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SOME HYDROLOGIC AND PHYSICAL PROPERTIES  
OF THE MAJOR SOIL TYPES ON THE PAWNEE  
INTENSIVE SITE

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## ABSTRACT

Experimental design of the soil water transect study is discussed in some detail. Soil bulk density, texture, and water retention data for the soil water transects and irrigation plots are presented. Preliminary water content data collected at the end of the 1970 growing season are also shown and discussed. Observations on snow accumulation and associated soil water recharge are introduced and discussed briefly. Results of the textural analysis, pore space distribution, and water desorption characteristics of the microwatershed soils are presented. The calibration results for the neutron probe are also given.

## INTRODUCTION

The primary goal of the Grassland Biome project is to quantitatively describe biological productivity and nutrient and energy cycling in a grassland ecosystem. The abiotic elements of this ecosystem serve as the driving and governing forces behind productivity. With an average annual precipitation of only 10 to 12 inches, water becomes one of the most important of these forces on the Pawnee Site. Soil, acting as a vast reservoir, receives precipitation and stores it, making water available for both plant and animal use. Infiltration, storage capacity, and evapotranspiration are the major factors that decide how much of the annual precipitation becomes available to biological systems in the grassland ecosystem. Soil physical properties play an important role in all three of these processes.

It is the objective of this study to examine the major hydrologic and physical characteristics of the representative soil types found on the Pawnee Site. This study, designed to augment the water balance study of the eight microwatersheds, has been divided into three areas. A series of soil water sampling transects serves as the basic experimental design for investigation of soil water recharge and depletion, soil water potential, and runoff characteristics for the major soils of the Pawnee Site.

## SOIL WATER TRANSECTS

Three sampling transects are located within the interior basin, one each in the heavy-, medium-, and light-grazed pastures. All three are straight-line transects and are oriented perpendicular to the main axis

of the drainage. In addition, another transect was located in a winter-grazed pasture of Section 22 as shown in Fig. 1. Also, a pair of small transects was located in Lynn Lake.

Sampling points were tentatively chosen for each of the transects and were spaced 25 m apart. Soils data were reviewed to determine representativeness of each sampling point. Final sampling points were chosen so that all combinations of grazing treatment, soil type, and "position on slope" were represented.

During the summer of 1970 a neutron probe access tube was installed at each sampling point and the A-, B-, and C-horizons were sampled for bulk density, texture, and water retention analyses.

Measurements of soil water content are taken at the 15, 30, 45, 60, 75, 90, 120, and 150 cm levels on a weekly basis during the growing season and on a monthly basis over the remainder of the year. Data are now available on a weekly basis from July 29, 1970, to September 2, 1970, and on a monthly basis from September 1970 to December 1970. These data are discontinuous due to equipment malfunction and failure.

#### IRRIGATION PLOT SOIL WATER MEASUREMENTS

Neutron probe access tubes have been installed in the ten stress plots on the irrigation site. Three tubes have been located in each of the ten plots and are arranged according to Fig. 2.

Sampling of these tubes will follow the same procedure as used on the microwatersheds and soil water transects. Bulk density, texture, and water retention samples were taken at each tube location using the same procedure as used on the soil water transects.

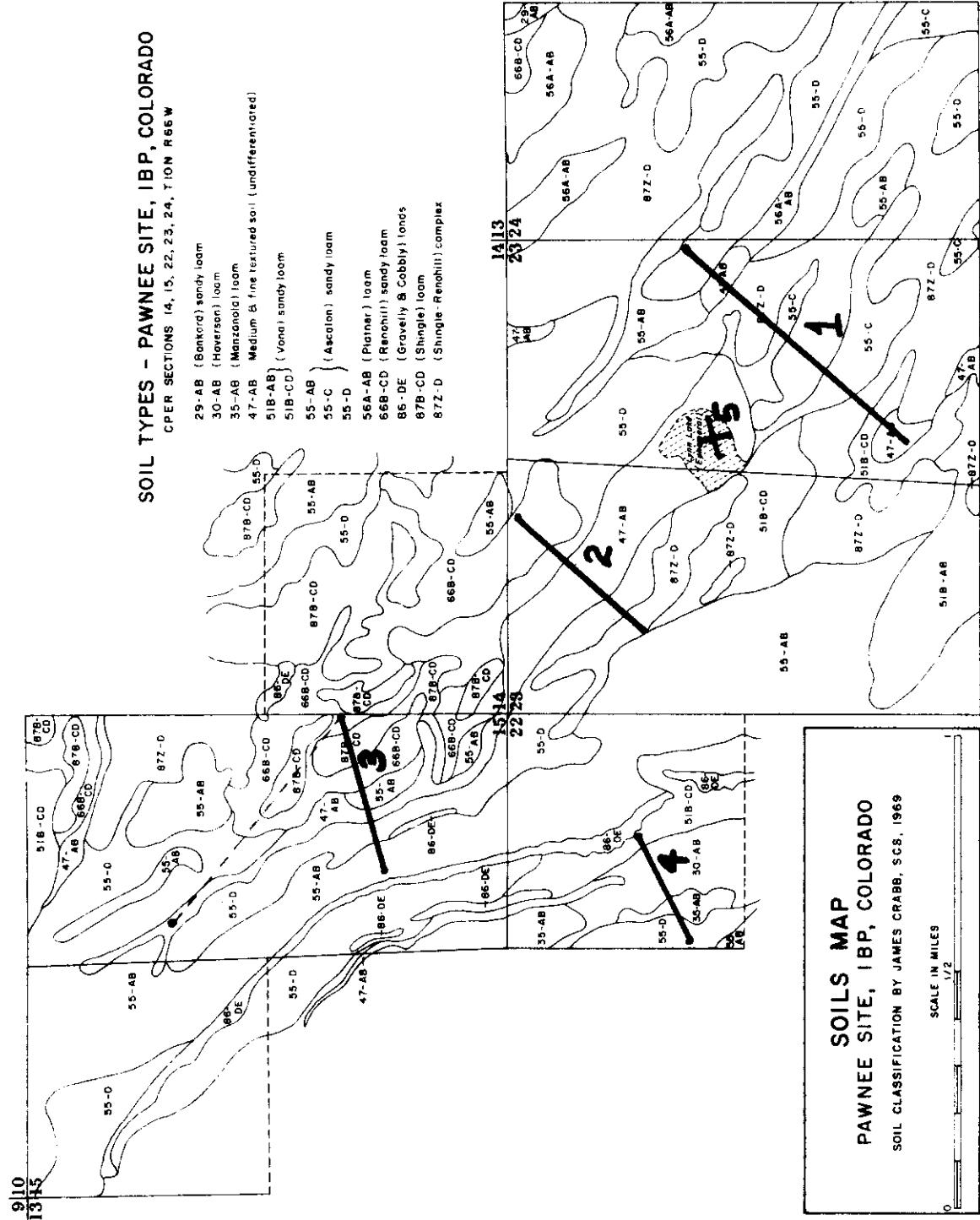


Fig. 1. Soils map showing the soil water transects.

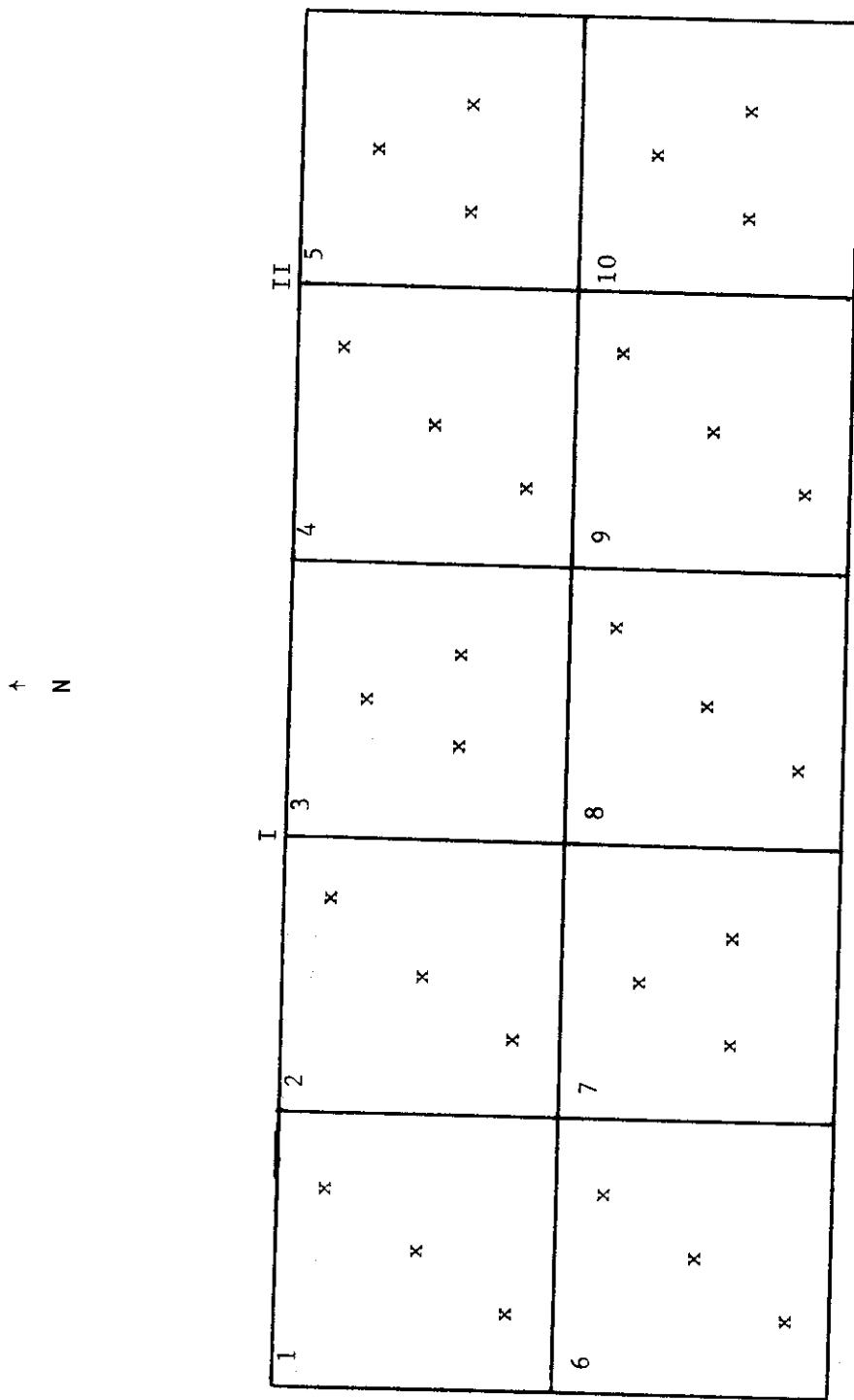


Fig. 2. Location of neutron probe access tubes in the irrigation plots.

Soil water content data for these plots are available on a limited basis for the month of August 1970.

#### SOIL WATER POTENTIAL MEASUREMENTS

In the past two decades a large amount of effort has been expended in investigating the free energy status of water in the soil-plant-atmosphere continuum. Recent developments in psychrometric instrumentation have enabled workers in this area of interest to accurately describe the free energy status of water in plant tissues and soils. A thermocouple psychrometer is now commercially available for use in soils. A later technical report will deal with the concept, theory, terminology, and instrumentation involved in measuring water potentials in porous media and biological systems.

Forty soil psychrometers are being included in the design of the soil water transects. Soil water potential measurements will be made at selected sampling points along the medium-grazed transect. Eight plastic access pipes, 15 cm in diameter and 75 cm in length, have been installed along the transect close to neutron probe access tubes. Two pipes are located in each of the four major soil types: Ascalon sandy loam, Renohill sandy loam, Shingle loam, and undifferentiated soil. Each access pipe will initially include a stack of five soil psychrometers at the 2, 10, 20, 40, and 60 cm levels. The sensors will protrude 10 to 15 cm horizontally from the outside wall of the access pipe. Extention of this study to the microwatersheds and irrigation plots is presently under consideration.

Soil water potential measurements will be taken weekly starting in April or May 1971. It is hoped that these data could be used to construct characteristic curves based on *in situ* measurement of both soil water content and soil water potential.

A laboratory calibration procedure is being developed whereby each soil psychrometer will be calibrated in aqueous sodium chloride solutions of known water potentials at various temperatures.

Also, this project, in cooperation with the photosynthesis group, is working on an improved design for soil psychrometers and is setting up a workshop for the construction and calibration of thermocouple psychrometers of all designs for intra-biome use.

#### RUNOFF STUDY -- DESIGN AND INSTRUMENTATION

At each access tube site on the light-, medium-, and heavy-grazed transects, two micro-runoff collectors have been installed to collect overland flow. The collectors consist of an ice cream container, 15 cm in diameter and 16 cm deep, fitted with a galvanized metal collecting funnel 7.5 cm wide. The micro-runoff collector is installed in the ground so that the collecting funnel is flush with the soil surface. These units will be in operation beginning in the spring of 1971.

Immediately following a rain event, each collector will be emptied, the total runoff volume measured, and the sample saved for analyses of sediment concentration and particle size distribution.

#### SOILS DESCRIPTION

Site descriptions, experimental design, instrumentation, and overall objectives for the hydrology study have been covered in previous

Table 1. Soil horizon depths in centimeters for the major soil types on the Pawnee Site.

Soil Type	Horizon	
	A	B
47	0-15	15-53
51	0-10	10-36
55	0-15	15-46
66	0-15	15-41
87B	0-13	13-46
87Z	0-15	15-53

technical reports (Smith and Striffler 1969; Galbraith 1969). In addition, a preliminary study on the mineralogy of the Pawnee Site has provided data on textural properties and background knowledge of the soils on the site (Franklin 1969).

Table 1 gives depths corresponding to each horizon for the major soil types. The C horizon was excluded because of the extreme variability found in the C horizons of all the soil types. Basically, the C horizon was composed of  $C_1$  and  $C_2$  sub-horizons having approximately the same texture. However, quite often a  $C_{ca}$  layer was sandwiched between the  $C_1$  and  $C_2$  or else was replacing one or the other. When the  $C_{ca}$  was present with either the  $C_1$  or  $C_2$ , it was found either above or below the  $C_1$  or  $C_2$ . Textural analysis of these  $C_{ca}$  layers showed them to be essentially clay sized material. Many of the  $C_{ca}$  samples appeared to hold no sand-sized material whatsoever. There seemed to be no pattern evident in the C-horizon makeup from one soil type to another. In general, the C horizon was made up of at least 60% sand-sized material. Usually, this sand was medium-grained and had the appearance of clean, white beach sand.

Tables 2a, 2b, and 2c give bulk density data for the soil water transects and the irrigation plot. In Table 2a the bulk densities are averaged within soil types for the A and B horizons. It was virtually impossible to obtain a bulk density sample for the C horizon anywhere on the site. The mean values for the A horizon all fell between 1.38 and  $1.45 \text{ g/cm}^{-3}$ , and the mean values for the B horizon fell between 1.40 and  $1.50 \text{ g/cm}^{-3}$ . It is doubtful that the differences between soil type means were significant because of the small sample sizes.

Table 2a. Bulk densities by soil type (g/cm<sup>-3</sup>).

Soil	Horizon	Mean B. D.	Maximum	Minimum	SD	n.
47	A	1.41	1.74	1.25	.173	10
	B	1.47	1.61	1.33	.109	7
51	A	1.40	1.44	1.33	.059	3
	B	1.50	1.52	1.48	.021	3
55	A	1.42	1.58	1.22	.101	20
	B	1.49	1.64	1.31	.089	17
66	A	1.39	1.51	1.25	.081	10
	B	1.45	1.65	1.05	.162	10
87B	A	1.38	1.41	1.31	.047	4
	B	1.40	1.46	1.32	.062	4
87Z	A	1.45	1.62	1.30	.106	11
	B	1.49	1.80	1.30	.126	12

Table 2b. Bulk densities by grazing treatment ( $\text{g/cm}^{-3}$ ).

Treatment	Horizon	Mean B. D.	Maximum	Minimum	SD	n.
Heavy	A	1.45	1.74	1.22	.132	21
	B	1.52	1.80	1.40	.097	17
Medium	A	1.42	1.55	1.25	.086	19
	B	1.46	1.65	1.32	.092	19
Light	A	1.37	1.58	1.25	.086	18
	B	1.43	1.61	1.05	.134	17

Table 2c. Bulk densities for irrigation plots ( $\text{g/cm}^{-3}$ ).

Horizon	Mean B. D.	Maximum	Minimum	SD	n.
A	1.35	1.50	1.20	.081	28
B	1.40	1.56	1.23	.077	15

Table 2b gives mean values by horizon of bulk density averaged within grazing treatments. For the A horizon, the difference between the mean bulk density of the heavy-grazed soil and that of the light-grazed soil was significant at the .99 level. However, in this case the heavy-grazed mean was only  $0.08 \text{ g/cm}^{-3}$  larger than the light-grazed mean. This difference does not seem large enough to substantiate the popular claim that grazing animals compact the soil surface, which gives rise to higher bulk densities. These bulk density data are now being run through an analysis of variance procedure to test soil type effect, grazing effect, and the interaction.

Particle size determinations for the sand, silt, and clay classes were made using the hydrometer method (Day 1965). Textural data for the major soils are shown in Table 3. From the textural triangle the following classifications are assigned:

47	A	sandy clay loam
	B	sandy clay loam
51	A	sandy loam
	B	sandy clay loam
55	A	sandy loam
	B	sandy clay loam
66	A	sandy clay loam
	B	sandy clay loam
87B	A	sandy loam
	B	sandy clay loam
87Z	A	sandy clay loam
	B	sandy clay loam
Irrig. Plot A		
	B	
	C	

Table 3. Percent sand, silt, and clay by soil type. All values except n are in percent.

Soil Horiz.	Mean % Sand	Mean % Silt						Mean % Clay						n.
		Max.	Min.	SD	Max.	Min.	SD	Max.	Min.	SD	Max.	Min.	SD	
47 A	62	77	38	12.3	17	25	5	6.7	21	41	15	8.3	11	
47 B	65	77	49	10.4	12	20	5	5.3	23	26	15	8.7	6	
51 A	75	85	57	11.6	9	15	5	4.2	16	38	5	13.5	5	
51 B	69	72	67	13.6	10	18	8	5.0	21	25	12	6.1	4	
55 A	69	82	47	10.3	14	26	2	6.0	17	35	5	6.8	22	
55 B	60	82	37	13.8	15	31	0	8.7	25	41	15	8.7	20	
66 A	64	77	47	10.8	14	20	8	3.8	22	38	8	11.9	10	
66 B	51	74	8	20.2	17	26	8	6.1	32	66	18	15.3	10	
87B A	65	80	52	11.6	16	28	8	9.1	19	28	12	7.0	4	
87B B	56	62	44	8.1	22	31	18	6.2	22	25	20	2.2	4	
87Z A	62	77	37	12.0	15	23	10	3.8	23	45	8	11.2	11	
87Z B	52	80	20	17.2	18	41	2	11.2	30	43	18	8.0	12	
Irrig. Plots (55)	A	61	75	37	9.6	23	53	13	7.2	16	38	10	7.6	29
	B	52	75	21	13.2	21	30	13	4.7	27	54	10	9.9	25
	C	62	75	36	8.6	17	36	7	6.3	20	35	10	5.3	28

#### WATER RETENTION CHARACTERISTICS OF THE MAJOR SOIL TYPES

The pressure plate method (Richards 1965) of determining absorption characteristics was employed in this analysis. Samples were subjected to applied pressures of 0.0, 0.1, 0.3, 1.0, 3.0, and 15.0 bars and the volumetric water content determined for each of the applied pressures. These data are tabulated in Table 4 and the associated desorption curves are shown in Fig. 3-9. With the exception of the undifferentiated soil (47), all the soil types showed the B horizon to have a higher water holding capacity than the A horizon. These results agree with the textural data for each soil type. The undifferentiated soil showed a higher water holding capacity for the A horizon and a corresponding heavier texture. For the other soil types, the texture of the A horizon was coarser than that for the B horizon and, consequently, had a lower water holding capacity than the B horizon. In considering the water holding capacity of the A and B horizons combined, it appears that the undifferentiated soil had the highest and Vona sandy loam the lowest water holding capacity. The desorption curves for the A and B horizons on the irrigation plots essentially matched those for the Ascalon sandy loam on the transects and, likewise, for the microwatersheds. The irrigation plot soils (mapped as Ascalon sandy loam) appeared to have a slightly higher water holding capacity. This may be due to aeolian deposits of silt-sized material, since there was a noticeably greater percentage of silt in the irrigation plot textured samples than in the other Ascalon samples.

Table 4. Desorption data for the soil water transects and irrigation plot soils.

Soil Type	Horizon	Matric Potential, -bars					
		0.0	0.1	0.3 (% H <sub>2</sub> O by volume)	1.0	3.0	15.0
47	A	66.0	48.0	28.8	25.2	15.5	13.9
	B	61.3	38.6	22.1	18.7	11.7	9.9
51	A	60.2	33.3	17.3	14.0	8.7	7.4
	B	61.4	33.8	19.5	16.4	10.7	9.4
55	A	65.2	45.9	20.0	16.2	10.0	9.0
	B	65.3	43.2	24.4	20.4	12.9	11.7
66	A	57.2	34.2	17.1	15.0	10.3	9.9
	B	67.6	46.4	27.6	24.5	15.3	13.4
87B	A	59.8	35.9	22.2	20.0	11.9	11.1
	B	62.9	45.0	25.5	22.1	11.9	11.0
87Z	A	57.2	37.4	23.9	21.3	13.7	12.9
	B	63.7	41.9	26.2	23.8	15.0	13.7
Irrig. Plots	A	66.7	36.4	19.0	15.6	8.9	7.4
	B	65.9	42.0	26.3	22.5	13.2	10.9

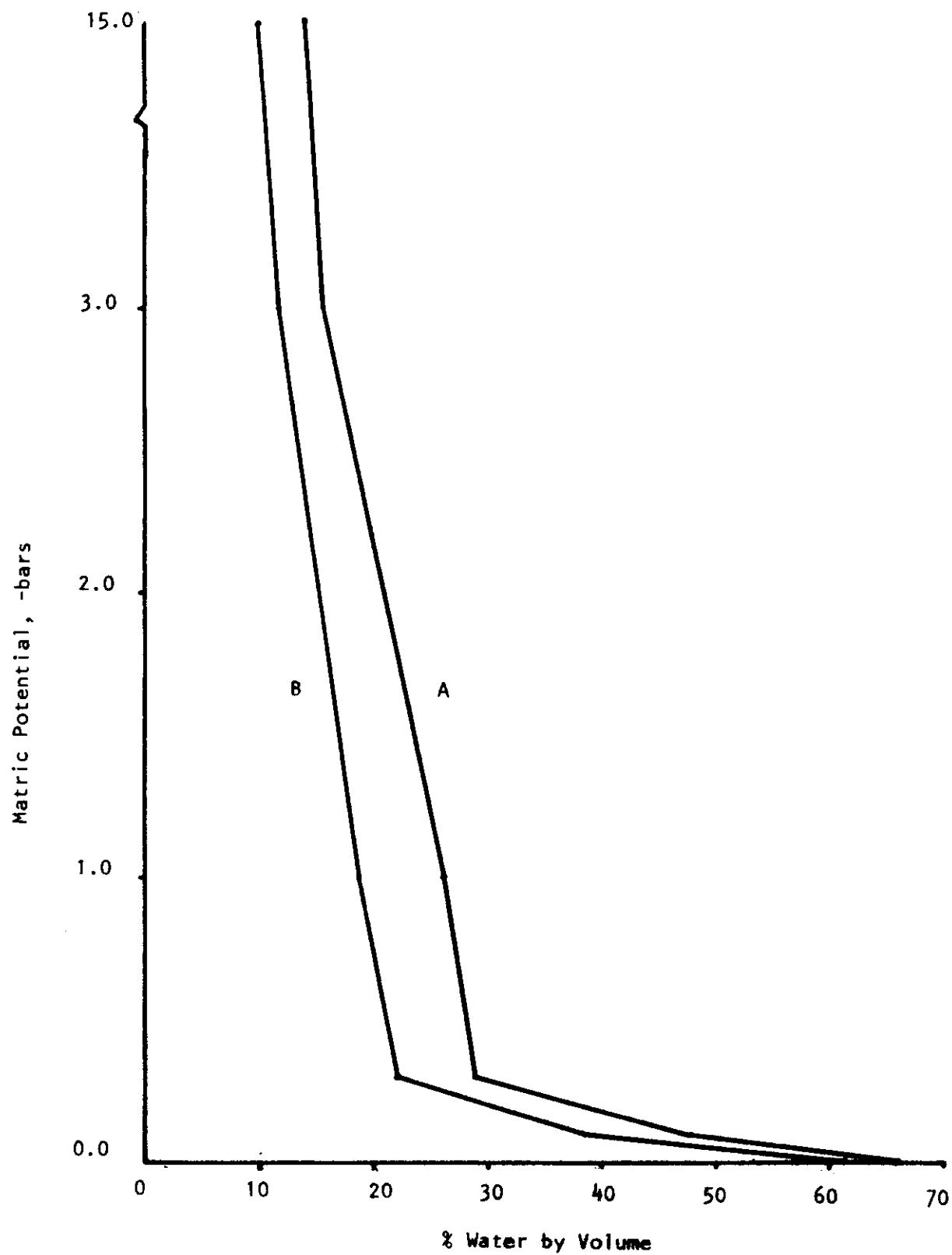


Fig. 3. Desorption curves for undifferentiated soil (47).

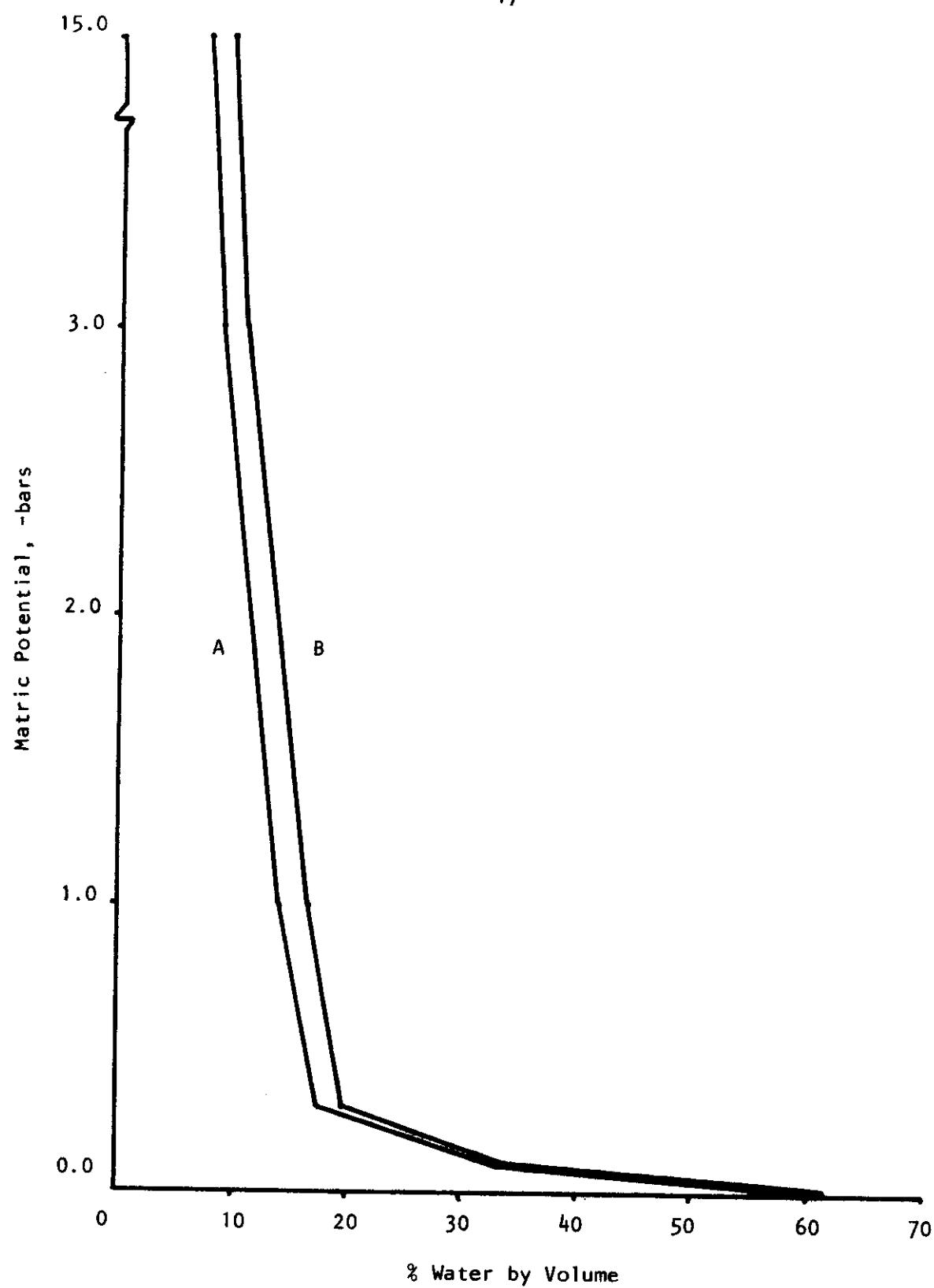


Fig. 4. Desorption curves for Vona sandy loam (51).

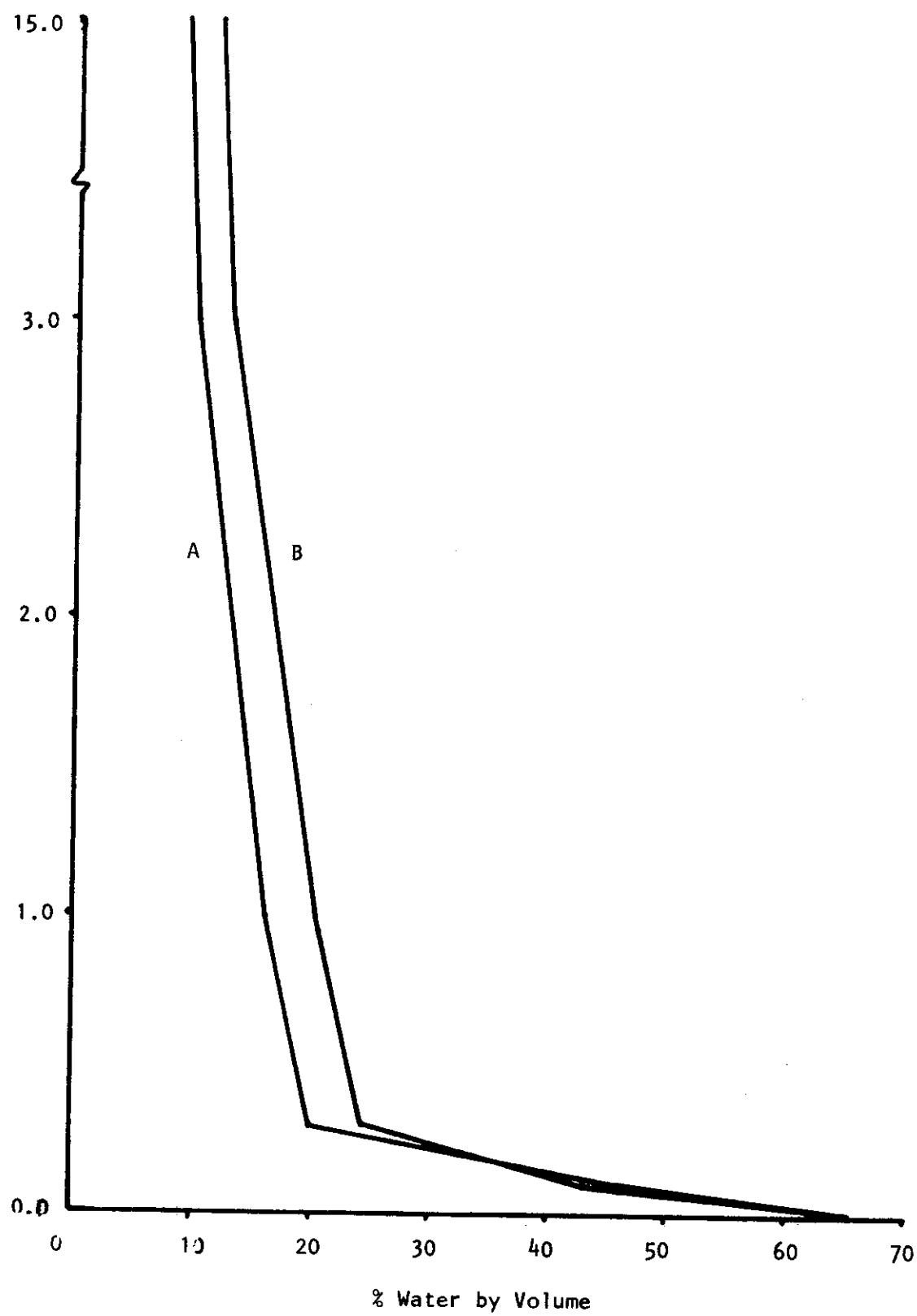


Fig. 5. Desorption curves for Ascalon sandy loam (55).

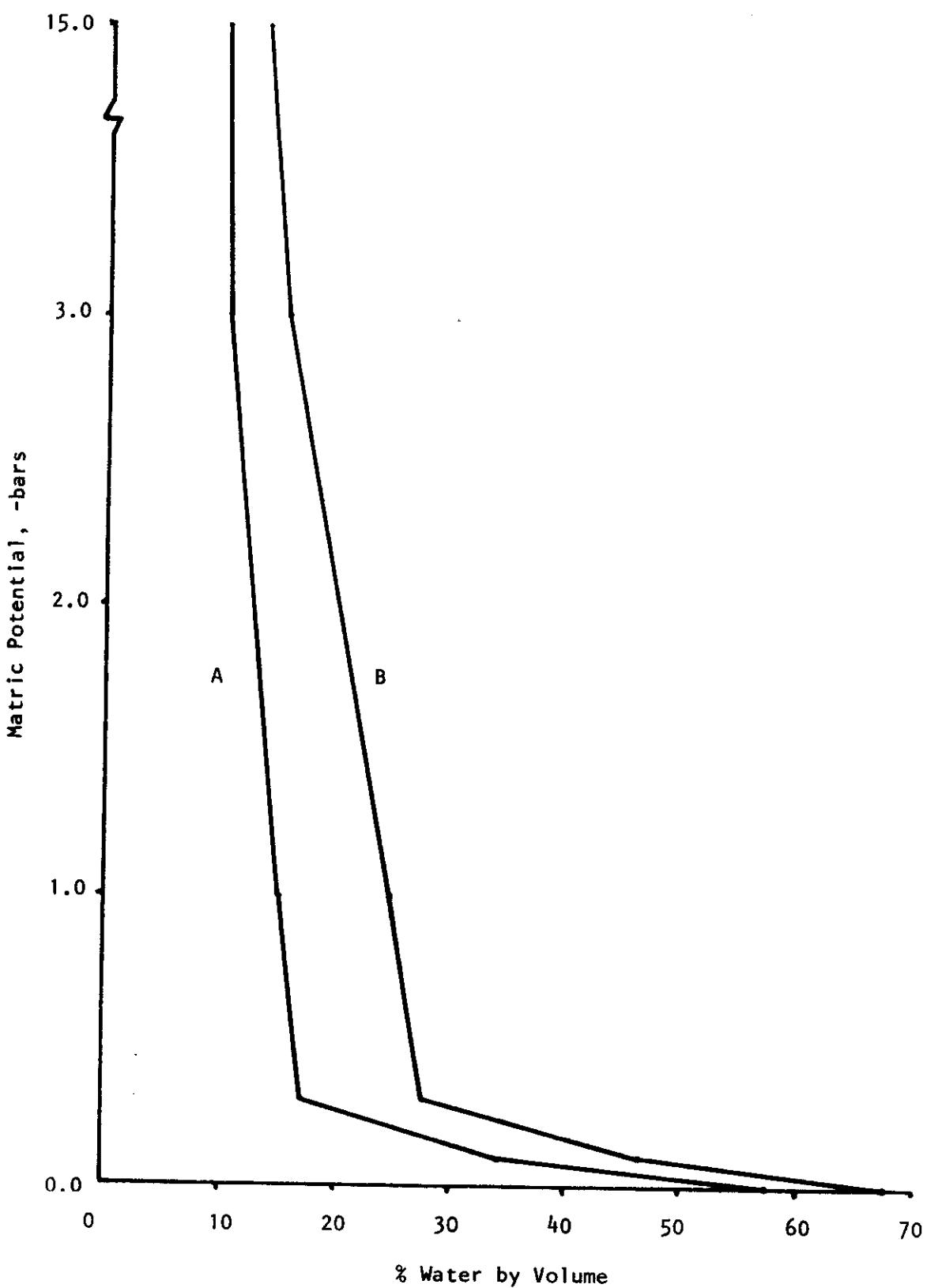


Fig. 6. Desorption curves for Renohill sandy loam (66).

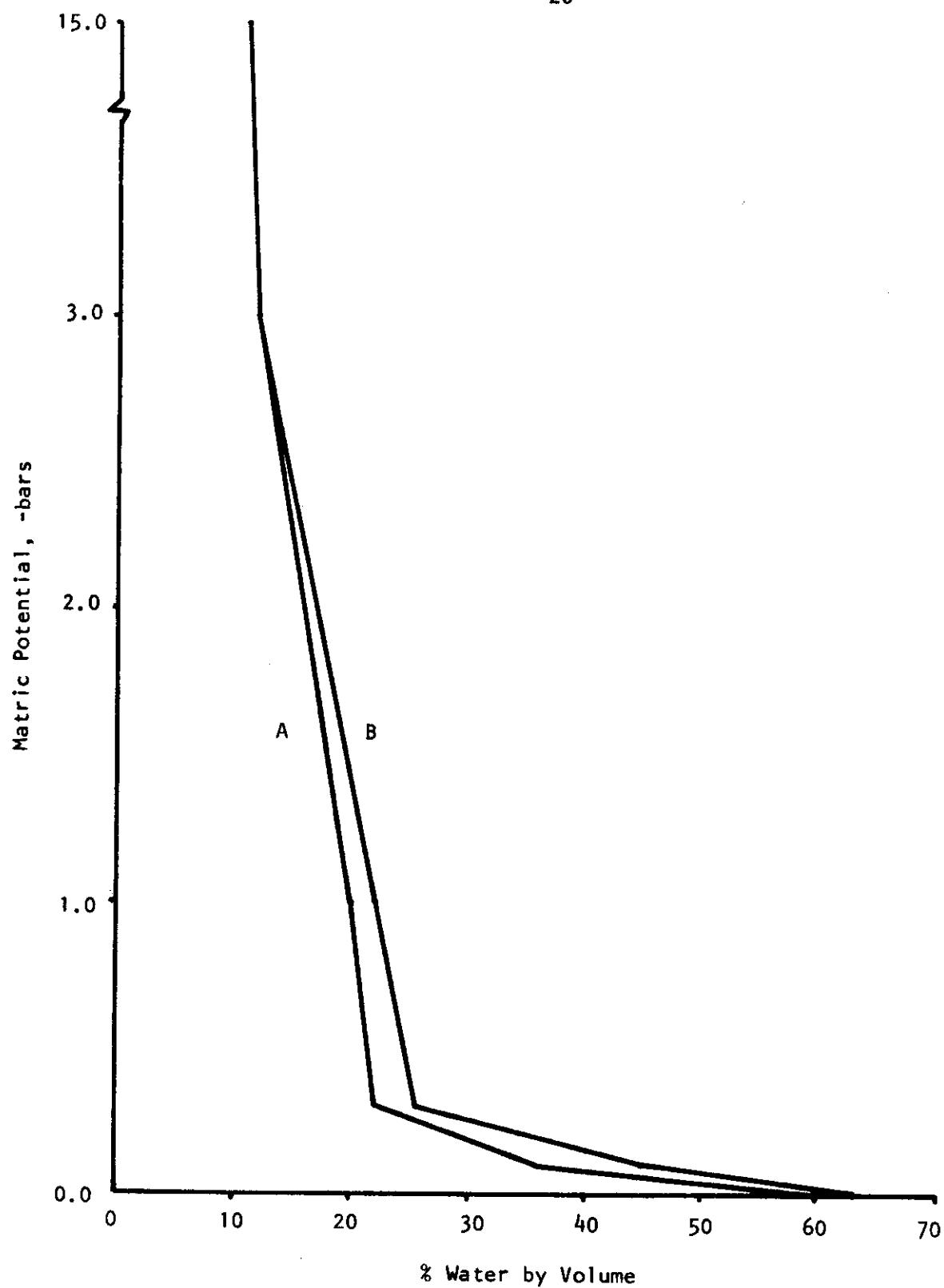


Fig. 7. Desorption curves for Shingle loam (87B).

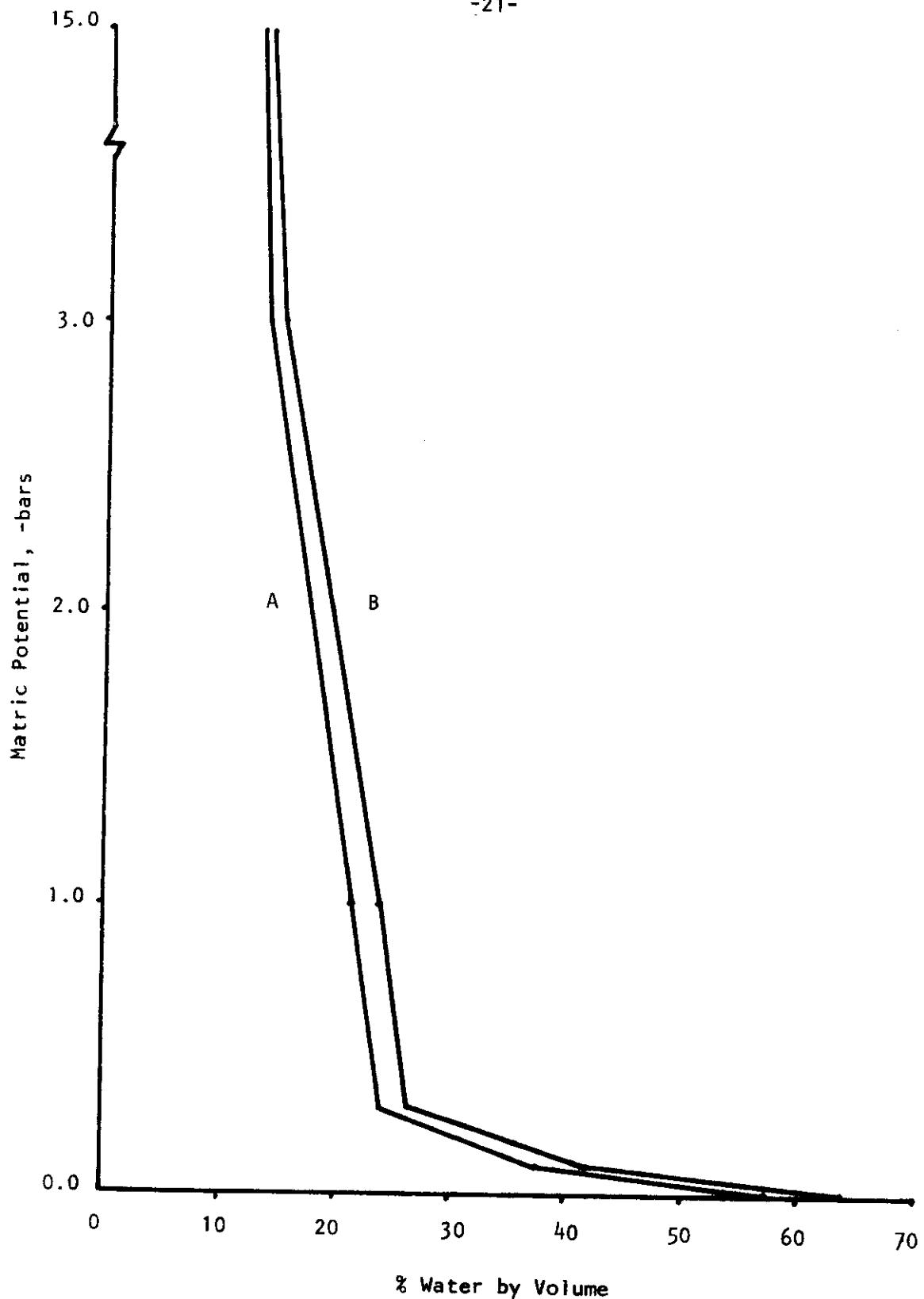


Fig. 8. Desorption curves for Shingle-Renohill Complex (87Z).

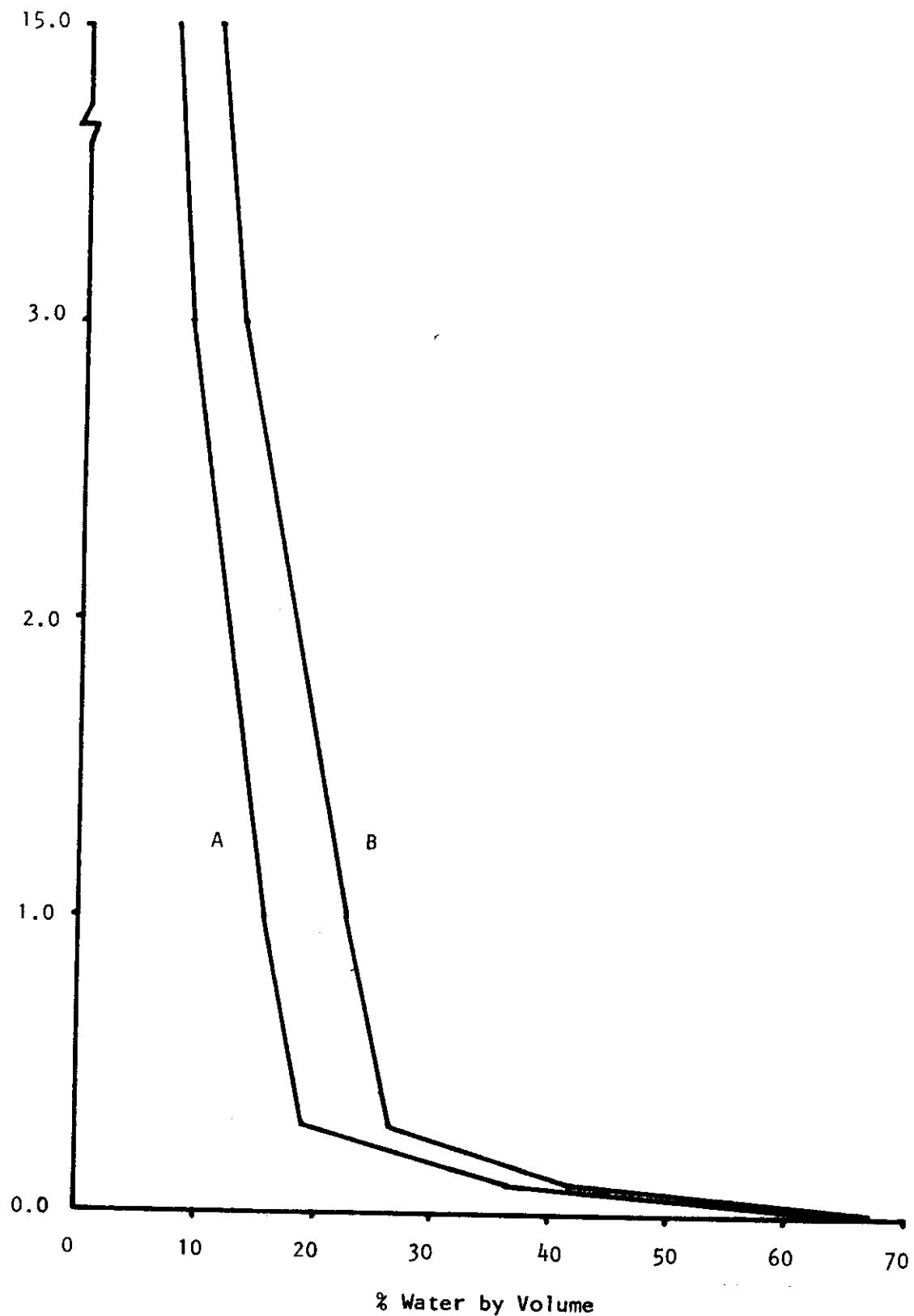


Fig. 9. Desorption curves for the irrigation plot soils (Ascalon sandy loam (55)).

## SOIL WATER CONTENT MEASUREMENT -- PRELIMINARY DATA

Fig. 10, 11, and 12 are topographical profiles of the soil water transects on each of the three grazing treatments. It is apparent from comparing the three graphs that the heavily-grazed pasture held more water than the light-grazed pasture and possibly more than the medium-grazed pasture. Differences in the water holding capacity of the various soils is noticeable in some cases. The Shingle-Renohill complex (87Z) seems to have a high water holding capacity. Differences between depletion rates were not readily apparent. It did seem, however, that the sampling sites at the tops of ridges were losing water more rapidly than other lowland sites. The reader must remember that at the time these measurements were begun, very little plant growth was occurring and the soil had dried considerably the previous two months. The low point of the soil water content curve was reached in the beginning of September. By the end of September, considerable recharge had occurred. A series of measurements taken recently showed that the soil is now recharged to about the level that existed in late July and early August.

## SOME OBSERVATIONS ON SNOW ACCUMULATION

Snow redistribution by wind may have an important effect on the recharge characteristics of prairie soils. Preliminary data and field observations for a storm in January 1971 suggest that mesotopography combined with high winds drastically influences snow redistribution on the ground. Initially, snowfall was evenly distributed on the ground.

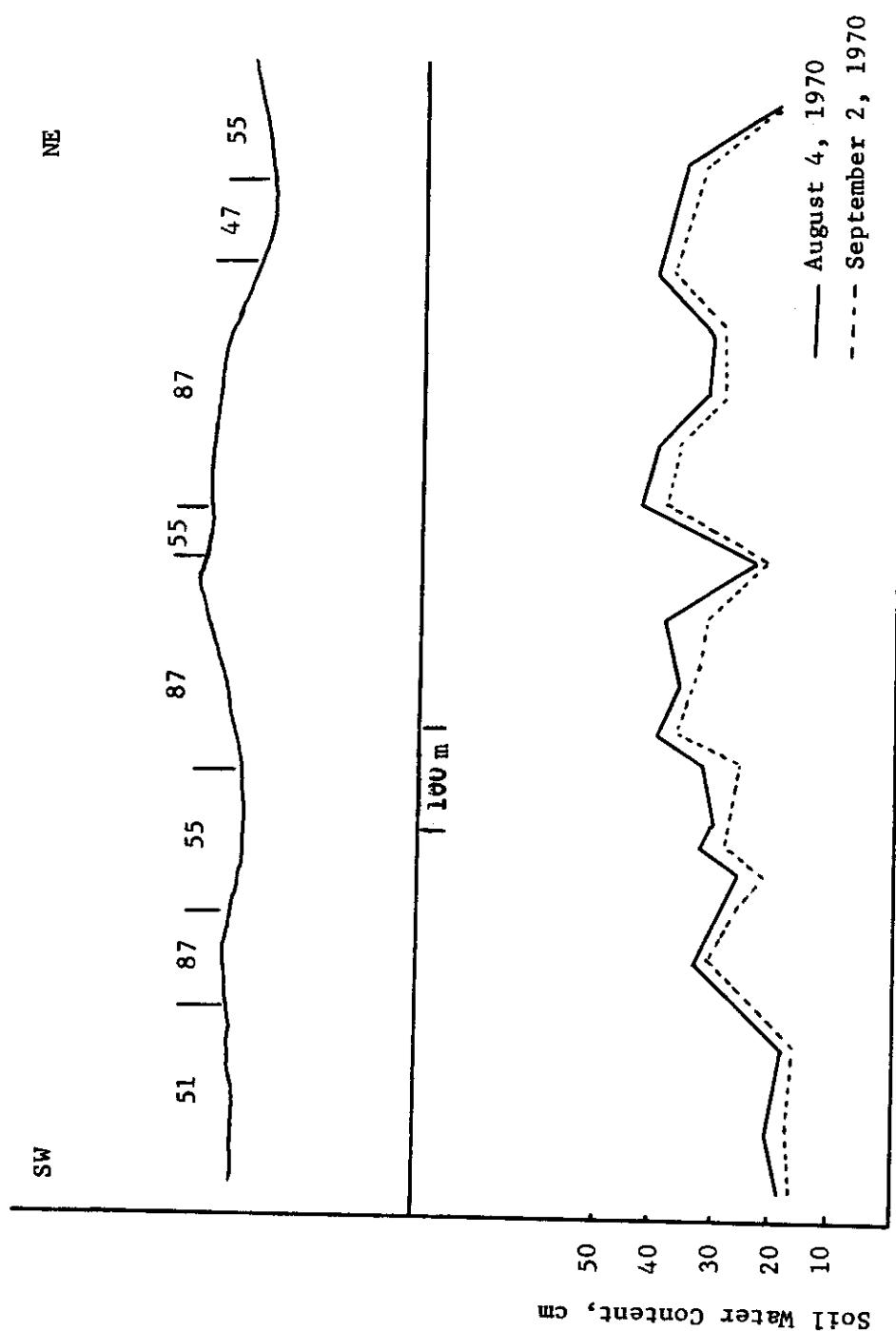


Fig. 10. Topographic profile and soil water contents for the heavy grazed soil transect.

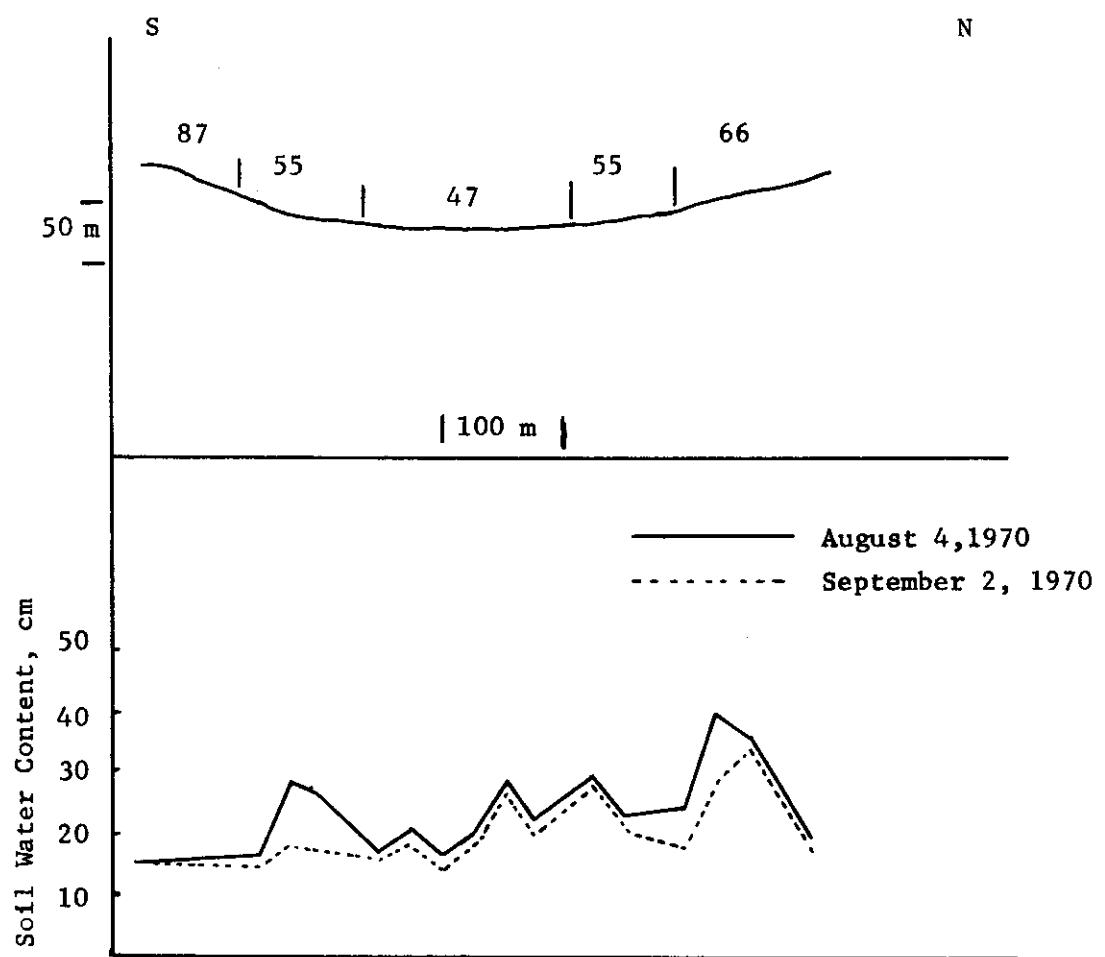


Fig. 11. Topographic profile and soil water content for the light grazed soil transect.

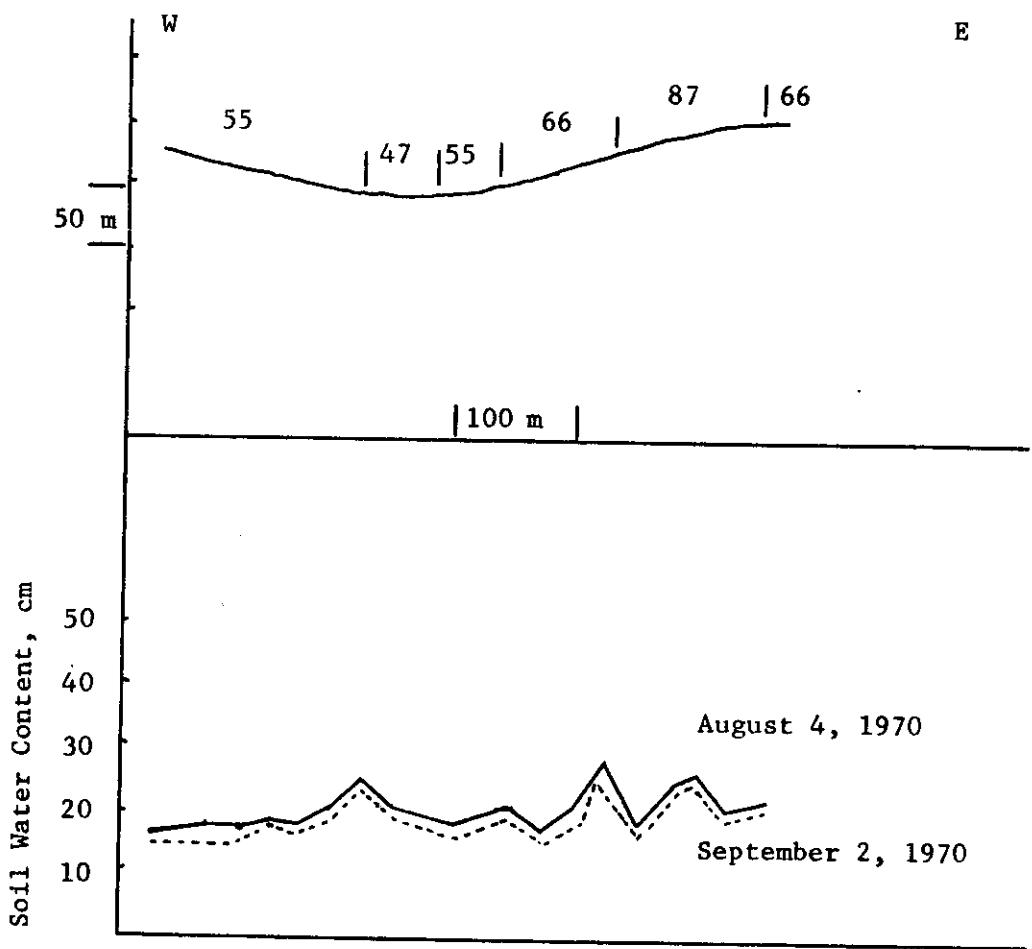


Fig. 12. Topographic profile and soil water content for the medium grazed soil transect.

Within 24 hours strong winds arose and began redistributing the snow. Redistribution appears to be a continuous process as long as the wind is blowing, but is finally abated when snow density and/or metamorphism reach some critical point. For the storm of January 1-3, gravimetric snow water equivalent samples taken at selected points along transects 1, 2, 3, and 5 showed depth water equivalents ranging from 1.0 to 1.4 cm with a mean of 1.2 cm before redistribution occurred. A series of samples taken 24 hours later revealed depth water equivalents ranging from 0.7 to 2.7 cm with a mean of 1.7 cm. The difference between the means is nonsignificant because of the small sample sizes. However, the data point out the extreme variability in water equivalents due to the redistribution of snow on the ground. Precipitation catch for this storm averaged 1.3 cm.

At two sampling points, an attempt was made to investigate the effect of vegetation on snow accumulation. In both cases, one sample was taken from a bare soil situation (i.e., devoid of vegetation) and a second was taken 6 m to the east of the first in a vegetated situation (vegetation at least 5 cm in height). Depth water equivalents for the two sampling points averaged 0.7 cm in the open and 2.1 cm in the vegetation. From these preliminary data it was hypothesized that grazing treatment might have an effect on snow accumulation because of wind transport and redeposition being influenced by vegetation height and density.

Two weeks later, on January 16, the snowpack was essentially stabilized and primed for melt. No visible melt had occurred in the previous 14 days. Samples of snow water equivalent were taken again at preselected snow sampling sites along transects 1, 2, 3, and 5. For the

heavy-, medium-, and light-grazed transects, average depth water equivalents were 1.0, 1.5, and 1.4 cm, respectively. Although there was considerable scatter in the data, at best these results indicate a possibility of grazing effect on the snow trapping efficiency of an area.

An attempt was made to classify sampling sites topographically. Classifications assigned were designated upland (or ridgeline), swale, or north, south, east, or west slope. Neutron probe access tube sites were classified in the same manner. Observations of volumetric soil water content were made with the neutron probe technique at the end of December and again at the end of January. Precipitation during this period included only that contributed by the storm discussed above. The average changes in volumetric soil water content,  $\bar{\Delta}\theta$ , for the designated topographic sites were computed and compared to the average depth water equivalents,  $D_w$ , obtained for the same sites. The results are shown in Table 5. Interestingly enough, the data suggest that this aspect might be influencing melt characteristics. Additions of water to the profile occurred on the traditionally warmer south and west slopes and in the bottoms or swales where ponded water may have had a better opportunity to infiltrate. The author had the opportunity to observe the melt characteristics on all the sites at the height of the melt period. Ponded water was evident in small depressions on all the sites. It was assumed this was due to frozen soil (soil on the Pawnee Site was frozen to depths as great as 20 cm in early January). The data of

Table 5. Comparison of average changes in volumetric soil water content ( $\Delta\theta$ ) to average depth water equivalents ( $D_w$ ).

Site	$\overline{\Delta\theta}$ a/	$D_w$ b/
----- cm -----		
Upland	-0.2	0.7
Swale	0.6	1.3
North	-0.3	1.9
South	1.5	1.7
East	-0.1	1.8
West	0.9	2.9

a/  $\overline{\Delta\theta}$  refers to the difference in volumetric soil water content between the end of December and the end of January.

b/  $D_w$  is the depth water equivalent measured gravimetrically.

Table 5 suggest very little penetration of melt water from the snowpack. The apparent differences between the depth water equivalents and the changes in soil water storage are assumed here to have been vapor losses to the atmosphere.

The drastic effect of mesotopography on snow redistribution was demonstrated quite clearly by the average depth water equivalent computed from samples intentionally taken from lee slopes. This average depth water equivalent was 4.3 cm as compared to the average of 1.3 cm for the entire site.

The preliminary data discussed above suggest that an investigation of snow accumulation should be considered in order to better understand the water balance on the Pawnee Site.

#### CHARACTERISTICS OF MICROWATERSHED SOILS

During 1970 all soil sample analyses for the microwatersheds were completed. Analyses included bulk densities (Galbraith 1969), texture analyses, and water retention characteristics. In addition, the calibration for the neutron probe was completed.

##### Soils Analysis

An analysis of the particulate size distribution for the sand, silt, and clay fractions was made for each of the microwatersheds. Textural classification (Table 6) shows that all soils are within the sandy loam to clay loam range. In general, the B horizon has the highest clay content and the C horizon the highest silt content. Variability in all three texture classes was the greatest in the parent material of the C horizon.

Table 6. Textural analysis of microwatershed soils.

Microwatershed	Horizon	Sand			Silt			Clay			Classification
		%	SD	Range	%	SD	Range	%	SD	Range	
1	A	69	2.2	6.0	12	2.4	7.0	19	1.6	5.0	Sandy Loam
	B	70	2.6	8.0	9	1.0	2.0	21	2.3	7.0	Sandy Clay Loam
	C	72	7.5	23.0	9	4.5	13.0	19	3.4	10.0	Sandy Loam
2	A	62	3.4	10.0	18	2.7	8.0	20	2.5	7.0	Sandy Loam
	B	49	7.5	28.0	21	4.7	14.0	30	3.9	14.0	Sandy Clay Loam
	C	51	8.3	25.0	20	5.4	18.0	29	3.8	10.0	Sandy Clay Loam
3	A	65	3.2	10.0	16	1.5	5.0	19	2.7	8.0	Sandy Loam
	B	42	10.1	32.0	21	5.7	18.0	37	5.2	16.0	Clay Loam
	C	45	6.7	21.0	24	4.6	15.0	31	4.4	15.0	Clay Loam
4	A	73	4.0	13.0	11	3.1	9.0	16	2.3	8.0	Sandy Loam
	B	68	3.5	11.0	9	3.3	8.0	23	2.9	8.0	Sandy Clay Loam
	C	62	9.3	28.0	16	5.3	18.0	22	4.5	11.0	Sandy Clay Loam
5	A	72	4.3	13.0	11	2.0	5.0	17	3.6	10.0	Sandy Loam
	B	71	3.1	8.0	9	2.4	8.0	20	2.5	8.0	Sandy Loam
	C	68	5.6	15.0	11	4.0	13.0	21	2.6	7.0	Sandy Clay Loam
6	A	58	5.0	17.0	18	4.4	15.0	24	2.3	9.0	Sandy Clay Loam
	B	57	5.7	19.0	16	2.2	6.0	27	4.9	16.0	Sandy Clay Loam
	C	52	12.4	30.0	19	8.8	21.0	29	5.1	14.0	Sandy Clay Loam
7	A	70	1.9	5.0	14	1.8	6.0	16	1.8	5.0	Sandy Loam
	B	60	2.4	8.0	15	1.4	3.0	25	1.9	5.0	Sandy Clay Loam
	C	57	15.8	47.0	19	8.1	24.0	24	8.2	25.0	Sandy Clay Loam
8	A	65	5.1	19.0	14	3.5	13.0	21	2.5	6.0	Sandy Clay Loam
	B	47	9.4	23.0	19	4.6	11.0	34	4.9	12.0	Sandy Clay Loam
	C	58	7.8	27.0	16	5.1	17.0	26	3.8	13.0	Sandy Clay Loam

#### Soil Water Desorption Analysis

Water retention characteristics were examined for each of the microwatershed soils. The amount of water retained by the soil from saturation to 15 bars applied pressure was determined by the standard pressure plate method. The results are given in Table 7. Microwatersheds 2, 3, and 8 appear to have higher water holding capacities for the B and C horizons than the other watersheds. The B and C horizons have higher water holding capacities than the A horizons as shown in Fig. 13. Micropore space, defined as water retained at 0.1 bars, exceeds macropore space generally by a factor of 2 for all horizons.

#### Neutron Probe Calibration

Calibration of the neutron probe was accomplished by a field gravitational method. The results of separate calibrations for the surface, 20 cm, and depths greater than 20 cm are shown in Fig. 14.

Table 7. Water retention characteristics of Ascalon soil.

Microwatershed	Horizon	Applied Pressure - Bars						Pore Space %		
		0.0 (Percent H <sub>2</sub> O by Volume)			-3.0 -3.4 <sup>23</sup> -3.6 <sup>22</sup>			-5.0 -5.3 <sup>23</sup> -5.6 <sup>22</sup>		
		A	B	C	A	B	C	Macro	Micro	%
1	A	43.5	32.5	17.0	12.3	8.6	6.9	11.0	32.5	A
	B	51.2	31.0	17.6	14.0	11.1	8.7	20.1	31.0	
	C	47.4	31.7	16.3	12.8	9.9	7.6	15.6	31.7	
2	A	50.4	33.8	15.0	11.6	7.9	5.5	16.5	33.8	F
	B	52.4	38.7	26.3	21.7	17.1	13.1	13.8	38.7	
	C	53.2	40.3	27.4	22.2	17.5	12.8	12.9	40.3	
3	A	42.4	27.1	16.5	14.1	9.6	6.2	15.1	27.1	
	B	52.7	34.1	27.7	22.7	18.3	13.6	18.5	34.1	
	C	54.2	35.9	27.6	21.0	16.4	12.1	18.3	35.9	
4	A	44.9	27.2	13.5	9.9	7.2	4.5	17.7	27.2	
	B	48.6	29.6	20.6	16.0	13.3	8.5	19.0	29.6	
	C	47.6	27.7	17.7	13.4	11.5	7.0	19.8	27.7	
5	A	45.6	24.9	13.6	--	7.5	6.4	19.9	24.9	
	B	48.9	27.9	19.1	15.5	11.3	10.5	21.1	27.9	
	C	50.6	30.5	19.3	15.0	11.3	10.5	20.0	30.5	
6	A	50.2	29.2	15.9	12.6	9.7	7.1	20.9	29.2	
	B	51.3	33.2	22.8	18.5	14.1	11.0	18.0	33.2	
	C	51.1	31.5	20.5	16.8	12.2	10.6	19.6	31.5	
7	A	48.3	22.3	16.4	11.2	8.8	6.0	25.8	22.3	
	B	52.3	31.7	22.6	16.1	11.6	9.2	20.5	31.7	
	C	55.7	37.7	30.7	22.3	17.3	11.2	17.9	37.7	
8	A	55.0	24.6	16.0	11.1	7.8	5.8	30.2	24.6	
	B	60.1	40.4	30.7	24.6	19.1	14.1	19.5	40.4	
	C	60.9	44.5	36.5	29.0	22.5	17.6	16.2	44.5	

<sup>3)</sup>/ Percentages are averages of four to five samples.

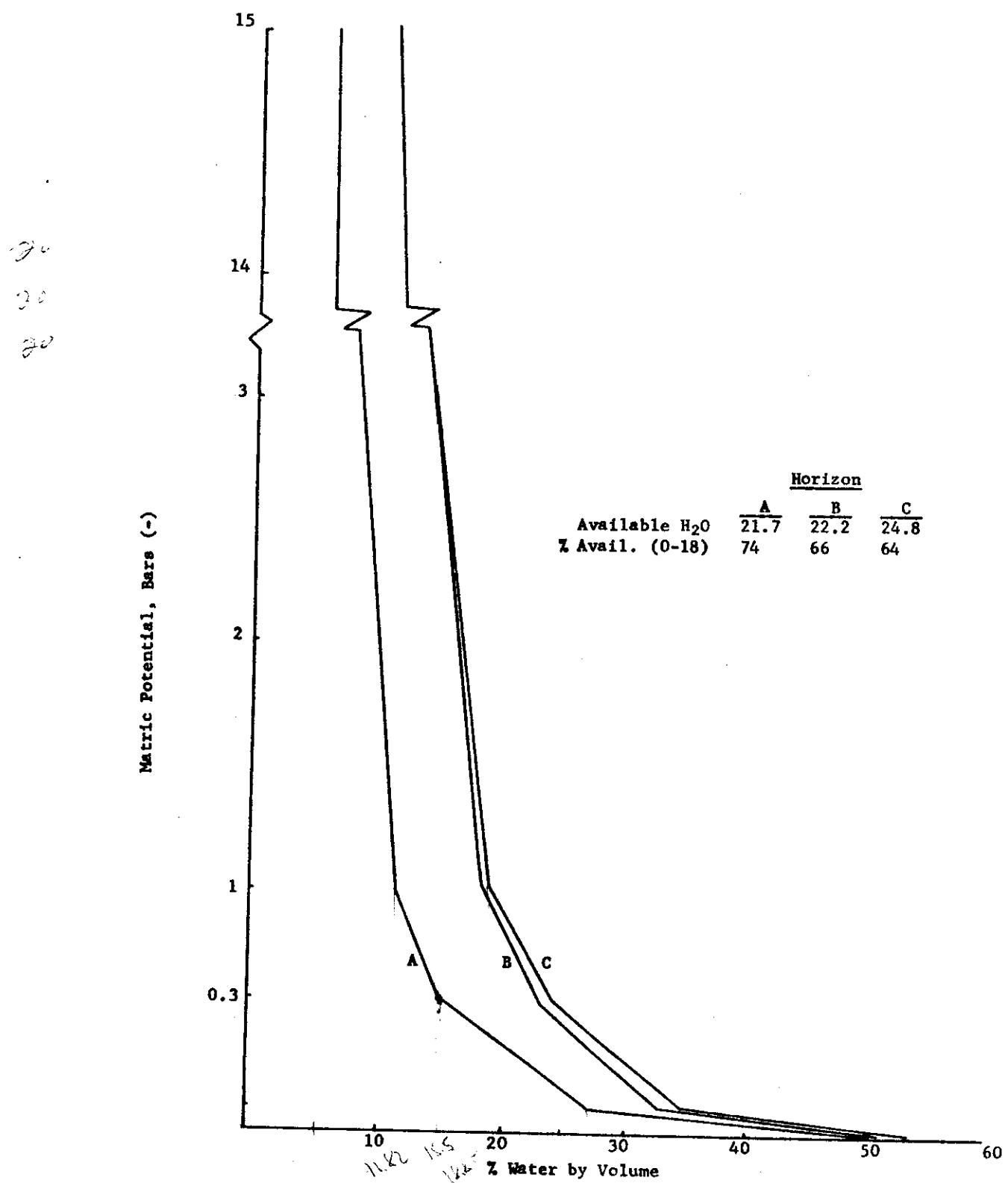


Fig. 13. Water retention of Ascalon soil.

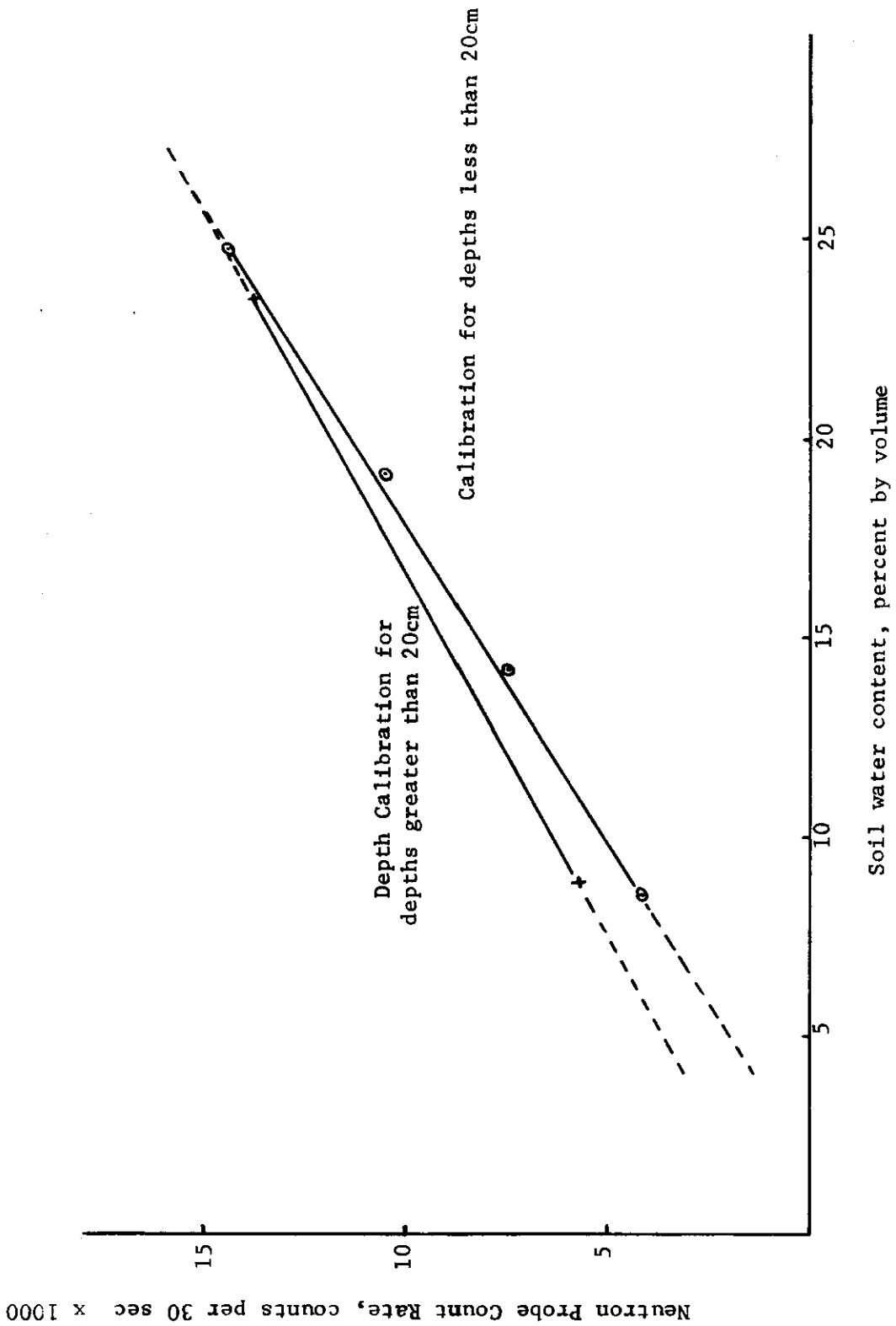


Fig. 14. Depth and surface calibrations for neutron probe.

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APPENDIX

FIELD DATA

Soil Moisture Transect Data

Soil Moisture Transect Data collected in 1970 at the Pawnee Site is Grassland Biome Data Set A2U703B. Data were collected on FORTRAN Coding Forms as follows:

---

Header Card (card 1)

---

Columns	Contents
2 - 9	Date (Month, Day, Year), eq., 08-05-70.
11	Number of this header card for this transect.
13 - 15	Maximum depth at which readings were made (cm).
17	N - defined as 10-J where J = the number of access tubes for which data are recorded on the following cards.

Sequence of Data Collection Card (card 2)

This card contains a pair of numbers for each of the J access tubes read. The first number of the pair is the tube number (a zero means tube 10) and the second number is the number of the neutron probe used on that tube. The format is (2X, 20I3).

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Data Cards (cards 3 through 10)

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Columns	Contents
1 - 3	Depth (cm) at which the readings on this card were taken.
4 - (3 + 6J)	Counts recorded in 1/2 minute for the J access tubes, in the order that tube numbers appeared on card 2.

---

This is followed by more sets of ten cards for this transect until data from the access tubes on the transect are recorded. Data for the remaining transects follow in the same fashion until all transects are recorded.

This data set will be changed to a more efficient format which will be compatible with the data which will be collected in 1971.

A listing of the existing 1970 field data follows.

\*\*\* FILED DATA \*\*\*

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12345678901234567890123456789012345678901234567890123456789012345678901234567890

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30	07105	16241 16367 10842 10245 10533 16834 13071 16814 12695
45	07524	14274 15957 11544 04740 08210 15920 17292 10658 11670
60	08367	11434 16094 13942 04170 10297 17779 17608 08812 10904
75	08258	13204 16786 14303 10111 13442 16141 17508 07331 11674
90	07474	16664 17585 15467 16559 14601 14748 17097 07913 15610
120	07627	15754 17645 17324 13699 17528 17865 17666 09998 17531
150	11675	14655 17685 18455 15486 10282 17885 21269 14095 19682
08-05-70	2	150
1	11	2 11
15	04383	06626 04674 13699 04286 03370 03539 03765 03423 04716
30	11246	10899 10773 07572 11671 06295 05005 06506 08490 08081
45	12974	12042 13285 12290 12413 06161 07642 07561 09073 07009
60	12231	11471 12118 15344 12119 08946 11300 07924 09082 06939
75	14767	14562 10374 13283 14655 07771 11691 08996 08921 07734
90	15286	15776 10340 09457 09443 09283 10653 14362 09406 10354
120	15502	17497 13243 12443 13042 13799 17383 17250 12406 08170
150	14274	19176 13518 13890 15199 10324 09977 19040 17650 07259
08-05-70	3	150
1	11	2 11
15	04226	03694
30	06869	07870
45	06329	04212
60	06271	06384
75	05165	05571
90	05624	05731
120	05163	07402
150	15216	07059
07-29-70	1	150
1	11	2 11
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45	07585	16220 14524 12023 09208 09184 14653 18092 10266 12816
60	08569	12070 15632 13575 09457 11152 16475 17655 08281 11592
75	08435	13298 16675 14564 10742 13754 16129 18016 06740 11774
90	07646	16573 17562 15744 16303 14923 14719 17029 07732 15669
120	07501	15645 17562 17612 13671 17053 17974 17678 09903 17503
150	11722	14563 17562 13567 15625 11223 17988 20958 13482 18431
07-29-70	1	150

07-29-70 2 150  
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09-02-70 1 150  
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08-04-70 1 150

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