be a function of slope,  $S_0$ , or relative roughness,  $k/Y$ . Using a function of  $S_0$  would bring about an approximation by assuming steady uniform flow, which is obviously not true when rainfall exists. On the other hand,  $K$  being a function of  $k/Y$ , as used by Phelps (1975) specifically for steady uniform flow over rough boundary, reflects the effect of non-uniformity of the flow with rainfall effect. As spatially varied flow moves on, the depth changes and the boundary resistance has to change accordingly to yield the relative roughness effect. Therefore, both friction factors due to boundary roughness and rainfall will be functions of distance, simply because depth and Reynolds number are not constant for sheet flow with rainfall :

$$f = \frac{\text{func}(k/Y) + 27.162i^{.407}}{(q_0 + ix)/v} \quad (32)$$

## 5.2 Turbulent Flow

Similar to the discussion for laminar flow with rainfall, the data provided by Yoon (1970) and Li (1972) are the most applicable and accurate compared to the other's data. Li first assumed that Blasius equation could be modified to accommodate the rainfall effect for turbulent flow over a smooth boundary :

$$f = \frac{\phi'(i)}{R_e^{0.25}} \quad (33)$$

which is valid for  $R_e > 2000$  where the turbulent flow begins. The regression analysis between Yoon's and Li's data showed that for available data  $\phi'$  was not a function of rainfall intensity but rather a constant. The results indicate that :

$$\begin{aligned} \phi' &= 0.262 & \text{for } 0.5 < i < 17.5 \text{ in/h} \\ \phi' &= 0.25 & \text{for } i = 0 \end{aligned} \tag{34}$$

The above results mean that the flow resistance begins to increase with rainfall intensity somewhat below 0.5 in/h. Once the flow resistance is increased, any further increase of rainfall intensity doesn't change the flow resistance at least for  $i < 17.5$  in/h. Since the major cause of increase in flow resistance due to rainfall is the creation of turbulence by rainfall impact, one should expect a little change in flow resistance when the flow is already turbulent.

As seen in Figs. 5 and 6, the  $f$  values decrease from that for the laminar range ending at  $R_e = 900$  to its value for the turbulent range starting at  $R_e = 2000$ . Li (1972) approximated the relation between  $\ln f$  and  $\ln R_e$  in transition range with a line and gave the following equation:

$$f = 0.0392 (R_e/2000)^a \tag{35}$$

in which  $a = -1.252 \ln(0.68 + 0.77i^{0.407})$ . The equation applies only for flow in the transition range,  $900 < R_e < 2000$ , over a smooth boundary.

## 6. VEGETATION EFFECT

Evaluation of vegetation resistance in sheet flow involves the most complicated experiments particularly for natural vegetation. So many interrelated variables contribute in flow resistance through vegetation that no test is able to separate the effect of each variable. The problem becomes more complex when the combined effects of vegetation, bottom roughness, and rainfall are present and yet no confirmed method of separation among those effects has been developed. Nevertheless, at least in case of laminar flow, it is believed that total resistance can be represented by the linear superposition of vegetation drag, bottom roughness, and rainfall effect. The last one is minor compared to vegetation drag and the natural bottom roughness of natural vegetated areas. The bottom effects due to roughness has been already discussed.

Although no unique equation in a general form has been derived to calculate the vegetation resistance, the following literature review and discussions will clarify, to some extent, the results of past studies.

### 6.1. Rigid Sparse Vegetation

The relationship between resistance to flow and hydraulic parameters of sheet flow through rigid sparse vegetation can be derived by applying momentum equation to a finite increment  $\Delta x$  along flow direction. For a steady flow in a wide channel one obtains :

$$F_g = F_b + F_D \quad (36)$$

where  $F_g$  = fluid weight component in flow direction per unit width approximately equal to  $\gamma Y S_0$  in case of sparse vegetation,  $F_b$  = boundary shear force per unit width, and  $F_D$  = total vegetation drag per unit width. The boundary shear force is equal to  $\gamma Y S_f$  or  $\rho f_b V^2/8$ , in which  $S_f$  = the friction slope due to boundary resistance, and drag force is equal to  $\int 0.5 C_D V_e^2 dA_e$  in which  $C_D$  = local drag coefficient, and  $dA_e$  = local area of vegetation projected normal to flow direction. If the vegetation system is composed of rigid uniform cylinders and local velocity can be approximated by mean velocity of the flow, then Eq.36 becomes :

$$\gamma Y S_0 = f_b \rho V^2/8 + 0.5 N C_D d Y V^2 \quad , \quad \text{for } h > Y \quad (37)$$

where  $N$  = the number of cylinders per unit area of bed,  $d$  = cylinder diameter, and  $h$  = cylinder height. When  $h < Y$ , then  $h$  should be substituted for  $Y$  in last term. In a more simplified form :

$$f_t = f_b + f_v = \frac{8 \gamma Y S_0}{V^2} \quad (38)$$

where  $f_v$  = friction factor due to vegetation equal to  $4 N C_D d Y$ . Hence, the contribution of vegetation effect,  $f_v$ , to total friction factor is dependent on flow depth as the hydraulic parameter, vegetation characteristics including number of single stems per unit area in a sparse pattern, stem diameter, and drag coefficient.

Li and Shen (1973) studied the drag coefficient for idealized vegetation, represented by rigid cylinders. As Fig.7 shows, the variation of mean drag coefficient in turbulent flow for second row cylinders in a staggered pattern is relatively small down to at least longitudinal spacing to diameter ratio of 5 at which  $C_D$  is only 8% higher than that of a single cylinder or that of first row cylinders. In case of a parallel pattern, however,  $C_D$  keeps continuously decreasing as the spacing is reduced for a given  $d$ , such that  $C_D$  equals only 60% of  $C_D$  for a single cylinder. Of course when the transverse spacing is changed, these ratios may change. Now, as long as  $C_D$  remains unchanged with the spacing, roughly down to  $10d$  in staggered pattern and  $50d$  in parallel pattern, the vegetation is considered sparse and  $C_D$  would be only function of element shape and Reynolds number, as has been classified by Hoerner (1965). Li and Shen recommend an average  $C_D$  being 1.2 for sparse cylinders. This value also has been reported in standard texts such as Schlichting (1968) for drag coefficient of a single cylinder in an idealized two-dimensional flow in cylinder Reynolds number,  $R_d = Vd/\nu$ , ranging from about  $8 \times 10^3$  to  $2 \times 10^5$ .

## 6.2. Dense Rigid Vegetation

Neglecting the free surface and flexibility effects, Kirsch and Fuchs (1967) studied the drag coefficient for pressure flow through parallel and staggered arrangements of dense rigid cylinders. They introduced a dimensionless coefficient of drag enhancement,  $F^*$ , which relates to  $C_D$  as the average drag coefficient for each cylinder in an array such that :

$$C_D = \frac{2F^*}{R_d} \quad (39)$$

where  $R_d$  = cylinder Reynolds number equal to  $Vd/\nu$ . If  $S_1$  and  $S_2$  represent the center to center spacing in the cross stream direction and in streamwise direction,  $F^*$  can be empirically evaluated as :

$$F^* = 4\pi \left[ -\ln\left(\frac{d}{2S_1} - 1.33\right) + \frac{\pi^2}{3} \left(\frac{d}{2S_1}\right)^2 \right], \text{ for } d/S_1 < 0.7$$

$$F^* = \frac{9\pi}{2\sqrt{2}} \left(1 - \frac{d}{S_1}\right)^{-2.5}, \text{ for } d/S_1 > 0.7 \quad (40)$$

Both above equations hold when  $S_2 > S_1$ . For  $S_2 < S_1$ ,  $F^*$  ratio decreased below unity with decrease in spacing between rows in a parallel arrangement. On the contrary, opposite relation was verified for staggered pattern in the case  $S_2 < S_1$ , depending on  $d/S_1$  and  $S_1/S_2$ . Kirsh and Fuchs also found that for nonuniform pattern of cylinders and for rotating rows of cylinders relative to one another,  $F^*$  showed less value than those of parallel and staggered patterns of equal density.

Chilton and Genereaux (1933) experimented pressure drop for the pressurized flow through staggered arrangement of cylinders presenting :

$$\frac{\Delta P}{L} = \frac{53V_{\max} \mu}{d_e^2} \quad \text{for laminar flow} \quad (41)$$

$$\frac{\Delta P}{L} = \frac{1.5\rho \cdot^8 V_{\max} \mu \cdot^2 \sqrt{N}}{G \cdot^2} \quad \text{for turbulent flow} \quad (42)$$

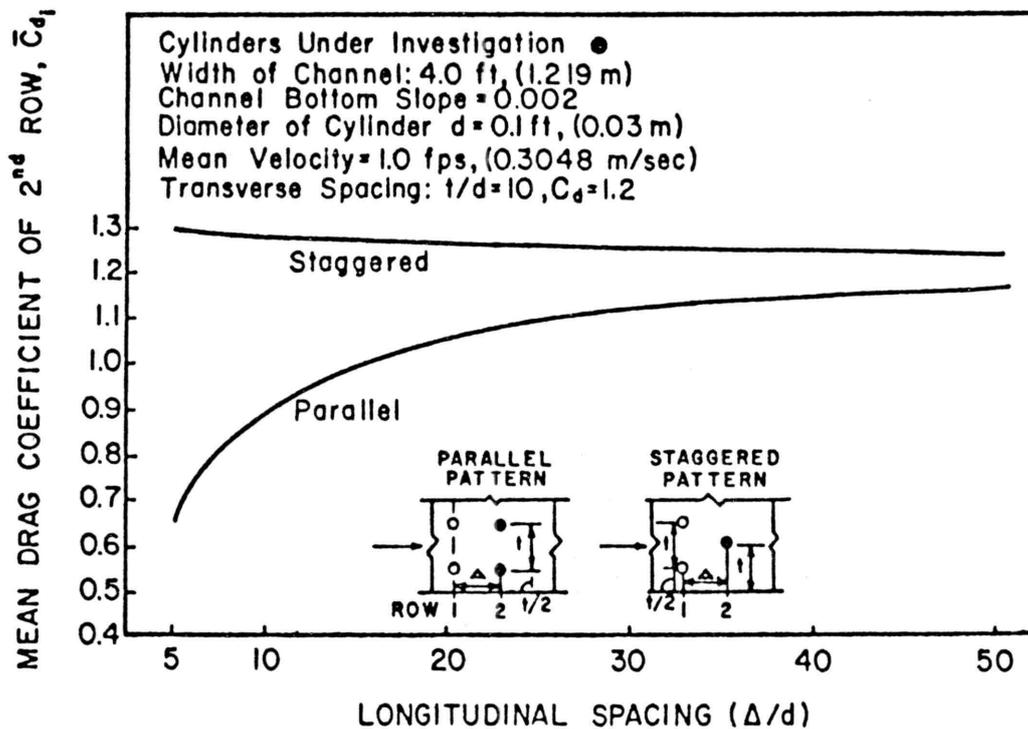


Fig. 7. The mean drag coefficient variation for staggered and parallel patterns, after Li and Shen (1973).

in which  $\Delta P$  = pressure drop over length  $L$ ,  $V_{\max}$  = maximum velocity through the gap or narrowest space between two adjacent cylinder elements,  $d_e$  = equivalent diameter equal to  $(4/\pi dN - d)$ ,  $d$  = cylinder diameter,  $N$  = number of elements per unit area of the bed,  $G$  = gap size.

Eq. 41 may be changed for the use in open channel with the aid of similarity between friction factor in open channel and pressure drop in pipes :

$$S_f = \frac{fV^2}{8gY} = \frac{53V_{\max}\mu}{\gamma d_e^2} = \frac{\Delta P}{\gamma L} \quad (43)$$

or by substituting  $V = GV_{\max}/S$  :

$$f = 424 \frac{Y S}{d_e G} (R_e)_e^{-1} \quad (44)$$

where  $(R_e)_e = V_{\max}d_e/\nu$ . This equation has not been verified experimentally for open channel flow. It confirms, however, the proportionality of friction factor directly with flow depth , and inversely with Reynolds number.

Similar modifications for turbulent flow relationship with recalling that  $N = 1/S^2$  yield :

$$f = \frac{12YS}{G^2} (R_e)_G^{-.2} \quad (45)$$

where  $(R_e)_G = V_{\max}G/\nu$ . Although the equation was primarily developed for pressure flow, it can confirm the linear dependence of  $f$  on flow depth,  $Y$ , in case of turbulent flow through rigid dense vegetation, similar to the relation for rigid sparse system. The small negative power of Reynolds number also satisfies the expectation for a turbulent flow.

Hartley (1980) tested the sheet flow on a smooth surface through 1/4 inch diameter cylinders representing ideal vegetation. He then measured the flow depths and velocities and used the following energy equation to evaluate friction slope :

$$S_f = (Y_1 - Y_2)/\Delta x + (V_1^2 - V_2^2)/2g\Delta x + S_0 \quad (46)$$

where subscripts 1 and 2 stand for upstream and downstream locations with the distance  $\Delta x$  apart. He reported that since the flow was close to a uniform flow, in most cases  $S_f$  showed values quite near  $S_0$ . Then, the total friction factor  $f$  could be calculated having  $S_f$ ,  $Y$ ,  $V$  and using Eq.12. Assuming linear superposition of drag, Hartley removed the sidewall effect applying the method by Vanoni and Brooks (1957) and then bottom resistance using  $f = 24/R_e$  for laminar flow and Blasius equation for turbulent flow. In case of smooth boundaries, the sidewall effect and bottom resistance showed quite minor values compared with the vegetation resistance.

Hartley assumed the following simple power model for laminar flow:  $f = A (Y/d)^B R_d^C$ , where  $A$  depends on density and pattern,  $Y/d$  is the depth diameter ratio to account for form drag effects, and  $R_d$  is diameter

Reynolds number equal to  $V_{\max} \cdot d / \nu$  . By performing regression, Hartley confirmed the general form  $f=K/R_d$  as:

$$f = A (Y/D) R_d^{-1.0} \quad (47)$$

Generally, having depth, instead of bed slope, as independent variable is advantageous because in case of non-uniform flow with rainfall the effect of change in depth would be included in flow resistance due to vegetation.

For turbulent flow, Hartley dropped the effect of Reynolds number, assuming negligible effect, and he allowed Froude number to enter the equation. Therefore, the power equation for turbulent flow became :

$$f = A (Y/d)^B F^E \quad (48)$$

where  $F$  = Froude number. By performing data regression, Hartley found the influence of Froude number to be marginal in its effect on resistance coefficient, even though the free surface effects were physically evident in some slopes. Also the exponent of  $Y/d$  turned out to be 1.

To account for density variation, Hartley introduced a correction factor being  $(d/S)^2$ . Therefore his resistance equation now becomes:

$$f = C \frac{dY}{S^2} R_d^P \quad (49)$$

in which p equals -1 for laminar flow and zero for turbulent flow. Constant C is dependent on the vegetation pattern as in the following table :

Table 2. Pattern Coefficient (C)

Pattern	Laminar Flow	Turbulent Flow	Relative C
Staggered	2995	11.4	1.0
Parallel	1366	5.2	0.46
Random	1576	6.0	0.53

Table 2 shows that the highest resistance is produced by staggered patterns for a given element density, whereas a random pattern yields somewhat more than half of that for staggered pattern. For the laminar flow, Hartley assumed that the relative pattern effect determined for turbulent flow was valid in the laminar range in order to avoid the lack of data in that range. However, no evidence has been provided to justify that assumption.

The conditions and restrictions on using Hartley's equations are as follows: (1) flow is laminar when  $R_d < 150$  and is turbulent otherwise -  $R_d$  may be replaced by  $(V_{max} \cdot d) / \nu = (S / (S-d)) \cdot (V \cdot d) / \nu$  in which  $(S-d)$  equals the gap size; (2) the vegetation surface is smooth and either no flexibility effect occurs or the flow is very shallow; (3) the vegetation pattern can be identified as one of staggered, parallel, or random; (4) the vegetation density is approximately constant along the height of stems; and (5) the equations only give the vegetation resistance.

### 6.3. Flexible Artificial Vegetation

The effect of flexibility of vegetation simulated by artificial turf on resistance to sheet flow was noticed by Fenzel (1964). He introduced a dimensionless deflection parameter,  $V^2Y^4/J$ , in which  $J = EI$ ,  $E =$  module of elasticity of the vegetation material, and  $I =$  moment of inertia of the turf cross section. For his particular studies on irrigation systems, Fenzel dropped this parameter from dimensional analysis because of no bending effect or other deflection of the vegetation in his experiments.

Hoerner (1965) modified the drag coefficient for a prismatic element by a factor equal to the cube of the cosine of the angle between the element and normal to the flow direction. This factor takes the degree of flexibility into account and implies that the drag coefficient for a flexible element is less than that of a rigid one. Obviously, the method can not be applied when the elements are semi-rigid which may be bent with varying angle and also the method holds for sparse vegetations.

More experiments on dense synthetic flexible turf were carried out by Phelps (1970). He did his experiments with artificial turf of raffia sewn to a jute fabric base. His procedure was to test the variation of  $f$  with  $R_e$  for different constant depths. This was accomplished for a series of depths by adjusting discharge to achieve these depths on a given slope. The reason for choosing constant depths with varying Reynolds number was to reflect the effect of decreasing vegetation density with the distance from the boundary, similar to natural grass. Phelps then found that the product of  $f.R_e$  was not a constant for

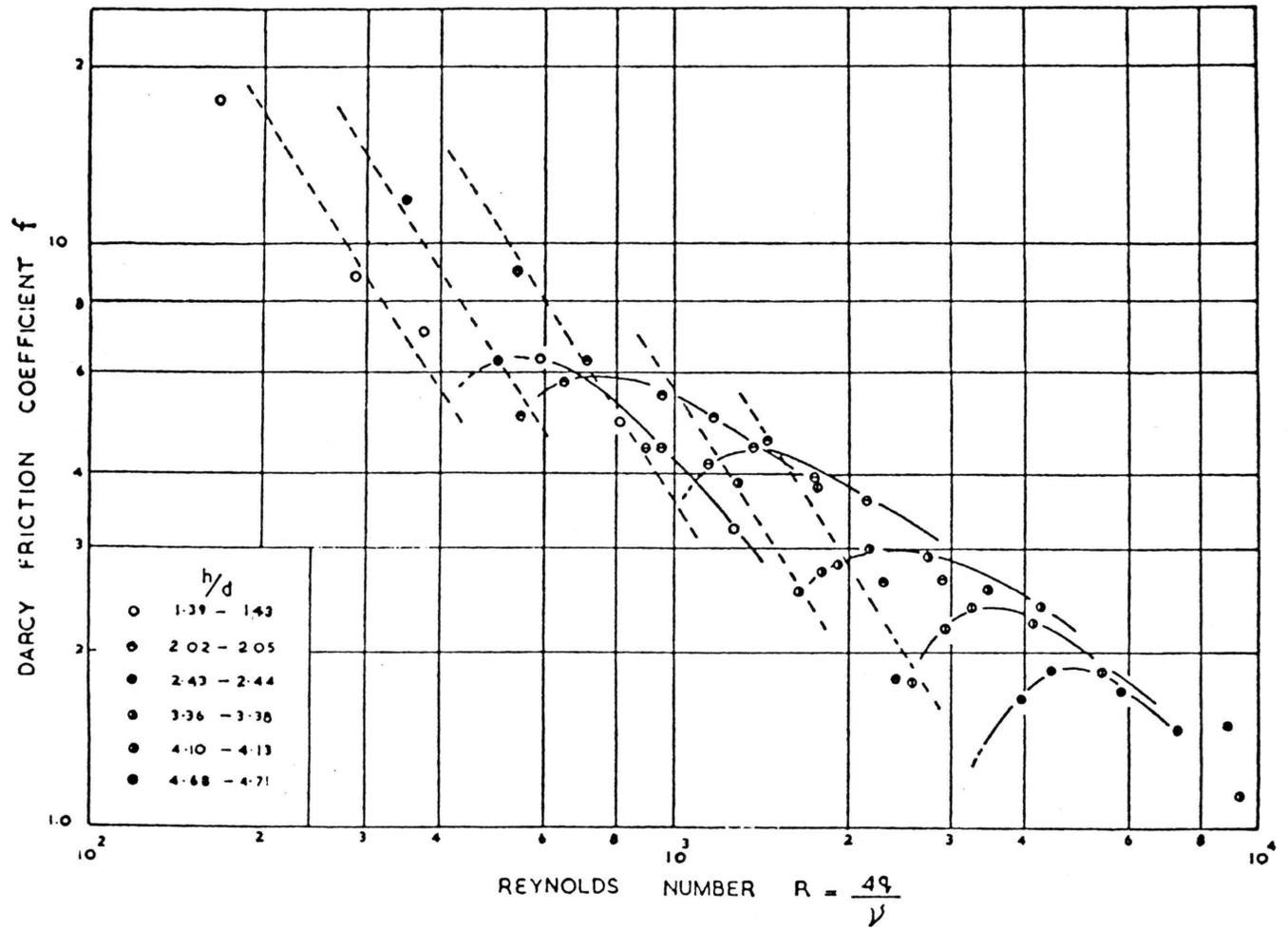


Fig. 8. The  $f$ - $R_e$  relationship for flexible artificial turf, after Phelps (1970).

laminar flow but rather decreasing with increase in  $R_e$  for every constant depth. This means a steeper slope than -1 on log-log paper which is the theoretical slope. Phelps (1970) explained this departure in terms of the flexibility of the synthetic turf in response to the flow condition. As the Reynolds number and velocity increased, the expansion of average pore size caused steeper decrease in resistance.

The data are depicted in Fig.8 illustrating  $f$  vs  $R_e$  for constant values of  $h/d$ , where  $h$  is flow depth and  $d$  is flow passage dimension which was set to .01 feet due to assumed similarity of flow through turf with groundwater flow through porous media, with convection  $d$  being .01. Therefore, constant lines of  $h/d$  represent constant depths. If one traces constant depth line in the direction of increasing  $R_e$  or discharge, he will find that the slope is increasing in that direction. As a result, the values of constant slope lines should decrease from the bottom to the top in direction of increasing  $f$ . Now, if for constant  $R_e$  or discharge the bed slope is reduced, the flow depth will increase and so will resistance. However, as will be indicated later, the same change in slope in Chen's data for natural vegetation causes less resistance. One may reason the difference in terms of the ability of contraction of pores due to lower velocity over the ability of the flow to find larger pores at higher depths in Phelps' tests. This is probably one difference between behavior of artificial turf and the natural one.

Although the adequacy of Phelps' data is in doubt particularly for higher depths, Phelps made three important conclusions for sheet flow through dense flexible artificial vegetation : (1) the varying density

of vegetation with depth has to be accounted for; (2) for constant depth, pore or flow tube size can expand as the velocity increases due to vegetation flexibility; and (3) the critical Reynolds number marking the limit of laminar flow decreases with the decrease in depth.

#### 6.4. Natural Vegetation

The early investigations of the flow resistance in a laminar flow through natural vegetation dates back to attempts to determine K value in Eq.16. As the first investigator, Izzard (1944) conducted a series of experiments on the laminar flow with the rainfall over a turf surface covered with Kentucky Blue grass. He found K to be as high as 10,000 for bed slope being .01 and with any rainfall intensity.

An extensive study on effect of specific natural vegetation on resistance to sheet flow was carried out by Chen (1976). Bermuda grass and Kentucky Blue grass were used as the typical vegetation in overland areas. Through a dimensional analysis with considering test results, Chen assumed Reynolds number, slope, relative roughness  $k/Y$ , and rainfall intensity as the independent variables in dimensional analysis. Chen concluded that the effect of the rainfall would decrease with increase in roughness size,  $k$ , and bottom slope and therefore it may be neglected for high roughness boundary of grassed area. Later, he dropped  $k$  from the analysis for sake of simplicity and difficulties involved in  $k$  measurement. Finally, the remaining variables became  $R_e$  and slope, i.e.  $f = \text{func}(R_e, S_o)$ . The regression analysis showed that K value for laminar flow through Bermuda grass began from 5000 up to 500,000 for slopes being .001 to .555 respectively. It was also found

that the upper limit of  $R_e$  for laminar flow decreased from  $10^4$  for  $S_o=.001$  to  $10^3$  for  $S_o=.555$ . The equation suggested by Chen to be applied for Bermuda grass and Kentucky Blue grass surfaces in the laminar range is:

$$f = \frac{510,000 S_o^{.662}}{R_e} \quad (50)$$

The increase in slope, if considered as an independent variable, would increase the friction factor of flow on a rough surface when discharge and other parameters held constant. The case of natural vegetation with higher density near the bed yields the same effect for bed slope. To reason such an effect, Kruse et al. (1965) explained the phenomena by considering the correspondence of increase in slope and decrease in depth for constant discharge and therefore higher average density opposing the flow. This trend resulted from Chen's tests on Bermuda grass.

Hartley (1980) superimposed the constant depth lines on Chen's data, as shown in Fig.9. Hartley confirmed the reason stated by Kruse et al. (1965) that for constant slope, resistance decreases as depth increases indicating lower average density of vegetation with increasing depth. Another trend in Fig.9 may be observed along constant depth lines. Generally, the friction factor grows along the path such that the tangent slope to the path starts from zero and increases toward infinity. This implies that constant depth at higher slope ranging from .001 to .164 and higher  $R_e$  up to some extent, corresponds to a higher friction factor. Obviously, the preceding conclusion is in

contradiction with the case of flow over a rough boundary in which friction factor decreases with slope and  $R_e$  with depth held constant. Hartley explains that the increase in resistance along constant depth lines in Chen's data could be due to either instability in free surface as velocity increases or flexibility effects. The former effect requires additional energy dissipation and the latter causes an increase in biomass brought down into the flow due to bending. Kouwen and Unny (1973) state that this effect of flexibility increases resistance as long as the vegetation is not totally overtopped or channelized by the flow.

In the second part of constant depth line in Chen's data,  $f$  tends to grow very rapidly with constant  $R_e$  and consequently discharge. The trend is true for depths being larger than 0.1 feet and when  $S_o > 0.164$ . This indicates that for steep slope with constant depth, the flow resistance becomes independent of  $R_e$  when  $R_e > 700$  and apparently flow enters the transition regime. Therefore, the upper limit for  $R_e$  for laminar regime in Chen's data would be probably close to 700 for slopes steeper than 0.164, whereas Chen extends it to 1100. One may reason the phenomenon for steep slope in terms of high free surface instability causing turbulence and making the flow exit from laminar regime. For practical purposes, however, a steeper slope ( $S_o > 0.164$ ) rarely occurs and the Chen's data on resistance to flow through Bermuda grass can be used for mild slope when  $R_e$  is as large as  $10^4$ .

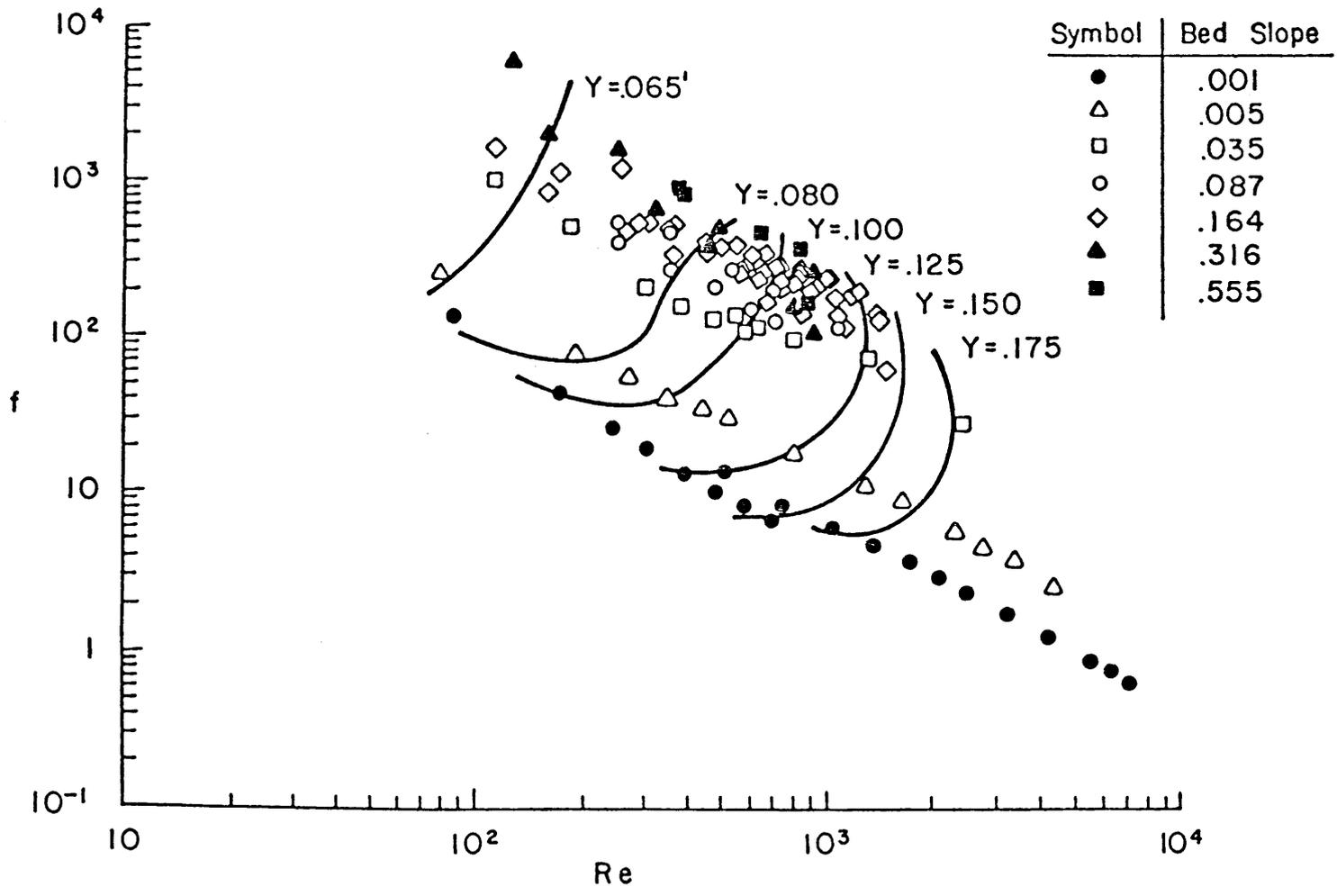


Fig. 9. The  $f-R_e$  restricted data for flow through Bermuda grass, after Chen (1976).

Even though there exist a debate concerning whether the bed slope can be an independent variable, Chen's data confirms a good agreement in laminar region with the equation  $f=K/R_e$ . Since Chen's equation directly computes the total resistance, there is no need to separate the boundary resistance and deal with it. Also, the equation comes from the experiments in which more similarity with natural situation occurs, particularly density variation with depth in addition to flexibility effect. The comparison of the data and the equation is shown in Fig.10.

Similar data on flow through Bermuda grass has been presented by Palmer (1945). Palmer data along with Chen data are plotted in Fig.10. Although most of the Palmer data fall within laminar range as indicated by Chen, it shows an almost constant  $f$  through the laminar range rather than decreasing  $f$  with  $R_e$ . Chen reasons the discrepancy in the results between his and Palmer's study in terms of high difficulties involved in depth measurements with such thin flows. Whatever the reason, the Palmer data in laminar range can not be trusted because showing nearly constant  $f$  in that range means the relative independency of resistance from Reynolds number that might be true for turbulent flows.

Ree and Palmer (1949) performed extensive experiments on resistance to turbulent flow through various grasses, particularly Bermuda grass, in two different channel cross sections, trapezoidal and rectangular, with channel slope ranging from 0.002 to 0.24. They plotted curves of Manning's  $n$  versus the product of velocity and hydraulic radius. Also the results of experiments identified three conditions of vegetal roughness system in terms of flexibility: (1) erect condition corresponding to low flows with high resistance, constant  $n$  until

partial submergence occurs; (2) deflected condition at intermediate flow, decreasing resistance with discharge, beyond complete submergence; and (3) prone condition at high flows and low resistance above the flattened vegetation, fully turbulent flow with constant  $n$ . Having Ree and Palmer data including the variation of  $n$  vs  $VR$  and the temperature at the time of experiments, Chen derived  $f$  vs  $R_e$  using the relation between  $f$  and  $n$  and then plotted the results along with his own data in Fig.10. Three interesting conclusions are revealed from Fig.10. First, the Ree and Palmer data falls mostly into transition and turbulent ranges, having a steep drop in resistance in transition range and terminating to, as Chen puts it, a fixed  $f$  when entering fully turbulent flow. The fixed  $f$  value is claimed to be 0.11 for  $R_e$  larger than  $10^6$ . However, almost all of the curves of  $n$  vs  $VR$  provided by Ree and Palmer terminates to a constant value for  $n$  indicating a fully turbulent flow independent of Reynolds number. Since  $n$  is proportional to  $f^{1/2}R^{1/6}$ , then constant  $n$  doesn't mean constant  $f$  while  $R_e$ , or discharge, is increasing. Therefore, referring to fixed  $f$  in  $f$ - $R_e$  diagrams, without having data in apparently constant  $f$  region, cannot be true and connection of two broken curves in Fig. 10 only indicates the independency of  $f$  from channel cross section for fully turbulent flow. In order to derive  $f$ - $n$  relationship and use it for fully turbulent region, one can use the Manning equation in addition to Eq.12 and then eliminate the depth parameter by introducing  $R_e$  into Eq.12. It yields:

$$f = 8(1.49)^{-1.8}n^{1.8}gS^{.1}R_e^{-.2}v^{-.2} \quad (52)$$

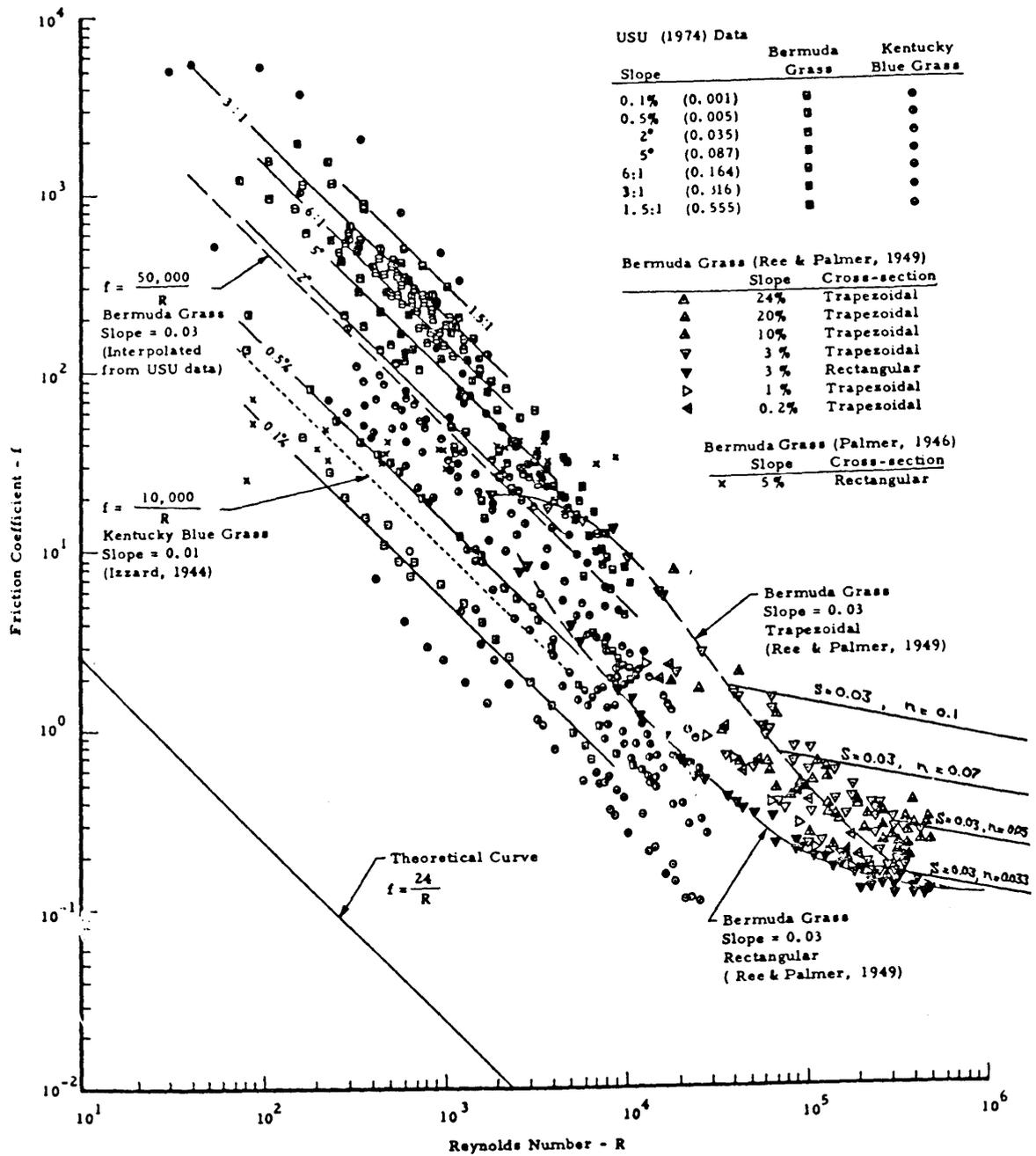


Fig. 10. The  $f$ - $R_e$  relationship for flow through Bermuda and Kentucky grasses, after Chen (1976).

For  $\nu = 1.5 \times 10^{-5}$  ft<sup>2</sup>/s,  $g = 32.2$  ft/s<sup>2</sup>, and specific slope being 0.03, the equation simplifies to :

$$f = 819.98n^{1.8}R_e^{-.2} \quad (53)$$

The Ree and Palmer's n-VR curves indicates a constant n being 0.033, corresponding to the line shown in Fig. 10, for fully turbulent flow when S=0.03. As it is seen that the f-n line extends the broken curves of f-R<sub>e</sub> from transition region into fully turbulent flow.

Second, the variation of f in the transition range may differ with the cross section shape for the same slope. Two broken curves in Fig.10 connect the data for trapezoidal and rectangular cross sections for 3% channel slope. The trapezoidal resistance curve represents larger f compared to that of a rectangular one for similar R<sub>e</sub>, or discharge. Equal discharge in rectangular and trapezoidal cross sections requires larger depth in rectangular channel, corresponding to less resistance. This trend was also derived from Chen's data in laminar region and was explained in terms of less vegetation density at higher depths in addition to lower resistance due to flexibility effects.

Third, both broken curves seem to meet at approximately R<sub>e</sub> = 2000 at a point that flow on the 3% slope starts to deviate from the laminar region to the transition. Interestingly, the point of intersection between two broken curves almost lies on the line representing f-R<sub>e</sub> relationship in Chen's equation for laminar flow on 3% channel slope. This indicates a good agreement between Chen's and Ree and Palmer's data.

### 6.5. Deep Flow over Flexible Vegetation

The importance of vegetation flexibility on relative roughness and flow resistance was suggested by Fenzel and Davis (1964) through a series of experiments on artificial turf. Element stiffness, spacing, and shape as well as fluid properties and flow parameters were realized to affect the flow resistance. Fenzel and Davis showed that the vegetation resistance was dominant over soil resistance, even though they couldn't evaluate the significance of flexibility parameters in their analysis due to lack of data. They also noticed the importance of soil resistance only at small depths in sparsely vegetated channels whereas it could be ignored for most deep flows in densely vegetated channels.

Probably, the most comprehensive analysis, which will be explained in details, of velocity profile and flow resistance in presence of flexible vegetation in deep flows was accomplished by Kouwen and his colleagues. Kouwen et al.(1969) and then Kouwen and Unny (1973) developed a semilogarithmic velocity profile equation by introducing a new relative roughness,  $Y/K$ , to account for the deflection effect of flexible vegetation.  $Y$  is simply flow depth and  $K$  stands for the deflected height of the vegetation. The equation is :

$$\frac{V}{V_*} = C_1 + C_2 \ln\left(\frac{Y}{K}\right) \quad (53)$$

where  $C_1$  and  $C_2$  are constants for a given vegetation type and density.  $C_1$  depends mainly on the flow through the vegetation and hence will be a function of its density. For small depths when  $Y < K$ , the equation

reduces to  $V/V_* = C_1$  or by substituting for  $V_*$ , it is obtained that  $\tau_0 = \rho V^2/C_1^2$  which looks like the familiar drag equation where  $C_D = 2/C_1^2$ . Since  $C_D$  is directly proportional to the number of stems per unit area, it becomes clear that  $C_1$  is dependent on the density of the vegetation.  $C_2$ , on the other hand, is related to vegetation stiffness.

Kouwen and Unny (1973) used flexible plastic strips to model and determine  $C_1$  and  $C_2$  for different conditions : prone and otherwise. The prone condition was found when the shear velocity exceeded a critical shear velocity as follows :

$$V_* > V_{*c} = 0.028 + 6.33 (\text{MEI})^2 \quad (54)$$

where MEI = a bulk stiffness parameter. The above relationship was primarily developed for elastic roughness which returns to its initial position after cessation of the flow. An analysis of Eastgate's (1966) data revealed that for tall natural grasses the critical shear velocity given by Eq.54 was too high. For natural long stiff grasses, which acts plastically under the flow, Eastgate's data indicated that :

$$V_{*c} = 0.23 (\text{MEI})^{.106} \quad (55)$$

Thus Eq.54 represents the shear velocity required to elastically bend the roughness to a prone condition and Eq.55 represents the plastic case. Both equations, which are not dimensionless, are in SI units. In practice, the smaller value between Eq.54 and Eq.55 is recommended to be used.

Assuming the validity of semilogarithmic velocity profile, the resistance coefficients,  $f$  and  $n$ , can be written in SI units as :

$$\frac{1}{\sqrt{f}} = a + b \log\left(\frac{Y}{K}\right) \quad (56)$$

and

$$n = \frac{Y^{1/6}}{\sqrt{8g} [ a + b \log(Y/K) ]} \quad (57)$$

Using the data on synthetic plastic roughness, Kouwen and Unny determined  $a$  and  $b$  as 0.15 and 1.85 for  $V_*/V_{*c} < 1.0$ ; 0.20 and 2.70 for  $1.0 < V_*/V_{*c} < 1.5$ ; 0.28 and 3.08 for  $1.5 < V_*/V_{*c} < 2.5$ ; and 0.29 and 3.50 for  $V_*/V_{*c} > 2.5$ .

Kouwen and Li (1980) established an equation to evaluate the deflected height of the vegetation, in SI units, as :

$$K = 0.14 h \left[ \left( \frac{MEI}{\rho V_*^2} \right)^{.25} / h \right]^{1.59} \quad (58)$$

The remaining difficulty is the value of MEI (in  $N.m^2$ ) for each grass type. Because there were no reported measurements of the deflected vegetation heights,  $K$ , for the experiments modeling the flow over natural vegetations, Kouwen and Li used a backward method to calibrate MEI values. They collected the experimental data of Chen (1975), Cox and Palmer (1948), Eastgate (1966), and Ree and Palmer (1949) including measurement of vegetation height,  $h$ , flow velocity,  $V$ , and effective slope,  $S_f$ . In their method, Kouwen and Li assumed an initial value for MEI for each grass. Then they calculated  $K$ ,  $n$ ,  $V_{cal}$ , and  $Q_{cal}$  for each individual experiment. That assumed value of MEI, which gave the smallest summation among the differences between calculated discharges

and corresponding measured discharges for all experiments with one grass, was tabulated as the value of MEI for that specific grass. The table was confirmed by computing retardance curves,  $n$  vs VR, and comparing with the measured retardance curves presented by Chen, Cox and Palmer, and the others. The good fit between the retardance curves was assumed to be an indication to justify the use of flexible plastic strips to model the flow over natural vegetation. Finally, Kouwen and Li proposed an iterative procedure for the design of a channel with vegetative lining. Kouwen (1969) classified flow through and over vegetation according to whether vegetation was erect and stationary, bent and waving, or prone. Shen and Li (1973) cited element waving as a possible mechanism increasing flow resistance. However the method by Kouwen and Li (1980) doesn't consider the element waving as a middle condition between erect and prone and only deflection effect contributes in the equations. Even though no report of applying Kouwen and Li's method is available, the method can be considered as a collection of existing data on turbulent flow resistance through various natural vegetations.

## 7. CONCLUSIONS

The following conclusions emerged from the discussion of the literature on resistance to sheet flows:

(1) total resistance in sheet flow can be represented by the sum of resistances due to rainfall, roughness, and vegetation;

(2) the relative roughness may represent a more general variable compared to bed slope, in flow resistance equation for laminar flow over a rough boundary. According to Phelps' paper, the friction factor equation in the form  $f = K/R_e$  has been verified.  $K$  is constant for a given relative roughness;

(3) the friction factor for turbulent flow depends on the condition of roughness related to the flow. Flow resistance under hydraulically smooth conditions is a function of Reynolds number whereas under fully rough condition the primary variable becomes the relative roughness;

(4) the friction factor, here defined as  $8\tau/\rho V^2$ , depends on Reynolds number and rainfall intensity for laminar flow over a smooth boundary and only on Reynolds number for turbulent flow. The resistance equation given by Li (1972) is recommended for the computation of flow resistance with rainfall;

(5) flow through vegetation is very complicated. Nevertheless, in limited number of cases several methods can be applied. Chen's equation is suggested for total friction factor due to laminar flow through Bermuda and Kentucky Blue grasses. For either flow through rigid vegetation with constant density along depth of flow, or very shallow flow through grass, Hartley's equations may be used to compute friction

factor for different vegetation patterns in both laminar and turbulent flow;

(6) in case of deep turbulent flow through natural vegetation, Ree and Palmer's resistance curves can provide Manning's  $n$ . Also in this case Kouwen and Unny's method is suitable for channel design with vegetative lining; and

(7) the relative magnitude of resistance to flow due to rainfall, roughness, and vegetation (represented by Bermuda grass) shows that rainfall resistance and roughness resistance for laminar flow are generally comparable whereas vegetation resistance drastically overcomes that of both rainfall and roughness combined.

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## APPENDIX II - LIST OF SYMBOLS

The following symbols are used in this paper:

- A = cross sectional area;
- C = concentration of roughness elements; also Chezy C;
- $C_D$  = drag coefficient of vegetation elements;
- D = average diameter;
- d = rainfall size; also diameter of vegetation elements;
- D = pipe diameter; also depth;
- EI = stiffness of vegetation;
- F = Froude number =  $V/\sqrt{gy}$ ;
- f = Darcy-Weisbach friction factor;
- g = gravitational acceleration;
- G = average gap size;
- h = vegetation height;
- $h_f$  = head loss in pipes;
- i = rainfall intensity;
- K = deflected height of the vegetation; also constant for  
description of f- $R_e$  relationship;
- k = mean boundary roughness height;
- N = number of cylinders per unit area;
- n = Manning's n;
- q = unit discharge;
- $q_0$  = unit base flow rate in case of rainfall;
- R = hydraulic radius;
- $R_e$  = Reynolds number =  $q/v$ ;

$R_d$  = diameter Reynolds number =  $VD/v$ ;  
 $S$  = average vegetation spacing;  
 $S_o$  = bed slope;  
 $S_f$  = friction or energy gradient;  
 $T$  = free surface width of the channel;  
 $U$  = velocity of raindrop entering main flow;  
 $V$  = mean flow velocity;  
 $Y$  = average flow depth;  
 $x$  = distance in the main flow direction;  
 $\beta$  = velocity distribution factor in momentum equation;  
 $\beta_i$  = regression coefficient in regression equation;  
 $\alpha$  = rainfall pattern dimensionless quantity;  
 $\gamma$  = Specific gravity of water;  
 $\epsilon$  = error in regression equation;  
 $\lambda$  = parameter describing raindrop shape; also soil roughness spacing;  
 $\rho$  = density of water;  
 $\tau$  = boundary shear stress;  
 $\theta$  = angle between main flow direction and horizontal; also  
cross sectional shape dimensionless quantity of vegetation  
elements;  
 $\mu$  = dynamic viscosity of water;  
 $\nu$  = kinematic viscosity of water;  
 $\psi$  = dimensionless vegetation pattern parameter;  
 $\phi$  = angle between the velocity  $U$  and  $x$ -direction;  
 $\sigma$  = soil roughness height;  
 $\delta$  = laminar sublayer thickness;

APPENDIX III. Tables of Data

f-Re Data Based on Relative Roughness, after Phelps (1975)

Data number	Channel slope	Relative roughness	Unit discharge (m <sup>2</sup> /s)	Mean depth (mm)	Reynold number	Darcy f
1	0.00083	0.35	0.000067	3.35	75	0.553
2	0.00083	0.28	0.000142	4.21	161	0.242
3	0.00083	0.23	0.000253	5.00	286	0.128
4	0.00083	0.52	0.000034	2.26	36	1.225
5	0.00154	0.35	0.000125	3.38	135	0.301
6	0.00154	0.28	0.000276	4.24	296	0.121
7	0.00154	0.23	0.000466	5.00	500	0.070
8	0.00238	0.54	0.000043	2.18	46	1.079
9	0.00238	0.35	0.000180	3.35	194	0.219
10	0.00238	0.28	0.000390	4.24	423	0.094
11	0.00238	0.23	0.000511	5.16	551	0.099
12	0.00199	0.52	0.000038	2.24	40	1.220
13	0.00199	0.27	0.000333	4.29	489	0.112
14	0.00458	0.53	0.000078	2.18	85	0.613
15	0.00458	0.35	0.000314	3.30	342	0.131
16	0.00458	0.27	0.000461	4.29	507	0.134
17	0.00458	0.23	0.000578	5.08	626	1.141
18	0.00302	0.35	0.000240	3.38	257	0.159
19	0.00302	0.27	0.000446	4.27	469	0.093
20	0.00048	0.27	0.000840	4.32	92	0.432
21	0.00048	0.23	0.000142	5.03	148	0.237
22	0.00048	0.35	0.000034	3.35	36	1.240
23	0.00120	0.23	0.000360	5.00	388	0.091
24	0.00120	0.28	0.000208	4.24	220	0.166
25	0.00120	0.35	0.000095	3.35	99	0.398
26	0.00612	0.54	0.000100	2.16	104	0.485
27	0.00612	0.35	0.000401	3.35	426	0.113
28	0.00612	0.28	0.000508	4.24	538	0.143
29	0.00612	0.23	0.000659	5.05	691	0.143
30	0.00761	0.55	0.000117	2.13	126	0.424
31	0.00761	0.28	0.000528	4.24	576	0.164
32	0.00761	0.23	0.000707	5.03	756	0.153
33	0.01000	0.55	0.000145	2.13	154	0.362
34	0.01000	0.28	0.000579	4.24	621	0.180
35	0.01000	0.23	0.000770	4.98	829	0.164
36	0.01490	0.53	0.000207	2.13	216	0.287
37	0.01490	0.27	0.000689	4.27	733	0.192
38	0.01980	0.53	0.000244	2.21	265	0.284
39	0.01980	0.35	0.000499	3.35	534	0.235
40	0.01980	0.23	0.001122	5.03	1204	0.158
41	0.02970	0.53	0.000214	2.21	229	0.550
42	0.02970	0.27	0.000953	4.32	1039	0.207
43	0.02970	0.23	0.001285	5.03	1393	0.180
44	0.04510	0.53	0.000271	2.18	297	0.507
45	0.04510	0.35	0.000659	3.33	716	0.301
46	0.04510	0.27	0.001107	4.29	1194	0.229
47	0.04510	0.23	0.001515	5.03	1600	0.197

f-Re . Data for Sand Surface, after Woo and Brater (1962)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Reynolds number	Darcy f
1	0.001	0.000786	0.145	66	0.7360
2	0.001	0.002345	0.195	195	0.2008
3	0.001	0.009700	0.310	789	0.0472
4	0.002	0.006140	0.210	547	0.0732
5	0.002	0.011980	0.295	992	0.0532
6	0.002	0.016600	0.375	1340	0.0570
7	0.002	0.021650	0.450	1730	0.0580
8	0.003	0.002017	0.145	166	0.3360
9	0.003	0.006730	0.190	553	0.0676
10	0.003	0.012600	0.300	1025	0.0762
11	0.003	0.020970	0.400	1687	0.0652
12	0.003	0.046550	0.695	3695	0.0694
13	0.004	0.002323	0.130	188	0.2432
14	0.004	0.004460	0.160	354	0.1232
15	0.004	0.006110	0.180	482	0.0932
16	0.004	0.023970	0.430	1870	0.0828
17	0.004	0.047700	0.650	3720	0.0720
18	0.006	0.003834	0.140	311	0.1672
19	0.006	0.008900	0.210	707	0.1048
20	0.006	0.017470	0.315	1384	0.0920
21	0.006	0.027560	0.430	2188	0.0936
22	0.006	0.048500	0.595	3795	0.0804
23	0.008	0.005800	0.145	469	0.1080
24	0.008	0.011500	0.230	919	0.1096
25	0.008	0.021060	0.320	1670	0.0880
26	0.008	0.022950	0.360	1811	0.1056
27	0.008	0.048100	0.550	3768	0.0856
28	0.010	0.002130	0.110	157	0.4380
29	0.010	0.005060	0.145	367	0.1780
30	0.010	0.007240	0.160	523	0.1164
31	0.010	0.009970	0.215	718	0.1496
32	0.010	0.009210	0.200	732	0.1404
33	0.010	0.014920	0.270	1080	0.1316
34	0.010	0.015580	0.260	1221	0.1080
35	0.010	0.018250	0.305	1330	0.1272
36	0.010	0.022400	0.340	1740	0.1168
37	0.010	0.047200	0.520	3664	0.0940
38	0.010	0.074100	0.685	5820	0.0872
39	0.015	0.002256	0.105	167	0.5100
40	0.015	0.004330	0.125	316	0.2328
41	0.015	0.004450	0.120	367	0.1952
42	0.015	0.007310	0.160	530	0.1716
43	0.015	0.008680	0.185	626	0.1880
44	0.015	0.014830	0.250	1070	0.1592
45	0.015	0.014370	0.230	1130	0.1320
46	0.015	0.019040	0.290	1393	0.1504
47	0.015	0.026650	0.355	2178	0.1408
48	0.015	0.052100	0.525	4075	0.1192
49	0.015	0.092100	0.690	7100	0.0868
50	0.020	0.001824	0.095	136	0.7720

f-Re Data for Sand Surface, after Woo and Brater (1962)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Reynolds number	Darcy f
51	0.020	0.004275	0.115	313	0.2480
52	0.020	0.007100	0.150	516	0.2000
53	0.020	0.010800	0.195	785	0.1900
54	0.020	0.011770	0.195	928	0.1600
55	0.020	0.014770	0.235	1076	0.1772
56	0.020	0.018350	0.265	1360	0.1648
57	0.020	0.020330	0.260	1570	0.1268
58	0.020	0.045600	0.445	3490	0.1264
59	0.020	0.066800	0.535	5110	0.1024
60	0.020	0.094800	0.640	7160	0.0868
61	0.040	0.001812	0.075	136	0.7660
62	0.040	0.004370	0.095	321	0.2684
63	0.040	0.006900	0.125	503	0.2444
64	0.040	0.009300	0.145	676	0.2108
65	0.040	0.011580	0.170	840	0.2188
66	0.040	0.016660	0.205	1220	0.1856
67	0.040	0.019700	0.175	1530	0.0828
68	0.040	0.023500	0.250	1780	0.1688
69	0.040	0.044250	0.360	3340	0.1424
70	0.040	0.063500	0.435	4800	0.1220
71	0.040	0.089500	0.555	6690	0.1272
72	0.060	0.001795	0.065	131	0.7620
73	0.060	0.007590	0.115	538	0.2368
74	0.060	0.012080	0.165	857	0.2756
75	0.060	0.016180	0.200	1048	0.2732
76	0.060	0.029600	0.285	2110	0.2368
77	0.060	0.028080	0.240	2165	0.1568
78	0.060	0.039240	0.315	2880	0.1820
79	0.060	0.043700	0.340	3300	0.1844
80	0.060	0.054350	0.385	4080	0.1728
81	0.060	0.080800	0.475	6020	0.1472
82	0.060	0.099100	0.530	7380	0.1356

## f-Re Data with Rainfall, Laminar Flow, after Yoon (1970)

Data number	Rainfall intensity (in/h)	Base flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
1	0.50	0.00254	0.00270	0.01000	249.0	0.16849
2	1.25	0.00237	0.00278	0.01060	247.8	0.16626
3	3.75	0.00150	0.00272	0.01167	245.3	0.25289
4	15.00	0.00141	0.00280	0.01467	249.9	0.41389
5	0.50	0.00368	0.00384	0.01092	354.1	0.11359
6	1.25	0.00339	0.00380	0.01208	348.8	0.15227
7	3.75	0.00248	0.00370	0.01308	345.4	0.19871
8	15.00	0.00182	0.00390	0.01625	349.4	0.29931
9	0.50	0.00471	0.00487	0.01250	449.0	0.10322
10	1.25	0.00463	0.00504	0.01308	443.2	0.11121
11	3.75	0.00366	0.00488	0.01517	444.4	0.17171
12	15.00	0.00153	0.00500	0.01733	449.4	0.23023
13	0.50	0.00585	0.00601	0.01300	548.0	0.07943
14	1.25	0.00554	0.00595	0.01383	547.9	0.09632
15	3.75	0.00479	0.00601	0.01625	544.9	0.14077
16	15.00	0.00182	0.00599	0.01792	539.3	0.18611
17	0.50	0.00693	0.00709	0.01433	648.3	0.07310
18	1.25	0.00666	0.00707	0.01505	647.6	0.08801
19	3.75	0.00585	0.00707	0.01683	644.1	0.11497
20	15.00	0.00218	0.00704	0.01933	630.4	0.16550
21	0.50	0.00853	0.00869	0.01542	798.9	0.06172
22	1.25	0.00829	0.00870	0.01600	797.0	0.07085
23	3.75	0.00750	0.00872	0.01758	798.8	0.08860
24	15.00	0.00385	0.00871	0.02058	779.9	0.13246
25	0.50	0.00962	0.00978	0.01667	899.1	0.05676
26	1.25	0.00938	0.00979	0.01867	896.9	0.07671
27	3.75	0.00853	0.00975	0.02000	893.2	0.09190
28	0.50	0.00260	0.00276	0.00792	248.4	0.16373
29	1.25	0.00235	0.00276	0.00825	248.9	0.18785
30	3.75	0.00150	0.00272	0.00900	245.3	0.24179
31	15.00	0.00000	0.00278	0.01075	249.8	0.36311
32	0.50	0.00371	0.00387	0.00883	348.8	0.11909
33	1.25	0.00345	0.00386	0.00942	347.2	0.14191
34	3.75	0.00262	0.00384	0.01000	344.9	0.17122
35	15.00	0.00000	0.00417	0.01183	373.0	0.22309
36	0.50	0.00481	0.00497	0.01017	447.9	0.10332
37	1.25	0.00456	0.00497	0.01067	447.3	0.11845
38	3.75	0.00372	0.00494	0.01150	444.6	0.14740
39	15.00	0.00000	0.00486	0.01317	434.0	0.22842
40	0.50	0.00592	0.00608	0.01092	547.9	0.08585
41	1.25	0.00565	0.00606	0.01158	547.0	0.10214
42	3.75	0.00481	0.00603	0.01250	544.3	0.12577
43	15.00	0.00113	0.00599	0.01433	531.1	0.18143
44	0.50	0.00701	0.00717	0.01158	647.9	0.07348
45	1.25	0.00677	0.00718	0.01208	646.4	0.08517
46	3.75	0.00594	0.00716	0.01317	643.3	0.10261
47	15.00	0.00229	0.00715	0.01508	634.0	0.15137
48	0.50	0.00869	0.00885	0.01233	797.5	0.05882
49	1.25	0.00841	0.00882	0.01267	796.3	0.06544
50	3.75	0.00759	0.00881	0.01408	793.3	0.08402
51	15.00	0.00395	0.00881	0.01608	781.1	0.11945

f-Re Data with Rainfall, Laminar Flow, after Yoon (1970)

Data number	Rainfall intensity (in/h)	Base flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
52	0.50	0.00977	0.00993	0.01292	897.2	0.05313
53	1.25	0.00951	0.00992	0.01325	895.7	0.06023
54	3.75	0.00867	0.00989	0.01458	893.0	0.07404
55	15.00	0.00497	0.00983	0.01650	880.1	0.10364

f-Re Data with rainfall, Laminar Flow, after Li (1972)

Data number	Rainfall intensity (in/h)	Base flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
1	7.5	0.00000	0.00148	0.00950	127.1	0.66496
2	7.5	0.00382	0.00530	0.01584	459.3	0.20148
3	7.5	0.00835	0.00983	0.01960	864.2	0.10132
4	7.5	0.00056	0.00204	0.01070	181.6	0.47095
5	7.5	0.00225	0.00373	0.01173	334.7	0.23783
6	15.0	0.00000	0.00295	0.01266	272.5	0.41859
7	15.0	0.00313	0.00608	0.01669	557.4	0.19564
8	15.0	0.00176	0.00471	0.01533	431.8	0.23227
9	15.0	0.00673	0.00968	0.01784	880.9	0.12677
10	17.5	0.00000	0.00334	0.01491	294.5	0.43650
11	17.5	0.00487	0.00831	0.01952	706.3	0.19197
12	10.5	0.00000	0.00207	0.01110	176.6	0.51200
13	10.5	0.00658	0.00865	0.01899	739.6	0.13498
14	17.5	0.00371	0.00715	0.01733	595.6	0.18171
15	10.5	0.00371	0.00578	0.01616	487.4	0.19630
16	17.5	0.00000	0.00304	0.01422	254.7	0.48261
17	17.5	0.00200	0.00504	0.01837	422.3	0.31544
18	17.5	0.00552	0.00856	0.02178	722.2	0.19090
19	10.5	0.00000	0.00182	0.01092	160.3	0.66452
20	10.5	0.00779	0.00961	0.01828	857.5	0.09497
21	10.5	0.00275	0.00457	0.01509	405.0	0.24792
22	12.5	0.00000	0.00217	0.01288	180.7	0.65689
23	12.5	0.00464	0.00681	0.01816	619.7	0.16169
24	12.5	0.00774	0.00991	0.02079	842.0	0.09710
25	17.5	0.00000	0.00952	0.02271	736.2	0.15090
26	17.5	0.00000	0.00344	0.01559	266.3	0.43960
27	12.5	0.00000	0.00680	0.01891	558.7	0.20812
28	12.5	0.00000	0.00246	0.01260	199.1	0.49223
29	12.5	0.00238	0.00918	0.02102	743.3	0.15252
30	12.5	0.00238	0.00484	0.01537	400.3	0.24377
31	7.5	0.00000	0.00148	0.00829	126.2	0.74339
32	7.5	0.00382	0.00530	0.01290	459.3	0.18719
33	7.5	0.00835	0.00983	0.01519	864.2	0.09660
34	7.5	0.00056	0.00204	0.00868	181.6	0.42328
35	7.5	0.00225	0.00373	0.00104	334.7	0.27483
36	15.0	0.00000	0.00295	0.01114	274.5	0.40388
37	15.0	0.00313	0.00608	0.01416	553.4	0.20530
38	15.0	0.00176	0.00471	0.01290	431.8	0.23961
39	15.0	0.00673	0.00968	0.01566	880.9	0.11916
40	17.5	0.00000	0.00344	0.01270	296.6	0.40490
41	17.5	0.00487	0.00831	0.01711	711.1	0.17250
42	10.5	0.00000	0.00207	0.01005	176.7	0.57594
43	10.5	0.00658	0.00865	0.01501	739.6	0.11721
44	17.5	0.00371	0.00715	0.01570	599.6	0.18398
45	10.5	0.00371	0.00578	0.01361	487.4	0.18901
46	12.5	0.00000	0.00217	0.01015	175.9	0.59854
47	12.5	0.00583	0.00800	0.01527	670.6	0.14492
48	12.5	0.00116	0.00333	0.01104	279.1	0.33646
49	17.5	0.00000	0.00304	0.01224	245.8	0.46890
50	17.5	0.00534	0.00838	0.01620	688.4	0.15486
51	10.5	0.00000	0.00182	0.00809	149.8	0.67114

f-Re Data with rainfall, Laminar Flow, after Li (1972)

Data number	Rainfall intensity (in/h)	Base flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
52	10.5	0.00246	0.00428	0.01139	351.9	0.25592
53	10.5	0.00537	0.00719	0.01434	602.9	0.15105
54	17.5	0.00000	0.00952	0.01806	742.0	0.15712
55	17.5	0.00000	0.00344	0.01197	268.4	0.36170
56	12.5	0.00000	0.00680	0.01487	555.1	0.18122
57	12.5	0.00000	0.00246	0.01104	200.8	0.51461
58	12.5	0.00238	0.00918	0.01592	746.9	0.14248
59	12.5	0.00238	0.00484	0.01347	403.0	0.27640

f-Re Data with Rainfall, Turbulent Flow, after Yoon (1970)

Data number	Rainfall intensity (in/h)	Base Flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
1	0.50	0.02690	0.02706	0.02833	2498.8	0.03840
2	1.25	0.02680	0.02721	0.02925	2493.6	0.04091
3	3.75	0.02596	0.02718	0.02883	2490.9	0.04024
4	15.00	0.02273	0.02759	0.03042	2479.0	0.04424
5	0.50	0.04325	0.04341	0.03583	4008.5	0.02992
6	1.25	0.04340	0.04381	0.03708	3993.2	0.03147
7	3.75	0.04280	0.04402	0.03767	3994.1	0.03247
8	15.00	0.03933	0.04419	0.03833	3974.0	0.02774
9	0.50	0.05923	0.05939	0.04333	5499.3	0.03109
10	1.25	0.05980	0.06021	0.04392	5488.2	0.02946
11	3.75	0.05863	0.05985	0.04458	5485.4	0.03128
12	15.00	0.05610	0.06096	0.04550	5477.2	0.03050
13	0.50	0.02743	0.02759	0.02300	2492.5	0.04060
14	1.25	0.02725	0.02766	0.02283	2191.4	0.03844
15	3.75	0.02643	0.02756	0.02308	2489.2	0.04101
16	15.00	0.02281	0.02767	0.02350	2477.3	0.03887
17	0.50	0.04399	0.04415	0.02892	3988.4	0.03270
18	1.25	0.04394	0.04435	0.02908	3987.9	0.03119
19	3.75	0.04290	0.04412	0.02925	3985.1	0.03296
20	15.00	0.03953	0.04439	0.02908	3974.1	0.03111
21	0.50	0.06088	0.06104	0.03525	5539.2	0.02967
22	1.25	0.06030	0.06071	0.03525	5483.7	0.03000
23	3.75	0.05956	0.06078	0.03567	5490.1	0.03089
24	15.00	0.05637	0.06123	0.03408	5467.1	0.02690

f-Re Data with Rainfall, Turbulent Flow, after Li ( 1972)

Data number	Rainfall intensity (in/h)	Base Flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
1	7.50	0.03951	0.04099	0.03362	3435.5	0.03910
2	7.50	0.10471	0.10619	0.05736	9209.5	0.02556
3	7.50	0.06676	0.06824	0.04401	6087.0	0.02865
4	7.50	0.04169	0.04317	0.03352	3850.6	0.03209
5	15.00	0.04415	0.04710	0.03481	4317.3	0.02931
6	15.00	0.02912	0.03207	0.02737	2939.6	0.03265
7	15.00	0.01999	0.02294	0.02397	2087.5	0.04131
8	17.50	0.05914	0.06258	0.04098	5281.3	0.03186
9	167.50	0.12076	0.12420	0.06260	10698.0	0.02363
10	17.50	0.03016	0.03360	0.02975	2894.3	0.03633
11	10.50	0.05283	0.05490	0.03869	4632.6	0.03422
12	10.50	0.03025	0.03232	0.02911	2802.8	0.03721
13	17.50	0.04019	0.04363	0.03254	3732.5	0.03265
14	17.50	0.11512	0.11816	0.06333	10038.9	0.02407
15	17.50	0.06538	0.06842	0.04409	5812.9	0.03134
16	17.50	0.02459	0.02763	0.02447	2363.4	0.03303
17	10.50	0.06765	0.06947	0.04231	6275.8	0.02765
18	10.50	0.04058	0.04240	0.03285	3677.6	0.03578
19	10.50	0.14381	0.14563	0.06679	12630.8	0.02173
20	10.50	0.03137	0.03319	0.03070	2878.8	0.04178
21	12.50	0.04715	0.04932	0.03556	4052.6	0.03341
22	12.50	0.12327	0.12544	0.06416	10307.3	0.02604
23	12.50	0.02586	0.02803	0.02647	2365.4	0.03329
24	12.50	0.08585	0.08802	0.05020	7328.9	0.02737
25	17.50	0.08585	0.08889	0.05158	7450.8	0.02595
26	10.50	0.08585	0.08767	0.05038	7300.0	0.02855
27	17.50	0.03515	0.04467	0.03619	3670.5	0.02933
28	17.50	0.03515	0.03859	0.03214	3171.2	0.03238
29	12.50	0.03677	0.04357	0.03463	3527.9	0.03668
30	12.50	0.03677	0.03923	0.03236	3223.5	0.03347
31	7.50	0.03951	0.04099	0.02615	3458.7	0.03225
32	7.50	0.10471	0.10619	0.04850	9339.1	0.02543
33	7.50	0.06676	0.06824	0.03635	6087.0	0.02724
34	7.50	0.04169	0.04317	0.02856	3850.6	0.03354
35	15.00	0.04415	0.04710	0.03006	4317.3	0.03346
36	15.00	0.02912	0.03207	0.02508	2939.6	0.03719
37	15.00	0.01999	0.02294	0.02039	2087.5	0.04333
38	17.50	0.05914	0.06258	0.03637	5281.3	0.03050
39	17.50	0.12076	0.12420	0.05411	10698.0	0.02572
40	17.50	0.03016	0.03360	0.02534	2864.3	0.03707
41	20.50	0.05283	0.05490	0.03372	4632.6	0.03406
42	10.50	0.03025	0.03232	0.02540	2802.8	0.04102
43	17.50	0.04019	0.04363	0.02939	3732.5	0.03646
44	12.50	0.04812	0.05029	0.03196	4202.0	0.03494
45	12.50	0.12466	0.12683	0.05273	10345.0	0.02193
46	12.50	0.06584	0.06801	0.03833	5547.3	0.03316
47	12.50	0.03009	0.03226	0.02565	2631.3	0.03827
48	17.50	0.05488	0.05792	0.03341	4567.7	0.03170
49	17.50	0.02682	0.02986	0.02267	2435.4	0.03835

f-Re Data with Rainfall, Turbulent Flow, after Li ( 1972)

Data number	Rainfall intensity (in/h)	Base Flow rate (cfs/ft)	Combined flow rate (cfs/ft)	Flow depth (ft)	Reynolds number	Darcy f
50	17.50	0.10979	0.11283	0.05431	9203.0	0.02997
51	10.50	0.05075	0.05257	0.03190	4288.2	0.03365
52	10.50	0.03369	0.03551	0.02690	2976.8	0.03858
53	10.50	0.09474	0.09656	0.04757	8184.8	0.02687
54	17.50	0.03515	0.04467	0.03004	3617.0	0.03332
55	17.50	0.03515	0.03859	0.02729	3150.5	0.03315
56	12.50	0.03677	0.04357	0.02832	3527.9	0.03182
57	12.50	0.03677	0.03923	0.02656	3223.5	0.03170

f-Re Data on Bermuda grass, after Chen (1976)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Mean velocity (fps)	Darcy f	Reynolds number
1	0.001	0.0105	1.717	0.073	6.813	696
2	0.001	0.0088	1.634	0.064	8.329	585
3	0.001	0.0073	1.550	0.056	10.396	483
4	0.001	0.0059	1.469	0.047	14.258	391
5	0.001	0.0046	1.406	0.039	18.942	309
6	0.001	0.0037	1.344	0.033	26.245	246
7	0.001	0.0026	1.262	0.025	41.314	178
8	0.001	0.0013	1.157	0.013	129.365	88
9	0.001	0.0077	1.769	0.052	13.567	515
10	0.001	0.0112	1.919	0.070	8.296	745
11	0.001	0.0156	2.154	0.087	6.068	1036
12	0.001	0.0205	2.370	0.104	4.679	1362
13	0.001	0.0260	2.575	0.121	3.758	1721
14	0.001	0.0313	2.694	0.139	2.975	2070
15	0.001	0.0375	2.831	0.159	2.391	2486
16	0.001	0.0485	3.027	0.192	1.754	3210
17	0.001	0.0636	3.250	0.235	1.261	4210
18	0.001	0.0829	3.445	0.288	0.886	5483
19	0.001	0.0955	3.587	0.319	0.753	6319
20	0.001	0.1069	3.662	0.350	0.639	7071
21	0.005	0.0012	0.787	0.019	216.076	86
22	0.005	0.0029	0.957	0.036	76.655	193
23	0.005	0.0041	1.076	0.046	53.173	277
24	0.005	0.0053	1.146	0.056	39.126	354
25	0.005	0.0067	1.278	0.063	33.947	448
26	0.005	0.0080	1.374	0.070	29.382	533
27	0.005	0.1030	1.571	0.078	27.196	681
28	0.005	0.0121	1.532	0.094	18.288	800
29	0.005	0.0195	1.790	0.131	11.196	1294
30	0.005	0.0248	1.951	0.153	8.931	1647
31	0.005	0.0350	2.116	0.198	5.759	2316
32	0.005	0.0420	2.209	0.218	4.962	2661
33	0.005	0.0514	2.384	0.258	3.818	3402
34	0.005	0.0660	2.477	0.320	2.593	4369
35	0.005	0.0909	2.618	0.416	1.616	6015
36	0.005	0.1157	2.551	0.544	0.923	7655
37	0.005	0.1459	2.704	0.647	0.691	9655
38	0.005	0.1706	2.871	0.713	0.605	11287
39	0.035	0.0096	1.281	0.090	117.206	639
40	0.035	0.0200	1.799	0.133	75.142	1328
41	0.035	0.0369	1.953	0.226	28.394	2443
42	0.035	0.0443	2.129	0.249	25.571	2930
43	0.035	0.0492	2.259	0.261	24.718	3258
44	0.035	0.0551	2.346	0.282	22.075	3648
45	0.035	0.0652	2.433	0.321	17.578	4316
46	0.035	0.0743	2.482	0.359	14.401	4914
47	0.035	0.1352	2.704	0.600	5.620	8945
48	0.035	0.1649	2.788	0.709	4.144	10907
49	0.035	0.0088	1.184	0.089	110.655	585

f-Re Data on Bermuda grass, after Chen (1976)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Mean velocity (fps)	Darcy f	Reynolds number
50	0.035	0.0072	1.096	0.078	131.609	477
51	0.035	0.0058	1.010	0.069	157.528	386
52	0.035	0.0046	0.946	0.058	207.550	305
53	0.035	0.0028	0.918	0.036	511.154	186
54	0.035	0.0017	0.825	0.025	954.938	116
55	0.035	0.0083	1.229	0.081	138.873	552
56	0.035	0.0121	1.408	0.103	98.920	801
57	0.087	0.0161	1.332	0.145	117.886	1067
58	0.087	0.0300	1.740	0.207	75.659	1990
59	0.087	0.0408	1.709	0.287	38.754	2705
60	0.087	0.0529	1.914	0.331	32.523	3499
61	0.087	0.0612	1.967	0.373	26.332	4054
62	0.087	0.0751	2.014	0.447	18.808	4970
63	0.087	0.0891	2.040	0.524	13.877	5896
64	0.087	0.1002	2.053	0.585	11.188	6628
65	0.087	0.1108	2.063	0.644	9.291	7328
66	0.087	0.1217	2.079	0.702	7.868	8054
67	0.087	0.0038	0.771	0.059	404.085	254
68	0.087	0.0054	0.855	0.076	274.946	359
69	0.087	0.0072	0.945	0.091	211.224	477
70	0.087	0.0092	1.009	0.109	156.253	612
71	0.087	0.0109	1.062	0.123	129.491	725
72	0.087	0.0038	0.851	0.054	542.208	254
73	0.087	0.0054	1.023	0.063	471.334	359
74	0.087	0.0081	1.119	0.087	274.288	539
75	0.087	0.0106	1.206	0.105	201.797	703
76	0.164	0.0260	1.415	0.220	102.622	1721
77	0.164	0.3699	1.626	0.272	77.184	2443
78	0.164	0.0536	1.893	0.340	57.718	1549
79	0.164	0.0768	1.976	0.466	31.999	5082
80	0.164	0.0918	1.822	0.604	17.566	6076
81	0.164	0.1147	1.825	0.754	11.316	7589
82	0.164	0.1331	1.821	0.877	8.343	8806
83	0.164	0.1459	1.815	0.964	9.426	10090
84	0.164	0.0211	1.374	0.184	142.454	1397
85	0.164	0.0017	0.575	0.036	1520.310	116
86	0.164	0.0041	0.694	0.072	469.603	257
87	0.164	0.0055	0.743	0.089	330.415	365
88	0.164	0.0086	0.929	0.111	263.367	571
89	0.164	0.0103	0.914	0.135	176.149	681
90	0.164	0.0129	0.997	0.155	144.779	857
91	0.164	0.0170	1.117	0.182	117.873	1126
92	0.164	0.0090	0.963	0.112	267.470	598
93	0.164	0.0103	1.124	0.109	328.209	681
94	0.164	0.0113	1.039	0.130	215.328	748
95	0.164	0.0135	1.136	0.143	195.210	898
96	0.164	0.0151	1.311	0.138	241.397	1000
97	0.164	0.0180	1.361	0.158	190.204	1192
98	0.164	0.0055	0.866	0.076	524.376	365

f-Re Data on Bermuda grass, after Chen (1976)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Mean velocity (fps)	Darcy f	Reynolds number
99	0.164	0.0076	0.974	0.094	384.754	508
100	0.164	0.0094	1.071	0.105	337.078	625
101	0.164	0.0111	1.135	0.118	286.119	740
102	0.164	0.0129	1.214	0.128	261.170	857
103	0.164	0.0152	1.299	0.140	231.115	1008
104	0.164	0.0187	1.413	0.159	196.888	1240
105	0.164	0.0024	0.585	0.049	836.494	160
106	0.164	0.0044	0.756	0.070	533.275	295
107	0.164	0.0053	0.837	0.076	501.479	354
108	0.164	0.0069	0.934	0.089	413.855	459
109	0.164	0.0093	0.998	0.112	279.087	618
110	0.164	0.0113	1.076	0.126	238.627	745
111	0.164	0.0137	1.160	0.141	203.873	906
112	0.164	0.0039	0.907	0.052	1155.106	263
113	0.164	0.0084	1.043	0.097	390.817	558
114	0.164	0.0107	1.106	0.116	287.433	711
115	0.164	0.0125	1.146	0.131	232.515	833
116	0.164	0.0160	1.236	0.155	179.348	1062
117	0.164	0.0214	1.344	0.191	128.849	1421
118	0.164	0.0026	0.680	0.046	1108.368	175
119	0.164	0.0047	0.795	0.071	544.609	314
120	0.164	0.0069	0.885	0.094	351.998	459
121	0.164	0.0097	0.981	0.119	242.364	646
122	0.164	0.0125	1.062	0.142	185.235	833
123	0.164	0.0164	1.153	0.171	138.627	1089
124	0.164	0.0224	1.090	0.246	63.002	1484
125	0.316	0.0096	0.687	0.168	163.614	639
126	0.316	0.0237	1.006	0.283	84.715	1574
127	0.316	0.0429	1.254	0.410	50.487	2839
128	0.316	0.0612	1.502	0.489	42.495	4054
129	0.316	0.0785	1.586	0.594	30.492	5196
130	0.316	0.0955	1.601	0.716	21.180	6319
131	0.316	0.1117	1.600	0.838	14.454	7393
132	0.316	0.1300	1.601	0.974	11.438	8598
133	0.316	0.1637	1.600	1.227	7.202	10832
134	0.316	0.1788	1.589	1.350	5.910	11826
135	0.316	0.0019	0.777	0.029	5936.334	128
136	0.316	0.0049	0.701	0.085	658.074	329
137	0.316	0.0074	0.838	0.107	495.101	495
138	0.316	0.0103	0.911	0.135	335.348	681
139	0.316	0.0138	1.015	0.163	257.697	914
140	0.316	0.0185	1.082	0.206	172.583	1230
141	0.316	0.0259	1.181	0.263	115.357	1717
142	0.316	0.0024	0.629	0.047	1912.662	164
143	0.316	0.0047	0.640	0.089	547.837	314
144	0.316	0.0075	0.608	0.149	184.899	501
145	0.316	0.0094	0.735	0.154	209.870	625
146	0.316	0.0122	0.804	0.182	164.023	809
147	0.316	0.0173	1.003	0.206	158.917	1144

f-Re Data on Bermuda grass, after Chen (1976)

Data number	Bed slope	Discharge (cfs/ft)	Depth (in)	Mean velocity (fps)	Darcy f	Reynolds number
148	0.316	0.0211	1.027	0.247	114.099	1400
149	0.316	0.0038	0.792	0.058	1583.515	254
150	0.316	0.0069	0.736	0.113	390.015	459
151	0.316	0.0100	0.825	0.146	259.998	667
152	0.316	0.0138	0.773	0.214	113.816	914
153	0.555	0.0184	0.528	0.419	35.720	1220
154	0.555	0.0266	0.500	0.639	14.547	1763
155	0.555	0.0059	0.708	0.100	838.192	391
156	0.555	0.0101	0.803	0.152	412.302	6784
157	0.555	0.0133	0.708	0.225	165.793	881
158	0.555	0.0157	0.786	0.240	161.431	1044
159	0.555	0.0170	0.845	0.241	172.175	1126
160	0.555	0.0227	0.966	0.282	143.691	1506
161	0.555	0.0327	0.823	0.477	42.911	2168
162	0.555	0.0295	0.968	0.366	85.865	1954
163	0.555	0.0216	0.905	0.286	130.917	1432
164	0.555	0.0171	1.020	0.201	297.692	1135
165	0.555	0.0127	0.905	0.168	379.661	841
166	0.555	0.0098	0.821	0.144	470.430	653
167	0.555	0.0057	0.706	0.097	880.628	381
168	0.555	0.0221	0.643	0.413	44.830	1466
169	0.555	0.0337	0.694	0.583	24.257	2233
170	0.555	0.0457	0.896	0.611	28.477	3023
171	0.555	0.0536	0.906	0.709	21.407	3549
172	0.555	0.0620	0.998	0.745	21.363	4106
173	0.555	0.0726	1.050	0.829	18.143	4803
174	0.555	0.0794	1.050	0.907	15.193	5253
175	0.555	0.0873	1.111	0.942	14.885	5777
176	0.555	0.0964	0.907	1.276	6.625	6380
177	0.555	0.1069	0.750	1.709	3.052	7071
178	0.555	0.1187	1.183	1.204	9.706	7854
179	0.555	0.1279	0.794	1.924	2.562	8461
180	0.555	0.1384	0.835	1.987	2.515	9156
181	0.555	0.1514	0.851	2.134	2.223	100178
182	0.555	0.1592	0.859	2.223	2.068	10533
183	0.555	0.1718	0.895	2.302	2.009	11363
184	0.555	0.1764	0.879	2.406	1.807	11671
185	0.555	0.1799	0.903	2.390	1.880	11904
186	0.555	0.0092	0.750	0.147	408.776	612
187	0.555	0.0187	0.584	0.385	46.884	1240
188	0.555	0.0264	0.530	0.599	17.586	1751
189	0.555	0.0337	0.932	0.434	58.832	2233

f-Re Data on Bermuda grass, after Ree and Palmer (1949)

Trapezoidal Shape, Bottom Width 1.5 ft, Bed Slope 24%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	0.950	3.090	0.2345	0.921	37300
2	1.850	4.300	0.2308	0.607	67200
3	2.900	5.300	0.2276	0.469	98600
4	3.750	5.580	0.2346	0.481	114000
5	4.900	6.200	0.1932	0.356	141000
6	2.900	5.320	0.2262	0.461	98500
7	5.020	6.620	0.2135	0.321	140000
8	3.030	5.660	0.2350	0.447	103000
9	5.320	7.540	0.2287	0.291	161000
10	7.320	7.730	0.2307	0.346	202000

Trapezoidal Shape, Bottom Width 1.5 ft, Bed Slope 20%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	4.200	5.010	0.1926	0.567	139000
2	6.500	6.660	0.1944	0.357	205000
3	9.850	7.770	0.1954	0.316	286000
4	13.400	8.640	0.1931	0.294	377000
5	17.300	9.480	0.1974	0.276	419000
6	21.600	9.880	0.1964	0.281	493000
7	21.300	10.000	0.2049	0.277	485000
8	27.300	9.310	0.1940	0.362	415000
9	9.510	2.120	0.1954	0.530	27400
10	30.200	4.120	0.1979	0.541	67500
11	4.680	5.040	0.1961	0.415	94000
12	9.400	6.720	0.1994	0.299	158000
13	14.260	7.980	0.2012	0.252	226000
14	19.170	8.850	0.1990	0.225	279000
15	23.650	9.890	0.2062	0.194	337000
16	29.310	10.080	0.1977	0.204	378000
17	4.570	4.080	0.1978	0.728	89800

Trapezoidal Shape, Bottom Width 1.5 ft, Bed Slope 10%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	4.650	4.090	0.0916	0.509	141000
2	7.120	4.970	0.0906	0.380	192000
3	10.000	5.660	0.0907	0.333	257000
4	13.500	6.400	0.0906	0.288	322000
5	17.900	7.070	0.0884	0.246	381000
6	23.000	7.800	0.0874	0.216	449000

Trapezoidal Shape, Bottom Width 1.5 ft, Bed Slope 10%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
7	28.100	8.060	0.0845	0.217	512000
8	26.100	8.510	0.0842	0.187	362000
9	25.900	8.740	0.0872	0.178	363000
10	26.300	8.700	0.0880	0.184	372000
11	26.300	8.820	0.0846	0.171	369000
12	26.100	8.900	0.0857	0.168	349000
13	1.040	0.940	0.1024	7.150	20300
14	2.960	1.940	0.1009	1.970	46200
15	4.940	2.650	0.0988	1.120	73900
16	9.840	3.900	0.0985	0.606	126000
17	15.120	4.980	0.0982	0.411	183000
18	20.820	5.930	0.0966	0.304	234000
19	25.840	6.560	0.0974	0.266	284000
20	30.440	7.070	0.0964	0.237	306000
21	35.460	7.510	0.0980	0.214	333000
22	0.979	1.580	0.1010	1.770	19300
23	2.820	2.960	0.1012	0.649	47100
24	4.710	3.900	0.1000	0.432	74300
25	9.930	5.670	0.0984	0.241	117000
26	14.700	6.560	0.0980	0.204	151000
27	19.800	7.440	0.0999	1.783	180000
28	24.600	8.060	0.0977	0.155	217000
29	29.800	8.740	0.1002	0.145	292000

Trapezoidal Shape, Bottom Width Varies , Bed Slope 3%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	4.090	1.630	0.0324	1.370	68300
2	4.090	1.690	0.0322	1.230	72600
3	6.870	2.340	0.0319	0.729	114000
4	9.760	2.790	0.0318	0.575	145000
5	14.000	3.500	0.0318	0.397	198000
6	18.800	4.100	0.0318	0.314	244000
7	23.300	4.540	0.0323	0.279	289000
8	29.000	5.030	0.0312	0.235	3580900
9	0.093	0.226	0.0319	18.400	1990
10	0.215	0.301	0.0320	16.000	4100
11	0.356	0.363	0.0327	14.700	6090
12	0.561	0.432	0.0323	12.600	9430
13	0.748	0.530	0.0314	8.650	11500
14	1.040	0.687	0.0308	5.320	17000
15	1.760	1.050	0.0312	2.510	28300
16	2.690	1.440	0.0321	1.470	42300
17	4.380	2.030	0.0329	0.845	66500
18	3.890	1.880	0.0337	0.949	62500
19	1.050	0.990	0.0322	1.940	19600
20	2.960	2.080	0.0328	0.560	53100

Trapezoidal Shape, Bottom Width Varies , Bed Slope 3%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
21	4.920	2.820	0.0332	0.362	75300
22	9.860	4.140	0.0352	0.223	128000
23	14.940	4.940	0.0356	0.187	163000
24	20.630	5.630	0.0355	0.161	226000
25	25.940	6.070	0.0354	0.151	272000
26	28.470	6.260	0.0354	0.148	294000
27	35.420	6.720	0.0348	0.137	337000
28	3.990	2.320	0.0359	0.753	93900
29	6.510	2.930	0.0361	0.546	141000
30	9.910	3.570	0.0370	0.426	193000
31	13.700	4.130	0.0365	0.347	253000
32	18.500	4.650	0.0354	0.288	313000
33	24.200	5.080	0.0352	0.264	377000
34	30.300	5.370	0.0352	0.260	439000
35	3.950	2.460	0.0350	0.618	91300
36	1.090	1.720	0.0318	0.693	38000
37	2.930	2.850	0.0335	0.346	82700
38	4.860	3.350	0.0346	0.312	118000
39	9.850	4.460	0.0340	0.217	205000
40	15.200	5.180	0.0341	0.192	279000
41	20.200	5.680	0.0344	0.179	341000
42	24.600	6.040	0.0342	0.169	389000
43	29.800	6.300	0.0348	0.170	379000
44	34.800	6.570	0.0348	0.166	433000
45	0.939	0.660	0.0314	5.180	17600
46	2.980	1.420	0.0310	1.400	458700
47	4.680	1.840	0.0312	0.919	64200
48	9.440	2.640	0.0315	0.539	109000
49	14.390	3.290	0.0312	0.384	153000
50	19.660	3.780	0.0312	0.315	174000

Trapezoidal Shape, Bottom Width 1.5 ft, Bed Slope 1%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	0.980	0.606	0.0098	2.080	13100
2	2.820	1.070	0.0097	0.845	29700
3	4.740	1.400	0.0103	0.604	46500
4	9.920	2.050	0.0108	0.359	74700
5	14.630	2.460	0.0111	0.285	97800
6	19.680	2.820	0.0101	0.211	116000
7	24.670	3.090	0.0102	0.190	140000
8	30.000	3.400	0.0102	0.163	189000

Rectangular Shape, Bottom Width Varies , Bed Slope 3%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	0.099	0.643	0.0297	2.850	6020
2	0.306	1.410	0.0301	0.850	15900
3	0.471	1.910	0.0294	0.512	29500
4	0.694	2.420	0.0290	0.364	44500
5	1.130	3.340	0.0278	0.218	71900
6	1.430	3.580	0.0278	0.222	91300
7	1.660	3.840	0.0279	0.210	106000
8	2.120	4.400	0.0278	0.179	137000
9	2.920	5.170	0.0276	0.150	193000
10	4.880	6.450	0.0267	0.126	328000
11	6.370	7.050	0.0264	0.124	405000
12	7.810	7.680	0.0270	0.121	505000
13	0.900	0.363	0.0298	7.340	2780
14	0.304	0.896	0.0302	1.660	9380
15	0.689	1.620	0.0304	0.644	22400
16	0.440	2.520	0.0300	0.354	46600
17	2.120	3.200	0.0298	0.254	69900
18	2.900	3.740	0.0298	0.217	97600
19	4.860	4.840	0.0303	0.169	167000
20	6.450	5.490	0.0304	0.154	211000
21	7.850	6.000	0.0309	0.145	261000
22	10.600	6.820	0.0310	0.138	345000
23	13.400	7.500	0.0307	0.129	435000
24	0.300	0.568	0.0298	3.610	5250
25	0.686	1.090	0.0296	1.160	12600
26	1.440	1.780	0.0296	0.560	26300
27	2.120	2.270	0.0295	0.398	39600
28	4.860	3.660	0.0294	0.217	93800
29	7.790	4.690	0.0300	0.168	148000
30	13.450	6.120	0.0300	0.131	248000
31	17.100	6.840	0.0294	0.117	313000
32	21.950	7.360	0.0284	0.117	424000
33	24.000	7.480	0.0270	0.115	473000
34	0.300	0.363	0.0292	7.910	3090
35	1.140	1.000	0.0294	1.390	11700
36	2.120	1.590	0.0293	0.669	22800
37	4.840	2.680	0.0294	0.315	55400
38	10.800	4.250	0.0305	0.184	115000
39	19.100	66.150	0.0332	0.118	207000
40	22.000	6.020	0.0316	0.137	251000
41	23.900	6.200	0.0312	0.135	277000

Trapezoidal Shape, Bottom Width 4.0 ft, Bed Slope .2%

Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
1	1.180	0.353	0.0188	2.150	19400
2	1.290	0.372	0.0174	1.780	17100

Trapezoidal Shape, Bottom Width 4.0 ft, Bed Slope .2%

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Data number	Discharge (cfs)	Velocity (fps)	Effective slope	Darcy f	Reynolds number
3	2.750	0.597	0.0202	0.957	39100
4	4.720	0.813	0.0208	0.624	61300
5	4.750	0.814	0.0192	0.565	51600
6	10.100	1.180	0.0213	3.369	114000
7	14.900	1.400	0.0199	0.282	157000
8	15.100	1.400	0.0185	0.269	124000
9	20.200	1.550	0.0178	0.229	192000
10	25.100	1.640	0.0141	0.177	214000
11	24.700	1.620	0.0144	0.179	173000
12	30.400	1.710	0.0127	0.158	251000
13	34.800	1.710	0.0111	0.144	272000
14	35.700	1.700	0.0108	0.149	214000

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