

THESIS

THE EFFECTS OF IRRIGATION RETIREMENT ON SOIL CARBON DYNAMICS OF A  
CONTINUOUS MAIZE AGROECOSYSTEM

Submitted by

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## ABSTRACT

### THE EFFECTS OF IRRIGATION RETIREMENT ON SOIL CARBON DYNAMICS OF A CONTINUOUS MAIZE AGROECOSYSTEM

Over half of the world's fresh water is used in crop production and, in some key agricultural regions, use far exceeds local water availability and recharge rates. With the increasing strain on freshwater resources caused by climate change and a growing population, agriculture is under pressure to reduce its water consumption and large areas of currently irrigated farmland across the Western U.S. will likely transition into dryland agriculture over the coming decades. The effects this will have on global soil carbon (C) dynamics, however, remain unclear. In 2016, a study was established in Northern Colorado to understand how stopping irrigation affects soil C turnover in a no-till, continuous maize agroecosystem. Earlier results showed limited responses of the soil microbial community to irrigation retirement, but differences in soil heterotrophic respiration ( $R_h$ ) rates were detected after two years of accumulated differences in plant residue inputs, thus suggesting a possible co-limitation of water and available C to the microbial community. We continued this experiment through 2022 to further explore the relationship between soil moisture and C inputs in shaping the soil microbial community under the new watering regimes and the consequential effects on soil respiration ( $R_s$ ) as an indicator of soil organic C (SOC) turnover rates. Two seasons of data collection in 2021 and 2022 showed decreases in available soil water, bacteria, fungi, protozoa and actinomycetes fatty acid methyl ester (FAME) biomarkers, activities of four extracellular enzymes and soil autotrophic respiration in response to both reductions in irrigation and plant inputs, with strong

interactive effects between the two factors. However, plots under dryland conditions had higher concentrations of dissolved organic carbon (DOC) and muted differences in soil  $R_h$  when compared to their irrigated counterparts; differences in  $R_h$  between fallow treatments with (YF) and without residue inputs (LTF), on the other hand, were more pronounced. Soil  $R_s$  in fallow plots was consistently, positively correlated with field soil temperature, while correlations with moisture were weak or even negative, thus suggesting soil moisture was not a strong direct driver of  $R_h$ . We investigated the direct and indirect influences of variables collected monthly across two seasons on soil  $R_h$  to test our hypothesized model using structural equation modeling. In contrast to the cumulative treatment level impacts of plant inputs and irrigation, monthly soil moisture measurements had a stronger, direct effect on  $R_h$  than substrate availability as estimated by water-extractable DOC. The final model only explained 24% of the variability in soil  $R_h$ . Changes in global C dynamics can be expected with transition of land areas from irrigated to dryland agriculture. However, focusing on soil health, resource conservation practices and the resiliency of the soil microbiome can be the key to minimize the potential negative impacts of this transition.

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## BACKGROUND AND RATIONALE

It is estimated that 70 percent of the world's fresh water is utilized to support crop production (FAO, 2011). Global crop production depends on water received as both precipitation and irrigation from surface water bodies and aquifers (Davis et al., 2017). In regions like the Great Plains, withdrawal rates to support agriculture far exceed local freshwater availability and aquifer recharge rates (McGuire (USGS SRI), 2011). Climate change studies predict shifts in precipitation timing and amount (Sherwood and Fu, 2014; Southworth et al., 2000). This, along with increasing global temperatures, could potentially further impact recharge rates and water availability. Moreover, due to rapid population growth, municipal and industrial water demands continue to increase, further putting pressure on limited freshwater sources. In some regions, this decline in resources is causing and/or will cause some farmers to transition from fully irrigated to deficit or complete dryland farming (Wheeler et al., 2008; Amosson et al., 2009; Rudnick et al., 2019).

The effects that these transitions will have on soil health and carbon (C) dynamics remain unclear. The retirement of irrigation will reduce crop productivity, thus reducing crop residue inputs to the soil with likely cascading impacts on multiple soil functions (Conant et al., 2004). Soil functions are regulated, in general, by soil biological communities and, thanks to their rapid response to environmental changes, they have proven to be effective predictors of soil health changes (Acosta-Martínez et al., 2011; Cano et al., 2018, Kallenbach et al., 2016; Lehmann et al., 2020; Fierer et al., 2021).

Although negative impacts on soil microbial communities can be expected in response to irrigation retirement (Siebelec et al., 2020; Hueso et al., 2011), the magnitude of these effects can vary. In climates with limited water availability, soil microbial biomass has been reported to decrease along with soil moisture (Carbone et al., 2011, Davidson et al., 1998; Wu et al., 2011; Manzoni et al., 2012; Sardans et al., 2008; Pérez-Guzmán et al., 2020), have no response (Canarini et al., 2016) or even increase when exposed to drought (Schaeffer et al., 2017). Further, drought drives shifts in the composition of the soil microbial community. Long-term drought has been reported to reduce bacterial and fungi abundance (Maestre et al., 2015), with bacteria being generally more sensitive to drought than fungi (Evans and Wallenstein, 2012).

The high temporal variability of available water in soils affects soil microbial activity (Porporato et al., 2004). Microorganisms may opt for dormancy (Jones and Lennon, 2010) or shift extra-cellular enzyme (EE) production to optimize C uptake when water is limited and conditions become unfavorable (Vetter et al., 1998; Allison, 2012; Moorhead et al., 2012). Soil EE activity tends to decrease under drought conditions, but the effects can be inconsistent (Sardans et al., 2008; Alster et al., 2013; Ren et al., 2017). Drought experiments have revealed decreases in activities of  $\beta$ -glucosidase (BGLUC) (Pohlon et al., 2013), leucine aminopeptidase (LAP) (Schreckinger et al., 2021),  $\beta$ -cellobiosidase (BCELL) and N-acetyl- $\beta$ -glucosaminidase (NAG) (Weitnraub et al., 2007) enzymes. On the other hand, an experiment which subjected a Central-European grassland to 100 years of extreme drought reported limited EE activity responses to reduced soil moisture (Kreyling et al., 2008), suggesting high adaptability of the soil microbial communities to drought (Sinsbaugh, 1994; Williamson and Wardle, 2007).

While soils may experience limited activity during drought periods, rewetting events (i.e., rainfall, irrigation) create large surges in activity (Austin et al., 2004; Borken and Matzner,

2009). Hence, even small increases in soil moisture after a long dry period may trigger large pulses in CO<sub>2</sub> efflux (Birch effect) (Birch, 1958; Reichstein et al., 2002; Carbone et al., 2011, Kim et al., 2012, Göransson et al., 2013). For this reason, soil respiration ( $R_s$ ) has the greatest level of uncertainty relative to other fluxes in global terrestrial C cycle estimates (Bond-Lamberty and Thompson, 2010). Soil respiration can be divided into autotrophic ( $R_a$ ) and heterotrophic ( $R_h$ ) components (Subke et al., 2006). Plant-derived  $R_a$  includes root and rhizomicrobial respiration, while  $R_h$  originates from microbial decomposition of litter and soil organic matter and has been reported to represent between 40 and 80% of  $R_s$  in maize cropping systems (Zhang et al., 2013; Li et al., 2019). Soil respiration generally increases with soil moisture (Cook and Orchard, 2008; Zornoza et al., 2016; Zhang et al., 2019; Zhou et al., 2019), therefore, a decrease in  $R_s$  can be expected after transitioning into a non-irrigated system. Since  $R_a$  does not contribute to soil C sequestration (Kuzyakov, 2006), the effect of the transition into dryland agriculture on SOC depends heavily on the changes in  $R_h$ . Soil temperature and moisture (Schipper et al., 2019), microbial community composition and activity (Troost et al., 2013; Ma et al., 2015; Calderon et al., 2016) and C availability (Orchard and Cook, 1983; Cook and Orchard, 2008; Canarini et al., 2017) all influence  $R_h$ , but it is still challenging to anticipate how  $R_h$  will respond to shifts in environmental conditions and C inputs (Falloon et al., 2011; Moyano et al., 2013; Bond-Lamberty et al., 2018).

A transition experiment was established in 2016 to determine the effect of irrigation retirement on SOC dynamics—more specifically C outputs—in a continuous maize agroecosystem in the High Plains of the United States (Nuñez et al., 2022). This experiment focused on comparing biomass production and soil respiration between irrigated and non-irrigated (dryland) continuous maize. Fallow subplots were installed to partition soil respiration

into its different components: a yearly fallow (YF) subplot where no new plants were allowed to grow but plant residue was retained from the previous year has served to represent  $R_h$  derived from recent and older residue and SOC sources, and a long-term fallow (LTF) subplot to estimate  $R_h$  from native SOC. After three years, a significant decrease in soil moisture and biomass production was observed in the plots retired from irrigation relative to the irrigated treatments. Despite lower biomass—hence reduced C returns into the soil—the irrigation retirement effect on soil microbial community size and composition was limited. Lower extra-cellular enzyme activity (EEA) was observed in the dryland maize system suggesting that irrigation retirement caused a reduction in soil microbial activity. Soil C outputs via  $R_h$  (YF plots), however, were affected by the change in irrigation regime only after two years of accumulated differences in litter inputs. Respiration of native SOC from LTF plots was not affected by irrigation. These observed effects, or the lack thereof, seem to point out that the irrigation treatment affected microbial respiration indirectly, due to changes in substrate (labile C) availability, rather than direct effects of soil moisture alone.

Drought affects the mechanisms of dissolution, diffusion and transport that allows microbial access to substrates (Schimel, 2018). In microbial dynamics, water acts as a resource, a solvent and a way of transport to produce complex biogeochemical processes (Tecon and Or, 2017; Moyano et al., 2013). With very few exceptions, the substrates used by microbes as energy sources are water soluble, often referred to as dissolved organic carbon (DOC) (Cleveland and Yavitt, 1997; Schimel, 2018). Because most substrates are water soluble, connected water-filled pores are vital to allow transport of resources to microbes (Skopp et al., 1990; Moldrup et al., 2001; Bailey et al., 2017; Tecon and Or, 2017). Schimel & Schaeffer (2012) argued that microbial physical access to soil C is the primary control on whether that carbon is decomposed

(and respired to CO<sub>2</sub>) rather than other factors such as microbial community composition or the chemical quality of the substrates.

A model proposed by Manzoni et al. (2014) and based on the structure proposed by Schimel and Weintraub (2003) aims to assess microbial responses to changes in moisture by coupling diffusion limitations due to low moisture in the system to osmoregulation, dormancy and extra-cellular enzyme synthesis in soil microbes. In this model, soil organic C receives a litterfall input that is decomposed by enzymatic reactions to soluble compounds that are then allocated by the same community for growth, enzyme production and synthesis of osmolytes, which together explain variability in heterotrophic soil respiration. Using the conceptual framework proposed in this model and based on results reported by Nuñez et al. (2021, 2022), we designed an experiment to explore the relationship between soil moisture and available C substrates in the shaping of the microbial community and their activity, measured as R<sub>h</sub> in the field and EEA in the lab.

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## CHAPTER 1

### LINKING SOIL MICROBIAL CO<sub>2</sub> RELEASE TO SUBSTRATE AND MOISTURE IN A DRYLAND AGROECOSYSTEM

#### **1 Introduction**

Global fresh water sources are under increasing stress. Agriculture consumes almost three quarters of global water withdrawals (FAO 2011) and in some key agricultural regions this consumption far exceeds local availability of water resources (McGuire (USGS SRI), 2011; Davis et al., 2018). As municipal and industrial demands increase (FAO 2011) and water availability decreases, the pressure to reduce agricultural water consumption grows and the number of farms transitioning to dryland agriculture in regions like the High Plains is expected to increase (Whittemore et al., 2016; Nuñez et al., 2022). Irrigated systems are far more productive than their dryland counterparts, therefore, the transition to dryland agriculture poses many challenges to not only crop yield, but also soil C dynamics and soil health.

Soil C plays a key role in many soil functions (Rawls et al., 2003; Wade et al., 2020; Kane et al., 2021) and is a good indicator of soil health (Cano et al., 2018; Lehmann et al., 2020). Unfortunately, SOC changes slowly and may take many years for changes in this pool to be quantifiable (Bradford et al., 2016). The soil microbial community, on the other hand, responds more quickly to changes in the environment (Acosta-Martinez et al., 2011) and microbial community abundance and composition has proven to be an effective benchmark of soil health, given its direct relationship to SOC formation, nutrient dynamics and other soil functions (Acosta-Martínez et al., 2011, Kallenbach et al., 2016; Cano et al., 2018; Fierer et al., 2021). In

addition, microbially produced soil EEAs are considered useful indicators of nutrient turnover and microbial activity, particularly under environmental stress (e.g., drought) (Kardol et al., 2010, Burns et al., 2013).

Global soils contain more C within the first meter than the atmosphere (Scharlemann et al., 2014) and a small change in soil C could have important effects on the atmospheric CO<sub>2</sub> concentrations (Köchy et al., 2015). Microbial respiration ( $R_h$ ) can account for up to 90% of the total soil CO<sub>2</sub> efflux depending on the ecosystem (Hanson et al., 2000; Curiel Yuste et al., 2004; Tang and Baldocchi, 2005; Churchland et al., 2013; Savage et al., 2013), and represents the C losses from the decomposition of litter and SOC by microorganisms (Subke et al., 2006). Studies suggest that seasonal and environmental factors drive soil  $R_h$  (Reichstein et al., 2003; Sperlich et al., 2016; de Dato et al., 2017), however, mechanisms at biome, regional and global scales remain uncertain. The study of the soil microbial community can provide insight into how the transition into dryland agriculture will affect soil processes, especially C turnover, and, in turn, expand current knowledge on the direct drivers of soil  $R_h$ . Decreases in soil moisture and plant productivity are a given when stopping irrigation, thus decreases in water and substrates (labile C) available to the soil microbiome are also expected and possible cascading effects in soil  $R_h$  will be observed.

The decline in soil moisture expected with the retirement of irrigation is likely to impact soil microbial abundance and community composition. Water acts as a resource, a solvent and a way of transporting substrates to soil microbes (Moyano et al., 2013; Tecon and Or, 2017). Thus, water availability directly and indirectly regulates soil communities and their function. Most of the changes in community structure are driven by death—organisms die from physiological stress, leaving other organisms dominating the community when activity and growth are limited

(Schimel, 2018). However, microbial communities have exhibited resilience by quickly returning to their original state after a disturbance, and, in certain cases, microbial composition may change but functionality remains unaffected (Allison and Martiny, 2008). It has been observed that microbial biomass often remains stable through month-long dry periods (Aponte et al., 2010; Landesman and Dighton, 2010; Parker and Schimel 2011; Schaeffer et al., 2017), with reductions in bacterial and fungal abundance (Maestre et al., 2015). Similarly, depending on what limits productivity in an ecosystem, changes in soil moisture may increase or decrease the ratio of bacteria and fungi (Schimel et al., 1999, Williams, 2007), as bacteria tend to be more sensitive to drought than fungi (Evans and Wallenstein, 2012; Göransson et al., 2013).

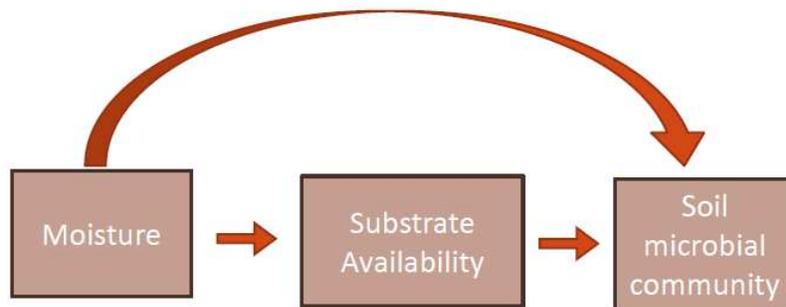
With changes in soil moisture, changes in microbial activity are often observed. To avoid losing water to their environment, microbes must accumulate solutes to match the water potential of their surroundings. The osmolyte demand is often not met in dry soils (Boot et al., 2013; Kakumanu et al., 2013) and microbes may opt for switching into a dormant state, consequently slowing down metabolic activity. Multiple studies have found linear decreases in microbial activity and respiration with decreasing water availability (e.g., Davidson et al., 1998; Jones and Lennon, 2010; Carbone et al., 2011; Wu et al., 2011; Manzoni et al., 2012). However, the magnitude of the decreases in respiration vary, and in some cases, even substantial reductions in soil moisture have only produced muted effects on respiration (Lu et al., 2017; Nuñez, 2021). It could be that microbes opt for a more efficient alternative to dormancy. The depolymerization of organic compounds catalyzed by EEA results in increase C availability (Schimel, 2018) and soil microbes may be able to optimize EE production to maximize C uptake under water limitation (Vetter et al., 1998; Allison, 2012; Moorhead et al., 2012). However, just like respiration, responses of EEAs to moisture are inconsistent (Sardans et al., 2008; Alster et al., 2013).

Microbial access to soil C is the primary control on whether C is decomposed (and respired to CO<sub>2</sub>) (Schimel and Schaeffer, 2012). With very few exceptions, the substrates used by microbes as energy sources are water soluble, often referred to as dissolved organic carbon (DOC) (Cleveland and Yavitt 1997; Schimel 2018). Because most substrates are water soluble, connected water-filled pores are vital to allow transport of resources to microbes (Skopp et al., 1990; Moldrup et al., 2001; Bailey et al 2017; Tecon and Or, 2017). Therefore, additional to moisture, substrate accessibility could limit soil microbial responses to irrigation retirement. While crop production and plant inputs are expected to decrease under dryland agriculture, studies have shown that drought increases plants' root to shoot ratio (Pallardy and Rhoads, 1993; Aspelmeyer and Leuschner, 2004; Zhou et al., 2018), which could potentially indicate that substrate availability for microorganisms may be less affected than other inputs.

A transition experiment was established in 2016 to determine the effect of the ceasing of irrigation on SOC dynamics in a continuous maize agroecosystem in the High Plains of the United States (Nuñez et al., 2022). This experiment has focused on comparing biomass production and soil respiration between irrigated and non-irrigated (dryland) continuous maize. Fallow subplots were installed to partition soil respiration into its different components: a yearly fallow (YF) subplot where no new plants were allowed to grow but plant residue provided cover has served to represent  $R_h$ , and a long-term fallow (LTF) subplot to estimate respiration from native SOC. In each of the first three years following irrigation retirement, soil moisture and biomass production decreased significantly. Despite lower biomass—hence C returns into the soil— the effect on soil microbial community size and composition was limited. Lower extra-cellular enzyme activity (EEA) was observed in the third year following irrigation retirement suggesting that retirement reduced soil microbial activity. Soil C outputs via  $R_h$  (YF plots),

however, were affected by the change in irrigation regime only in the third year of accumulated differences in litter inputs. Respiration of native SOC from LTF plots was not affected by irrigation. These observed effects, or the lack thereof, suggested that the irrigation treatment affected microbial respiration indirectly, due to changes in substrate (labile C) availability, rather than changes in soil moisture alone.

Thus, this experiment was designed to explore the influence of different factors on C turnover under drought in the same agroecosystem by relating soil moisture, available substrates and microbial (heterotrophic) respiration ( $R_h$ ). We define substrate as the dissolved organic carbon pool. We also explored how soil moisture and substrate availability—in conjunction—affected soil microbial EEA and community composition, and their consequent effects on  $R_h$ . We hypothesized that the soil microbiome is co-limited by moisture and plant residue inputs (as the main source of substrates), rather than available water alone (Fig 1). Therefore, we expected to see a reduced water-soluble C availability in plots under non-irrigated conditions that, in turn, would result in lower enzymatic activity and respiration. This shift in activity may also be associated with a shift in microbial community composition.



**Figure 1.** We hypothesized that soil microbial activity is co-limited by substrate and soil moisture rather than water availability alone.

## 2 Materials and Methods

### 2.1 Site description

Our research was conducted at a field experiment established in 2016 at Colorado State University's Agricultural Research Development and Education Center (ARDEC) near Fort Collins, CO (40°39'16"N, 104°59'55"W; 1555 m a.s.l). The site is characterized by an average annual precipitation of 408 mm and annual mean temperature of 10.2°C (1981-2010 average, <https://usclimatedata.com/>). The soil is classified as a Fort Collins loam (Aridic Haplustalfs) (USDA NRCS, 2019). As described by Nuñez et al. (2021, 2022), the selected field was previously used for irrigated crop production and managed under conventional tillage but transitioned to a no-till system in 2016 when the present experiment was established. This factorial experiment has four replicate blocks of a continuous maize (*Zea mays*) system under two contrasting irrigation practices (full irrigation and no irrigation/dryland conditions) and two subplots within each main irrigation treatment plot representing a gradient of plant and plant residue inputs: long-term fallow subplot where all vegetation has been removed since the establishment of the experiment and a yearly fallow subplot where each year after planting a location is randomly selected and new vegetation is removed. Irrigated treatments are watered once per week from May to September. Weekly irrigation amounts were managed at the farm level and were defined based on moisture conditions, crop phenology, and weather forecast. As a long-term experimental site, our current analysis presented here covers the 2021 and 2022 growing seasons in continuation of results from previous seasons (2016-2019) that have already been reported (Nuñez et al., 2021, 2022).

## 2.2 Aboveground biomass

Aboveground maize biomass was measured in 2021 and 2022 at physiological maturity via hand harvest. To estimate C inputs for each treatment, we selected two representative areas (9.12 m<sup>2</sup>) at the center of each plot and collected all plants, cutting about 8 cm from the soil

surface. We weighed the entire sample and separated ears and grain from stover. Dry matter was estimated by oven-drying a subsample of each portion at 55-60 °C for 72hrs.

### *2.3 Soil sampling and analysis*

To evaluate the relationship between soil available C and water, and their collective importance in shaping an agroecosystem's microbiome, we measured soil gravimetric water content (GWC), available dissolved C (DOC), soil microbial activity and community composition across the growing season. Starting in August 2021, ten (August through October 2021 and May through October 2022) monthly surface soil samples (0-10 cm) were collected using a 2-cm diameter soil probe (Tate, 1979; Fang et al., 2005; Canizales-Paredes et al., 2012; Chen et. al., 2021). Triplicate soil cores were collected in linear transects from each subplot and main planted plot, homogenized using a 2mm soil sieve, and separated into subsamples for the various analyses within hours after collection. Soil subsamples for EEA and DOC/DON analyses were stored at 4°C for no more than 48hrs until ready to be analyzed and/or extracted. To measure GWC (lab GWC), approximately 10g of soil were weighed, dried in an oven at 105°C for 48hrs, and reweighed.

Soil microbial community composition was approximated via ester-linked FAME (EL-FAME) fingerprinting method. Microbial biomass and community composition can be estimated via fatty acid methyl ester (FAME) analyses since the abundance of individual fatty acids varies among soil microorganisms and membrane lipids have a quick turnover during metabolism (Balsler et al., 2019). These measures, along with soil microbial extracellular enzyme activity (EEA), are often used as proxies for soil health as they are sensitive to changes in management systems (Cano et al., 2018; Lehmann et al., 2020). Following Shutter and Dick (2000), ester-linked fatty acids from field-moist soil samples were extracted and analyzed as described in Li et

al. (2020) by the Acosta-Martínez's soil microbiology lab at USDA-ARS in Lubbock, TX. Out of a total of 116 named FAME biomarkers, 46 selected fatty acids were used as microbial biomarkers according to previous research (Olsson et al., 1995; Zelles, 1999; Willers et al., 2015). These included gram positive (i13:0, a13:0, i14:0, a14:0, i15:0, a15:0, i15:1  $\omega$ 6c, a15:1 $\omega$ 9c, i16:0, a16:0, i17:0, a17:0, i17:1 $\omega$ 9c, i18:0, i19:0), gram negative (cy17:0 $\omega$ 7c, i17:0 3OH, cy19:0 $\omega$ 7c, 14:1 $\omega$ 9c, 14:1 $\omega$ 8c, 15:1 $\omega$ 9c, 15:1 $\omega$ 8c, 15:1 $\omega$ 6c, 16:1 $\omega$ 9c, 16:1 $\omega$ 7c, 16:0 2OH, 17:1 $\omega$ 8c, cy17:0 $\omega$ 7c, 18:1 $\omega$ 7c, 18:1 $\omega$ 6c, 18:1 $\omega$ 5c, 19:1 $\omega$ 9c, 19:1 $\omega$ 8c, 19:1 $\omega$ 7c), Actinobacteria (10Me16:0, 10Me17:1 $\omega$ 7c, 10Me17:0, 10Me18:1 $\omega$ 7c, 10Me18:0, 10Me19:1 $\omega$ 7c) and Protozoa (20:4  $\omega$ 6c). Fungal biomarkers included Saprophytic Fungi (SF) (18:3 $\omega$ 6c, 18:4 $\omega$ 3c, 18:2 $\omega$ 6c, 18:1 $\omega$ 9c), and Arbuscular mycorrhizal fungi (AMF) (16:1 $\omega$ 5c). Marker notations are represented as methyl group (Me), cyclopropane group (cy), and iso- (i) and anteiso- (a) branched fatty acids. Abundances were calculated to concentration (nmol g<sup>-1</sup> soil). The bacterial sum was calculated based on gram-positive and gram-negative bacterial FAMES, while the fungal sum was calculated by summation of SF and AMF biomarkers.

We quantified four soil EEAs related to glucose, cellulose, protein and chitin degradation (Table 1). Following the protocol by Saiya-Cork et al. (2002) we homogenized 1 g of fresh soil in 125 mL of 50 mM, pH 8.1, tris buffer with a blender. We then transferred 200  $\mu$ L of each slurry into black, 96-well microplates and mixed with 50  $\mu$ L of each substrate. Slurries were also mixed with buffer only or with the corresponding standards (10 mM 4 methylumbelliferone, or 7-amino-4-methyl coumarin) as negative and quenching controls, respectively. Samples in microplates were incubated for 4 h at 25 °C and the developed fluorescence read on a microplate reader (Cytation 5, BioTek, Vermont, USA) at 365 nm excitation and 450 nm emission wavelengths.

**Table 1.** Enzyme name, abbreviation, function in the soil (nutrient cycle indicator), and final product for the four soil hydrolytic enzymes assayed.

| Enzyme Name                            | Abbreviation | Function in Soil   |
|--|--------------|--|
| $\beta$ -1, 4-n-acetyl-glucosaminidase | NAG          | Hydrolysis of chitin   |
| Leucine amino peptidase                | LAP          | Hydrolysis of amino acid residues (N- terminus of peptides and proteins) |
| $\beta$ -D-1,4-cellobiosidase          | BCELL        | Hydrolysis of cellulose and cellotetraose                                |
| $\beta$ -1, 4-glucosidase              | BGLUC        | Hydrolysis of cellulose  |

To quantify water-soluble C (DOC) available in the soil, duplicate water extractions in a 1:5 fresh soil to water ratio were performed. Soil slurries were shaken for an hour, centrifuged at 3400 rpm for 15 min and filtered through 0.2  $\mu$ m nylon syringe filters (J.T. Baker® syringe filter; VWR International, LLC, Radnor, PA, USA). Extracts were analyzed for total DOC content using a Shimadzu TOC/TN analyzer (TOC-L CPH/CPN; Shimadzu Scientific Instruments Inc., Columbia, MD, USA).

#### *2.4 Soil respiration*

To understand the effect of irrigation and substrate availability on different components of soil respiration we measured soil CO<sub>2</sub> flux from the main planted area and the yearly and long-term fallow subplots weekly from May to November in 2021 and 2022. A 20-cm diameter PVC collar was installed at the center of each fallowed subplot at approximately 10-cm depth as attachments 20 cm diameter survey chamber (LI-COR, Nebraska, USA). Two collars were placed in the main plot planted area, one within a planted row and another one in-between rows. Soil CO<sub>2</sub> flux measurements were conducted the morning before weekly irrigations using a portable infrared gas analyzer connected to a chamber (LI-8100, LI-COR, Nebraska, USA). We

used the SoilFluxPro software (2015, LI-COR Biosciences, Inc.) to estimate the rate of soil respiration by fitting an exponential equation to the CO<sub>2</sub> concentration increase during a 3-minute measurement period.

Soil respiration measured in the fallow subplots was used to approximate heterotrophic respiration ( $R_h$ ) (Subke et al., 2006). The measurements in the yearly fallow were assumed to include microbial respiration of native SOC and the decomposition of crop residues from previous seasons. We also assumed that soil respiration from the long-term fallow represented microbial respiration of native SOC, as these subplots received no new crop residues or living roots starting in 2016. Soil temperature and moisture (field VWC) to 10cm depth were also recorded during the 3-minute CO<sub>2</sub> measurement period at the collar location using the probe connected to the LI-COR chamber and analyzer.

## *2.5 Data Analysis*

We analyzed data using both the factorial experimental design as well as exploring relationships between variables using correlation analyses and structural equation modeling. To assess the effect of irrigation and plant input treatments on response variables, the experimental data was analyzed as a replicated, randomized split plot design with irrigation management (irrigated or dryland) as the main plot and fallow as the subplot (planted, yearly fallow, or long-term fallow).

The effect of irrigation and fallow treatment on the microbial components (EEAs, FAME biomarker abundances), available substrate (DOC) and collective  $R_h$  CO<sub>2</sub> efflux was assessed by creating individual repeated measures model for each with irrigation, fallow treatment and their interaction as fixed categorical factors and block, sampling and interaction of block:sampling as random effects. Fixed effects were tested by anova with Kenward-Roger, and means were

compared using Tukey adjusted pairwise comparisons. All data analysis was done with R version 4.2.3 (R Core Team, 2023). Most data was log-transformed to meet normality assumptions (Shapiro-Wilk test), except for temperature and measures of water content (lab GWC, field VWC).

The response of soil respiration to soil temperature and moisture was analyzed across all sampling dates, with block, sampling date and their interactions as random effects. Relationships between individual variables were studied by creating linear regression models with block, sample ID and their interaction as random effects. Sample ID is a unique number assigned to each sample based on location in the field that remained constant across the duration of the experiment. For the temperature sensitivity of soil respiration, we used the van't Hoff equation (Lloyd and Taylor, 1994; Borken et al., 2006; Suseela et al., 2012; Nuñez et al., 2021):

$$R_s = a e^{bT} \quad (1)$$

where  $R_s$  is the measured CO<sub>2</sub> flux rate ( $\mu\text{mol m}^{-2} \text{s}^{-1}$ ),  $T$  is soil temperature ( $^{\circ}\text{C}$ ) at 10 cm,  $a$  is the basal respiration, and  $b$  is the temperature sensitivity of  $R_s$ . The response to moisture was evaluated by fitting a linear equation between respiration and field volumetric water content (%). We used only the  $R_h$  measured on the week preceding the date of soil collection to relate  $R_h$  efflux to the monthly measured variables, including lab GWC.

We used structural equation modeling (SEM) to explore the direct and indirect factors regulating soil  $R_h$  efflux (Ding et al., 2016). An SEM tests hypothetical pathways of influence among multiple variables using covariance among them (Grace, 2006; Miao et al., 2009; Shipley, 2000). This technique allowed us to quantitatively test our hypothesized model of direct and indirect effects of moisture and substrate available on  $R_h$  as mediated by soil microbial variables. To obtain the final SEM (Fig 7), first we established a base model based on empirical knowledge

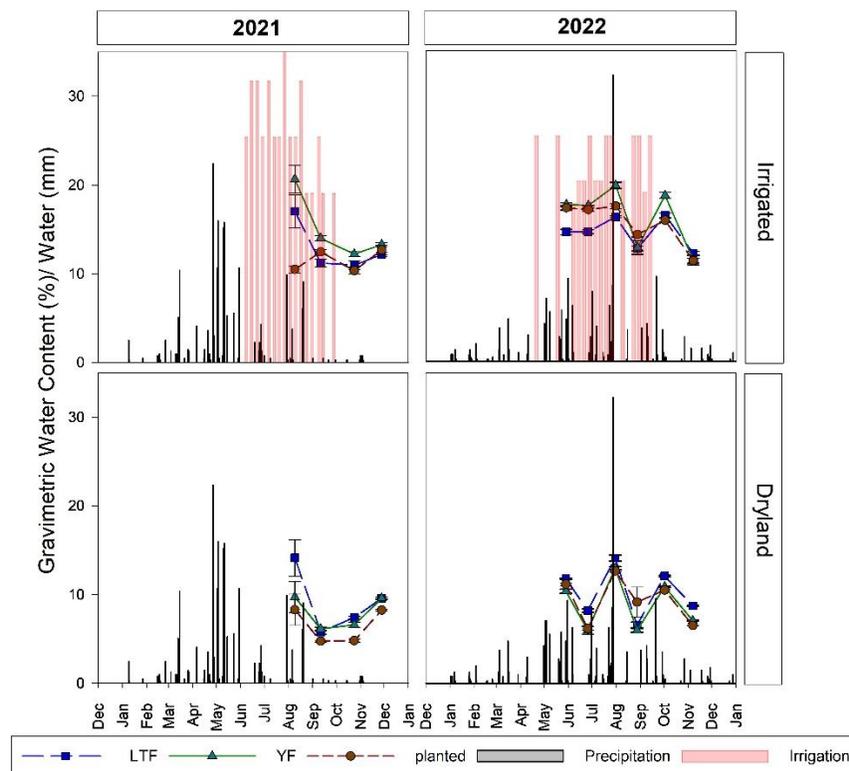
and proposed mechanisms (Fig 1 and Fig S8). We assumed a direct relationship between soil moisture and DOC, microbial community composition, EEAs, and soil  $R_h$ , since available water plays a key role in soil substrate availability and microbial processes. We assumed the microbial community would have a direct effect on soil  $R_h$  through community composition and an indirect effect through EE production. We created latent variables for each of these sets of microbial measures. Moreover, available substrates defined as DOC were theorized to play direct roles in soil  $CO_2$  release and indirect roles through regulating microbial community composition and EEAs.

Once the base model was defined, we analyzed result statistics to optimize it. All data were tested for normality (Shapiro-Wilk) and non-normal data was log-transformed. First, we examined modification indices to ensure that no important paths were left out of the model. Non-significant paths were removed to improve model parsimony without removing pathways central to our core hypotheses. The iterative model optimization continued until the model predictions—chi-square ( $\chi^2$ ) statistic, whole model p-value, the root-mean-square error of approximation (RMSEA), comparative fit index (CFI) and Akaike information criterion (AIC) — fit well with the observed values (Grace, 2006; Ding et al., 2016). Low values of  $\chi^2$ , RMSEA and AIC and high p-value and CFI suggest that there is a small difference between modeled and observed values (Grace, 2006; Shipley, 2000; Ding et al., 2016). The “lavaan” package in R version 4.2.3 was used to perform SEM analyses (R Core Team, 2015).

### **3 Results**

#### ***3.1 Soil moisture***

Soil moisture under dryland and irrigated water treatments followed similar temporal dynamics across both years (Fig 2). The irrigated treatment received nearly three times as much water (1160 mm vs 400 mm) than dryland across both 2021 and 2022 growing seasons, resulting in overall higher soil moisture when compared to dryland (Fig 2 and 3A,  $p < 0.01$ ). There was also an interactive effect of plant residue inputs and watering regime on soil moisture (Fig 3A,  $p < 0.01$ ) as planted subplots had the lowest moisture across both watering regimes (Fig 3A, Table S2). However, there was no difference between planted, YF, and LTF subplots within the same irrigation treatment (Fig 3A, Table S2).



**Figure 2.** Monthly average (mean  $\pm$  SE) soil (0-10 cm) gravimetric water content (shapes), weekly irrigation (red lines) and daily precipitation (black lines) dynamics of a continuous maize no-till agroecosystem under two irrigation regimes— irrigated and dryland agriculture—and three plant and residue input treatments— long-term fallow (LTF), yearly fallow (YF), and planted— in Northern Colorado.

### 3.2 C inputs

Our experimental design created a gradient of plant inputs across different water treatments. The LTF plots received no plant or plant residue inputs since the experiment initiation in 2016. The YF plots, on the other hand, received plant inputs each year except for the measurement year, thus included accumulated residues from harvest of the previous years. Overall, plots under irrigation, except for LTF, received plant residue inputs nearly three times larger in magnitude than their dryland counterparts (Table 2). Two years of crop failure resulted in no grain harvest from the dryland maize plots, but plant residue was retained (Table 2).

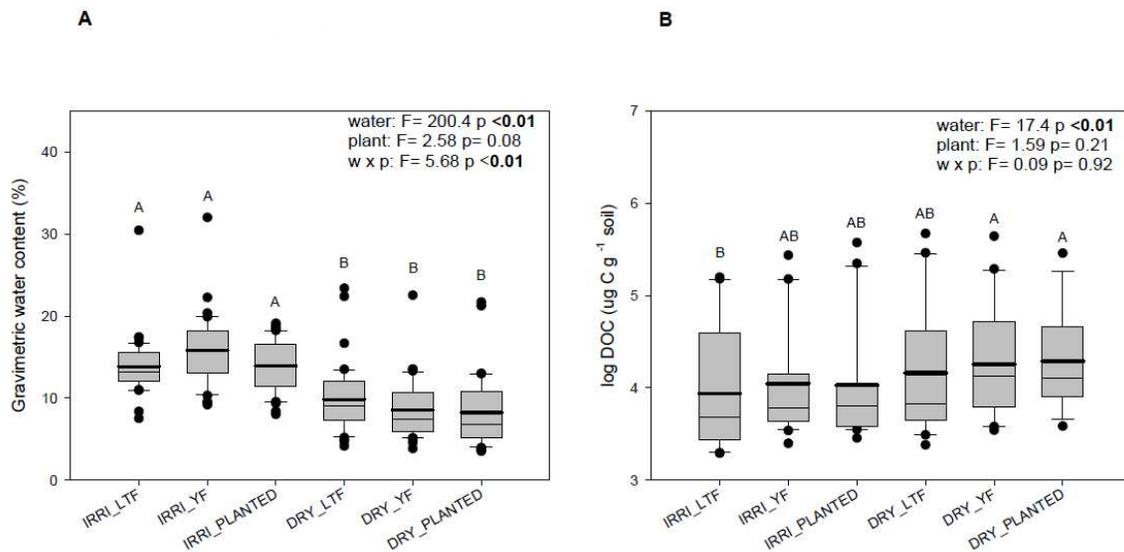
**Table 2.** Biomass inputs (Mg dry matter ha<sup>-1</sup>) in irrigated and dryland maize during three growing seasons. Values represent the mean  $\pm$  1 SE with n = 4.

| Season      | Irrigated maize                       |                  | Dryland maize                         |                 |
|-------------|---------------------------------------|------------------|---------------------------------------|-----------------|
|             | Aboveground Biomass (Grain + residue) | Plant residue    | Aboveground Biomass (Grain + residue) | Plant residue   |
| *2020       | 15.78 $\pm$ 1.50                      | 7.46 $\pm$ 0.79  | 0.95 $\pm$ 0.3                        | 0.95 $\pm$ 0.3  |
| 2021        | 17.93 $\pm$ 0.61                      | 8.16 $\pm$ 1.44  | 2.60 $\pm$ 0.16                       | 1.58 $\pm$ 0.25 |
| *2022       | 13.44 $\pm$ 0.30                      | 6.40 $\pm$ 0.52  | 4.49 $\pm$ 0.53                       | 4.49 $\pm$ 0.53 |
| Accumulated | 47.15 $\pm$ 1.65                      | 22.02 $\pm$ 1.72 | 8.04 $\pm$ 0.63                       | 7.02 $\pm$ 0.66 |

\* Years of dryland maize crop failure which resulted in total biomass being equivalent to plant residue as no grain was harvested.

### 3.3 Soil DOC

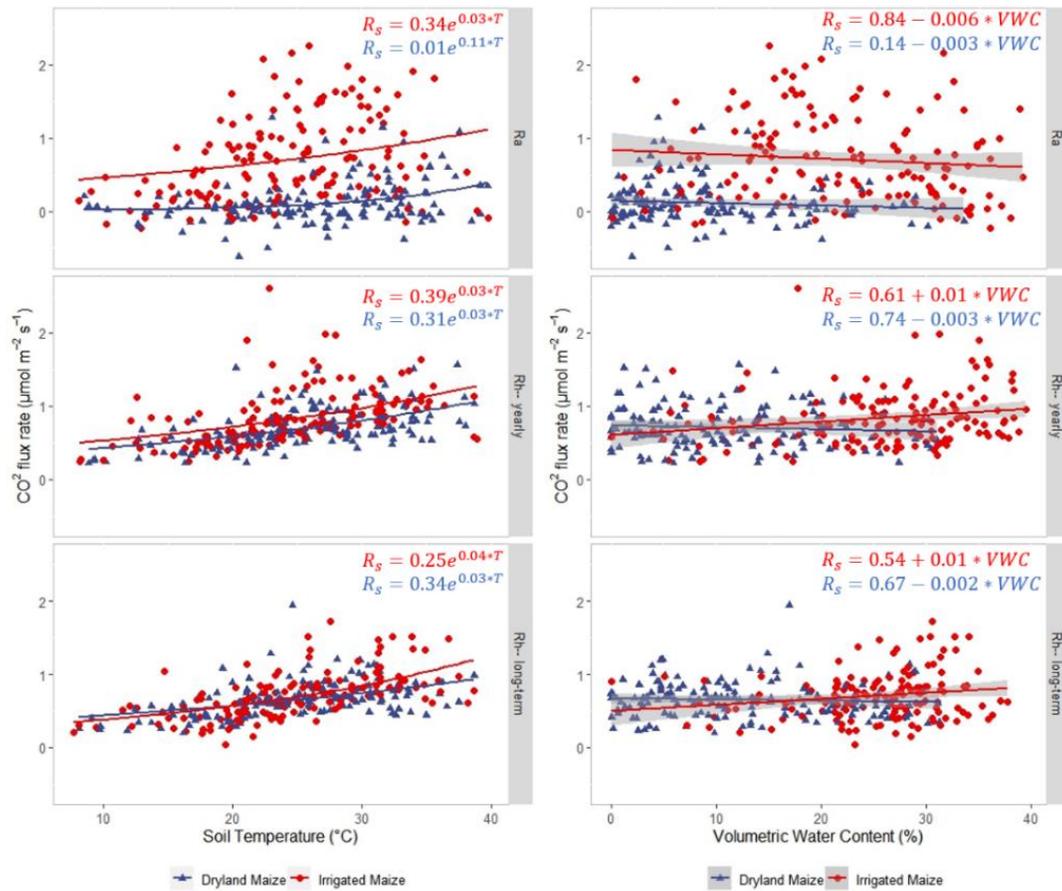
We observed a strong irrigation effect ( $p < 0.01$ ) on soil water-extractable C (Fig 3B). Despite the larger C inputs in irrigated systems (Table 2), dryland plots had, on average, slightly higher concentrations of water-extractable C (Fig 3B, Table S2), with the long-term fallow plot under irrigation being the lowest average of all six unique combination treatments and dryland planted being the highest (Fig 3B, Table S2). Subplots under the same irrigation regime did not differ in water-extractable C (Fig 3B, Table S2).



**Figure 3.** Gravimetric water content (A) and water-extractable C (B) from top-soil (0-10 cm) of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and planted— in Northern Colorado. Overall significance for a two-way ANOVA with the main factors of watering regime (water) and plant residue input (plant) are shown ( $\alpha = 0.05$ ). Different letters represent significant differences between unique treatments ( $p < 0.05$ ). Horizontal lines represent mean (bold) and median (light).

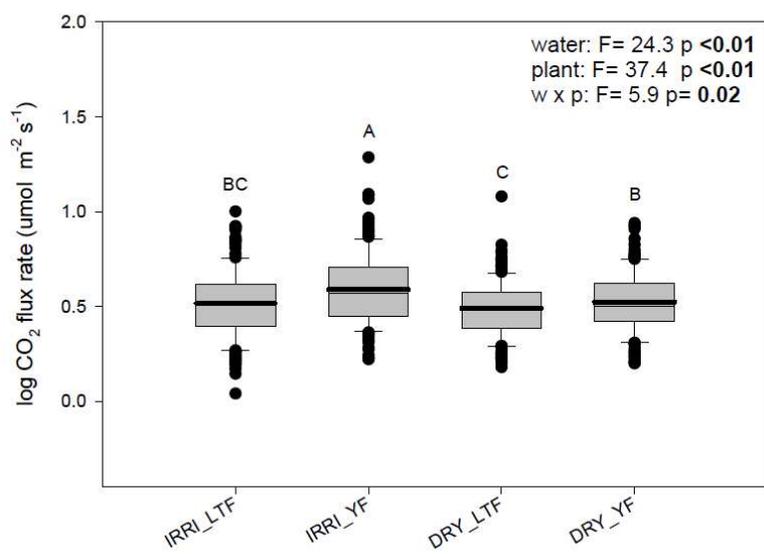
### 3.4 Temperature and moisture as drivers of soil respiration

Soil respiration increased exponentially with increases in soil temperature (Fig 4), with the autotrophic component exhibiting greater differences between irrigation treatments than the heterotrophic component. Increases in soil moisture, on the other hand, resulted in almost negligible increases heterotrophic respiration from irrigated fallow plots (IRRI YF and LTF: slope= 0.01,  $p = 0.01$  and  $0.03$ , respectively) and decreases in dryland fallow plots (DRY YF: slope= -0.003,  $p = 0.35$ ; LTF: slope= -0.002,  $p = 0.50$ ). Autotrophic respiration decreased with increases in soil moisture under both irrigation treatments (IRRI: slope= - 0.006,  $p = 0.24$ ; DRY: slope= -0.003,  $p = 0.27$ ).



**Figure 4.** Response of individual components of soil respiration— autotrophic ( $R_a$ ) and heterotrophic ( $R_h$ )— to field soil temperature (left) and moisture (right) in a Northern Colorado continuous, no-till maize agroecosystem under two irrigation regimes. Points represent individual measurements and lines represent fitted curves based on assumed exponential response to temperature and linear response to moisture.

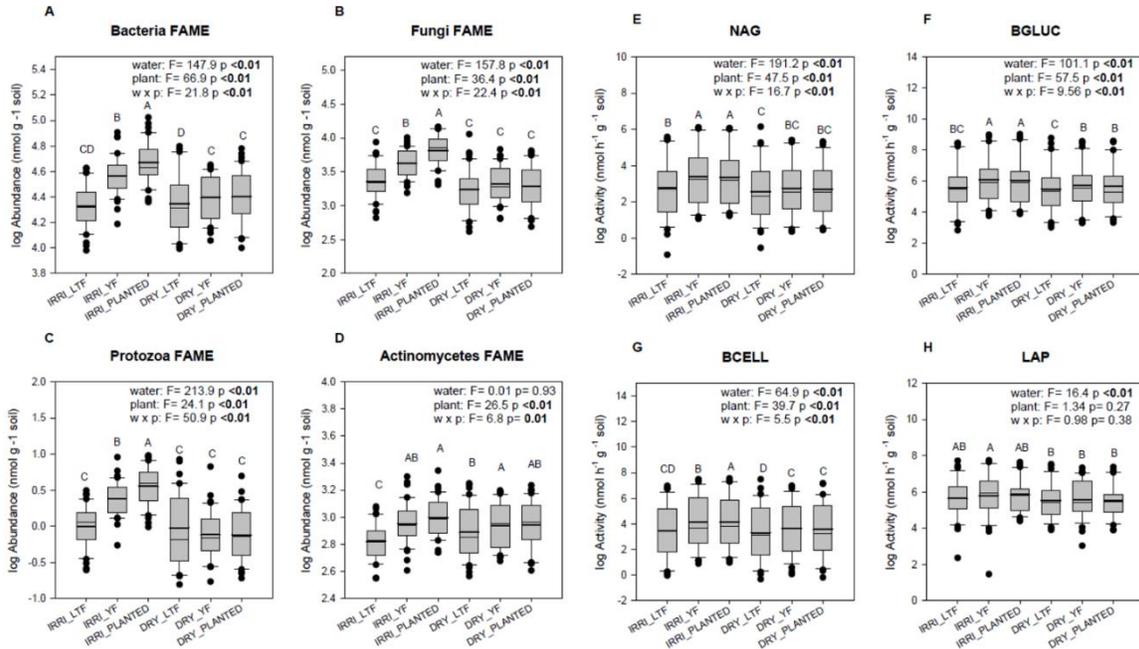
Irrigation treatment and plant inputs—as well as their interactions— had strong effects on overall soil respiration (cumulative  $R_s$  data not shown,  $p < 0.01$ ). Looking at just the heterotrophic component of soil respiration from the fallow subplots (Fig 5), we observed a response to irrigation ( $p < 0.01$ ) and plant inputs ( $p = 0.02$ ), and a weaker interactive effect of irrigation and plant inputs ( $p = 0.05$ ), with the yearly fallow plot under regular irrigation exhibiting higher rates of  $\text{CO}_2$  efflux than the other subplots (Fig 5, Table S2).



**Figure 5.** Soil  $R_h$  of irrigated (IRRI) and dryland (DRY) maize under two fallow treatments— long-term fallow (LTF), yearly fallow (YF), across two seasons. Overall significance for a two-way ANOVA with the main factors of watering regime (water) and plant residue input (plant) are shown ( $\alpha = 0.05$ ). Letters represent significant difference between the four unique treatments ( $p < 0.05$ ). Horizontal lines represent mean (bold) and median (light).

### 3.5 Microbial responses

Differences in plant inputs resulted in shifts in soil microbial community abundance with smaller effects of irrigation treatment. All selected FAME biomarkers for the microbial groups of interest showed strong responses to differences in plant inputs (Fig 5A-D; bacteria, fungi, protozoa and actinomycetes FAME:  $p < 0.01$ ), with planted plots generally showing the highest abundances in most microbial groups under both irrigation treatments, followed by YF and LTF, respectively (Fig 5; Table S2). Irrigation treatment, on the other hand, resulted in decreases in only three out of the four microbial groups for the dryland relative to irrigated plots (bacteria, fungi, protozoa FAME  $p < 0.01$ ; actinomycetes FAME  $p = 0.93$ ). However, interactive effects of irrigation and plant input treatments were observed in all four microbial groups (bacteria, fungi, and protozoa FAME  $p < 0.01$ ; actinomycetes FAME  $p = 0.01$ ). These interactions were driven by the joint effect of soil moisture and labile C in shaping the composition of the soil microbial community.

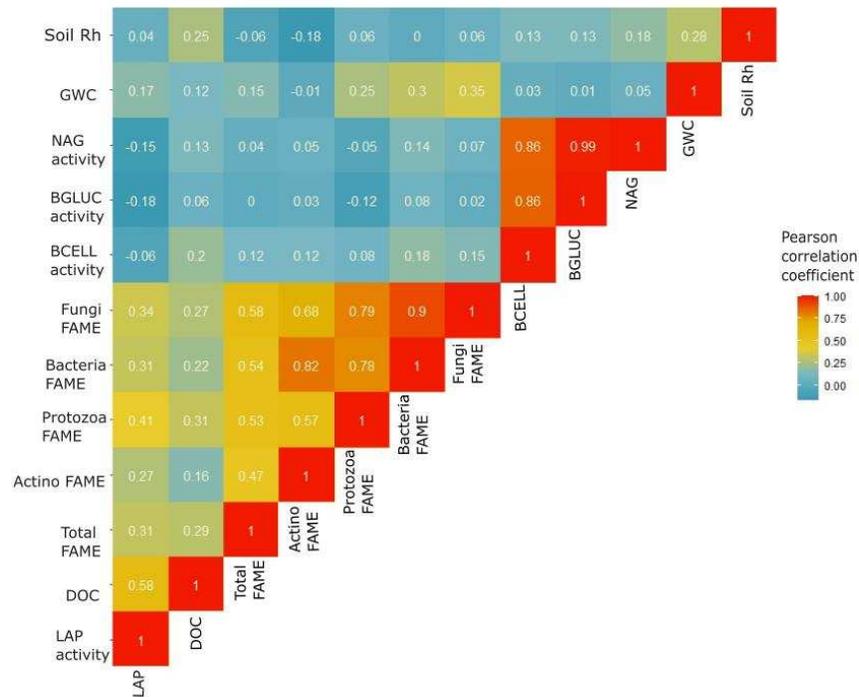


**Figure 6.** Fatty acid methyl ester biomarker abundances (A-D) and extracellular enzyme activities (E-H) from top-soil (0-10 cm) of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and regular planted— in Northern Colorado. Overall significance for a two-way ANOVA with the main factors of watering regime (water) and plant residue input (plant) are shown ( $\alpha = 0.05$ ). Letters represent significant differences between unique treatments ( $p < 0.05$ ). Horizontal lines represent mean (bold) and median (light).

Similarly, the irrigation treatments resulted in differences in activities of all four measured EEAs (Fig 3E-H; NAG, BGLUC, BCELL, and LAP =  $p < 0.001$ ) and were generally higher in plots under irrigation treatment (Table S2). Three out of the four EEAs also showed strong positive responses to plant inputs (Fig 5E-G, NAG, BGLUC and BCELL  $p < 0.01$ ) and the interaction between both water and plant input treatments ( $p < 0.001$ ), except for LAP (Fig 5H,  $p = 0.27$ ,  $p = 0.38$ , respectively). Under the irrigated treatment, planted plots showed higher NAG (Fig 5E,  $p < 0.01$ ), BGLUC (Fig 5F,  $p < 0.01$ ) and BCELL (Fig 5G,  $p < 0.01$ ) activities when compared to long-term fallow, but LAP activity did not differ by irrigated subplot (Fig 5H,  $p = 0.32$ ). The yearly fallow plots had intermediary EEAs between planted and long-term fallow in the irrigated system (Fig 5H, Table 3). A similar trend was observed under dryland conditions: planted plots showed higher activities on three out of the four measured EEAs when compared to LTF (Fig 5E-H, Table 3), but exhibited no difference from YF (Fig 5E-H, Table 3).

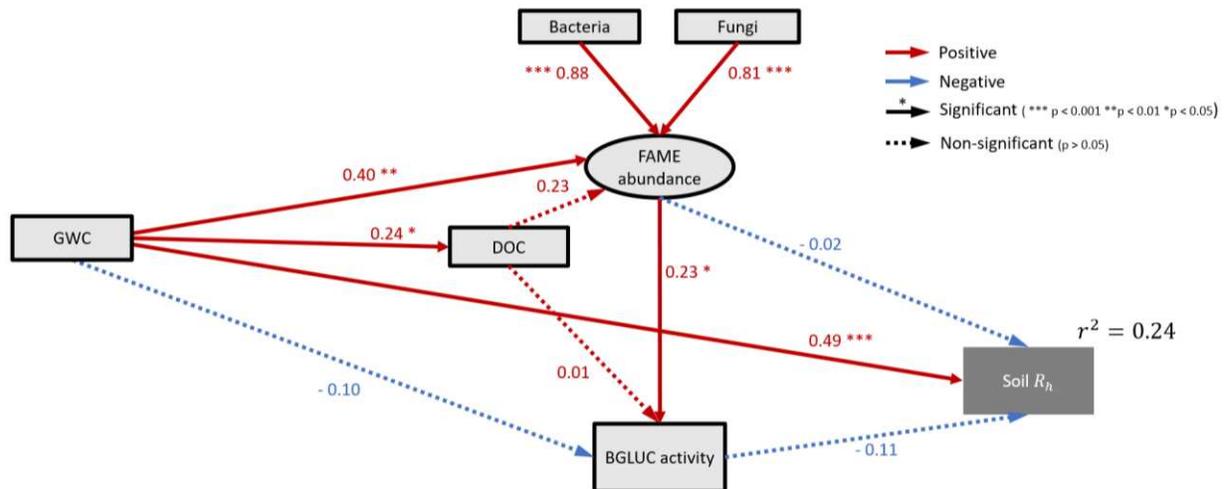
### 3.6 Exploring relationships

We explored simple, two-way correlations between soil moisture (GWC), substrate (DOC) and microbial properties (FAME, EEAs, soil  $R_h$ ) using data across all ten sampling time points (Fig S4-S7; Fig 7). Soil moisture was not correlated with DOC ( $r = 0.12$ ,  $p = 0.17$ ) or the measured EEAs (NAG:  $r = 0.05$ ,  $p = 0.45$ ; BGLUC:  $r = 0.01$ ,  $p = 0.92$ ; BCELL:  $r = 0.03$ ,  $p = 0.61$ ) except for LAP ( $r = 0.17$ ,  $p = 0.01$ ). Soil moisture was positively correlated with some FAME marker abundances (fungi:  $r = 0.35$ ,  $p = <0.01$ ; bacteria:  $r = 0.30$ ,  $p = 0.04$ ), but not others (protozoa:  $r = 0.25$ ,  $p = 0.88$ ). Similar to soil moisture, soil DOC was positively correlated with LAP ( $r = 0.58$ ,  $p = <0.01$ ), and total FAME ( $r = 0.29$ ,  $p = <0.01$ ), bacteria ( $r = 0.22$ ,  $p = <0.01$ ), and fungi ( $r = 0.27$ ,  $p = <0.01$ ), as well as protozoa ( $r = 0.31$ ,  $p = <0.01$ ) biomarker abundances. While soil  $R_h$  was not strongly correlated with field-measured soil moisture relative to soil temperature as shown in Figure 4,  $R_h$  showed a stronger correlation with GWC ( $r = 0.28$ ,  $p = <0.01$ ) than DOC ( $r = 0.25$ ,  $p = 0.12$ ). Both DOC and GWC were more strongly correlated with  $R_h$  than with the selected FAME biomarker abundances and EEAs.



**Figure 7.** Pearson correlation coefficient heatmap exploring the relationship among all measured variables. All data (except GWC) has been log transformed for normality.

To test our initial theoretical framework (Fig 1), an SEM of soil heterotrophic CO<sub>2</sub> efflux rate was constructed, including all measured variables regardless of correlation strength with the other factors (Fig 8). The soil microbial community abundance was best represented by a combination of soil fungal and bacterial FAME biomarker abundances while soil EEAs were best represented by the single BGLUC activity due to the high degree of correlation amongst the EEAs (Fig 8). Soil GWC had direct effects on DOC, FAME biomarker abundance and was the only factor with direct influence on soil R<sub>h</sub>. Soil moisture influenced soil EEA indirectly via its effect on FAME abundance, but EEA did not explain the variability in R<sub>h</sub>. Together, these variables predicted only 24% of the variance in the soil heterotrophic respiration (Fig 8).



**Figure 8.** Structural equation model (SEM) with variables (boxes) and potential causal relationships (arrows) for soil heterotrophic CO<sub>2</sub> release rate ( $R_h$ ) across a continuous maize system under two irrigation treatments. Arrows indicate the hypothesized direction of causation. The numbers are standardized path coefficients, which can reflect the importance of the variables within the model [Colman and Schimel, 2013]. The model had  $\chi^2 = 5.62$ , p-value = 0.23, CFI= 0.99, RMSEA = 0.07, and AIC = 1202.20.

#### 4 Discussion

We followed the changes in SOC dynamics of a continuous maize agroecosystem, five years after its transition from irrigated to a non-irrigated system, across two growing seasons. We quantified C inputs in the form of crop residues as well as C outputs in the form of heterotrophic soil respiration. We hypothesized that soil moisture and plant residue inputs (as the main source of substrates), in combination, control the abundance and activity of the soil microbial community, rather than available water alone. Our findings indicate that  $R_h$  was seemingly regulated by both available substrates and soil moisture; differences in soil moisture were responsible for seasonal changes in  $R_h$ , while cumulative variability in  $R_h$  could be explained by available substrates and/or the interaction between both factors.

Relative to the irrigated treatments, non-irrigated plots had reduced soil moisture, aboveground biomass, microbial biomass, and microbial activity, but higher DOC and similar rates of heterotrophic respiration. Dryland plots had lower plant residue returns, and

correspondingly lower soil C inputs, than irrigated plots. We observed shifts in microbial community abundance with decreases in available bacteria, fungi, protozoa and actinomycetes FAME biomarkers under dryland conditions and with reduced C inputs. Similarly, decreases in microbial activity—quantified as activities of four extracellular enzymes—were observed in the dryland system. On the other hand, plots under dryland conditions showed higher DOC concentrations and soil  $R_h$  fluxes that are comparable to those from their irrigated counterparts.

We investigated all possible correlations between factors and used structural equation modeling to explore the direct and indirect influences on soil  $R_h$  in the system by building causal relationships between all measured variables. Across all 10 sampling points, soil moisture was weakly correlated with  $R_h$ . While soil moisture and DOC had stronger correlations with  $R_h$  than other soil variables measured, only soil moisture explained a significant proportion of the variability in our final model.

#### *4.1 Effects on soil moisture and C inputs*

Higher soil moisture content in irrigated plots than dryland was expected, given that the irrigated treatment received nearly three times as many water inputs than dryland across both growing seasons. Productivity under dryland conditions suffered a six-fold decrease from irrigated plots, and received only a third of estimated plant residue returns, which was again expected given maize's high irrigation requirements and high water-stress susceptibility (Sherrod et al., 2014; Cano et al., 2018). The transition to dryland agriculture likely has had cumulative effects on substrate availability and accessibility to microbes, as the difference in biomass and water inputs increased with each growing season.

#### *4.2 Effects on the soil microbial community*

After two seasons, the size of three major microbial groups (bacteria, fungi, protozoa) in the system, assessed as EL-FAME, was affected by irrigation, while actinomycetes were mostly affected by plant inputs, with strong interactive effects between these factors for all groups. Differences in plant inputs, and hence likely substrate availability, was much greater between planted and fallow plots under irrigation than those under dryland. Thus, differences in microbial group abundances was greater between irrigated subplots than dryland, further emphasizing the importance of both substrate availability and soil moisture in sustaining microbial development. A higher biomass production corresponded with increased litter and root inputs, which are an important microbial source of C (Finzi et al., 2015; Calderon et al., 2016). This remains consistent with what was observed in previous years of this study (Nuñez, 2021), where, under irrigation, soil moisture had greater effects in community size (measured as PLFAs) in the yearly fallow subplots, but now these differences have magnified given the cumulative differences in plant inputs.

Similarly, three out of four soil EEAs (BGLUC, NAG, BCELL) showed strong responses to soil moisture, plant inputs and their interaction, while LAP was mostly influenced by moisture. Long-term fallow plots under both treatments showed similar levels of potential enzymatic activity, and differences between irrigated planted and yearly fallow plots were muted, further emphasizing the importance of substrate availability in influencing microbial activity, not just water availability. Although measured under saturated substrate concentrations (German et al., 2011), microbial activity measured as EEAs in the lab was consistent with overall trends in measurements of *in situ* microbial activity ( $R_h$ ).

#### *4.4 Influences in soil CO<sub>2</sub> efflux and SOC turnover*

Lower  $R_s$  fluxes were observed in non-irrigated plots. These differences were mostly driven by  $R_a$ , the component of respiration directly associated with plant growth (Hopkins et al., 2013). Variations in  $R_s$  related to plant productivity, even when  $R_h$  was the main component of  $R_s$ , have been reported (Xu and Shang, 2016; Gao et al., 2017). Plots differing only in soil moisture (LTF IRR v DRY) showed negligible differences in  $R_h$ , and the magnitude of differences in  $R_h$  observed in yearly fallow plots (IRRI v DRY) do not reflect the magnitude of differences in water and plant residue inputs that the systems have received.

While plant inputs were used as a proxy for substrate availability at the cumulative annual scale, we used DOC as the proxy for monthly differences in substrate availability (Fang et al., 2005; Dannenmann et al., 2009), given that most substrates used by microbes as energy sources are water soluble (Cleveland and Yavitt, 1997; Schimel, 2018). Our monthly sampling aimed to account for the rapid changes in this pool. Consistent with other drought studies in agroecosystems (e.g., Bloor and Bardgett, 2012; Niboyet et al., 2017; Peng et al., 2023), our results indicate a greater DOC concentration in the dryland system, where we know C inputs are lower. This could be interpreted as strong evidence that microbial physical access to soil C is the primary control on whether that carbon is decomposed (and respired to  $CO_2$ ) (Schimel and Schaeffer, 2012) and that the microbial community in these plots is water-limited and therefore unable to access substrates, thus allowing it to accumulate (Zeglin et al., 2013). This is supported by the slightly higher DOC in the planted and yearly fallow dry plots, and the observed lower potential enzymatic activity under dryland.

Using the conceptual framework proposed by previous models (Schimel and Weintraub, 2003; Manzoni et al., 2014), our initial theoretical framework depicted soil moisture and available substrates as co-limiting the microbial community and consequently soil  $R_h$ . The final

SEM revealed soil moisture as a direct driver of soil  $R_h$ , and a direct and indirect driver of soil DOC and microbial community abundance. Microbial effects on soil  $CO_2$  release rates could be enhanced only when aridity is alleviated (Wu et al., 2014; Ding et al., 2016). Thus, the microbial community shows a stronger moisture limitation, given accumulation of labile C available to microbes under dryland conditions. Soil moisture also influenced EE production indirectly as mediated by soil microbial community abundance. This may indicate that the soil microbial community has opted for dormancy awaiting better moisture conditions.

#### *4.5 Limitations in explaining soil $CO_2$ efflux variability*

The final model created based on our theorized framework only explained 24% of the variability in soil  $R_h$ . We aimed to relate enzymatic activity and activity in the form of *in situ* heterotrophic respiration to soil moisture and DOC. For this, we included a combination of field and laboratory analyses in our experimental design, and thus some of the missing relationships could be attributed to the disconnection of field and laboratory conditions. Similarly, unlike other models linking soil  $CO_2$  release rates to microbial properties (e.g., Allison, 2005; Davidson et al., 2012; Manzoni et al., 2014; Ding et al., 2016), we failed to include some of the same physical and environmental factors, such as pH and bulk density, into our initial framework. Another possible explanation to the low score in our model could lie in our choice of microbial group biomarker analysis. With lower cost and simpler methodology, EL-FAME has the potential to be a reliable method in agricultural management soil evaluations (Li et al., 2020). However, the EL-FAME method extracts all of the lipid groups present in the sample (Sharma and Buyer, 2015), this includes fatty acids originating from sources other than living microbes (Li et al., 2020). Thus, it is impossible to distinguish between active and inactive microbes just from EL-FAME analysis alone.

## ***5 Conclusion***

In summary, this study analyzed the effects of the transition into dryland agriculture on SOC dynamics and attempted to find links between soil moisture, soil DOC and the microbial community to explain variations in soil CO<sub>2</sub> fluxes. Our results revealed that soil moisture was a stronger driver of soil R<sub>h</sub> than other measured variables, with direct effects on soil DOC and microbial FAME biomarker abundance. Soil DOC, microbial FAME biomarker abundance and EEAs were of minor importance in the tested SEM. However, this does not mean that they should be disregarded, as factorial analyses showed that the retirement of irrigation and the gradient of soil plant residue inputs created by our experimental design affected these ecosystem elements and R<sub>h</sub>. Overall, our results help illustrate that changes in global C dynamics can be expected with transition of land areas from irrigated to dryland agriculture. However, focusing on resource-conserving practices and further exploring the resiliency of the soil microbiome can be the key to minimize the potential negative impacts of the transition into non-irrigated cropping systems.

## LITERATURE CITED

1. Acosta-Martínez, V., Lascano, R., Calderón, F., Booker, J. D., Zobeck, T. M., and Upchurch, D. R.: Dryland cropping systems influence the microbial biomass and enzyme activities in a semiarid sandy soil, *Biol. Fertil. Soils*, 47, 655–667, <https://doi.org/10.1007/s00374-011-0565-1>, 2011.
2. Allison, S. D.: A trait-based approach for modelling microbial litter decomposition, *Ecol. Lett.*, 15, 1058–1070, <https://doi.org/10.1111/j.1461-0248.2012.01807.x>, 2012.
3. Allison, S. D. and Martiny, J. B. H.: Resistance, resilience, and redundancy in microbial communities, *Proc. Natl. Acad. Sci. U. S. A.*, 105, 11512–11519, <https://doi.org/10.1073/pnas.0801925105>, 2008.
4. Alster, C. J., German, D. P., Lu, Y., and Allison, S. D.: Microbial enzymatic responses to drought and to nitrogen addition in a southern California grassland, *Soil Biol. Biochem.*, 64, 68–79, <https://doi.org/10.1016/j.soilbio.2013.03.034>, 2013.
5. Amosson, S., Almas, L., and Golden, B.: Economic Impacts of Selected Water Conservation Policies in the Ogallala Aquifer, *Ogallala Aquifer ...*, 1–50, 2009.
6. Aponte, C., Marañón, T., and García, L. V.: Microbial C, N and P in soils of Mediterranean oak forests: Influence of season, canopy cover and soil depth, *Biogeochemistry*, 101, 77–92, <https://doi.org/10.1007/s10533-010-9418-5>, 2010.
7. Aspelmeier, S. and Leuschner, C.: Genotypic variation in drought response of silver birch (*Betula pendula*): Leaf water status and carbon gain, *Tree Physiol.*, 24, 517–528, <https://doi.org/10.1093/treephys/24.5.517>, 2004.
8. Austin, A. T., Yahdjian, L., Stark, J. M., Belnap, J., Porporato, A., Norton, U., Ravetta, D. A., and Schaeffer, S. M.: Water pulses and biogeochemical cycles in arid and semiarid ecosystems, *Oecologia*, 141, 221–235, <https://doi.org/10.1007/s00442-004-1519-1>, 2004.
9. Bailey, V. L., Smith, A. P., Tfaily, M., Fansler, S. J., and Bond-Lamberty, B.: Differences in soluble organic carbon chemistry in pore waters sampled from different pore size domains, *Soil Biol. Biochem.*, 107, 133–143, <https://doi.org/10.1016/j.soilbio.2016.11.025>, 2017.
10. Balsler, T. C., Liang, C., and Gutknecht, J. L. M.: Linking microbial community analysis and ecosystem studies: A rapid lipid analysis protocol for high throughput, *Soil Ecol. Lett.*, 1, 22–32, <https://doi.org/10.1007/s42832-019-0003-0>, 2019.
11. Birch, H. F.: The effect of soil drying on humus decomposition and nitrogen availability, *Plant Soil*, 10, 9–31, <https://doi.org/10.1007/BF01343734>, 1958.

12. Bloor, J. M. G. and Bardgett, R. D.: Stability of above-ground and below-ground processes to extreme drought in model grassland ecosystems: Interactions with plant species diversity and soil nitrogen availability, *Perspect. Plant Ecol. Evol. Syst.*, 14, 193–204, <https://doi.org/10.1016/j.ppees.2011.12.001>, 2012.
13. Bond-Lamberty, B., Bailey, V. L., Chen, M., Gough, C. M., and Vargas, R.: Globally rising soil heterotrophic respiration over recent decades, *Nature*, 560, 80–83, <https://doi.org/10.1038/s41586-018-0358-x>, 2018.
14. Bond-Lamberty, B. and Thomson, A.: Temperature-associated increases in the global soil respiration record, *Nature*, 464, 579–582, <https://doi.org/10.1038/nature08930>, 2010.
15. Boot, C. M., Schaeffer, S. M., and Schimel, J. P.: Static osmolyte concentrations in microbial biomass during seasonal drought in a California grassland, *Soil Biol. Biochem.*, 57, 356–361, <https://doi.org/10.1016/j.soilbio.2012.09.005>, 2013.
16. Borken, W. and Matzner, E.: Reappraisal of drying and wetting effects on C and N mineralization and fluxes in soils, *Glob. Chang. Biol.*, 15, 808–824, <https://doi.org/10.1111/j.1365-2486.2008.01681.x>, 2009.
17. Borken, W., Savage, K., Davidson, E. A., and Trumbore, S. E.: Effects of experimental drought on soil respiration and radiocarbon efflux from a temperate forest soil, *Glob. Chang. Biol.*, 12, 177–193, <https://doi.org/10.1111/j.1365-2486.2005.001058.x>, 2006.
18. Bradford, M. A., Wieder, W. R., Bonan, G. B., Fierer, N., Raymond, P. A., and Crowther, T. W.: Managing uncertainty in soil carbon feedbacks to climate change, *Nat. Clim. Chang.*, 6, 751–758, <https://doi.org/10.1038/nclimate3071>, 2016.
19. Burns, R. G., DeForest, J. L., Marxsen, J., Sinsabaugh, R. L., Stromberger, M. E., Wallenstein, M. D., Weintraub, M. N., and Zoppini, A.: Soil enzymes in a changing environment: Current knowledge and future directions, *Soil Biol. Biochem.*, 58, 216–234, <https://doi.org/10.1016/j.soilbio.2012.11.009>, 2013.
20. Calderón, F. J., Nielsen, D., Acosta-Martínez, V., Vigil, M. F., and Lyon, D.: Cover Crop and Irrigation Effects on Soil Microbial Communities and Enzymes in Semiarid Agroecosystems of the Central Great Plains of North America, *Pedosphere*, 26, 192–205, [https://doi.org/10.1016/S1002-0160\(15\)60034-0](https://doi.org/10.1016/S1002-0160(15)60034-0), 2016.
21. Canarini, A., Carrillo, Y., Mariotte, P., Ingram, L., and Dijkstra, F. A.: Soil microbial community resistance to drought and links to C stabilization in an Australian grassland, *Soil Biol. Biochem.*, 103, 171–180, <https://doi.org/10.1016/j.soilbio.2016.08.024>, 2016.
22. Cano, A., Núñez, A., Acosta-Martínez, V., Schipanski, M., Ghimire, R., Rice, C., and West, C.: Current knowledge and future research directions to link soil health and water conservation in the Ogallala Aquifer region, *Geoderma*, 328, 109–118, <https://doi.org/10.1016/j.geoderma.2018.04.027>, 2018.

23. Carbone, M. S., Still, C. J., Ambrose, A. R., Dawson, T. E., Williams, A. P., Boot, C. M., Schaeffer, S. M., and Schimel, J. P.: Seasonal and episodic moisture controls on plant and microbial contributions to soil respiration, *Oecologia*, 167, 265–278, <https://doi.org/10.1007/s00442-011-1975-3>, 2011.
24. Churchland, C., Grayston, S. J., and Bengtson, P.: Spatial variability of soil fungal and bacterial abundance: Consequences for carbon turnover along a transition from a forested to clear-cut site, *Soil Biol. Biochem.*, 63, 5–13, <https://doi.org/10.1016/j.soilbio.2013.03.015>, 2013.
25. Cleveland, C. C. and Yavitt, J. B.: Consumption of atmospheric isoprene in soil, *Geophys. Res. Lett.*, 24, 2379–2382, <https://doi.org/10.1029/97GL02451>, 1997.
26. Conant, R. T., Dalla-Betta, P., Klopatek, C. C., and Klopatek, J. M.: Controls on soil respiration in semiarid soils, *Soil Biol. Biochem.*, 36, 945–951, <https://doi.org/10.1016/j.soilbio.2004.02.013>, 2004.
27. Cook, F. J. and Orchard, V. A.: Relationships between soil respiration and soil moisture, *Soil Biol. Biochem.*, 40, 1013–1018, <https://doi.org/10.1016/j.soilbio.2007.12.012>, 2008.
28. Curiel Yuste, J., Janssens, I. A., Carrara, A., and Ceulemans, R.: Annual Q10 of soil respiration reflects plant phenological patterns as well as temperature sensitivity, *Glob. Chang. Biol.*, 10, 161–169, <https://doi.org/10.1111/j.1529-8817.2003.00727.x>, 2004.
29. Dannenmann, M., Simon, J., Gasche, R., Holst, J., Naumann, P. S., Kögel-Knabner, I., Knicker, H., Mayer, H., Schloter, M., Pena, R., Polle, A., Rennenberg, H., and Papen, H.: Tree girdling provides insight on the role of labile carbon in nitrogen partitioning between soil microorganisms and adult European beech, *Soil Biol. Biochem.*, 41, 1622–1631, <https://doi.org/10.1016/j.soilbio.2009.04.024>, 2009.
30. Davidson, E. A., Belk, E., and Boone, R. D.: Soil water content and temperature as independent or confounded factors controlling soil respiration in a temperate mixed hardwood forest, *Glob. Chang. Biol.*, 4, 217–227, <https://doi.org/10.1046/j.1365-2486.1998.00128.x>, 1998.
31. Davis, K. F., Rulli, M. C., Garrassino, F., Chiarelli, D., Seveso, A., and D’Odorico, P.: Water limits to closing yield gaps, *Adv. Water Resour.*, 99, 67–75, <https://doi.org/10.1016/j.advwatres.2016.11.015>, 2017.
32. de Dato, G., Lagomarsino, A., Lellei-Kovacs, E., Liberati, D., Abou Jaoudé, R., Marabottini, R., Stazi, S. R., Guidolotti, G., Kovacs-Lang, E., Kroel-Dulay, G., and De Angelis, P.: The response of soil CO<sub>2</sub> efflux to water limitation is not merely a climatic issue: The role of substrate availability, *Forests*, 8, 1–22, <https://doi.org/10.3390/f8070241>, 2017.
33. Ding, J., Chen, L., Zhang, B., Liu, L., Yang, G., Fang, K., Chen, Y., Li, F., Kou, D., Ji, C., Luo, Y., and Yang, Y.: Linking temperature sensitivity of soil CO<sub>2</sub> release to substrate,

- environmental, and microbial properties across alpine ecosystems, *Global Biogeochem. Cycles*, 30, 1310–1323, <https://doi.org/10.1002/2015GB005333>, 2016.
34. Evans, S. E. and Wallenstein, M. D.: Soil microbial community response to drying and rewetting stress: does historical precipitation regime matter?, *Biogeochemistry*, 109, 101–116, <https://doi.org/10.1007/s10533-011-9638-3>, 2012.
  35. Falloon, P., Jones, C. D., Ades, M., and Paul, K.: Direct soil moisture controls of future global soil carbon changes: An important source of uncertainty, *Global Biogeochem. Cycles*, 25, 1–14, <https://doi.org/10.1029/2010GB003938>, 2011.
  36. Fang, C., Smith, P., Moncrieff, J. B., and Smith, J. U.: Similar response of labile and resistant soil organic matter pools to changes in temperature, *Nature*, 433, 57–59, <https://doi.org/10.1038/nature03138>, 2005.
  37. FAO: The State of Food and Agriculture, 2011. Fierer, N., Wood, S. A., and Bueno de Mesquita, C. P.: How microbes can, and cannot, be used to assess soil health, *Soil Biol. Biochem.*, 153, 108111, <https://doi.org/10.1016/j.soilbio.2020.108111>, 2021.
  38. Finzi, A. C., Abramoff, R. Z., Spiller, K. S., Brzostek, E. R., Darby, B. A., Kramer, M. A., and Phillips, R. P.: Rhizosphere processes are quantitatively important components of terrestrial carbon and nutrient cycles, *Glob. Chang. Biol.*, 21, 2082–2094, <https://doi.org/10.1111/gcb.12816>, 2015.
  39. Gao, W., Reed, S. C., Munson, S. M., Rui, Y., Fan, W., Zheng, Z., Li, L., Che, R., Xue, K., Du, J., Cui, X., Wang, Y., and Hao, Y.: Responses of soil extracellular enzyme activities and bacterial community composition to seasonal stages of drought in a semiarid grassland, *Geoderma*, 401, 115327, <https://doi.org/10.1016/j.geoderma.2021.115327>, 2021.
  40. German, D. P., Weintraub, M. N., Grandy, A. S., Lauber, C. L., Rinkes, Z. L., and Allison, S. D.: Optimization of hydrolytic and oxidative enzyme methods for ecosystem studies, *Soil Biol. Biochem.*, 43, 1387–1397, <https://doi.org/10.1016/j.soilbio.2011.03.017>, 2011.
  41. Göransson, H., Godbold, D. L., Jones, D. L., and Rousk, J.: Bacterial growth and respiration responses upon rewetting dry forest soils: Impact of drought-legacy, *Soil Biol. Biochem.*, 57, 477–486, <https://doi.org/10.1016/j.soilbio.2012.08.031>, 2013.
  42. Grace, J. B.: *Structural Equation Modeling and Natural Systems*, Cambridge University Press, <https://doi.org/10.1017/CBO9780511617799>, 2006.
  43. Hanson, P. J., Edwards, N. T., Garten, C. T., and Andrews, J. A.: Separating Root and Soil Microbial Contributions to Soil Respiration: A Review of Methods and Observations, *Biogeochemistry*, 48, 115–146, 2000.
  44. Hopkins, F., Gonzalez-Meler, M. A., Flower, C. E., Lynch, D. J., Czimczik, C., Tang, J., and Subke, J.: Ecosystem-level controls on root-rhizosphere respiration, *New Phytol.*, 199, 339–

- 351, <https://doi.org/10.1111/nph.12271>, 2013.
45. Hueso, S., Hernández, T., and García, C.: Resistance and resilience of the soil microbial biomass to severe drought in semiarid soils: The importance of organic amendments, *Appl. Soil Ecol.*, 50, 27–36, <https://doi.org/10.1016/j.apsoil.2011.07.014>, 2011.
  46. Joergensen, R. G. and Wichern, F.: Alive and kicking: Why dormant soil microorganisms matter, *Soil Biol. Biochem.*, 116, 419–430, <https://doi.org/10.1016/j.soilbio.2017.10.022>, 2018.
  47. Jones, S. E. and Lennon, J. T.: Dormancy contributes to the maintenance of microbial diversity, *Proc. Natl. Acad. Sci. U. S. A.*, 107, 5881–5886, <https://doi.org/10.1073/pnas.0912765107>, 2010.
  48. Kakumanu, M. L., Cantrell, C. L., and Williams, M. A.: Microbial community response to varying magnitudes of desiccation in soil: A test of the osmolyte accumulation hypothesis, *Soil Biol. Biochem.*, 57, 644–653, <https://doi.org/10.1016/j.soilbio.2012.08.014>, 2013.
  49. Kallenbach, C. M., Frey, S. D., and Grandy, A. S.: Direct evidence for microbial-derived soil organic matter formation and its ecophysiological controls, *Nat. Commun.*, 7, <https://doi.org/10.1038/ncomms13630>, 2016.
  50. Kane, D. A., Bradford, M. A., Fuller, E., Oldfield, E. E., and Wood, S. A.: Soil organic matter protects US maize yields and lowers crop insurance payouts under drought, *Environ. Res. Lett.*, 16, <https://doi.org/10.1088/1748-9326/abe492>, 2021.
  51. Kardol, P., Cregger, M. A., Campanv, C. E., and Classen, A. T.: Soil ecosystem functioning under climate change: Plant species and community effects, *Ecology*, 91, 767–781, <https://doi.org/10.1890/09-0135.1>, 2010.
  52. Köchy, M., Hiederer, R., and Freibauer, A.: Global distribution of soil organic carbon – Part 1: Masses and frequency distributions of SOC stocks for the tropics, permafrost regions, wetlands, and the world, *Soil*, 1, 351–365, <https://doi.org/10.5194/soil-1-351-2015>, 2015.
  53. Kreyling, J., Beierkuhnlein, C., Elmer, M., Pritsch, K., Radovski, M., Schloter, M., Wöllecke, J., and Jentsch, A.: Soil biotic processes remain remarkably stable after 100-year extreme weather events in experimental grassland and heath, *Plant Soil*, 308, 175–188, <https://doi.org/10.1007/s11104-008-9617-1>, 2008.
  54. Kuzyakov, Y.: Sources of CO<sub>2</sub> efflux from soil and review of partitioning methods, *Soil Biol. Biochem.*, 38, 425–448, <https://doi.org/10.1016/j.soilbio.2005.08.020>, 2006.
  55. Landesman, W. J. and Dighton, J.: Response of soil microbial communities and the production of plant-available nitrogen to a two-year rainfall manipulation in the New Jersey Pinelands, *Soil Biol. Biochem.*, 42, 1751–1758,

<https://doi.org/10.1016/j.soilbio.2010.06.012>, 2010.

56. Lehmann, J., Bossio, D. A., Kögel-Knabner, I., and Rillig, M. C.: The concept and future prospects of soil health, *Nat. Rev. Earth Environ.*, 1, 544–553, <https://doi.org/10.1038/s43017-020-0080-8>, 2020.
57. Li, C., Cano, A., Acosta-Martinez, V., Veum, K. S., and Moore-Kucera, J.: A comparison between fatty acid methyl ester profiling methods (PLFA and EL-FAME) as soil health indicators, *Soil Sci. Soc. Am. J.*, 84, 1153–1169, <https://doi.org/10.1002/saj2.20118>, 2020.
58. Li, L., Zheng, Z., Biederman, J. A., Xu, C., Xu, Z., Che, R., Wang, Y., Cui, X., and Hao, Y.: Ecological responses to heavy rainfall depend on seasonal timing and multi-year recurrence, *New Phytol.*, 223, 647–660, <https://doi.org/10.1111/nph.15832>, 2019.
59. Li, X., Xie, J., Zhang, Q., Lyu, M., Xiong, X., Liu, X., Lin, T., and Yang, Y.: Substrate availability and soil microbes drive temperature sensitivity of soil organic carbon mineralization to warming along an elevation gradient in subtropical Asia, *Geoderma*, 364, <https://doi.org/10.1016/j.geoderma.2020.114198>, 2020.
60. Liu, H. S., Li, L. H., Han, X. G., Huang, J. H., Sun, J. X., and Wang, H. Y.: Respiratory substrate availability plays a crucial role in the response of soil respiration to environmental factors, *Appl. Soil Ecol.*, 32, 284–292, <https://doi.org/10.1016/j.apsoil.2005.08.001>, 2006.
61. Lloyd J., Taylor, J. A.: On the Temperature Dependence of Soil Respiration, *British Ecological Society Stable*, <http://www.jstor.org/stable/2389824>, *Funct. Ecol.*, 8, 315–323, 1994.
62. Ma, Z., Zhao, W., Liu, M., and Liu, Q.: Responses of soil respiration and its components to experimental warming in an alpine scrub ecosystem on the eastern Qinghai-Tibet Plateau, *Sci. Total Environ.*, 643, 1427–1435, <https://doi.org/10.1016/j.scitotenv.2018.06.243>, 2018.
63. Maestre, F. T., Delgado-Baquerizo, M., Jeffries, T. C., Eldridge, D. J., Ochoa, V., Gozalo, B., Quero, J. L., García-Gómez, M., Gallardo, A., Ulrich, W., Bowker, M. A., Arredondo, T., Barraza-Zepeda, C., Bran, D., Florentino, A., Gaitán, J., Gutiérrez, J. R., Huber-Sannwald, E., Jankju, M., Mau, R. L., Miriti, M., Naseri, K., Ospina, A., Stavi, I., Wang, D., Woods, N. N., Yuan, X., Zaady, E., and Singh, B. K.: Increasing aridity reduces soil microbial diversity and abundance in global drylands, *Proc. Natl. Acad. Sci. U. S. A.*, 112, 15684–15689, <https://doi.org/10.1073/pnas.1516684112>, 2015.
64. Manzoni, S., Schaeffer, S. M., Katul, G., Porporato, A., and Schimel, J. P.: A theoretical analysis of microbial eco-physiological and diffusion limitations to carbon cycling in drying soils, *Soil Biol. Biochem.*, 73, 69–83, <https://doi.org/10.1016/j.soilbio.2014.02.008>, 2014.
65. Manzoni, S., Schimel, J. P., and Porporato, A.: Responses of soil microbial communities to water stress: Results from a meta-analysis, *Ecology*, 93, 930–938, <https://doi.org/10.1890/11-0026.1>, 2012.

66. McGuire, V. L.: Water-level changes in the high plains aquifer, change in water in storage, U.S. Dep. Inter. U.S. Geol. Surv., 13, 2011. Miao, S., Carstenn, S., and Nungesser, M.: Real World Ecology, 308 pp., 2009. Moldrup, P., Olesen, T., Komatsu, T., Schjønning, P., and Rolston, D. E.: Tortuosity, Diffusivity, and Permeability in the Soil Liquid and Gaseous Phases, *Soil Sci. Soc. Am. J.*, 65, 613–623, <https://doi.org/10.2136/sssaj2001.653613x>, 2001.
67. Moorhead, D. L., Lashermes, G., and Sinsabaugh, R. L.: A theoretical model of C- and N-acquiring exoenzyme activities, which balances microbial demands during decomposition, *Soil Biol. Biochem.*, 53, 133–141, <https://doi.org/10.1016/j.soilbio.2012.05.011>, 2012.
68. Moyano, F. E., Manzoni, S., and Chenu, C.: Responses of soil heterotrophic respiration to moisture availability: An exploration of processes and models, *Soil Biol. Biochem.*, 59, 72–85, <https://doi.org/10.1016/j.soilbio.2013.01.002>, 2013.
69. Niboyet, A., Bardoux, G., Barot, S., and Bloor, J. M. G.: Elevated CO<sub>2</sub> mediates the short-term drought recovery of ecosystem function in low-diversity grassland systems, *Plant Soil*, 420, 289–302, <https://doi.org/10.1007/s11104-017-3377-8>, 2017.
70. Núñez, A.: When the wells run dry: soil organic carbon dynamics during the transition from irrigated to dryland cropping systems, Colorado State University, 2021.
71. Núñez, A., Ball, R., and Schipanski, M.: Plant and soil microbial responses to irrigation retirement in semiarid cropping systems, *Environ. Res. Commun.*, 4, <https://doi.org/10.1088/2515-7620/ac59c3>, 2022.
72. Olsson, P. A., Bååth, E., Jakobsen, I., and Söderström, B.: The use of phospholipid and neutral lipid fatty acids to estimate biomass of arbuscular mycorrhizal fungi in soil, *Mycol. Res.*, 99, 623–629, [https://doi.org/10.1016/S0953-7562\(09\)80723-5](https://doi.org/10.1016/S0953-7562(09)80723-5), 1995.
73. Orchard, V. A. and Cook, F. J.: Relationship between soil respiration and soil moisture, *Soil Biol. Biochem.*, 15, 447–453, [https://doi.org/10.1016/0038-0717\(83\)90010-X](https://doi.org/10.1016/0038-0717(83)90010-X), 1983.
74. Pallardy, S. G. and Rhoads, J. L.: Morphological adaptations to drought in seedlings of deciduous angiosperms, *Can. J. For. Res.*, 23, 1766–1774, <https://doi.org/10.1139/x93-223>, 1993.
75. Parker, S. S. and Schimel, J. P.: Soil nitrogen availability and transformations differ between the summer and the growing season in a California grassland, *Appl. Soil Ecol.*, 48, 185–192, <https://doi.org/10.1016/j.apsoil.2011.03.007>, 2011.
76. Peng, Y., Xu, H., Wang, Z., Shi, J., Lv, J., and Wang, X.: Responses of the content and spectral characteristics of dissolved organic matter in intercropping soil to drought in northeast China, *Plant Soil*, <https://doi.org/10.1007/s11104-023-05931-w>, 2023.
77. Pérez-Guzmán, L., Acosta-Martínez, V., Phillips, L. A., and Mauget, S. A.: Resilience of the microbial communities of semiarid agricultural soils during natural climatic variability

- events, *Appl. Soil Ecol.*, 149, 103487, <https://doi.org/10.1016/j.apsoil.2019.103487>, 2020.
78. Pohlen, E., Fandino, A. O., and Marxsen, J.: Bacterial community composition and extracellular enzyme activity in temperate streambed sediment during drying and rewetting, *PLoS One*, 8, <https://doi.org/10.1371/journal.pone.0083365>, 2013.
79. Porporato, A., Daly, E., and Rodriguez-Iturbe, I.: Soil water balance and ecosystem response to climate change, *Am. Nat.*, 164, 625–632, <https://doi.org/10.1086/424970>, 2004.
80. Rawls, W. J., Pachepsky, Y. A., Ritchie, J. C., Sobecki, T. M., and Bloodworth, H.: Effect of soil organic carbon on soil water retention, *Geoderma*, 116, 61–76, [https://doi.org/10.1016/S0016-7061\(03\)00094-6](https://doi.org/10.1016/S0016-7061(03)00094-6), 2003.
81. Reichstein, M., Rey, A., Freibauer, A., Tenhunen, J., Valentini, R., Banza, J., Casals, P., Cheng, Y., Grünzweig, J. M., Irvine, J., Joffre, R., Law, B. E., Loustau, D., Miglietta, F., Oechel, W., Ourcival, J. M., Pereira, J. S., Peressotti, A., Ponti, F., Qi, Y., Rambal, S., Rayment, M., Romanya, J., Rossi, F., Tedeschi, V., Tirone, G., Xu, M., and Yakir, D.: Modeling temporal and large-scale spatial variability of soil respiration from soil water availability, temperature and vegetation productivity indices, *Global Biogeochem. Cycles*, 17, <https://doi.org/10.1029/2003gb002035>, 2003.
82. Reichstein, M., Tenhunen, J. D., Roupsard, O., Ourcival, J. M., Rambal, S., Miglietta, F., Peressotti, A., Pecchiari, M., Tirone, G., and Valentini, R.: Severe drought effects on ecosystem CO<sub>2</sub> and H<sub>2</sub>O fluxes at three Mediterranean evergreen sites: Revision of current hypotheses, *Glob. Chang. Biol.*, 8, 999–1017, <https://doi.org/10.1046/j.1365-2486.2002.00530.x>, 2002.
83. Ren, C., Zhao, F., Shi, Z., Chen, J., Han, X., Yang, G., Feng, Y., and Ren, G.: Differential responses of soil microbial biomass and carbon-degrading enzyme activities to altered precipitation, *Soil Biol. Biochem.*, 115, 1–10, <https://doi.org/10.1016/j.soilbio.2017.08.002>, 2017.
84. Rudnick, D. R., Irmak, S., West, C., Chávez, J. L., Kisekka, I., Marek, T. H., Schneekloth, J. P., Mitchell McCallister, D., Sharma, V., Djaman, K., Aguilar, J., Schipanski, M. E., Rogers, D. H., and Schlegel, A.: Deficit Irrigation Management of Maize in the High Plains Aquifer Region: A Review, *J. Am. Water Resour. Assoc.*, 55, 38–55, <https://doi.org/10.1111/1752-1688.12723>, 2019.
85. Saiya-Cork, K. R., Sinsabaugh, R. L., and Zak, D. R.: The effects of long-term nitrogen deposition on extracellular enzyme activity in an *Acer saccharum* forest soil, *Soil Biol. Biochem.*, 34, 1309–1315, [https://doi.org/10.1016/S0038-0717\(02\)00074-3](https://doi.org/10.1016/S0038-0717(02)00074-3), 2002.
86. Sardans, J., Peñuelas, J., and Estiarte, M.: Changes in soil enzymes related to C and N cycle and in soil C and N content under prolonged warming and drought in a Mediterranean shrubland, *Appl. Soil Ecol.*, 39, 223–235, <https://doi.org/10.1016/j.apsoil.2007.12.011>, 2008.

87. Savage, K., Davidson, E. A., and Tang, J.: Diel patterns of autotrophic and heterotrophic respiration among phenological stages, *Glob. Chang. Biol.*, 19, 1151–1159, <https://doi.org/10.1111/gcb.12108>, 2013.
88. Schaeffer, S. M., Homyak, P. M., Boot, C. M., Roux-Michollet, D., and Schimel, J. P.: Soil carbon and nitrogen dynamics throughout the summer drought in a California annual grassland, *Soil Biol. Biochem.*, 115, 54–62, <https://doi.org/10.1016/j.soilbio.2017.08.009>, 2017.
89. Scharlemann, J. P. W., Tanner, E. V. J., Hiederer, R., and Kapos, V.: Global soil carbon: Understanding and managing the largest terrestrial carbon pool, *Carbon Manag.*, 5, 81–91, <https://doi.org/10.4155/cmt.13.77>, 2014.
90. Schimel, J. P.: Life in dry soils: Effects of drought on soil microbial communities and processes, *Annu. Rev. Ecol. Evol. Syst.*, 49, 409–432, <https://doi.org/10.1146/annurev-eolsys-110617-062614>, 2018.
91. Schimel, J. P., Gullledge, J. M., Clein-Curley, J. S., Lindstrom, J. E., and Braddock, J. F.: Moisture effects on microbial activity and community structure in decomposing birch litter in the Alaskan taiga, *Soil Biol. Biochem.*, 31, 831–838, [https://doi.org/10.1016/S0038-0717\(98\)00182-5](https://doi.org/10.1016/S0038-0717(98)00182-5), 1999.
92. Schimel, J. P. and Schaeffer, S. M.: Microbial control over carbon cycling in soil, *Front. Microbiol.*, 3, 1–11, <https://doi.org/10.3389/fmicb.2012.00348>, 2012.
93. Schimel, J. P. and Weintraub, M. N.: The implications of exoenzyme activity on microbial carbon and nitrogen limitation in soil: A theoretical model, *Soil Biol. Biochem.*, 35, 549–563, [https://doi.org/10.1016/S0038-0717\(03\)00015-4](https://doi.org/10.1016/S0038-0717(03)00015-4), 2003.
94. Schipper, L. A., Petrie, O. J., O’Neill, T. A., Mudge, P. L., Liáng, L. L., Robinson, J. M., and Arcus, V. L.: Shifts in temperature response of soil respiration between adjacent irrigated and non-irrigated grazed pastures, *Agric. Ecosyst. Environ.*, 285, 106620, <https://doi.org/10.1016/j.agee.2019.106620>, 2019.
95. Schreckinger, J., Mutz, M., Mendoza-Lera, C., and Frossard, A.: Attributes of Drying Define the Structure and Functioning of Microbial Communities in Temperate Riverbed Sediment, *Front. Microbiol.*, 12, 1–15, <https://doi.org/10.3389/fmicb.2021.676615>, 2021.
96. Schutter, M. E. and Dick, R. P.: Comparison of Fatty Acid Methyl Ester (FAME) Methods for Characterizing Microbial Communities, *Soil Sci. Soc. Am. J.*, 64, 1659–1668, <https://doi.org/10.2136/sssaj2000.6451659x>, 2000.
97. Sharma, M. P. and Buyer, J. S.: Comparison of biochemical and microscopic methods for quantification of arbuscular mycorrhizal fungi in soil and roots, *Appl. Soil Ecol.*, 95, 86–89, <https://doi.org/10.1016/j.apsoil.2015.06.001>, 2015.

98. Sherrod, L. A., Ahuja, L. R., Hansen, N. C., Ascough, J. C., Westfall, D. G., and Peterson, G. A.: Soil and rainfall factors influencing yields of a dryland cropping system in Colorado, *Agron. J.*, 106, 1179–1192, <https://doi.org/10.2134/agronj13.0520>, 2014.
99. Sherwood, S. and Fu, Q.: A drier future?, *Science* (80-. ), 343, 737–739, <https://doi.org/10.1126/science.1247620>, 2014. Shipley, B.: *Cause and Correlation in Biology: A User's Guide to Path Analysis, Structural Equations and Causal Inference with R*, 2nd ed., Cambridge University Press, <https://doi.org/10.1017/CBO9781139979573>, 2016.
100. Siebielec, S., Siebielec, G., Klimkiewicz-Pawlas, A., Gałazka, A., Grzadziel, J., and Stuczynski, T.: Impact of water stress on microbial community and activity in sandy and loamy soils, *Agronomy*, 10, <https://doi.org/10.3390/agronomy10091429>, 2020.
101. Sinsabaugh, R. S.: Enzymic analysis of microbial pattern and process, *Biol. Fertil. Soils*, 17, 69–74, <https://doi.org/10.1007/BF00418675>, 1994.
102. Skopp, J., Jawson, M. D., and Doran, J. W.: Steady-State Aerobic Microbial Activity as a Function of Soil Water Content, *Soil Sci. Soc. Am. J.*, 54, 1619–1625, <https://doi.org/10.2136/sssaj1990.03615995005400060018x>, 1990.
103. Southworth, J., Randolph, J. C., Habeck, M., Doering, O. C., Pfeifer, R. A., Rao, D. G., and Johnston, J. J.: Consequences of future climate change and changing climate variability on maize yields in the midwestern United States, *Agric. Ecosyst. Environ.*, 82, 139–158, [https://doi.org/10.1016/S0167-8809\(00\)00223-1](https://doi.org/10.1016/S0167-8809(00)00223-1), 2000.
104. Sperlich, D., Barbeta, A., Ogaya, R., Sabaté, S., and Peñuelas, J.: Balance between carbon gain and loss under long-term drought: Impacts on foliar respiration and photosynthesis in *Quercus ilex* L., *J. Exp. Bot.*, 67, 821–833, <https://doi.org/10.1093/jxb/erv492>, 2016.
105. Subke, J. A., Inglima, I., and Cotrufo, M. F.: Trends and methodological impacts in soil CO<sub>2</sub> efflux partitioning: A metaanalytical review, *Glob. Chang. Biol.*, 12, 921–943, <https://doi.org/10.1111/j.1365-2486.2006.01117.x>, 2006.
106. Suseela, V., Conant, R. T., Wallenstein, M. D., and Dukes, J. S.: Effects of soil moisture on the temperature sensitivity of heterotrophic respiration vary seasonally in an old-field climate change experiment, *Glob. Chang. Biol.*, 18, 336–348, <https://doi.org/10.1111/j.1365-2486.2011.02516.x>, 2012.
107. Tang, J. and Baldocchi, D. D.: Spatial-temporal variation in soil respiration in an oak-grass savanna ecosystem in California and its partitioning into autotrophic and heterotrophic components, *Biogeochemistry*, 73, 183–207, <https://doi.org/10.1007/s10533-004-5889-6>, 2005.
108. Tecon, R. and Or, D.: Biophysical processes supporting the diversity of microbial life in soil, *FEMS Microbiol. Rev.*, 41, 599–623, <https://doi.org/10.1093/femsre/fux039>, 2017.

109. Trost, B., Prochnow, A., Drastig, K., Meyer-Aurich, A., Ellmer, F., and Baumecker, M.: Irrigation, soil organic carbon and N<sub>2</sub>O emissions. A review, *Agron. Sustain. Dev.*, 33, 733–749, <https://doi.org/10.1007/s13593-013-0134-0>, 2013.
110. Vetter, Y. A., Deming, J. W., Jumars, P. A., and Krieger-Brockett, B. B.: A predictive model of bacterial foraging by means of freely released extracellular enzymes, *Microb. Ecol.*, 36, 75–92, <https://doi.org/10.1007/s002489900095>, 1998.
111. Wade, A. M., Richter, D. D., Cherkinsky, A., Craft, C. B., and Heine, P. R.: Limited carbon contents of centuries old soils forming in legacy sediment, *Geomorphology*, 354, 107018, <https://doi.org/10.1016/j.geomorph.2019.107018>, 2020.
112. Weintraub, M. N., Scott-Denton, L. E., Schmidt, S. K., and Monson, R. K.: The effects of tree rhizodeposition on soil exoenzyme activity, dissolved organic carbon, and nutrient availability in a subalpine forest ecosystem, *Oecologia*, 154, 327–338, <https://doi.org/10.1007/s00442-007-0804-1>, 2007.
113. Wheeler, E., Golden, B., Johnson, J., and Peterson, J.: Economic Efficiency of Short-Term Versus Long-Term Water Rights Buyouts, *J. Agric. Appl. Econ.*, 40, 493–501, <https://doi.org/10.1017/S1074070800023786>, 2008.
114. Whittemore, D. O., Butler, J. J., and Wilson, B. B.: Assessing the major drivers of water-level declines: new insights into the future of heavily stressed aquifers, *Hydrol. Sci. J.*, 61, 134–145, <https://doi.org/10.1080/02626667.2014.959958>, 2016.
115. Willers, C., Jansen van Rensburg, P. J., and Claassens, S.: Phospholipid fatty acid profiling of microbial communities-a review of interpretations and recent applications, *J. Appl. Microbiol.*, 119, 1207–1218, <https://doi.org/10.1111/jam.12902>, 2015.
116. Williams, M. A.: Response of microbial communities to water stress in irrigated and drought-prone tallgrass prairie soils, *Soil Biol. Biochem.*, 39, 2750–2757, <https://doi.org/10.1016/j.soilbio.2007.05.025>, 2007. Williamson, W. M. and Wardle, D. A.: The soil microbial community response when plants are subjected to water stress and defoliation disturbance, *Appl. Soil Ecol.*, 37, 139–149, <https://doi.org/10.1016/j.apsoil.2007.05.003>, 2007.
117. Wu, X., Fang, H., Zhao, L., Wu, T., Li, R., Ren, Z., Pang, Q., and Ding, Y.: Mineralisation and Changes in the Fractions of Soil Organic Matter in Soils of the Permafrost Region, Qinghai-Tibet Plateau, China, *Permafr. Periglac. Process.*, 25, 35–44, <https://doi.org/10.1002/ppp.1796>, 2014.
118. Wu, Z., Dijkstra, P., Koch, G. W., Peñuelas, J., and Hungate, B. A.: Responses of terrestrial ecosystems to temperature and precipitation change: A meta-analysis of experimental manipulation, *Glob. Chang. Biol.*, 17, 927–942, <https://doi.org/10.1111/j.1365-2486.2010.02302.x>, 2011.

119. Xu, M. and Shang, H.: Contribution of soil respiration to the global carbon equation, *J. Plant Physiol.*, 203, 16–28, <https://doi.org/10.1016/j.jplph.2016.08.007>, 2016.
120. Zeglin, L. H., Bottomley, P. J., Jumpponen, A., Rice, C. W., Arango, M., Lindsley, A., McGowan, A., Mfombep, P., and Myrold, D. D.: Altered precipitation regime affects the function and composition of soil microbial communities on multiple time scales, *Ecology*, 94, 2334–2345, <https://doi.org/10.1890/12-2018.1>, 2013.
121. Zelles, L.: Fatty acid patterns of phospholipids and lipopolysaccharides in the characterisation of microbial communities in soil: A review, *Biol. Fertil. Soils*, 29, 111–129, <https://doi.org/10.1007/s003740050533>, 1999.
122. Zhang, K., Ali, A., Antonarakis, A., Moghaddam, M., Saatchi, S., Tabatabaenejad, A., Chen, R., Jaruwatanadilok, S., Cuenca, R., Crow, W. T., and Moorcroft, P.: The Sensitivity of North American Terrestrial Carbon Fluxes to Spatial and Temporal Variation in Soil Moisture: An Analysis Using Radar-Derived Estimates of Root-Zone Soil Moisture, *J. Geophys. Res. Biogeosciences*, 124, 3208–3231, <https://doi.org/10.1029/2018JG004589>, 2019.
123. Zhang, Z. S., Li, X. R., Nowak, R. S., Wu, P., Gao, Y. H., Zhao, Y., Huang, L., Hu, Y. G., and Jia, R. L.: Effect of sand-stabilizing shrubs on soil respiration in a temperate desert, *Plant Soil*, 367, 449–463, <https://doi.org/10.1007/s11104-012-1465-3>, 2013.
124. Zhou, L., Liu, Y., Zhang, Y., Sha, L., Song, Q., Zhou, W., Balasubramanian, D., Palingamoorthy, G., Gao, J., Lin, Y., Li, J., Zhou, R., Zar Myo, S. T., Tang, X., Zhang, J., Zhang, P., Wang, S., and Grace, J.: Soil respiration after six years of continuous drought stress in the tropical rainforest in Southwest China, *Soil Biol. Biochem.*, 138, 107564, <https://doi.org/10.1016/j.soilbio.2019.107564>, 2019.
125. Zornoza, R., Rosales, R. M., Acosta, J. A., de la Rosa, J. M., Arcenegui, V., Faz, Á., and Pérez-Pastor, A.: Efficient irrigation management can contribute to reduce soil CO<sub>2</sub> emissions in agriculture, *Geoderma*, 263, 70–77, <https://doi.org/10.1016/j.geoderma.2015.09.003>, 2016.



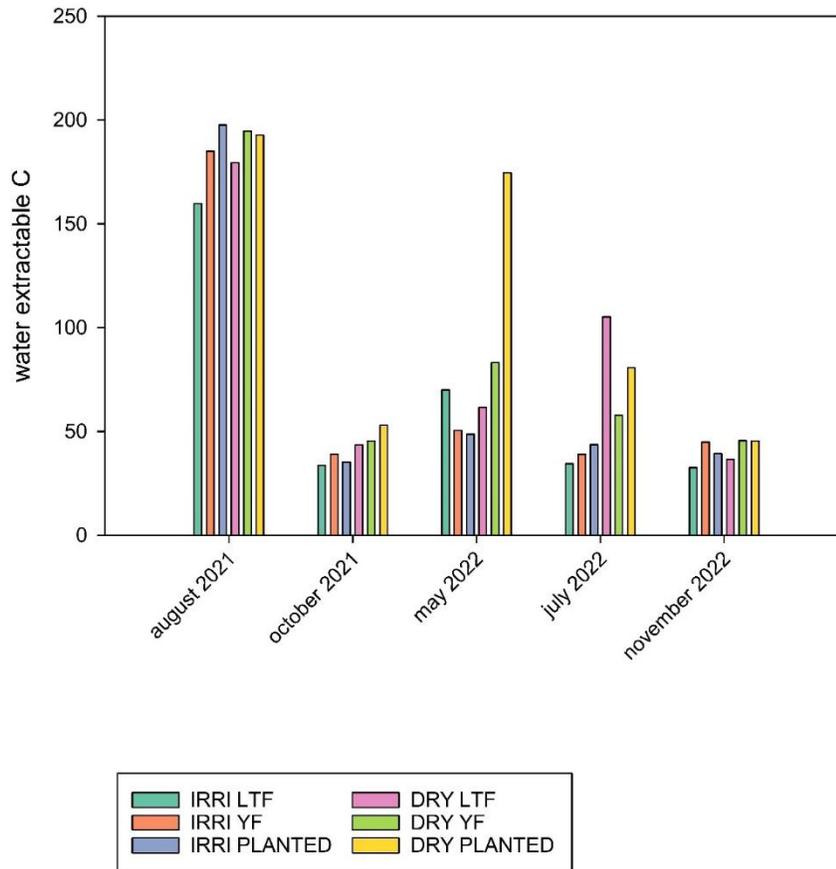
**Table S1.** Results of a permutational analysis of variance (PERMANOVA) between six unique irrigation and plant input treatment combinations across ten soil samplings (0-10 cm) in 2021 and 2022 of a continuous maize agroecosystem in Northern Colorado. *df*: irrigation= 1, plant input= 1, irri x plant= 1.

|                 | irrigation            |          |              | plant input           |          |              | irri x plant          |          |             |
|-----------------|-----------------------|----------|--------------|-----------------------|----------|--------------|-----------------------|----------|-------------|
|                 | <i>R</i> <sup>2</sup> | <i>F</i> | <i>p</i>     | <i>R</i> <sup>2</sup> | <i>F</i> | <i>p</i>     | <i>R</i> <sup>2</sup> | <i>F</i> | <i>p</i>    |
| GWC             | 0.38                  | 99.5     | <b>0.001</b> | 0.003                 | 0.88     | 0.37         | 0.023                 | 5.93     | <b>0.02</b> |
| DOC             | -                     | -        | -            | -                     | -        | -            | -                     | -        | -           |
| soil Rh         | 0.01                  | 1.90     | 0.155        | 0.008                 | 1.26     | 0.27         | 0.005                 | 0.85     | 0.38        |
| <b>FAME</b>     |                       |          |              |                       |          |              |                       |          |             |
| bacteria        | 0.72                  | 15.28    | <b>0.002</b> | 0.16                  | 35.07    | <b>0.001</b> | 0.03                  | 6.94     | <b>0.01</b> |
| fungi           | 0.13                  | 25.74    | <b>0.001</b> | 0.09                  | 18.11    | <b>0.001</b> | 0.02                  | 4.56     | <b>0.04</b> |
| protozoa        | -0.67                 | -340.15  | 0.93         | 0.68                  | 346.36   | 0.15         | 0.68                  | 347.99   | 0.15        |
| actino          | 0.006                 | 1.09     | 0.28         | 0.07                  | 12.61    | <b>0.001</b> | 0.02                  | 2.54     | 0.11        |
| <b>EEAs</b>     |                       |          |              |                       |          |              |                       |          |             |
| BGLUC           | -                     | -        | -            | -                     | -        | -            | -                     | -        | -           |
| NAG             | -                     | -        | -            | -                     | -        | -            | -                     | -        | -           |
| BCELL           | -                     | -        | -            | -                     | -        | -            | -                     | -        | -           |
| LAP             | -                     | -        | -            | -                     | -        | -            | -                     | -        | -           |
| <b>Together</b> | 0.20                  | 40.74    | <b>0.001</b> | 0.01                  | 2.81     | <u>0.05</u>  | 0.02                  | 3.37     | <b>0.03</b> |

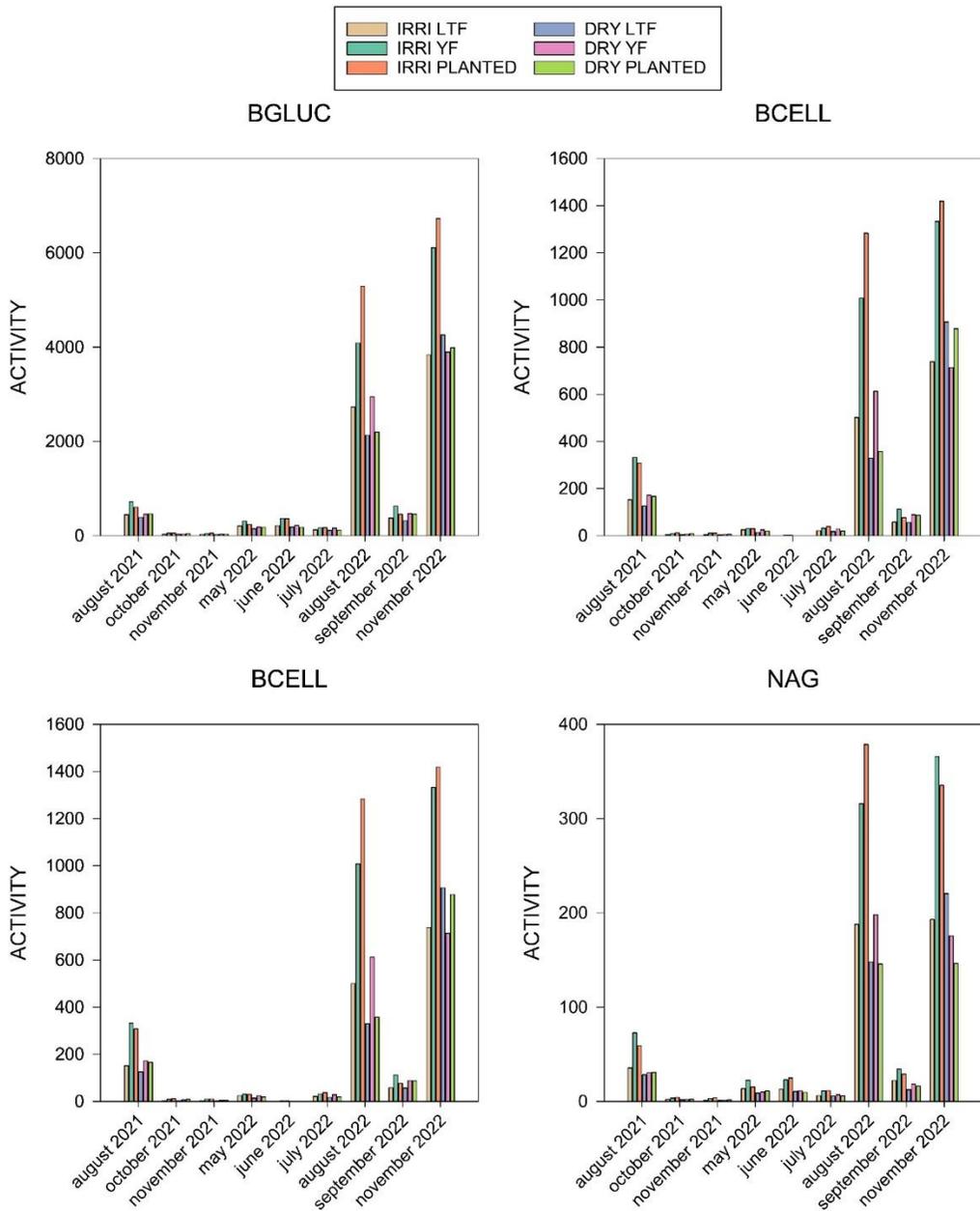
**Table S2.** Means comparisons between six unique irrigation and plant input treatment combinations across ten soil samplings (0-10 cm) in 2021 and 2022 of a continuous maize agroecosystem in Northern Colorado, across eleven categories. *df* for: gravimetric water content (GWC): 195; dissolved organic carbon (DOC): 94.7; Soil Rh: 117; fatty acid methyl ester biomarkers (FAME): 195; extra-cellular enzyme activities (EEAs): 175.

| Contrast                   | GWC            |                 | DOC            |                 | Soil Rh        |                 | Bacteria FAME  |                 | Fungi FAME     |                 | Protozoa FAME  |                 | Actino FAME    |                 |
|----------------------------|----------------|-----------------|----------------|-----------------|----------------|-----------------|----------------|-----------------|----------------|-----------------|----------------|-----------------|----------------|-----------------|
|                            | <i>t.ratio</i> | <i>p.value</i>  |
| IRRI LTF - IRRI YF         | -2.86          | 0.05            | -1.13          | 0.87            | -3.09          | <b>0.01</b>     | -8.42          | <b>&lt;0.01</b> | -6.22          | <b>&lt;0.01</b> | -7.97          | <b>&lt;0.01</b> | -5.53          | <b>&lt;0.01</b> |
| IRRI LTF - IRRI PLANTED    | -0.17          | 1.00            | -0.94          | 0.94            |                |                 | -11.82         | <b>&lt;0.01</b> | -10.63         | <b>&lt;0.01</b> | -11.79         | <b>&lt;0.01</b> | -7.46          | <b>&lt;0.01</b> |
| IRRI LTF - DRY LTF         | -5.75          | <b>&lt;0.01</b> | 2.34           | 0.19            | -0.56          | 0.94            | -1.16          | 0.85            | -2.66          | 0.09            | -0.51          | 1.00            | 2.91           | <b>0.05</b>     |
| IRRI LTF - DRY YF          | -7.65          | <b>&lt;0.01</b> | 3.29           | <b>0.02</b>     | -0.21          | 1.00            | 1.97           | 0.36            | -0.77          | 0.97            | -2.52          | 0.12            | 4.97           | <b>&lt;0.01</b> |
| IRRI LTF - DRY PLANTED     | -8.09          | <b>&lt;0.01</b> | 3.63           | <b>0.01</b>     |                |                 | 1.86           | 0.43            | -1.49          | 0.67            | -2.55          | 0.11            | 5.26           | <b>&lt;0.01</b> |
| IRRI YF - IRRI PLANTED     | -2.70          | 0.08            | -0.19          | 1.00            |                |                 | 3.40           | <b>0.01</b>     | 4.41           | <b>&lt;0.01</b> | 3.82           | <b>&lt;0.01</b> | 1.93           | 0.39            |
| IRRI YF - DRY LTF          | -8.62          | <b>&lt;0.01</b> | 1.21           | 0.83            | -3.66          | <b>&lt;0.01</b> | -9.59          | <b>&lt;0.01</b> | -8.87          | <b>&lt;0.01</b> | -8.47          | <b>&lt;0.01</b> | -2.62          | 0.10            |
| IRRI YF - DRY YF           | -10.52         | <b>&lt;0.01</b> | 2.17           | 0.26            | -3.31          | <b>0.01</b>     | -6.45          | <b>&lt;0.01</b> | -6.99          | <b>&lt;0.01</b> | -10.49         | <b>&lt;0.01</b> | -0.57          | 0.99            |
| IRRI YF - DRY PLANTED      | -10.95         | <b>&lt;0.01</b> | 2.52           | 0.13            |                |                 | -6.56          | <b>&lt;0.01</b> | -7.71          | <b>&lt;0.01</b> | -10.52         | <b>&lt;0.01</b> | -0.27          | 1.00            |
| IRRI PLANTED - DRY LTF     | -5.92          | <b>&lt;0.01</b> | 1.40           | 0.73            |                |                 | -12.98         | <b>&lt;0.01</b> | -13.28         | <b>&lt;0.01</b> | -12.29         | <b>&lt;0.01</b> | -4.55          | <b>&lt;0.01</b> |
| IRRI PLANTED - DRY YF      | -7.82          | <b>&lt;0.01</b> | 2.36           | 0.18            |                |                 | -9.85          | <b>&lt;0.01</b> | -11.39         | <b>&lt;0.01</b> | -14.31         | <b>&lt;0.01</b> | -2.49          | 0.13            |
| IRRI PLANTED - DRY PLANTED | -8.26          | <b>&lt;0.01</b> | 2.71           | 0.08            |                |                 | -9.96          | <b>&lt;0.01</b> | -12.12         | <b>&lt;0.01</b> | -14.34         | <b>&lt;0.01</b> | -2.20          | 0.24            |
| DRY LTF - DRY YF           | 1.90           | 0.41            | -0.95          | 0.93            | -0.35          | 0.99            | -3.13          | <b>0.02</b>     | -1.89          | 0.41            | 2.02           | 0.34            | -2.06          | 0.32            |
| DRY LTF - DRY PLANTED      | 2.34           | 0.18            | -1.33          | 0.77            |                |                 | -3.02          | <b>0.03</b>     | -1.16          | 0.85            | 2.04           | 0.32            | -2.35          | 0.18            |
| DRY YF - DRY PLANTED       | -0.44          | 1.00            | 0.39           | 1.00            |                |                 | -0.11          | 1.00            | -0.73          | 0.98            | -0.03          | 1.00            | 0.30           | 1.00            |
| Contrast                   | NAG            |                 | BGLUC          |                 | BCELL          |                 | LAP            |                 |                |                 |                |                 |                |                 |
|                            | <i>t.ratio</i> | <i>p.value</i>  | <i>t.ratio</i> | <i>p.value</i>  | <i>t.ratio</i> | <i>p.value</i>  | <i>t.ratio</i> | <i>p.value</i>  |                |                 |                |                 |                |                 |
| IRRI LTF - IRRI YF         | -9.68          | <b>&lt;0.01</b> | -9.56          | <b>&lt;0.01</b> | -7.35          | <b>&lt;0.01</b> | -1.54          | 0.64            |                |                 |                |                 |                |                 |
| IRRI LTF - IRRI PLANTED    | -9.25          | <b>&lt;0.01</b> | -8.82          | <b>&lt;0.01</b> | -7.53          | <b>&lt;0.01</b> | -2.05          | 0.32            |                |                 |                |                 |                |                 |
| IRRI LTF - DRY LTF         | -3.28          | <b>0.02</b>     | -2.24          | 0.22            | -2.00          | 0.35            | -1.26          | 0.81            |                |                 |                |                 |                |                 |
| IRRI LTF - DRY YF          | -0.36          | 1.00            | 2.16           | 0.26            | 1.86           | 0.43            | -1.00          | 0.92            |                |                 |                |                 |                |                 |
| IRRI LTF - DRY PLANTED     | -1.39          | 0.74            | 1.04           | 0.90            | 1.06           | 0.90            | -1.16          | 0.86            |                |                 |                |                 |                |                 |
| IRRI YF - IRRI PLANTED     | -0.43          | 1.00            | -0.74          | 0.98            | 0.18           | 1.00            | 0.51           | 1.00            |                |                 |                |                 |                |                 |
| IRRI YF - DRY LTF          | -12.95         | <b>&lt;0.01</b> | -11.80         | <b>&lt;0.01</b> | -9.35          | <b>&lt;0.01</b> | -2.80          | 0.06            |                |                 |                |                 |                |                 |
| IRRI YF - DRY YF           | -10.04         | <b>&lt;0.01</b> | -7.40          | <b>&lt;0.01</b> | -5.50          | <b>&lt;0.01</b> | -2.54          | 0.12            |                |                 |                |                 |                |                 |
| IRRI YF - DRY PLANTED      | -11.06         | <b>&lt;0.01</b> | -8.51          | <b>&lt;0.01</b> | -6.29          | <b>&lt;0.01</b> | -2.70          | 0.08            |                |                 |                |                 |                |                 |
| IRRI PLANTED - DRY LTF     | -12.52         | <b>&lt;0.01</b> | -11.06         | <b>&lt;0.01</b> | -9.53          | <b>&lt;0.01</b> | -3.31          | <b>0.01</b>     |                |                 |                |                 |                |                 |
| IRRI PLANTED - DRY YF      | -9.61          | <b>&lt;0.01</b> | -6.66          | <b>&lt;0.01</b> | -5.68          | <b>&lt;0.01</b> | -3.05          | <b>0.03</b>     |                |                 |                |                 |                |                 |
| IRRI PLANTED - DRY PLANTED | -10.63         | <b>&lt;0.01</b> | -7.78          | <b>&lt;0.01</b> | -6.47          | <b>&lt;0.01</b> | -3.21          | <b>0.02</b>     |                |                 |                |                 |                |                 |
| DRY LTF - DRY YF           | -2.92          | <b>0.05</b>     | -4.40          | <b>0.00</b>     | -3.85          | <b>&lt;0.01</b> | -0.26          | 1.00            |                |                 |                |                 |                |                 |
| DRY LTF - DRY PLANTED      | -1.89          | 0.41            | -3.29          | <b>0.02</b>     | -3.06          | <b>0.03</b>     | -0.10          | 1.00            |                |                 |                |                 |                |                 |
| DRY YF - DRY PLANTED       | -1.03          | 0.91            | -1.12          | 0.87            | -0.79          | 0.97            | -0.16          | 1.00            |                |                 |                |                 |                |                 |

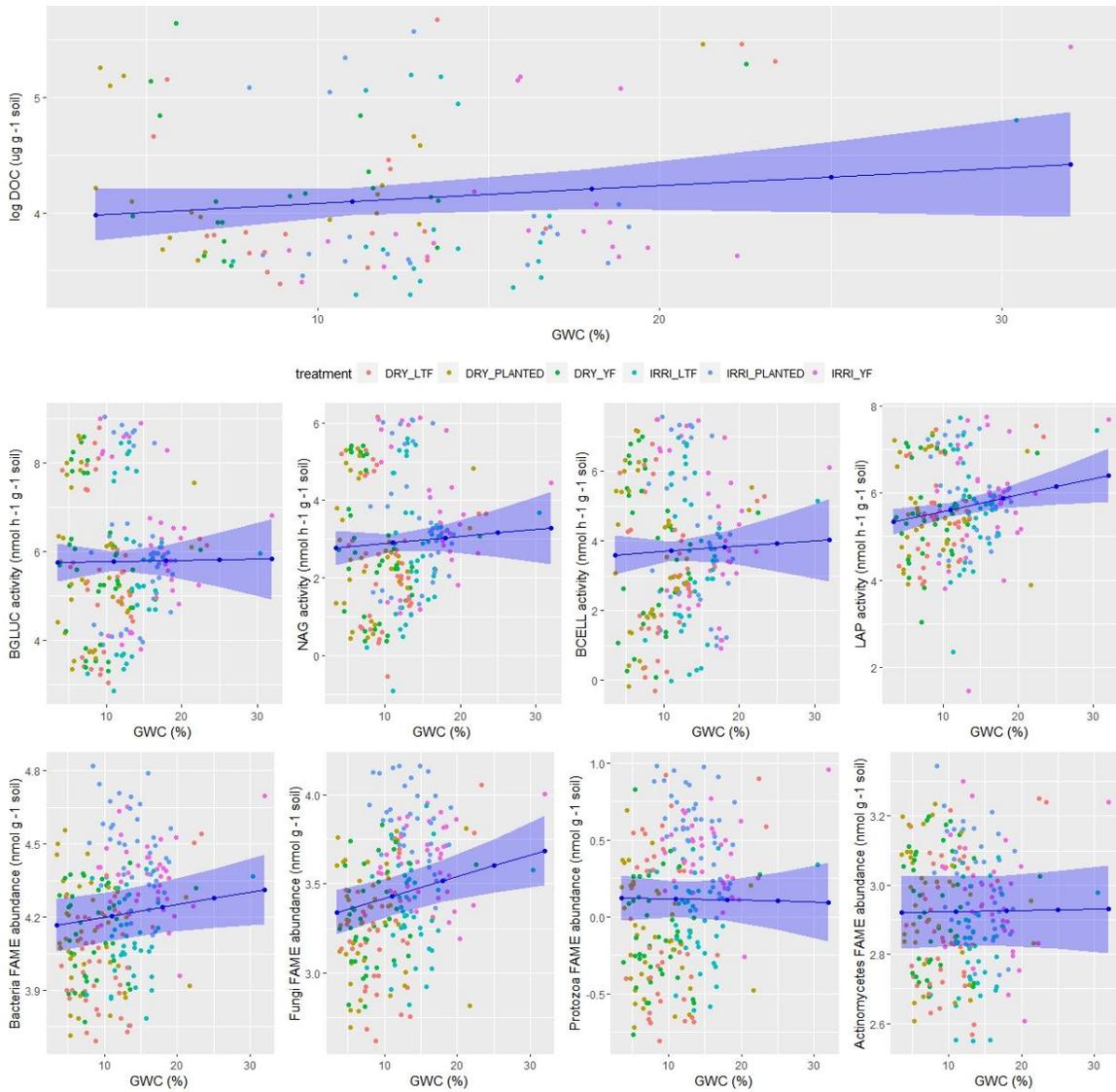
## DOC



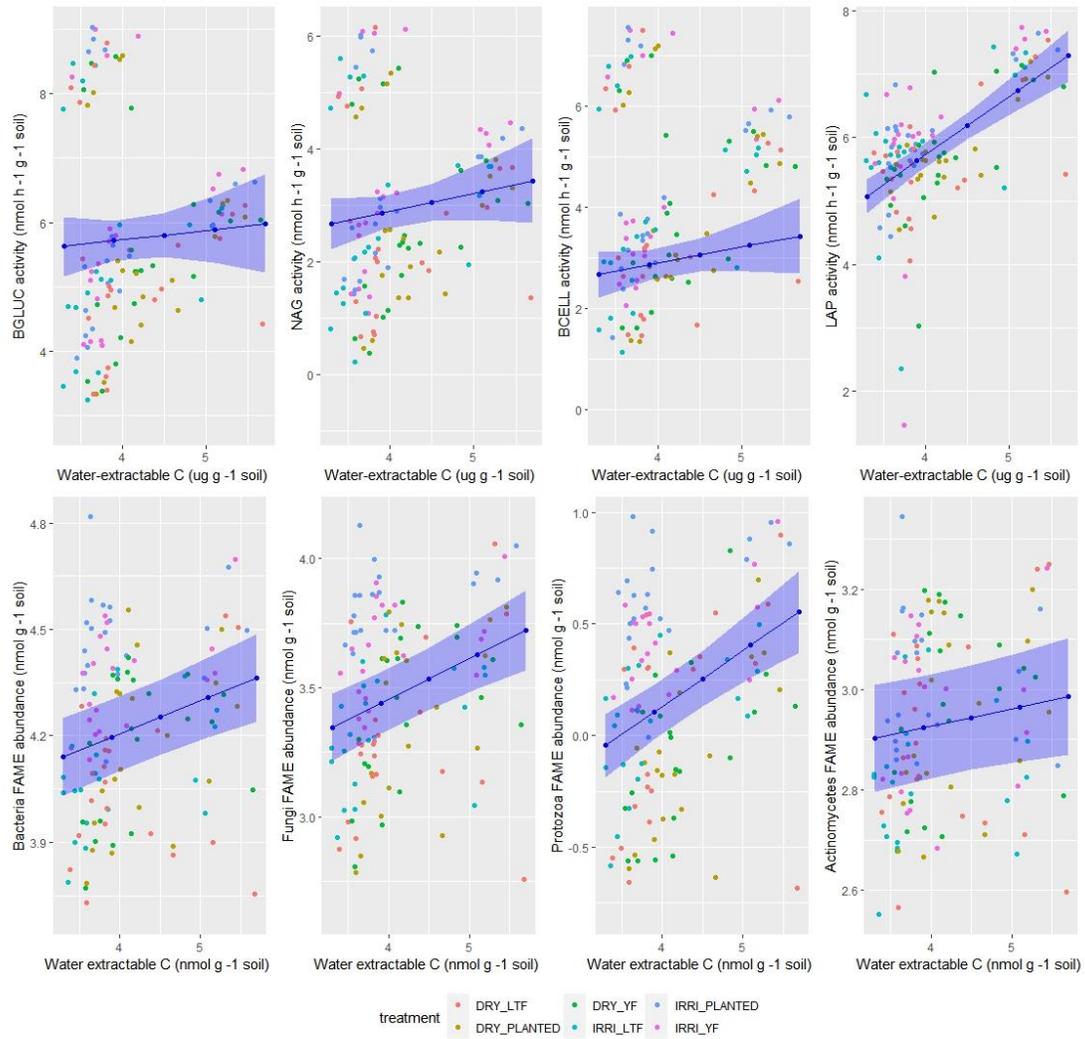
**Figure S2.** Temporal changes in dissolved organic carbon concentrations (DOC; ug C g<sup>-1</sup> soil) from top-soil (0-10 cm) of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and planted— in Northern Colorado.



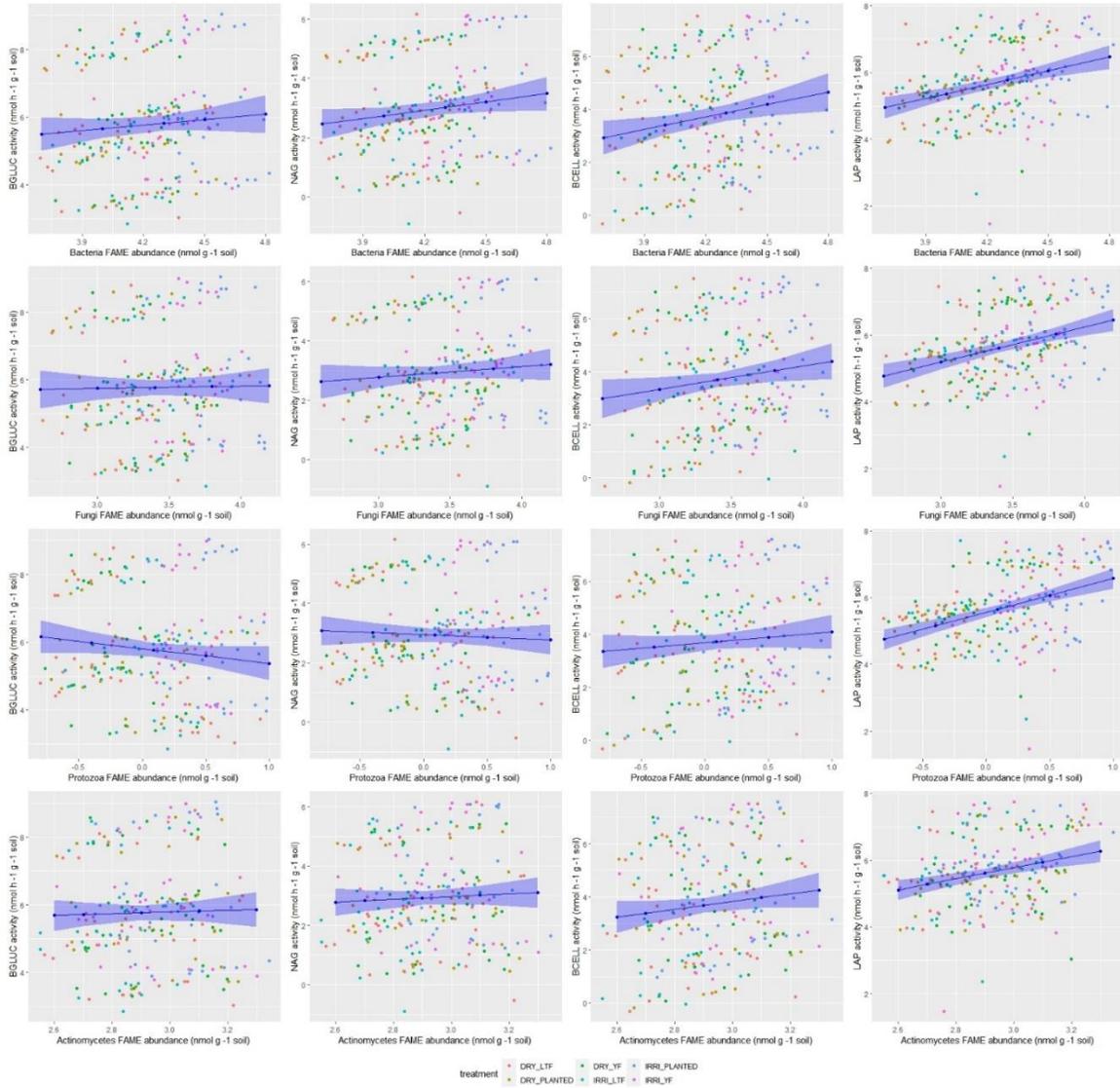
**Figure S3.** Temporal changes in four extra-cellular enzyme activities (EEAs; nmol h<sup>-1</sup> g<sup>-1</sup> soil) from top-soil (0-10 cm) of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and planted— in Northern Colorado.



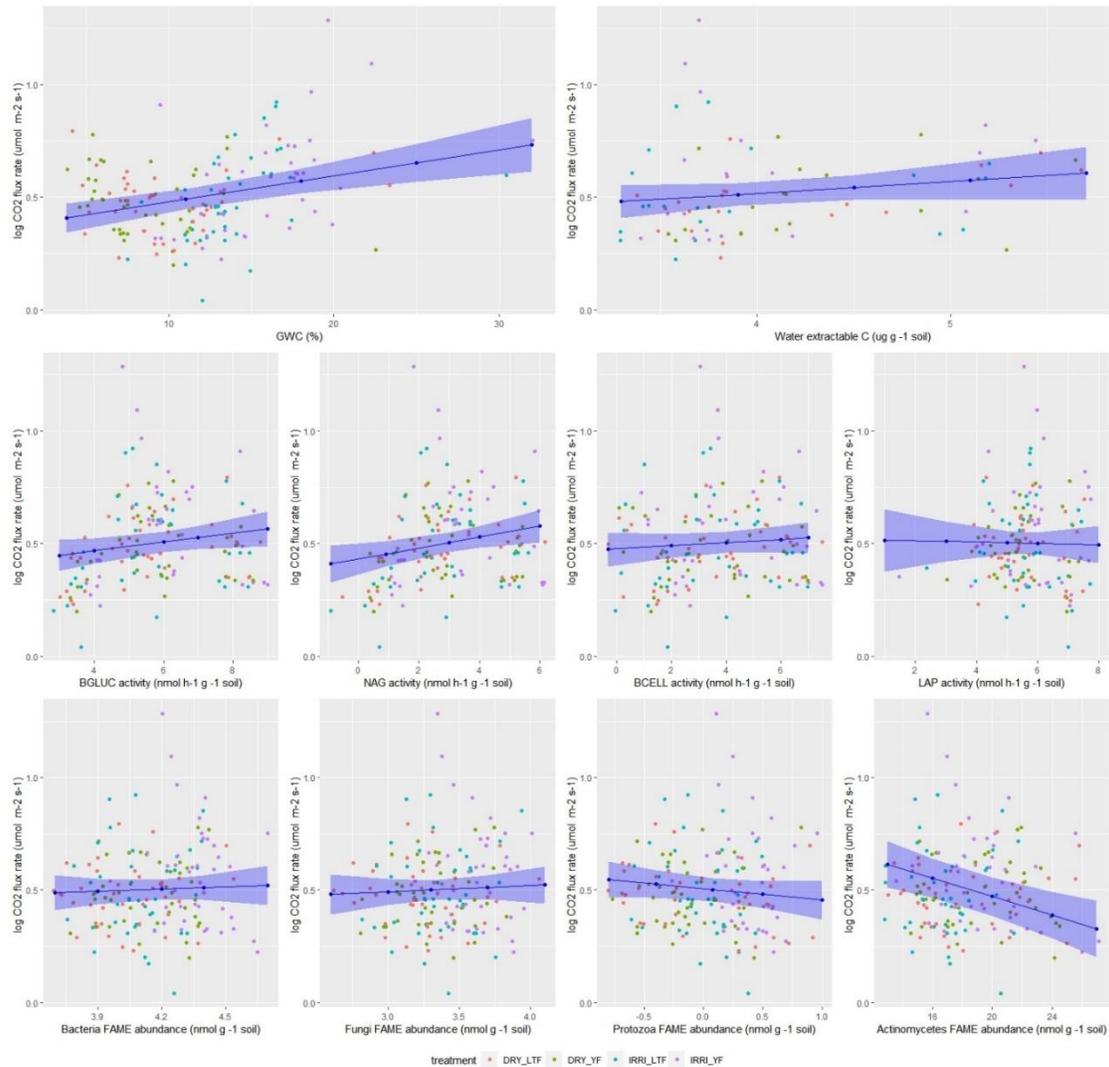
**Figure S4.** Relationship between soil moisture, represented by topsoil (0-10 cm) gravimetric water content of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and regular planted— in Northern Colorado. All data has been log transformed (except GWC).



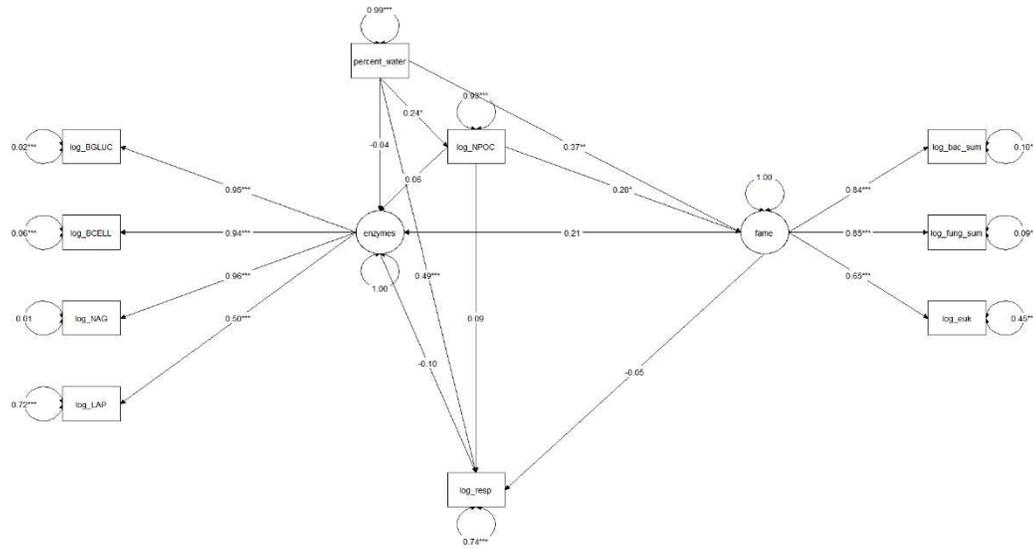
**Figure S5.** Relationship between available soil substrates, represented by topsoil (0-10 cm) water-extractable soil C content of irrigated (IRRI) and dryland (DRY) maize, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and regular planted— in Northern Colorado. All data has been log transformed.



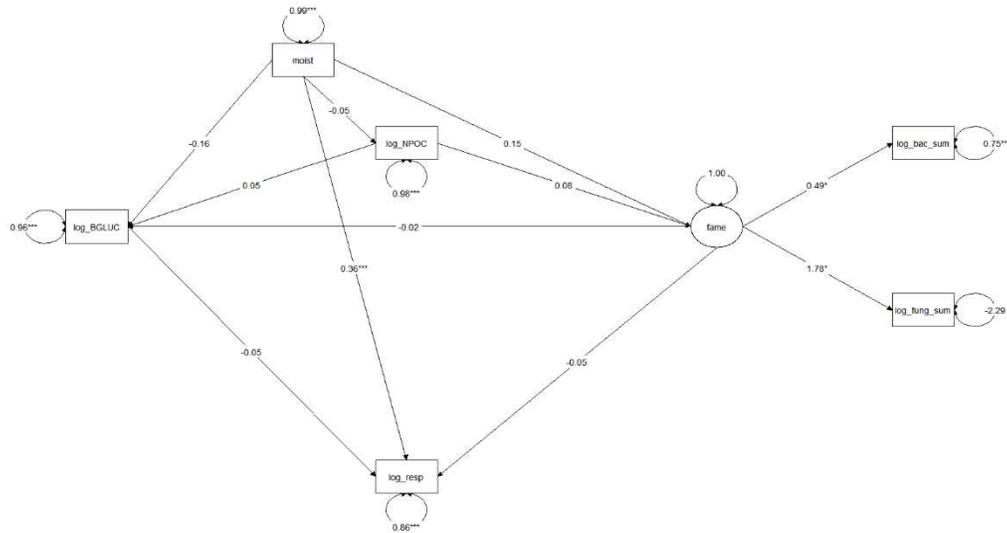
**Figure S6.** Relationship between bacteria, fungi, protozoa and actinomycetes (rows) fatty acid methyl-ester biomarkers and soil extracellular enzyme activities (columns), represented by  $\beta$ -1, 4-glucosidase (BGLUC),  $\beta$ -1, 4-n-acetyl-glucosaminidase (NAG),  $\beta$ -D-1,4-cellobiosidase (BCELL) and Leucine amino peptidase (LAP) activities in top soil (0-10 cm) collected from a maize agroecosystem undergoing irrigated (IRRI) and dryland (DRY) watering regimes, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and regular planted—in Northern Colorado. All data has been log transformed.



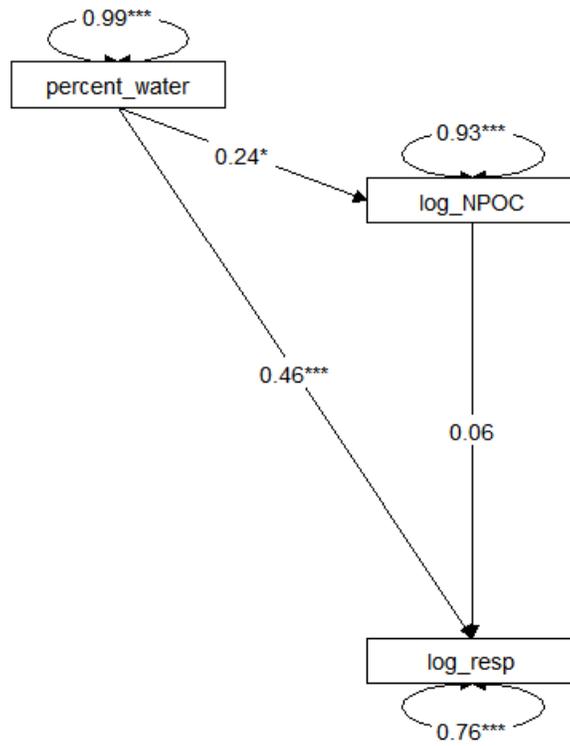
**Figure S7.** Relationship between soil heterotrophic respiration (soil Rh) and multiple soil characteristics (in order, top to bottom, left to right): gravimetric water content, water-extractable C (DOC),  $\beta$ -1, 4-glucosidase (BGLUC),  $\beta$ -1, 4-n-acetyl-glucosaminidase (NAG),  $\beta$ -D-1,4-cellobiosidase (BCELL) and Leucine amino peptidase (LAP) extracellular activities and bacteria, fungi, protozoa and actinomycetes (rows) fatty acid methyl-ester biomarkers and in top soil (0-10 cm) collected from a maize agroecosystem undergoing irrigated (IRRI) and dryland (DRY) watering regimes, under three plant and residue inputs— long-term fallow (LTF), yearly fallow (YF), and regular planted— in Northern Colorado. All data has been log transformed (except GWC).



**Figure S8.** First iteration of the structural equation model (SEM) with variables (boxes) and potential causal relationships (arrows) for soil heterotrophic CO<sub>2</sub> release rate (R<sub>h</sub>) across a continuous maize system under two irrigation treatments. FAME abundance and enzymes (ovals) are latent variables created from the combination of bacterial, fungal and protozoa FAME abundances and BGLUC, BCELL, NAG, LAP EEAs, respectively. Arrows indicate the hypothesized direction of causation. The numbers are standardized path coefficients, which can reflect the importance of the variables within the model [Colman and Schimel, 2013]. The model had  $\chi^2 = 137.41$ , p-value = 0.00, CFI= 0.87, RMSEA = 0.22, and AIC = 1550.15.



**Figure S9.** Iteration of the structural equation model (SEM) using field moisture measurements (LICOR soil probe) with variables (boxes) and potential causal relationships (arrows) for soil heterotrophic CO<sub>2</sub> release rate (R<sub>h</sub>) across a continuous maize system under two irrigation treatments. FAME abundance and enzymes (ovals) are latent variables created from the combination of bacterial, fungal and protozoa FAME abundances and BGLUC, BCELL, NAG, LAP EEs, respectively. Arrows indicate the hypothesized direction of causation. The numbers are standardized path coefficients, which can reflect the importance of the variables within the model [Colman and Schimel, 2013]. The model had  $\chi^2 = 14.35$ , p-value = 0.01, CFI= 0.94, RMSEA = 0.18, and AIC = 1211.61.



**Figure S10.** Iteration of the structural equation model (SEM) using factors with highest correlation coefficients related to  $R_h$  with variables (boxes) and potential causal relationships (arrows). Arrows indicate the hypothesized direction of causation. The numbers are standardized path coefficients, which can reflect the importance of the variables within the model [Colman and Schimel, 2013]. The model had  $\chi^2 = 0.00$ , p-value = 0.00, CFI= 1.00, RMSEA = 0.00, and AIC = 664.15.